

Integrating *in situ* Geochronology and Metamorphic Petrology:
An Example from the Gruf Complex, European Central Alps

By
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**Integrating *in situ* Geochronology and Metamorphic Petrology:
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ABSTRACT

Understanding the thermal evolution (i.e. the timing, rate, duration, and magnitude of thermal events) within mountain belts has important implications for the geodynamic evolution of both ancient and modern orogenies. Ultra-high temperature (UHT) metamorphism requires geodynamic or tectonic processes that bring excess heat to the lower crust. Therefore, dating UHT metamorphism can shed light on the geodynamic evolution of the geological settings in which UHT rocks are exposed. In recent years, many researchers have used accessory mineral U-Pb geochronology to date (U)HT metamorphic events. However, it is not always clear to what part of the pressure-temperature (P-T) path the ages relate.

Using an *in situ* approach, this study combines accessory mineral U-Pb geochronology with single mineral thermometry, thermobarometric modeling, and trace element geochemistry to elucidate the P-T-time (P-T-t) evolution of sapphirine-bearing granulites from the Gruf Complex in the Central Alps. Two main questions are addressed: (1) When did the Gruf Complex experience UHT metamorphism? (2) What parts of the P-T evolution of high-grade metamorphic rocks can be dated using U-Pb geochronology of different accessory phases?

Equilibrium phase diagrams calculated from whole rock and microdomain compositions and Zr-in-rutile thermometry indicate that the Gruf granulites underwent UHT metamorphism at 900–1000°C and 7.0–9.5 kbar after decompressing from ca. 800°C and 9–12 kbar. This decompression-heating event resulted in the breakdown of garnet to form orthopyroxene, sapphirine, and cordierite. A lack of inherited monazite and presence of young (34–30 Ma) monazite within UHT textures is interpreted to indicate that UHT metamorphism was the last main metamorphic event the Gruf granulites experienced, thus precluding a Permian UHT event followed by a lower temperature (700–750°C) Alpine event. Textural observations and Ti-in-

zircon thermometry reveal that minor zircon growth occurred in equilibrium with garnet at 34.8 ± 1.1 Ma, and zircon was not growing, but resorbing during UHT metamorphism. Therefore, the youngest zircon rims can only be used to date post-UHT melt crystallization and cooling at 32.7 ± 0.7 Ma. The U-Pb zircon ages of variable deformed felsic dikes indicate that the lower crustal UHT rocks were juxtaposed against the midcrustal migmatites between 30 and 27 Ma and contractional deformation ceased by 25.6 ± 0.3 Ma in the Gruf Complex. Finally, U-Pb rutile ages indicate that the amalgamated Gruf Complex cooled from 700–420°C over an 11 m.y. period from 30 to 19 Ma.

These results indicate that different accessory minerals can be used to date different stages of the evolution of UHT rocks. However, depending on the reactions in the rock volume, dateable accessory minerals may be crystallizing, resorbing, or not reacting at a given P-T condition. Therefore, combining accessory mineral ages with textural, geochemical, and petrological information is necessary to elucidate the P-T-t evolution of a particular rock package.

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CHAPTER 1

Introduction

Background and Significance

Determining the ages of metamorphic and igneous events preserved in the exposed roots of orogens has important implications for the thermal and tectonic evolutions of mountain belts. In the past 15 to 20 years, advances have been made in U-Pb geochronology and trace element thermometry of accessory minerals (e.g. Spear and Pyle, 2002; Hanchar and Hoskin, 2003; Harley and Kelly, 2007; Meinhold, 2010). However, tying the ages of accessory minerals to a specific pressure-temperature (P-T) stage can be ambiguous, especially in high-grade metamorphic rocks from terranes that have experienced multiple high-temperature events (e.g. Degeling et al., 2001; Rubatto, 2002; Kelly et al., 2006; Kelsey et al., 2008).

Ultra-high temperature (UHT) metamorphism is defined as metamorphism occurring in excess of 900°C at 7–13 kbar pressure (Harley, 1998; Brown, 2006; Harley, 2008). Attaining UHT conditions requires tectonic or geodynamic processes that provide excess heat to the lower crust (Collins, 2002; Kelsey, 2008; Clark et al., 2011). Therefore, the timing of UHT metamorphism has important implications for the timing of important processes within the regions UHT rocks are found.

The sapphirine-bearing granulites of the Gruf Complex of the Central Alps are the primary subjects of this study. Although these rocks are quite scarce, they have well-preserved reaction textures that are useful in constraining P-T conditions at different stages in their history. These rocks are also unique in that they are the only known UHT rocks within the Central Alps. Therefore, constraining the pressure-temperature-time (P-T-t) history of these rocks has important implications for the geodynamic evolution of the Central Alps.

By combining U-Pb ages of rutile, zircon, and monazite with pressure-temperature estimates determined by accessory phase thermometry and equilibrium phase diagram modeling, this study addresses two main questions: (1) What was the P-T-t path taken by the sapphirine-bearing granulites and related charnockites from the Gruf Complex? In particular, when did UHT metamorphism occur? (2) What geologic processes can be dated by accessory mineral U-Pb geochronology in high-grade metamorphic rocks? Do the ages constrain the timing of prograde, peak, or retrograde metamorphism or other events?

Summary of Chapters

This dissertation comprises four studies on the UHT rocks from the Gruf Complex that will be published in peer-review journals. Chapters 2, 3, and 4 each focus on a different accessory mineral. In Chapter 5, the major minerals are the main concern.

Chapter 2, which was submitted to the journal *Chemical Geology*, focuses on U-Pb dating and trace element geochemistry of rutile. The main conclusion of this chapter is that rutile U-Pb ages and Zr-in-rutile temperatures are decoupled. The ages indicate that the Gruf complex underwent cooling from 700°C to 420°C between 30 and 19 Ma. The temperatures, which are strongly correlated with textural setting, record the conditions during prograde (500–890°C), peak (840–1000°C), and retrograde metamorphism (650–850°C).

Zircon U-Pb dating and trace element geochemistry is the focus of Chapter 3. Based on zircon textures and Ti-in-zircon thermometry, we conclude that zircon was not growing but resorbing during UHT metamorphism. Therefore, the zircon dates that are preserved constrain the timing of three main events: (1) Permian protolith crystallization, (2) a Jurassic to Cretaceous magmatic or thermal event, and (3) cooling and melt crystallization after UHT metamorphism at ca. 32.7 ± 0.7 Ma. The U-Pb zircon ages of variably deformed felsic dikes constrain the timing of

amalgamation of the lower crustal and mid crustal units of the Gruf Complex to between 30 and 27 Ma.

Monazite U-Pb ages are the focus of Chapter 4. All monazite grains and domains have U-Pb dates between 34 and 31 Ma, and inherited monazite was not observed. Numerous studies have shown that monazite can record multiple high-temperature metamorphic events unless near complete recrystallization occurs, and diffusional resetting of U-Pb monazite ages is unlikely at even the highest grades of metamorphism. Therefore, we conclude that UHT metamorphism must have completely recrystallized the monazite (and most of the other minerals) in the granulites, and the UHT event occurred during Cenozoic Alpine orogenesis. This conclusion is supported by the UHT P-T estimates of monazite-bearing textures that is presented in Chapter 5.

In Chapter 5, we present the results from thermobarometric modeling of microdomains and whole rock compositions to constrain the P-T evolution of the Gruf granulites and migmatitic paragneisses. The Gruf granulites and charnockites underwent decompression from ca. 800°C and 10–13 kbar to UHT conditions of 900–1000°C and 7.0–9.5 kbar followed by exhumation into the midcrustal migmatites, which equilibrated at 675–750°C and 5–7 kbar. This P-T path and the timing constraints presented in the other chapters indicates that UHT metamorphism occurred after exhumation of (U)HP units in the Central Alps and was most likely driven by slab breakoff.

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CHAPTER 2

Decoupling of Zr-thermometry and U-Pb thermochronology in rutile from the UHT granulites and charnockites of the Gruf Complex (Central Alps)

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Abstract

In situ U-Pb dating and trace element analysis of rutile is used to constrain the thermal history of UHT granulites and charnockites in the Gruf Complex, European Central Alps. Zirconium-in-rutile temperatures are coupled with textural settings providing temperature estimates for prograde (495–895°C), peak (840–1000°C), and retrograde (610–890°C) metamorphism. Even at the highest temperatures, Zr concentrations are not affected by diffusion unless a Zr sink (e.g. zircon) is present or nucleates at or near the rutile grain boundary. Zr and Hf concentrations are strongly correlated, but other high field strength elements (HFSE) are not correlated, indicating that Zr and Hf are more robust to modification during retrograde metamorphism and cooling. Some trace elements (e.g. W) may provide constraints on fluid compositions during retrograde metamorphism. U-Pb cooling dates are decoupled from Zr-in-rutile temperatures and textural settings, indicating that Pb diffusion is controlled by the presence or absence of localized fast diffusion pathways. Dispersion of concordant ²⁰⁶Pb/²³⁸U dates between

30 Ma and 19 Ma is interpreted to record ca. 11 m.y. of crustal processes in the range of 420–700°C, constraining the average cooling rate to ca. 25°C/m.y. for this time period. Assuming that UHT metamorphism in the Gruf Complex occurred at ca. 34 Ma, these rocks initially cooled at ca. 63°C/m.y. between 34 and 30 Ma. These results are consistent with a tectonic model in which asthenospheric upwelling after slab breakoff provided heat for magmatism, migmatization, and UHT metamorphism in the Central Alps.

Introduction

Ultra-high temperature (UHT) rocks are defined as metamorphic rocks that experienced peak temperatures in excess of 900°C within the sillimanite stability field (Harley, 1998; Brown, 2007; Harley, 2008). These extreme crustal conditions require an elevated geothermal gradient $\gg 20^\circ\text{C}/\text{km}$ (e.g. Kelsey, 2008). Therefore, thermal studies of UHT terranes are important in understanding lower crustal rheology, mantle to crust heat flow, and geodynamic processes in the tectonic settings where they are found.

Determining peak pressure and temperature conditions in UHT granulites and polymetamorphic terranes presents a major challenge. Peak mineral assemblages and phase compositions commonly become overprinted during retrogression or subsequent metamorphic events (e.g. Carswell and O'Brien, 1993; Fitzsimons and Harley, 1994; Harley, 1998; Pattison et al., 2003). Due to their resistance to recrystallization and diffusional loss of certain trace elements, accessory phases such as rutile can record metamorphic conditions that major minerals do not preserve (e.g. Zack et al., 2004; Luvizotto and Zack, 2009; Meinhold, 2010; Kooijman et al., 2012). Zirconium-in-rutile thermometry is particularly robust to resetting even in slowly cooled terranes (e.g. Kooijman et al., 2012; Pape et al., 2016), and thus rutile has the ability to record the temperatures of different metamorphic reactions and recrystallization events. Other trace elements

(e.g., Nb, Ta, U, W) can record the fluid composition and degree of elemental fractionation at different temperatures (e.g. Meyer et al., 2011). The high U and low common-Pb concentrations in rutile make it a useful chronometer (e.g. Mezger et al., 1989; Mezger et al., 1991), and it is well established that rutile records cooling in upper amphibolite to granulite facies rocks (e.g. Mezger et al., 1989; Möller et al., 2000; Vry and Baker, 2006). Other studies have used rutile U-Pb thermochronology to constrain the timing of cooling and cooling rates in UHT rocks (Zack et al., 2004; Jiao et al., 2011; Meyer et al., 2011; Kooijman et al., 2012). These characteristics make rutile a useful mineral for determining thermal histories of complex metamorphic terranes from prograde metamorphism through cooling at upper crustal conditions.

This study presents rutile trace element and U-Pb isotopic data collected *in situ* (i.e. directly from thin sections) by laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS). The goal is to constrain the temperature conditions and cooling ages recorded by different textural domains within the UHT granulites and charnockites of the Gruf Complex in the Central Alps. This provides a thermal history used to interpret the specific geodynamic setting and evolution of UHT metamorphism during continental collision.

Geological Setting

The Central Alps consist of nappes derived from the European paleomargin and the Penninic Units (the Valais and Piemonte-Liguria oceans plus the Briançonnais microcontinent) that were thrust together during the late stages of the Cenozoic Alpine orogeny (Stampfli et al., 1998; Figure 2.1; Bousquet et al., 2012b). Prior to amalgamation, some of these units experienced various degrees of (ultra) high-pressure metamorphism related to south-directed subduction of Europe beneath the Adriatic microplate, which is now represented by the Austroalpine nappes and the Southern Alps (Becker, 1993; Gebauer, 1996; Brouwer et al., 2005; Hermann et al., 2006). A

large part of the Central Alps, known as the Lepontine Dome, is characterized by a Barrovian metamorphic field gradient in which metamorphic grade increases from greenschist facies in the north to upper amphibolite facies in the south toward the Insubric fault zone (Frey and Ferreiro Mählmann, 1999; Rubatto et al., 2009; Berger et al., 2011). Isotherms and mineral isograds crosscut nappe boundaries, indicating that temperature-dominated metamorphism occurred after nappe amalgamation (Niggli and Niggli, 1965; Trommsdorff, 1966; Todd and Engi, 1997; Frey and Ferreiro Mählmann, 1999). Burri et al. (2005) divided the Lepontine dome into three zones of migmatites based on the style of melting. The most northern zone consists of rocks intruded by pegmatitic and aplitic dikes and did not undergo *in situ* melting. Migmatites in the intermediate zone underwent *in situ*, fluid induced, anatexis. The southernmost migmatite zone, the Southern Steep Belt (SSB) was formed by muscovite dehydration melting. Rubatto et al. (2009) interpreted the SSB to have undergone protracted or repeated migmatization between 32 and 22 Ma.

The Gruf Complex is bounded by Penninic Units (the Chiavenna ophiolite and the Tambo Nappe) to the north, structurally overlain by the 32–30 Ma (von Blanckenburg, 1992; Samperton et al., 2015) Bergell intrusion in the east (Berger et al., 1996; Davidson et al., 1996), intruded by the 24 ± 1 Ma (Liati et al., 2000) Novate S-type leucogranite (Chappell and White, 1974) in the south, and bordered by the European plate-derived (Schmid et al., 2004) Adula Nappe to the west (Figure 2.1). The northern and western contacts are covered by valley fill. The Adula Nappe consists of rock packages with distinct P-T histories and has been interpreted as a lithospheric *mélange* (Trommsdorff, 1990) that formed by tectonic accretion in a subduction channel (e.g. Engi et al., 2001). Some researchers (e.g. Frey and Ferreiro Mählmann, 1999; Liati and Gebauer, 2003) have interpreted the Gruf Complex as part of the Adula Nappe, which contains rocks that experienced ultrahigh pressure metamorphism, but lack evidence for UHT metamorphism. In

contrast, Galli et al. (2013) proposed that the Gruf Complex is not equivalent to the Adula Nappe, but is a separate migmatitic unit exhumed to the middle crust together with the Bergell intrusion. This interpretation is based in part on the lack of evidence for ultrahigh pressure metamorphism in the Gruf Complex. Some researchers (e.g., Ciancaleoni and Marquer, 2006; Galli., 2010) have concluded that the Gruf Complex and the Adula Nappe are separated by the normal-sense Forcola fault (Figure 2.1) and represent distinct pieces of European lower crust with different metamorphic histories.

Migmatitic ortho- and paragneisses are the volumetrically dominant rock types in the Gruf Complex (Galli, 2010; Galli et al., 2011; Figure 2.2). The migmatitic layering defines the main NE-SW striking foliation (Figure 2.2), and the changing dip direction from predominately SE in the south to NW in the north gives the Gruf Complex a dome-like shape. Mineral lineations, defined by aligned biotite or sillimanite, plunge moderately to the E-NE. The migmatites underwent peak metamorphism at 700–750°C and 7 kbar (Oalmann et al., in preparation-a).

Two large charnockite bodies occur within a narrow, east-west striking zone in the center of the Gruf Complex (Figure 2.2) at the highest elevations, and smaller charnockitic boudins are present within the lower grade migmatitic gneisses. Scarce sapphirine-bearing granulites preserve UHT mineral assemblages and occur as enclaves and schollen within charnockite and less commonly within orthogneiss (Galli, 2010; Galli et al., 2011). These are the only known UHT rocks in the Central Alps (Bousquet et al., 2012a). In the charnockite bodies, foliation and lineation orientations vary significantly from the main foliation in the other units. Mylonites along the charnockite–migmatite contacts consist of a zone of <50 cm-wide mylonitic layers that anastomose around the less deformed charnockitic sheets and schollen (Figure 2.3F). Galli *et al.* (2011; 2012) concluded that the charnockites and granulites underwent peak metamorphism at 920–940°C and

8.5–9.5 kbar during Permian rifting at 282–260 Ma. They interpreted orthopyroxene-, sapphirine-, cordierite-, and spinel-bearing symplectites and coronae to have formed at 720–740°C and 7.0–7.5 kbar subsequently during the Alpine orogeny 34–29 Ma (Galli et al., 2011; 2012). In contrast, evidence from a U-Pb zircon rim age of 32.7 ± 0.5 Ma (Liati and Gebauer, 2003) and a texturally controlled monazite U-Th-Pb chemical age of 33 ± 4 Ma (Schmitz et al., 2009) from the sapphirine-bearing granulites indicates that UHT metamorphism occurred during the Alpine orogeny. Thus, the timing of UHT metamorphism in the Gruf Complex is still a matter of debate. Regardless of the timing of UHT metamorphism, the sapphirine-bearing granulites as well as the charnockites were subjected to upper amphibolite facies conditions, equivalent to the peak conditions in the migmatitic units, at some point after the UHT event. This indicates that the juxtaposition of the UHT and lower grade units occurred prior to or coevally with the upper amphibolite facies metamorphism.

Abundant felsic dikes, ranging from hornblende-bearing granitoids to biotite \pm muscovite \pm garnet \pm tourmaline \pm beryl pegmatites, occur throughout the Gruf Complex. The least evolved dikes are commonly ductily deformed whereas the most evolved dikes crosscut migmatitic foliations. Some of these dikes are deformed within the mylonitic shear zones separating the UHT units from the lower grade migmatites (Galli et al., 2013). The youngest zircon populations from the deformed dikes have U-Pb ages between ca. 30 and 27 Ma. (Oalman et al., in preparation-c) constraining the timing of amalgamation of these units to 30–27 Ma. The crosscutting, most evolved pegmatites have an age 25.6 ± 0.3 Ma (Oalman et al., in preparation-c).

Petrographic Descriptions and Major Mineral Compositions

Residual Granulite

Sample G1A-2009 (Figure 2.3A) is a massive, nonfoliated, mafic, residual granulite that was collected from a talus slope east of Bresciadega village (Figure 2.2). The peak metamorphic assemblage consists of garnet + orthopyroxene + sapphirine + biotite + sillimanite porphyroblasts. Leucosomes containing intergrown quartz + plagioclase + K-feldspar occur as discrete pods. Anhedral garnet porphyroblasts (up to 1 cm diameter) have multiphase embayments that consist of orthopyroxene ± sapphirine ± rutile ± cordierite ± minor biotite (Figure 2.5A). Garnet-replacing symplectites, which consist of elongate orthopyroxene and scarce sapphirine grains as well as blocky plagioclase, are also present (Figure 2.5A, B). Prismatic to bladed sapphirine porphyroblasts are 0.4-0.7 mm long. Orthopyroxene porphyroblasts forms subhedral to anhedral porphyroblasts, some of which are intergrown with sapphirine porphyroblasts. Primary porphyroblasts are invariably surrounded by cordierite coronae or multiphase symplectites. Fine-grained sapphirine grains are the dominant mineral within fine-grained, polymineralic aggregates that also contain sillimanite, spinel, cordierite, and/or quartz (Figure 2.5B). Some of these aggregates surround or emanate from the larger sapphirine porphyroblasts and are commonly surrounded by cordierite moats (Figure 2.5B). Biotite primarily forms subrounded aggregates of bladed to tabular grains that also contain scarce sapphirine and rutile (Figure 2.5B). Biotite also occurs as discrete grains (up to 1 mm diameter), within symplectites, and replacing orthopyroxene porphyroblasts. Prismatic sillimanite porphyroblasts (0.4 mm long on average) are commonly surrounded by cordierite moats, leucosomes, or fine-grained, sapphirine-dominated aggregates.

Anhedral rutile grains (0.02-0.80 mm) mostly occur as part of the multiphase embayments in garnet (Figure 2.5A). Outside of garnet, rutile grains are only found within biotite-dominated aggregates where they are commonly in contact with sapphirine (Figure 2.5B). These textures

suggest that rutile was stable during prograde and peak metamorphism but was consumed by retrograde reactions.

Stromatic Granulites

Samples Bras10-5 and Bras10-6 are stromatic, migmatitic granulites that consist of charnockitic leucosomes and garnet- + orthopyroxene- + sapphirine- ± cordierite- ± sillimanite bearing melanosomes. Alternating leucosomes and melanosomes give this rock type its layered (gneissic) appearance. In one thin section (Bras10-5b-1), coarse-grained and fine-grained textural domains with similar mineralogies can be distinguished (Figure 2.4A). Quartzofeldspathic layers, which are interpreted to represent crystallized melt, consist of subhedral quartz + mesoperthitic alkali feldspar + plagioclase + minor orthopyroxene (Figure 2.5C). Subhedral garnet porphyroblasts (0.8-1.2 mm) contain small proportions of fine-grained rutile, quartz, spinel + plagioclase, and opaque inclusions and are commonly rimmed by fine-grained orthopyroxene + sapphirine symplectites, which contain minor biotite and rutile (Figure 2.5D). Subhedral to anhedral orthopyroxene grains (1-3 mm on average) commonly contain embayments filled with cordierite and/or biotite (Figure 2.5E). In some cases, coarse-grained orthopyroxene + sapphirine intergrowths and minor plagioclase fill embayments in garnet (Figure 2.5F). In addition to the symplectites and intergrowths with orthopyroxene, sapphirine occurs as distinct prismatic to bladed porphyroblasts (Figure 2.5D). Sample Bras10-6 has a higher proportion of cordierite and lower proportion of sapphirine compared to sample Bras10-5. Anhedral to tabular biotite grains (0.1-1.8 mm) are primarily intergrown with and in contact with orthopyroxene, suggesting that much of the biotite is secondary. Bladed biotite is the dominant mineral in some mafic layers (Figure 2.5G).

Subhedral to euhedral rutile grains (0.01-0.22 mm) occur primarily in the more mafic layers of the granulites and as single-phase inclusions in garnet. Rutile is also present in sapphirine + orthopyroxene symplectites and multi-phase, coarse grained embayments around garnet (Figure 2.5A, D, F, 7E). In the quartzofeldspathic layers, rutile grains have similar shapes and sizes as those in the mafic layers but are less abundant. Rarely, rutile also occurs as fine-grained needles within orthopyroxene porphyroblasts (Figure 5E). In some cases, there are small zircon grains along the edges of rutile (Figure 2.5C, G).

Leucogranulite

Sample GR11-35 is felsic, sapphirine-bearing granulite schlieren that was collected from one of the large charnockite bodies (Figure 2.2). It consists of ≤ 4 mm-long quartz ribbons, subhedral plagioclase and mesoperthite, subhedral to anhedral, inclusion-poor garnet porphyroblasts (0.1-0.7 mm), bladed to tabular biotite, and orthopyroxene. Minor, prismatic sillimanite and aggregates of fine-grained sapphirine are also present. Aggregates of green spinel \pm opaque minerals \pm plagioclase \pm biotite \pm orthopyroxene + sillimanite \pm sapphirine commonly surround or fill embayments in garnet (Figure 2.5H). Compositional layering and shape-preferred orientations of quartz ribbons, biotite grains, and sapphirine aggregates define a weakly developed foliation.

Subhedral to anhedral rutile grains mostly occur within the matrix of this sample and range from 0.05 to 0.5 mm in size. Some rutile grains are included in garnet, and others are in contact with sapphirine aggregates (Figure 2.7F).

Charnockite

Sample GR11-43 is a garnet-biotite-alkali feldspar granitic gneiss (retrogressed charnockite) collected from a small charnockite body near Val Codera (Figure 2.2). A hypidiomorphic granular matrix consists mainly of quartz + mesoperthitic, alkali feldspar. Scarce, blocky plagioclase grains (≤ 0.1 mm) occur mainly as inclusions in and adjacent to garnet and, more rarely, within the matrix. Anhedral to subhedral garnet porphyroblasts (0.4-1.75 mm) contain embayments filled with quartz, plagioclase, alkali feldspar, biotite, and spinel (Figure 2.5I) and form atoll structures in some cases. Tabular biotite grains (≤ 1.75 mm) most commonly occur adjacent to garnet porphyroblasts and form finger-like intergrowths with plagioclase, quartz, and opaque minerals in some cases. Bladed biotite grains (≤ 1.5 mm long) are restricted to fractures and microschlieren.

Subhedral to euhedral rutile grains (0.05-0.15 mm) occur mainly within the matrix and as inclusions in quartz and K-feldspar (Figure 2.5I). Rare rutile inclusions in garnet are also present.

Retrogressed Granulites

Two samples (GR11-1A and GR11-1C) of retrogressed granulite were collected from a large boulder in Val Codera near the village of Bresciadega (Figure 2.2). These samples have the assemblage quartz + feldspar + biotite + pinnitized cordierite + orthopyroxene + garnet + minor sillimanite + abundant rutile (Figure 2.5J, K). Cordierite + orthopyroxene symplectites (Figure 2.5J, K) commonly surround garnet. Within garnet embayments, sample GR11-1C contains some relict sapphirine porphyroblasts, which are rimmed by cordierite coronae (Figure 2.5L). Most of the orthopyroxene in these samples is surrounded by fine-grained masses of chlorite + cordierite (Figure 2.5K).

Sample GR11-1A contains two textural domains, based on grain size, that have similar mineralogy (Figure 2.4B), and both domains contain rutile. The majority of rutile grains in sample GR11-1C are included in garnet (Figure 2.5L).

Analytical Methods

Trace element compositions and U-Pb isotopic data were collected by laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) using a Thermo Scientific Element 2 ICP-MS coupled with a Photon Machines 193 nm ArF excimer laser at The University of Kansas, Isotope Geochemistry Laboratories (IGL). Rutile grains were ablated directly in 40 μ m-thick, polished section to preserved textural context. Trace element and U-Pb data were collected in separate analytical sessions. See Table 2.1 for details on the analytical set-up, procedures, and data reduction.

Zirconium-in-rutile temperatures were calculated with the calibration of (Tomkins et al., 2007) using the equation for the β -quartz field and a pressure of 9 kbar. This corresponds to the maximum pressure of UHT metamorphism (Oalman et al., in preparation-a). Therefore, the reported temperatures are maximum estimates. Decreasing the pressure to 7 kbar, which is the minimum pressure estimated for the UHT assemblages, decreases the apparent temperature by ca. 5–10°C. The analytical uncertainty of Zr compositions is 2–7%, corresponding to a maximum temperature uncertainty of 63°C at 900°C and 49°C at 700°C. Therefore, the temperature difference due to a pressure difference of 2 kbar is well within analytical uncertainty.

The legacy mode of ET_Redux (McLean et al., 2016) was used to plot concordia diagrams and calculate ages. Because high Hg backgrounds made it impossible to perform a ^{204}Pb -based common Pb correction, lower intercept ages are the preferred ages, but weighted mean $^{206}\text{Pb}/^{238}\text{U}$ ages of concordant analyses are also reported in some cases.

U-Pb Ages

Regardless of rock type, textural setting, and grain size, many of the U-Pb rutile ages are similar, but there is a high degree of dispersion in $^{206}\text{Pb}/^{238}\text{U}$ dates within the populations of each sample. All analyzed samples contain rutile grains with observable amounts of common Pb as seen on the concordia diagrams (Figure 2.6). Combining all analyses from all samples yields a lower intercept age of 22.79 ± 0.05 Ma (MSWD = 32, n = 358) and a common-Pb upper intercept. The $^{206}\text{Pb}/^{238}\text{U}$ dates of concordant analyses range from 18.8 ± 1.5 to 32.2 ± 1.5 Ma (Figure 2.6 and Table 2.2). Lower intercept ages from individual samples range from 20.8 ± 0.2 to 25.8 ± 0.3 Ma, and all have common-Pb upper intercepts (Figure 2.6). Spot analyses along traverses in the largest rutile grains do not show any detectable core to rim age variation (Figure 2.7).

Residual granulite sample G1A-2009 (Figure 2.6B) contains rutile inclusions in garnet (Figure 2.7A) with scarce matrix rutile (Figure 2.7B). All analyses combined yield a lower intercept age of 23.7 ± 0.6 Ma (MSWD = 3.5, n = 28). A single rutile inclusion in garnet yields the only two concordant analyses with $^{206}\text{Pb}/^{238}\text{U}$ dates of 20.2 ± 2 and 22.1 ± 2 Ma. Excluding these two concordant analyses results in a lower intercept age of 24.6 ± 0.7 Ma (MSWD = 2.4, n = 26).

Sample GR11-1A (Figure 2.6C) is a retrogressed residual granulite that contains abundant rutile as inclusions in garnet and as matrix grains. The analyzed grains yield a lower intercept age of 25.9 ± 0.3 Ma (MSWD = 6.3, n = 27). The $^{206}\text{Pb}/^{238}\text{U}$ dates of concordant grains included in the age calculation range from 23.0 ± 1.0 to 27.8 ± 1.3 Ma. Two matrix grains, which are not included in the age calculation but are concordant, yield $^{206}\text{Pb}/^{238}\text{U}$ dates of 18.8 ± 1.5 and 19.8 ± 0.9 Ma.

Another retrogressed residual granulite (sample GR11-1C, Figure 2.6D) contains rutile inclusions in garnet that yield a lower intercept age of 24.0 ± 0.5 Ma (MSWD = 1.8, n = 18). Concordant analyses have $^{206}\text{Pb}/^{238}\text{U}$ dates between 22.4 ± 2.6 and 25.7 ± 2.6 Ma.

Rutile grains are restricted to the matrix of charnockite sample GR11-43 (Figure 2.6E). The lower intercept age calculated for this sample is 23.5 ± 0.5 Ma (MSWD = 0.95, n = 12). Concordant analyses have $^{206}\text{Pb}/^{238}\text{U}$ dates ranging from 22.2 ± 2.2 to 25.9 ± 3.1 Ma.

Leucogranulite sample GR11-35 (Figure 2.6F) contains only matrix rutile grains. Two different common-Pb discordia can be interpreted from the analyses. The younger group, which includes more analyses, yields a lower intercept age of 25.5 ± 0.4 Ma (MSWD = 2.0, n = 11), and the older group yields a lower intercept age of 27.6 ± 0.6 Ma (MSWD = 0.42, n = 4). One discordant analysis ($^{206}\text{Pb}/^{238}\text{U}$ date = 21.9 ± 1.4 Ma) was not included in either of the calculations.

The following stromatic granulite samples (Figure 2.6G–J) contain rutile included in garnet and within the UHT matrix assemblages except where otherwise noted. These samples are the most texturally complex, and they all show excessive dispersion of concordant dates and complex patterns of discordant results.

Sample Bras10-5b-1 (Figure 2.6G) is a stromatic granulite sample that yields a lower intercept rutile age of 20.5 ± 0.2 Ma (MSWD = 3.4, n = 15). The $^{206}\text{Pb}/^{238}\text{U}$ dates of concordant analyses range from 19.8 ± 0.7 to 23.9 ± 1.2 Ma. Two nearly concordant analyses ($^{206}\text{Pb}/^{238}\text{U}$ dates = 23.5 ± 0.9 and 23.9 ± 1.2 Ma) were excluded from the age calculation.

Stromatic granulite sample Bras10-5b-2 (Figure 2.6H) contains only matrix rutile and yields the largest dispersion of concordant analyses with $^{206}\text{Pb}/^{238}\text{U}$ dates ranging from 21.3 ± 0.6 Ma to 32.2 ± 1.5 Ma. Constructing a discordia through all analyses results in a lower intercept at 24.4 ± 0.2 Ma (MSWD = 42, n = 24) coinciding with a cluster of concordant analyses (weighted

mean $^{206}\text{Pb}/^{238}\text{U}$ age = 24.6 ± 0.3 Ma, MSWD = 1.2, n = 4). The youngest group of concordant analyses from this sample has a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 22.2 ± 0.2 Ma (MSWD = 5.4, n = 7).

The rutile dates obtained from stromatic granulite sample Bras10-6-1 (Figure 2.6I) also exhibit a high degree of dispersion along concordia, with $^{206}\text{Pb}/^{238}\text{U}$ dates ranging from 18.7 ± 0.7 to 28.2 ± 0.7 Ma. A lower intercept age calculated for this sample is 22.2 ± 0.1 Ma (MSWD = 32, n = 131), but the discordant results do not lie on the same common Pb array. The implications of such U-Pb patterns for age interpretation are discussed below.

Stromatic granulite sample Bras10-6-2 (Figure 2.6J) yields a lower intercept age of 25.0 ± 0.2 Ma (MSWD = 13, n = 57). Concordant analyses have $^{206}\text{Pb}/^{238}\text{U}$ dates between 21.2 ± 1.5 and 28.3 ± 1.0 Ma.

Figure 2.7 shows the dates of rutile grains from various textures in different rock types. The dates of rutile grains within a single microdomain vary in some cases, but these variations do not correlate with grain size, and typical intragrain diffusion profiles are not observed (Figure 2.7).

Rutile Trace Element Compositions

Most samples display a wide range of compositional variation in trace elements from grain to grain (Figure 2.8). There is a higher degree of correlation between trace element content and textural setting than for rock type. In general, matrix rutile grains have higher trace element abundances than inclusions in garnet. Matrix grains from peak and lower temperature assemblages in all samples show similar concentration ranges for most trace elements (Figure 2.8). However, grains from peak assemblages contain higher concentrations of Zr, Hf, and U than grains from lower temperature domains (Figure 2.8A–B). Retrogressed granulite sample GR11-1A and the finer-grained parts of stromatic granulite samples Bras105b-1 and Bras10-5b-2 contain matrix

rutile grains with elevated concentrations of W (Figure 2.8C). The two grains with the highest Nb concentrations (ca. 5.0 and 6.8 wt. %) are from a sapphirine-bearing domain in residual granulite sample G1A_2009, and several grains from all textural domains have Nb concentrations between 2.7 and 4.0 wt. % (Figure 2.8D).

Only two pairs of elements show strong correlations, Zr-Hf (Figure 2.8A) and Nb-Ta (Figure 2.8E). Zirconium and Hf are positively correlated, with an average Zr-Hf ratio of 29.2 ($y = 0.034x + 6.26$) and a correlation factor r of 0.97 (Figure 2.8A). Niobium and Ta are also positively correlated, but show a wider range of trends (Figure 2.8E) than Zr and Hf. Figures 2.8F, G, and H show that, on ratio plots, there is a high degree of scatter for the high field strength elements (HFSE). Grains with high Ta concentrations tend to have less scatter in the Nb-Ta ratio (Figure 2.8F), but this may be a result of higher analytical uncertainties for lower concentrations. Figure 2.8G shows that there is a higher degree of scatter in the Zr-Hf ratios of grains included in garnet (blue symbols) and from retrogressed domains (green symbols) than for rutile grains from UHT domains (red symbols). The Nb/Ta vs. Zr/Hf plot (Figure 2.8H) shows that there is high degree of variability, but no correlation with granulite type or textural setting.

Zr-in-rutile Thermometry

Zirconium concentrations of rutile grains and thus calculated temperatures are highly correlated with their textural setting in all investigated samples (Figure 2.7). Intragrain Zr concentration profiles are not observed (Figure 2.7). There is considerable overlap of temperature ranges obtained from similar textural settings in the different rock types. The details of these intrasample variations are discussed in this section (see also Table 2.3).

Massive, Mg- and Al-rich Granulite (Residual Granulite)

Rutile grains occur mostly within multiphase embayments (e.g. orthopyroxene + sapphirine + rutile; Figure 2.7A) and as single phase inclusions in the large garnet porphyroblasts of the residual granulite sample (G1A-2009) and record a wide range of temperatures ranging from 588–880°C (Figure 2.7A). Scarce grains outside of the garnet porphyroblasts occur primarily within large biotite clots and in contact with sapphirine (Figure 2.7B). These grains record UHT conditions (950–1000°C).

Stromatic Granulites

Rutile inclusions in garnet have the lowest trace element abundances and record the lowest Zr-in-rutile temperatures, ranging from 540–640°C and 490–700°C for samples Bras10-5b-1 (Figure 2.7C) and Bras10-6, respectively. Rutile, which records UHT temperatures (900–1000°C), occurs within some multiphase embayments in garnet that contain plagioclase + orthopyroxene + biotite + sapphirine (Figure 2.7D).

Outside of garnet, within the matrix domains of the stromatic granulite samples Bras10-6 (Figure 2.7E) and in the coarser-grained domain of sample Bras10-5b-1 (Figure 2.4A), rutile grains record mainly UHT conditions, ranging from 870–1070°C; however, only a few grains record temperatures higher than 1000°C. Rutile grains within orthopyroxene + sapphirine symplectites surrounding garnet porphyroblasts invariably record UHT conditions (Figure 2.7C). Rutile grains with small (<10 µm) zircon grains at their boundaries (Figure 2.5G) as well as rutile in contact with larger zircon grains (Figure 2.5C) record temperatures between 710 and 750°C. Rutile grains within the finer-grained matrix domain (Domain B) of sample Bras10-5b-1 (Figure 2.4A) record intermediate temperatures, ranging from 710–850°C.

Leucocratic Granulite

In the leucocratic granulite sample (GR11-35), matrix rutile grains (Figure 2.7F) record temperatures of 790–1040°, and rutile inclusions in garnet were too small to be analyzed. Seven of the 45 rutile grains that were analyzed record extremely low temperatures (435–570°C). This is possibly the result of analyzing exsolution lamellae of ilmenite or inclusions of another Zr-poor phase, and thus these temperatures are not considered to be geologically relevant.

Charnockite

Matrix rutile grains within the charnockite sample (GR11-43) are commonly surrounded by quartz + mesoperthite, which represents crystallized melt. These grains record temperatures between 715 and 960°C (Figure 2.5I). Two grains from this sample record very low temperatures (402 and 545°C), probably related to ablation of micro inclusions or exsolution lamellae, and are therefore not considered for discussion.

Retrogressed Granulites

Rutile grains were analyzed from three textural domains in samples GR11-1A and GR11-1C, which came from the same boulder: grains in the 1) coarse- and 2) fine-grained matrix domains of sample GR11-1A (Figure 2.5K) and 3) inclusions within garnet from sample GR11-1C (Figure 2.5J). Zr-in-rutile temperatures are similar in all three domains. The coarse-grained (Domain A) and fine-grained (Domain B) matrix domains of sample GR11-1A contain rutile grains recording temperatures ranging from 510–890°C and 650–890°C, respectively. Four matrix grains have Zr concentrations corresponding to Zr-in-rutile temperatures $\leq 460^\circ\text{C}$, likely caused by ablating micro inclusions or exsolution lamellae. Rutile inclusions in garnet (sample GR11-1C) record temperatures between 610 and 895°C.

Discussion

Significance of Trace Element Compositions

The trace element patterns presented here are similar to those discussed by Meyer et al. (2011), who studied rutile from UHT granulites in NW South Africa. The good positive correlation between Zr and Hf (Figure 2.8A) indicates that these elements were strongly coupled throughout peak metamorphism and lower temperature retrogression and cooling. The correlation is not as good for Nb and Ta (Figure 2.8B), and in agreement with Meyer et al. (2011), we conclude that Zr and Hf are decoupled from the other HFSE. The high degree of correlation between Zr and Hf indicate that these elements exhibited similar behavior during the lower temperature overprint and cooling. The high degree of scatter in the Zr-Hf ratios for rutile grains included in garnet and from the retrogressed domains (Figure 2.8G) indicates that Zr and Hf are more easily fractionated at lower temperatures and during recrystallization. The wide scatter between concentrations and ratios of other elements (e.g. Zr/Hf vs. Hf and Nb/Ta vs. Hf; Figure 2.8) indicates that the primary concentrations of these elements were modified during recrystallization. Scatter in some elements is also possibly due to diffusional loss. The elevated W concentrations in rutile grains from low temperature domains (Figure 2.8C) of the retrogressed and stromatic granulite samples (GR11-1A, Bras105b-1, and -2) is interpreted to be the result of recrystallization in the presence of a W-rich fluid, which was likely sourced from the abundant pegmatitic dikes (Graham and Morris, 1973) within the Gruf Complex.

Zr-in-rutile Thermometry

Temperatures estimated by Zr-in-rutile thermometry are well correlated with textural setting within the samples (inclusions, matrix, recrystallized domains) indicating that Zr diffusion is negligible in most cases, and thus the temperatures represent crystallization temperatures, supporting the results of previous studies (Cherniak et al., 2007; Luvizotto and Zack, 2009;

Kooijman et al., 2012). This allows estimation of temperature at different stages in the metamorphic history of the Gruf granulites. Because the Zr-in-rutile temperatures are better correlated with textural setting than rock type, the results from all samples are combined and organized based on textural setting for this discussion. Mean temperatures of rutile from the different textural domains were estimated using histograms (Figure 2.9), and uncertainties are given as 2 SE.

Rutile included in garnet shows the widest range of temperatures (ca. 495–895°C) and the most non-Gaussian histogram (Figure 2.9), which is skewed toward higher temperatures. Because of the wide temperature range and the highest temperatures being slightly lower than the peak temperatures (see below), rutile inclusions in garnet are interpreted to have formed during prograde metamorphism. These grains likely crystallized continuously along the prograde path and were subsequently included in garnet. However, a detailed account of the garnet growth history is outside the scope of this study.

The highest average Zr-in-rutile temperatures were obtained from grains within matrix assemblages indicative of the metamorphic peak temperature in the granulites and charnockites. Temperature estimates for all except 9 analyses range from ca. 715–1075°C. The majority of these grains (219 of 282) record UHT conditions, and the mean temperature is $908 \pm 83^\circ\text{C}$ (Figure 2.9). This mean temperature from the peak assemblages corresponds well with peak temperature estimates of 900–1000°C at 7.5–9.5 kbar (Oalman et al., in preparation-a) obtained by thermodynamic modeling techniques. The occurrence of UHT rutile grains within multiphase embayments and symplectites around garnet porphyroblasts indicates that garnet breakdown, likely due to decompression, began at temperatures in excess of 900°C. This is consistent with textural evidence for UHT decompression (e.g., sapphirine-bearing symplectites around garnet in

stromatic granulites and spinel- and sapphirine-bearing garnet breakdown textures in the leucogranulites). The absence of UHT rutile inclusions in garnet and presence of UHT rutile within garnet-breakdown textures suggests that the UHT conditions occurred late in the garnet-growth history of the granulites or after garnet growth.

Some rutile grains are part of multiphase embayments in garnet consisting of orthopyroxene + sapphirine + rutile in the residual granulite sample G1A_2009 (Figure 2.7A). We interpret these embayments to have formed as the result of garnet breakdown at UHT conditions based on the presence of orthopyroxene and sapphirine. However, the apparent Zr-in-rutile temperature estimates from these grains are quite low, at ca. 750°C (Figure 2.7A), which is likely the result of low Zr activity in the local reaction environment (i.e. absence of zircon and melt/fluid to transport Zr). In contrast, a sapphirine-bearing embayment in garnet from the stromatic granulite sample Bras10-6-2 (Figure 2.7C) contains rutile grains that record UHT conditions (900–1000°C). This embayment also contains biotite and an apatite grain with a zircon inclusion, indicating that fluid was present and Zr activity was high. Therefore, we interpret the temperatures recorded by rutile in this texture to accurately represent the conditions of garnet breakdown.

Experimental studies (e.g. Cherniak et al., 2007) indicate that primary Zr concentrations should not be preserved in UHT rutile unless very rapid cooling ($>100^{\circ}\text{C}/\text{m.y.}$) occurred. The majority of the rutile grains analyzed in this study preserve Zr concentrations corresponding to UHT conditions, but lack intragrain diffusion profiles or grain size–temperature relationships (Figure 2.7). Some rutile grains within UHT assemblages record temperatures between ca. 700–750°C, which corresponds to the closure temperature of Zr diffusion in a 200 μm rutile grain cooling from a peak temperature of 900°C at $10^{\circ}\text{C}/\text{m.y.}$ (Cherniak et al., 2007). These grains are commonly in contact with zircon grains (Figure 2.5C, G). Based on the position and size of the

small (<10 μ m) zircon grains (Figure 2.5G), we interpret that they formed as a result of Zr-exsolution from the UHT rutile during cooling. Therefore, we propose that grain boundary diffusion of Zr from rutile occurs only if a Zr-rich sink (e.g., zircon) is already present (Figure 2.5C) or can nucleate along the rutile grain boundary (Figure 2.5G). This is in good agreement with Taylor-Jones and Powell (2014) and Pape et al. (2016), who suggested that grain boundary diffusion of Zr from rutile is sluggish under the dry conditions experienced by many UHT rocks.

Within matrix domains that contain major mineral assemblages indicative of recrystallization at lower than peak temperatures (stromatic granulite samples Bras10-5b-1 and -2 and retrogressed granulite sample GR11-1A), rutile grains record temperatures ranging from ca. 610–890°C with a mean of $740 \pm 72^\circ\text{C}$. The Gaussian distribution of this population (Figure 2.9) can possibly be explained by (re)crystallization of rutile in response to a single thermal event and over a limited temperature range. The recrystallization of these domains was likely driven by fluid or melt influx, supported by the abundance of biotite relative to orthopyroxene in these assemblages and the elevated W concentrations of some of these rutile grains (Figure 2.8C). The mean temperature corresponds to the upper range of peak temperature estimates of 700–750°C at ≤ 7.5 kbar (Oalman et al., in preparation-a) for migmatitic paragneisses, which lack textural evidence for ever experiencing UHT metamorphism. Therefore, we interpret it to be likely that this recrystallization occurred within the granulites after they were juxtaposed against the lower grade units (Galli et al., 2013; Oalman et al., in preparation-c). This temperature range is also consistent with temperatures recorded by migmatites elsewhere in the southern part of the Leopontine Dome (e.g., Todd and Engi, 1997; Engi et al., 2004; Burri et al., 2005).

In summary, the crystallization temperatures of rutile from different textural settings record three stages of the metamorphic evolution of the Gruf granulites and charnockites (Figure 2.10):

1) prograde metamorphism, which was accompanied by garnet growth and did not exceed 900°C; 2) peak metamorphism and garnet breakdown at UHT conditions; 3) retrograde (Leptontine) metamorphism and/or cooling in the range of ca. 700–750°C during amalgamation of the UHT and migmatitic units in the Gruf complex. The good agreement between the results presented here and those of other studies, which utilized other techniques to estimate temperature, indicate that Zr-in-rutile thermometry is a robust method for constraining the thermal evolution of high-grade crustal rocks that experienced multiple high temperature events.

Rutile U-Pb Cooling Ages

The most striking result from *in situ* rutile dating in these samples is that there is no evident correlation between the obtained U-Pb ages and the textural setting or specific trace element characteristics. This decoupling indicates that textural setting has no bearing on Pb diffusivity in rutile here. The analyzed grains do not exhibit intragrain age zoning, and there is not a clear correlation between grain size and age. We therefore interpret that Pb diffusivity is most likely controlled by the presence or absence of localized fast diffusion pathways such as grain boundary melts or fluids, exsolution lamellae, and intragrain fractures (e.g. Kooijman et al., 2010).

Some samples are characterized by clusters of concordant analyses and common Pb arrays with different trajectories. These clusters and arrays may represent distinct events that affected some grains but not others. For example, the vast majority of analyses from samples G1A_2009 (residual granulite) and GR11-1A (retrogressed granulite) lie along common Pb arrays with similar lower intercept ages (24.6 ± 0.7 and 25.8 ± 0.3 Ma, respectively). These samples also contain scarce grains with concordant $^{206}\text{Pb}/^{238}\text{U}$ dates of 19–22 Ma. Assuming that most of the grains were closed to Pb diffusion by ca. 25 Ma (i.e. the sample was below the Dodson closure

temperature of Pb diffusion in rutile), the grains that give younger dates may have been subjected to melts or fluids that allowed Pb to diffuse out at lower temperatures.

Core to rim age variations have been observed in rutile grains from other (ultra)high temperature granulites (e.g. Kooijman et al., 2010), which cooled slowly (on the order of $0.1^{\circ}\text{C}/\text{m.y.}$). Because no intragrain age variations along traverses of large rutile grains were observed in this study (Figure 2.7), the Gruf granulites and charnockites must have cooled faster than $0.1^{\circ}\text{C}/\text{m.y.}$ as a first order approximation.

The large range in concordant dates and common Pb arrays is interpreted to represent ca. 11 m.y. of low- to mid-crustal processes (e.g. melt crystallization, fluid flow, cooling, and deformation) between the onset of retrograde metamorphism and final closure to Pb diffusion. Assuming that peak metamorphic conditions (ca. 950°C and 8 kbar) occurred by ca. 34 Ma (oldest U-Pb rutile dates; interpretation of earliest Alpine U-Pb zircon ages (Liatì and Gebauer, 2003; Oalman et al., in preparation-c); U-Pb monazite ages (Oalman et al., in preparation-b); model of (Beltrando et al., 2010)) and the Gruf Complex reached the upper limit of the Pb closure temperature range for rutile (ca. 700°C , e.g. Cherniak, 2000) by ca. 30 Ma, we estimate the cooling rate to have been ca. $63^{\circ}\text{C}/\text{m.y.}$ between 34 and 30 Ma (Figure 2.10). Based on the range of concordant U-Pb rutile dates, we interpret that the Gruf Complex remained in the temperature window of Pb partial retention in rutile ($700\text{--}420^{\circ}\text{C}$, e.g. Mezger et al., 1989; Cherniak, 2000) from 30–19 Ma, which yields an average cooling rate of ca. $25^{\circ}\text{C}/\text{m.y.}$ for this time period (Figure 2.10). However, the temperature-time path for the Gruf Complex after juxtaposition of the UHT and lower grade units (between 30 and 19 Ma) may have been more complex. The crystallization ages of underformed dikes (ca. 25 Ma; Oalman et al., in preparation-c) and the Novate granite (ca. 24 Ma; Liatì et al., 2000) indicate that the Gruf Complex was above 650°C until at least 24

Ma. Therefore, a likely scenario involves slow cooling (ca. 8°C/m.y.) between 30 and 24 Ma followed by more rapid cooling (ca. 46°C/m.y.) between 24 and 19 Ma (Figure 2.10). This scenario is in good agreement with the conclusions of Berger et al. (2011) who used lower temperature thermochronology (zircon and apatite fission track ages) to constrain the cooling history of the Central Alps.

Decoupling of Zr-in-rutile temperatures and U-Pb ages

Figures 7 and 11 and Table 2.4 show that U-Pb ages and Zr-in-rutile temperatures are strongly decoupled. Grains with Zr concentrations indicating UHT crystallization have $^{206}\text{Pb}/^{238}\text{U}$ dates anywhere from 19 to 30 Ma, and lower temperature (700–900°C) grains also span this age range. We interpret these results to indicate that Pb and Zr behave significantly different during retrograde metamorphism and cooling.

We interpret that zirconium concentrations represent crystallization temperatures in most cases, and that Zr is only mobilized by recrystallization or diffusion if there is a transport medium and pathway (see *Zr-in-rutile Thermometry* section above). The lowest Zr-in-rutile temperatures recorded by matrix rutile grains are ca. 700°C. Lower temperatures are recorded by rutile grains included in garnet, but these grains likely formed during prograde metamorphism (see *Zr-in-rutile Thermometry* section above). We interpret these results to indicate that rutile does not release Zr by diffusion below ca. 700°C, which is consistent with empirical and experimental studies (e.g. Cherniak et al., 2007; Kooijman et al., 2012). Based on the lack of correlation between ages and temperatures, we conclude that Pb is completely lost from rutile at temperatures above 700°C, and Pb is retained only below 700°C at the high cooling rates experienced by the UHT rocks of the Gruf complex. This is consistent with results of other studies that determined the closure temperature for Pb diffusion in rutile to be between ca. 420 and 700°C (Mezger et al., 1989;

Mezger et al., 1991; Cherniak, 2000; Möller et al., 2000; Vry and Baker, 2006). In summary, these results support the notion that the temperature range for Zr retention is higher than that for Pb retention. Diffusion behavior is controlled by availability of fluids/melts as transport media and grain boundaries as diffusion pathways.

Regional significance of Zr-in-rutile thermometry and U-Pb thermochronology

The results from texturally controlled Zr-in-rutile thermometry of this study indicate that the UHT event occurred late in the evolution of the granulites and charnockites. However, from this study alone, the precise timing of this event cannot be constrained. Our interpretation that initial decompression leading to garnet breakdown and symplectite formation occurred at UHT conditions is in contrast to the results of Galli et al. (2011), who estimated that symplectite formation occurred at ca. 740°C and 6.5–7.5 kbar. However, the results of Galli et al. (2011) are inconsistent with the first-order petrographic observation of UHT assemblages in the garnet breakdown textures (Figure 2.5A, D, F) and the Zr-in-rutile temperatures presented here. Therefore, we argue that this short-lived UHT event occurred during Alpine orogenesis. The lack of Zr-diffusion profiles in rutile grains supports the hypothesis that UHT metamorphism and the subsequent upper amphibolite to lower granulite facies overprint occurred within a very short timeframe (ca. 5–6 Ma), during a single metamorphic cycle.

There is good agreement between the mean Zr-in-rutile temperature from lower temperature matrix domains in the UHT granulites and peak temperature estimates of the paragneisses of 700–750 °C (Oalman et al., in preparation-a). We interpret this as evidence for coeval recrystallization of some granulite domains and peak metamorphism of the paragneisses in the Gruf Complex at the same crustal level. We conclude that this happened after the Alpine juxtaposition of the granulite/charnockite and lower grade gneiss units as documented by Galli et

al. (2013). The juxtaposition and exhumation of the composite Gruf Complex likely occurred during and as a result of emplacement of the Bergell intrusion at ca. 30 Ma as discussed by (Galli et al., 2013).

The range of rutile ages (30–19 Ma) from all samples centers broadly on the crystallization age of the Novate leucogranite at 24 ± 1 Ma (Liati et al., 2000) and the latest crosscutting dikes within the Gruf Complex at 25.6 ± 0.3 Ma (Oalmann et al., in preparation-c). These combined results indicate that the Gruf granulites and charnockites cooled to below the closure temperature of Pb diffusion in rutile coevally with the latest magmatic crystallization events in the immediate vicinity (Figure 2.10, 12). The rutile cooling ages also coincide with the latest crystallization of zircon in migmatites throughout the Lepontine Dome at ca. 22 Ma (Figure 2.12; Rubatto et al., 2009). It can therefore be proposed that after the UHT rocks of the Gruf Complex were exhumed to mid-crustal levels, upper amphibolite facies conditions persisted between ca. 30 and 24 Ma in the Southern Central Alps, and rapid cooling ensued between 24 and 19 Ma (Figure 2.10, 12).

Conclusions

The results presented in this study have some general implications for the use of rutile as a geothermochronometer in high-grade metamorphic rocks. Hafnium is the only HFSE that is strongly correlated with Zr. Concentrations of the other HSFES do not appear to be a function of crystallization temperature and have most likely been modified during retrogression and cooling. Some trace elements (e.g. W) may be used to constrain fluid compositions. Zirconium-in-rutile thermometry is a robust method for determining prograde, peak, and retrograde temperatures in UHT rocks when specific attention is given to the textural setting of rutile. However, diffusional loss of Zr from rutile may occur if a Zr sink is present at or near the grain boundary. In the case of a short-lived UHT event with a very high cooling rate as in the Gruf Complex, Zr-in-rutile

temperatures generally represent crystallization temperatures. However, in terranes with slow post-peak metamorphic cooling, some modification of Zr-in-rutile temperatures may occur, and retrieving peak-T conditions can be difficult (e.g. Kooijman et al., 2012; Pape et al., 2016).

U-Pb rutile cooling ages are not correlated with texture or Zr-in-rutile temperatures. This indicates that Pb diffusion is controlled by the presence or absence of localized fast diffusion pathways such as fluids, melts, or intragrain defects. Above 700°C rutile does not effectively retain Pb, whereas Zr is fully retained by rutile below ca. 700°C. The temperature range for Pb-retention is likely cooling rate dependent.

The granulites of the Gruf Complex most likely reached UHT conditions during Alpine orogenesis, and some recrystallization of the granulites occurred during and after juxtaposition of these lower-crustal rocks with mid-crustal migmatites. The Gruf complex underwent fast cooling on the order of 63°C/m.y. between 34 and 30 Ma; cooling rates slowed to ca. 8°C/m.y. between 30 and 24 Ma and then accelerated to ca. 46°C/m.y. between 24 and 19 Ma. The U-Pb rutile age range of 30–19 Ma coincides with emplacement of the Novate leucogranite at 24 ± 1 Ma (Liati et al., 2000) and migmatization within the Southern Steep Belt between 32 and 22 Ma (e.g. Rubatto et al., 2009). The most likely geodynamic scenario for UHT conditions being reached during decompression is considered to be slab breakoff, which resulted in heat transfer from asthenospheric upwelling to the crust by ca. 34 Ma.

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assistance, and C. Fischer at the University of Potsdam and W.C. Dickerson at the University of Kansas for preparing thin sections.

Figure Captions

Figure 2.1. Tectonic map of the Central Alps showing the location of the Gruf Complex and surrounding units (color coded by paleogeographic affinity) and major faults (modified from Bousquet et al., 2012b). SSB = Southern Steep Belt. Inset shows location of map area within Europe. Ages of Bergell and Novate Intrusions are from von Blanckenburg (1992) and Liati et al. (2000), respectively.

Figure 2.2. Geologic map of the Gruf Complex and surrounding units with strike and dip of foliations and locations of samples used in this study (modified from Galli et al., 2013). Inset shows location of map area within Europe. Ages of Bergell and Novate Intrusions are from von Blanckenburg (1992) and Liati et al. (2000), respectively.

Figure 2.3. Photos of rock types. (A) Residual granulite (sample G1A_2009). (B) Stromatic granulite (sample Bras10-5). (C) Leucogranulite (upper part of photo; sample GR11-35). (D) Charnokite (sample GR11-43). (E) Large granulite boulder from Val Codera with multiple mineralogical/textural domains; 1 and 2 refer to samples GR11-1A and GR11-1C, which are retrogressed granulites. (F) Example of mylonite zone separating charnokite from orthogneiss.

Figure 2.4. Scans of thin sections of selected granulites showing Zr-in-rutile temperature ranges for different textural domains. (A) Stromatic granulite sample BRAS10-5b showing coarser-

grained Domain A and finer-grained Domain B. (B) Retrogressed granulite GR11-1C showing coarser-grained Domain A and finer-grained Domain B.

Figure 2.5. Photomicrographs with Zr-in-rutile temperatures (in red). (A) Residual granulite (sample G1A_2009): Embayments in large garnet porphyroblast filled with orthopyroxene + sapphirine + rutile + minor biotite. Also note the orthopyroxene + cordierite sillimanite + sapphirine symplectite, which surrounds biotite, on the left side of the image. (B) Residual granulite (sample G1A_2009): Large biotite clot with inclusions of sapphirine + rutile. Left side of photo shows a relict garnet porphyroblast with sapphirine + cordierite + orthopyroxene-filled embayments and orthopyroxene + cordierite + sapphirine symplectites between the garnet and the biotite clot. Right side of photo shows quartz + K-feldspar leucosome with fine-grained sapphirine aggregates and biotite. (C) Stromatic granulite (sample Bras10-5b-2): Quartz + K-feldspar + plagioclase leucosome surrounding orthopyroxene + several rutile grains. Note the low apparent temperature for the rutile grain in contact with zircon compared. (D) Stromatic granulite (sample Bras10-5b-1): Sapphirine + orthopyroxene + biotite + rutile intergrowths and symplectites at the edge of a large garnet porphyroblasts. Quartz + K-feldspar + plagioclase leucosomes and large orthopyroxene porphyroblasts surround the garnet and its finer-grained reaction products. (E) Stromatic granulite (Bras10-6-2): Large orthopyroxene porphyroblasts surrounded by cordierite + biotite with minor quartz and plagioclase. Note the very fine-grained needles (red arrow) and small grains (green arrow) of rutile within the orthopyroxene. (F) Stromatic granulite (Bras10-6-2): Orthopyroxene + sapphirine intergrowths, rutile, biotite, and plagioclase within an embayment in garnet. (G) Stromatic granulite (Bras10-6-1): Rutile with small zircon grains along grain boundaries within a biotite rich layer. (H) Leucogranulite (sample GR11-35): Relict garnet

porphyroblast breaking down to plagioclase + spinel + opaques surrounded by quartz + K-feldspar + minor plagioclase leucosome. A small garnet porphyroblast is intergrown with orthopyroxene. (I) Charnockite (sample GR11-43): Anhedral garnet + biotite + rutile within quartz + K-feldspar + plagioclase leucosome. (J) Retrogressed granulite (sample GR11-1A): Large garnet porphyroblast with a rutile inclusion surrounded by orthopyroxene + cordierite symplectite. Biotite (with exsolved rutile needles) + orthopyroxene + sillimanite + cordierite occurs outside of the symplectite. (K) Retrogressed granulite (sample GR11-1A): Garnet porphyroblast with orthopyroxene + cordierite symplectite surrounded by orthopyroxene + biotite + cordierite + rutile. (L) Retrogressed granulite (sample GR11-1C): Garnet porphyroblast with chlorite-filled fractures and spinel, rutile, opaque, and biotite inclusions. Sillimanite and cordierite are present between the garnet and an orthopyroxene porphyroblast. Relict sapphirine grains fringed by spinel aggregates are also present. Black boxes in panels A, B, and D denote areas shown in Figure 2.7 where Zr-in-rutile temperatures are shown. Mineral abbreviations after Whitney and Evans (2010).

Figure 2.6. Concordia plots: (A) All samples combined. (B)–(J) Individual plots for each sample. Gray ellipses represent analyses not used in age calculations. L.I. = lower intercept age. In panel (F), the two lower intercept ages shown are calculated for the yellow and purple ellipses, respectively. In panel (H), ages near the inset refer to weighted mean $^{206}\text{Pb}/^{238}\text{U}$ ages for each group of ellipses (blue and purple).

Figure 2.7. Photomicrographs with plots of $^{206}\text{Pb}/^{238}\text{U}$ dates (upper), and Zr-in-Rt temperatures (lower) for selected textures from different granulite types. (A) Rutile within garnet breakdown texture in garnet embayment from residual granulite sample G1A_2009. (B) UHT rutile +

sapphirine with biotite-rich clots from residual granulite sample G1A_2009. (C) UHT rutile-bearing sapphirine + orthopyroxene symplectite around garnet with lower temperature rutile inclusions in garnet from stromatic granulite sample Bras10-5b-1. (D) UHT rutile from stromatic granulite sample Bras10-6-2. (E) UHT rutile from stromatic granulite sample Bras10-6-1. (F) Near UHT rutile with sapphirine + spinel aggregate in leucogranulite sample GR11-35. Error bars on date plots represent 2SE internal plus propagated uncertainty. Error bars on temperature plots represent 2SE internal uncertainty. Mineral abbreviations after Whitney and Evans (2010). See Figure 2.5 for locations of photos A, B, and C.

Figure 2.8. Rutile trace element concentrations for granulites and charnockites. Points represent individual spot analyses. Symbol colors represent textural domains (blue: inclusions in garnet; red: peak matrix domains; green: retrogressed matrix domains), and symbol shapes represent sample. (A) Hf versus Zr with calculated linear regression line. (B) U versus Zr. (C) W versus Zr. (D) Nb versus Zr. (E) Ta versus Nb. (F) Nb/Ta versus Ta. (G) Zr/Hf versus Hf. (H) Nb/Ta versus Zr/Hf.

Figure 2.9. Composite histogram showing Zr-in-rutile temperatures color coded by textural setting (blue: inclusions in garnet; green: retrogressed matrix domains; red: peak matrix domains). Temperatures given in the figure are means calculated from each histogram. Rextal. = recrystallized.

Figure 2.10. Schematic temperature versus time diagram for evolution of Gruf granulites. Boxes represent time/temperature ranges for the ages listed in the figure. U-Pb zircon ages for granulites and dikes are from (Oalman et al., in preparation-c). Cooling rates are also shown for various

segments of the temperature-time path. For the time period between 30 and 19 Ma, both a continuous and two-stage cooling path are shown. The diagrams across the top of the Figure 2 are schematic cross sections showing stages of dike emplacement and deformation at given times. The pluton symbols across the bottom are scaled to show U-Pb crystallization ages of the Bergell intrusion (von Blanckenburg, 1992) and Novate leucogranite (Liati et al., 2000).

Figure 2.11. Zr-in-rutile temperature versus $^{206}\text{Pb}/^{238}\text{U}$ date for rutile analyses from representative stromatic and retrogressed granulite samples. Rutile from all textural domains are included. Error bars are 2SE internal plus propagated uncertainties on the $^{206}\text{Pb}/^{238}\text{U}$ date. Uncertainty of the Zr-in-rutile temperatures are less than or equal to the height of the symbols.

Figure 2.12. Summary of metamorphic and plutonic ages in the Central Alps (going from north to south) roughly color coded for temperature of event. Numbers refer to the following ages: 1) U-Pb staurolite: ca. 28 Ma (Nagel, 2002); 2) U-Pb zircon (rims): 33–32 Ma (Liati et al., 2009); 3) Ar-Ar phengite: 33–30 Ma (Talerico, 2000); 4) U-Pb zircon: 37.1 ± 0.9 (Liati et al., 2003); 5) U-Pb zircon: undeformed dikes = 25.6 ± 0.3 Ma; deformed dikes = 30–27; zircon rims in UHT rocks = 32.9 ± 0.5 Ma (Oalman et al., in preparation-c); 6) U-Pb monazite in UHT rocks: 34–31 Ma (Oalman et al., in preparation-b); 7) U-Pb zircon: 24.0 ± 1.2 Ma (Liati et al., 2000); 8) U-Pb zircon (tonalite): 31.88 ± 0.09 Ma, U-Pb allanite (granodiorite): 30.1 ± 0.25 Ma (von Blanckenburg, 1992); 9) U-Pb zircon: 32–22 Ma (Rubatto et al., 2009); 10) U-Pb zircon: 34.2 ± 0.2 and 32.9 ± 0.3 Ma (Hermann et al., 2006); 11) Lu-Hf garnet: 70–36 Ma (Brouwer et al., 2005); 12) U-Pb zircon: 43 ± 2 and 35.4 ± 0.5 Ma (Gebauer, 1996). Height of black boxes is scaled to age plus uncertainty or age range. SSB = Southern Steep Belt.

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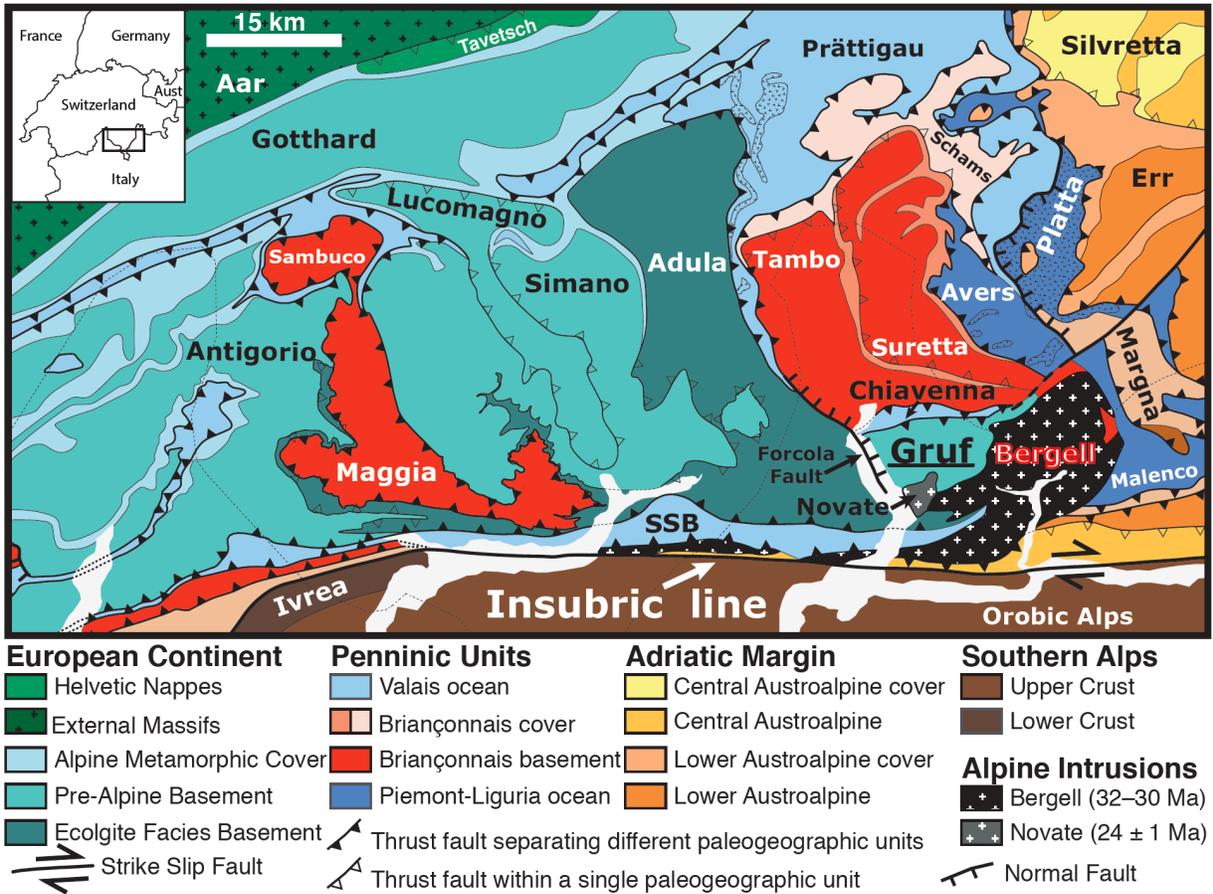


Figure 2.1

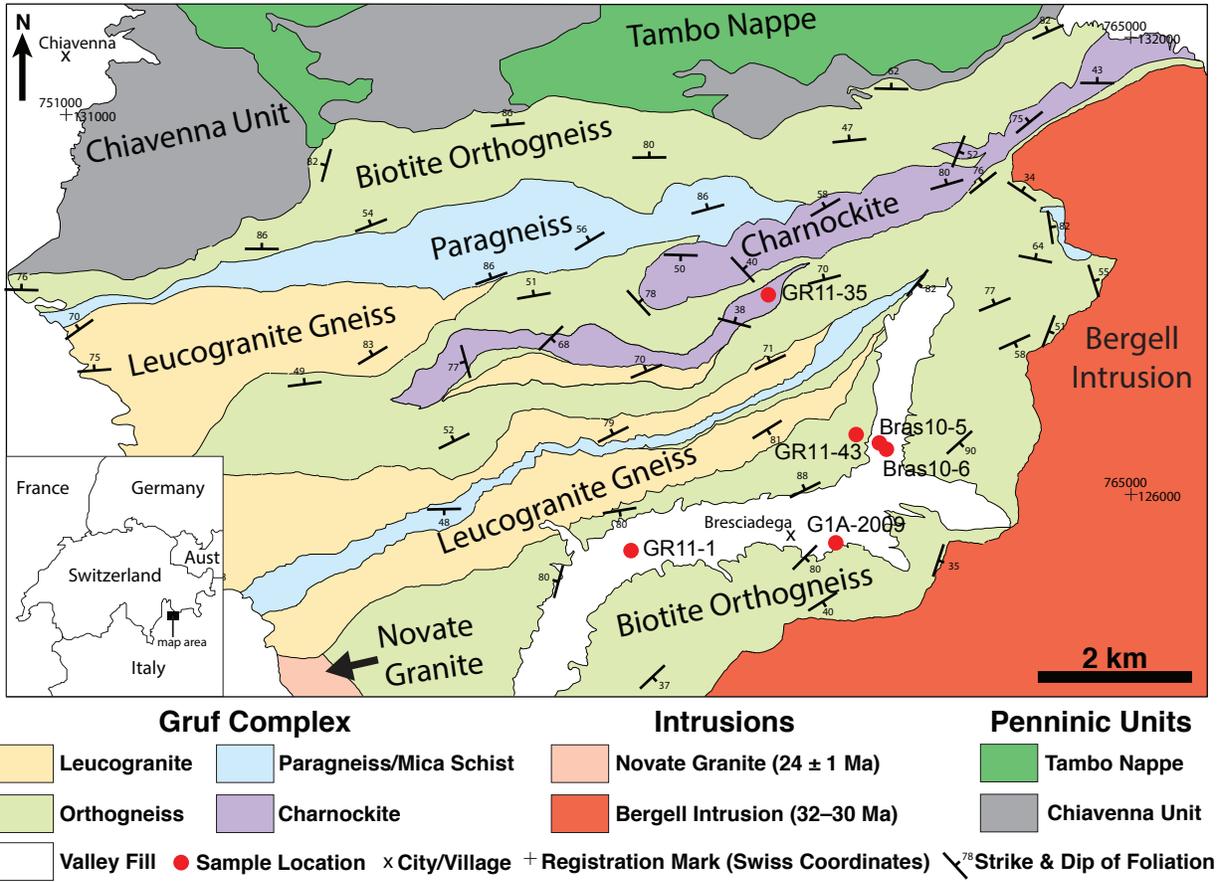


Figure 2.2

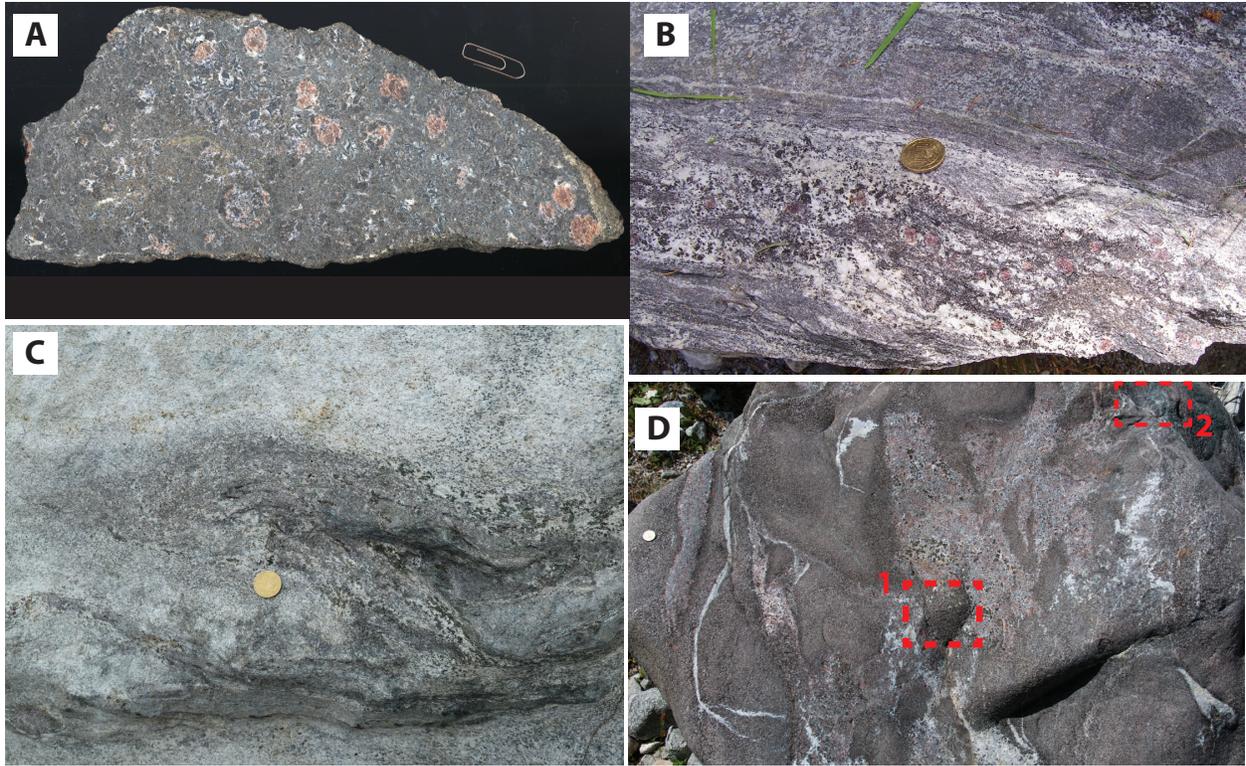


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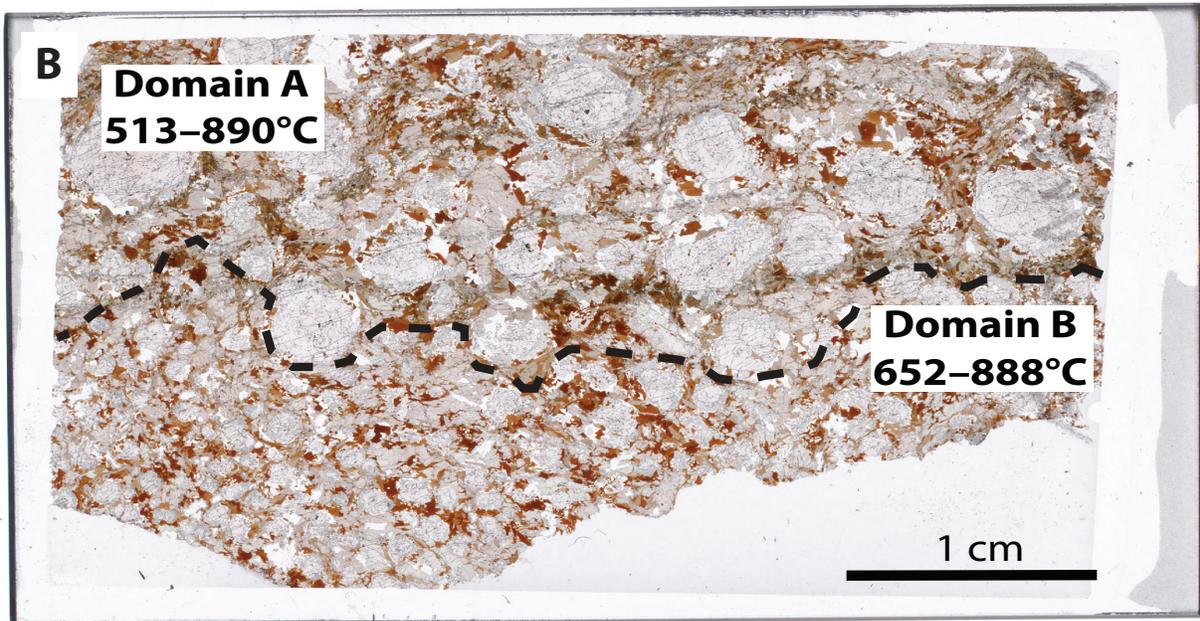
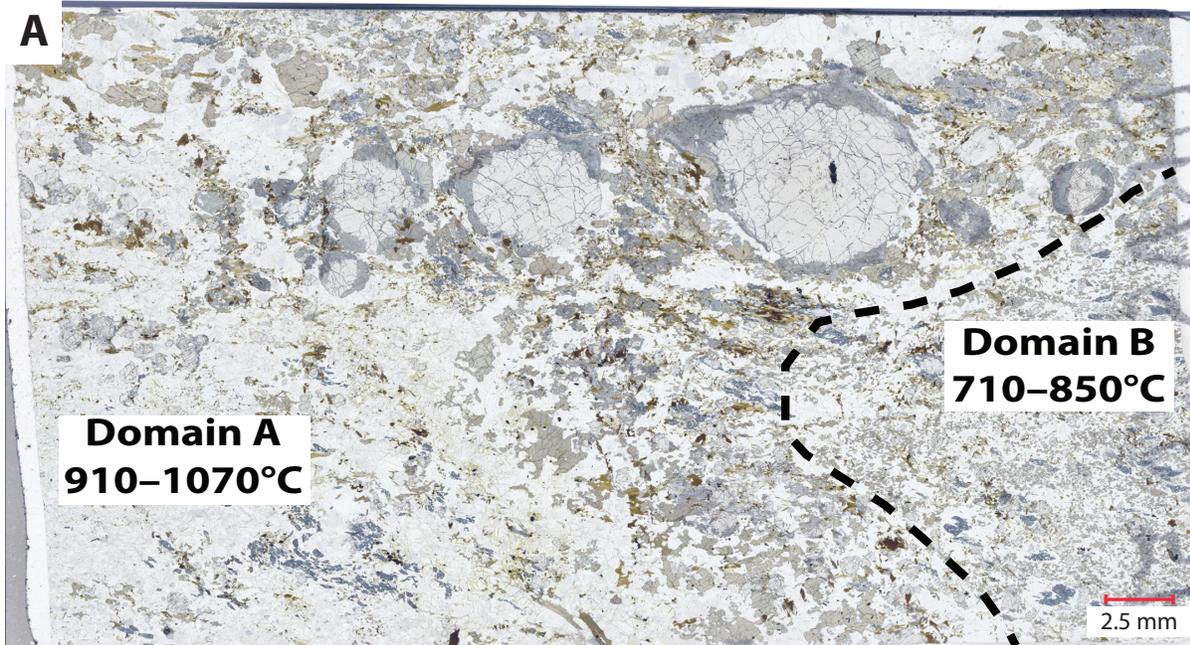


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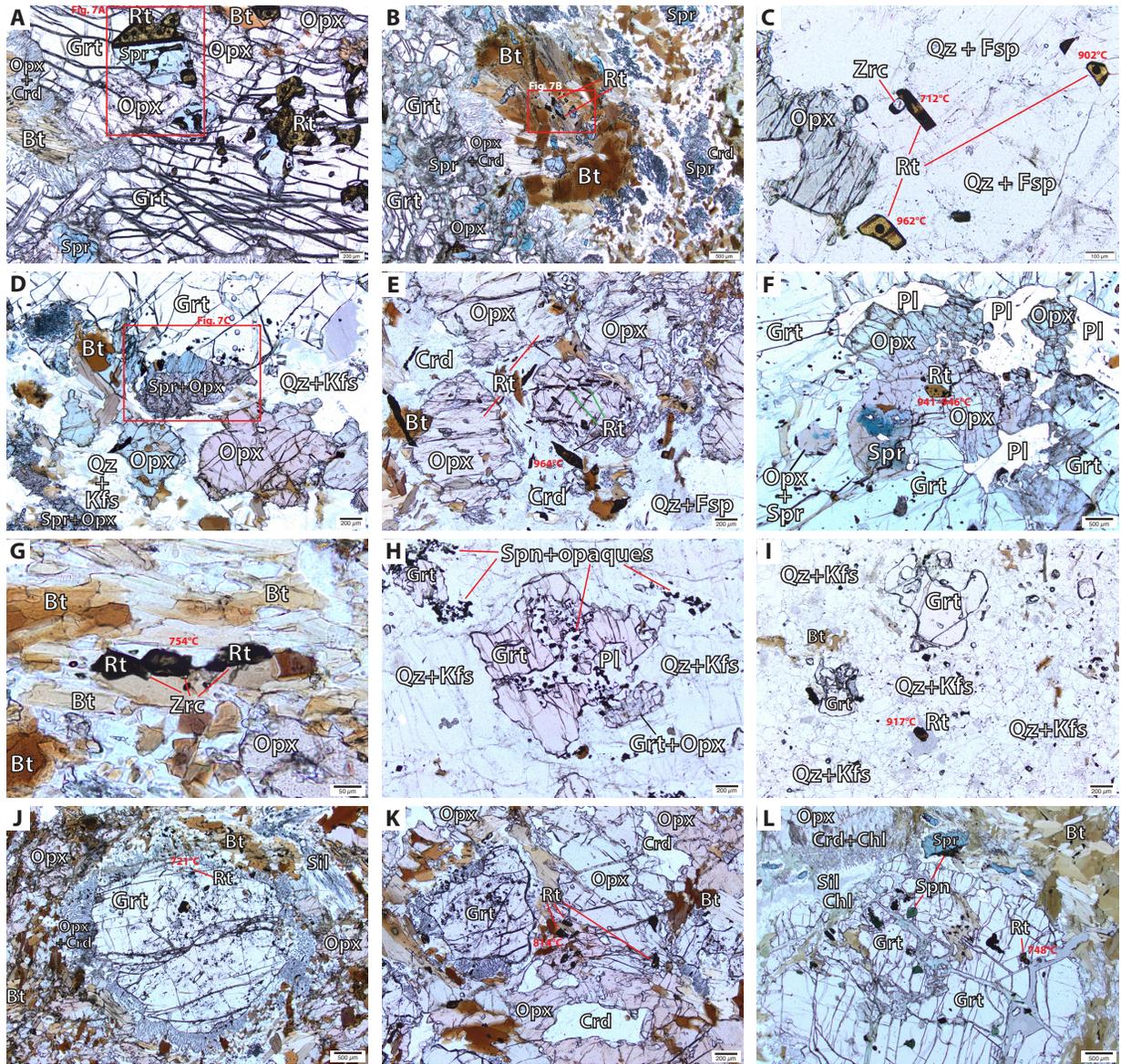


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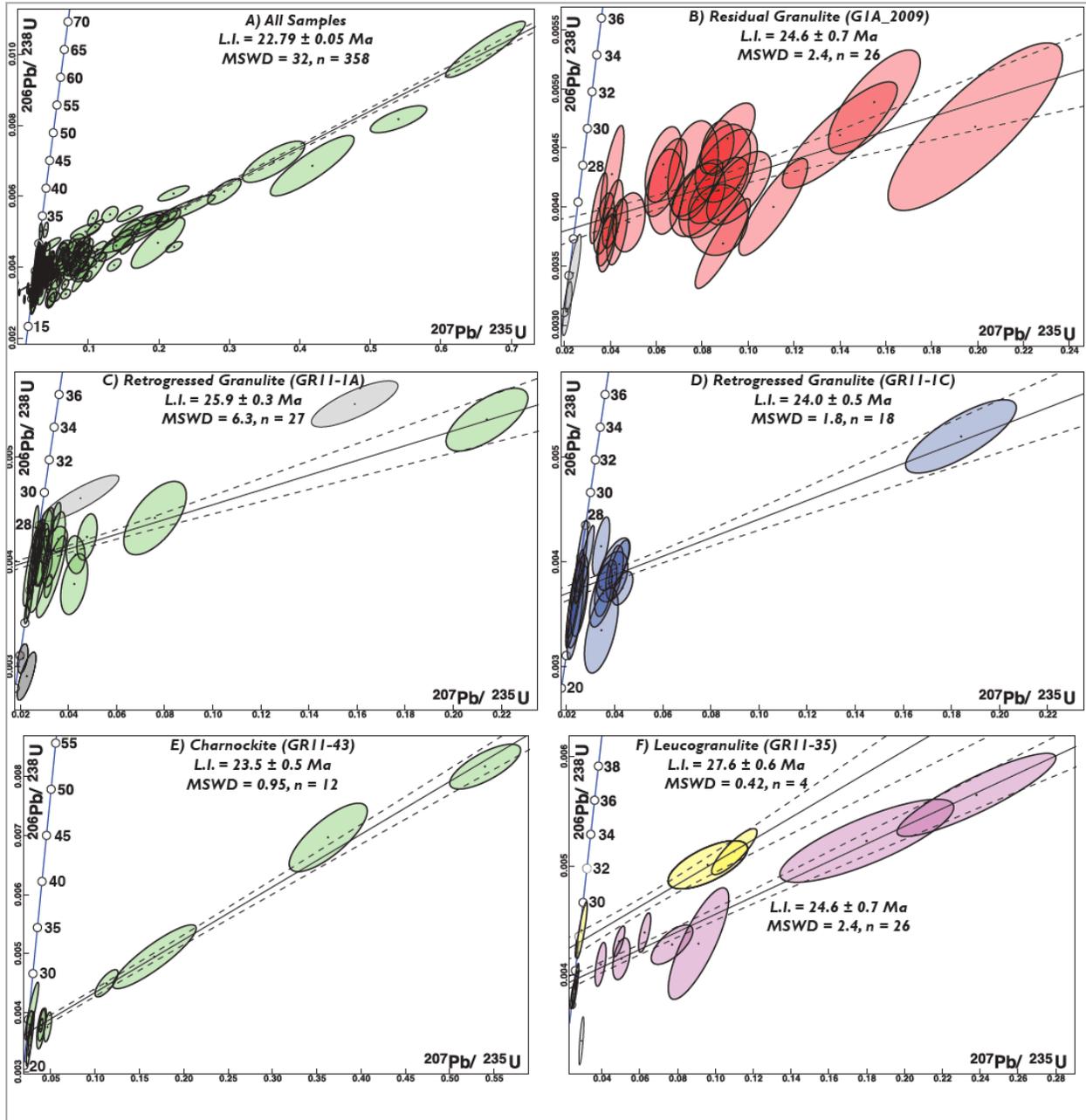


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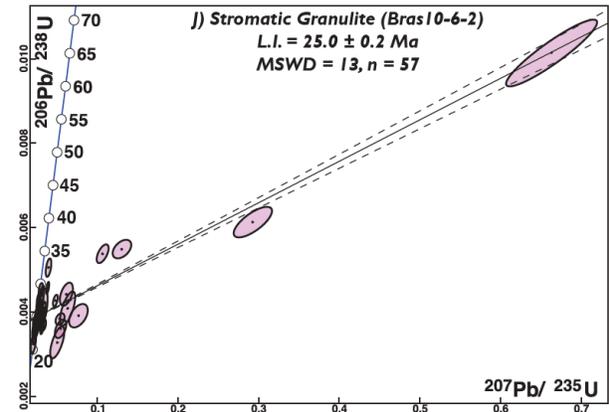
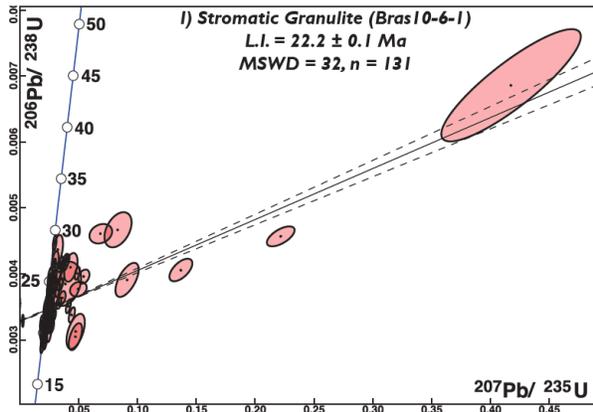
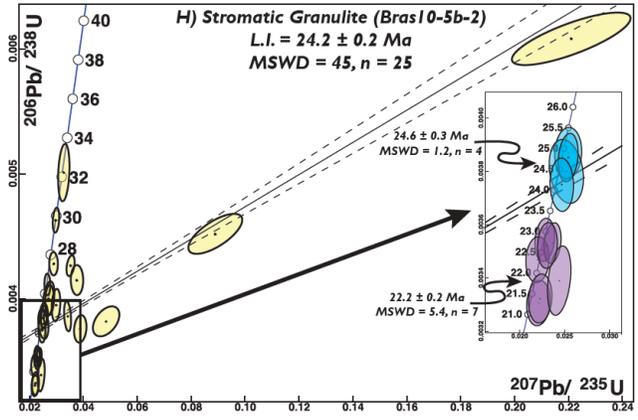
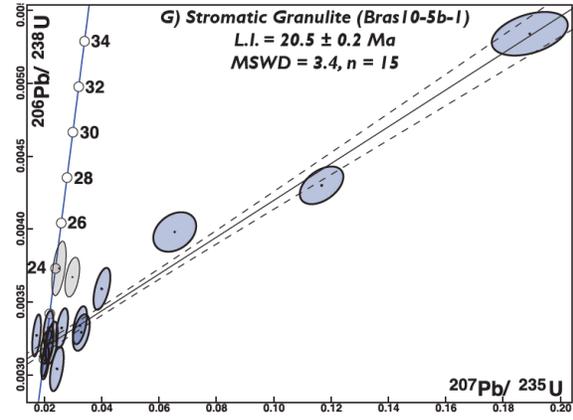


Figure 2.6 (continued)

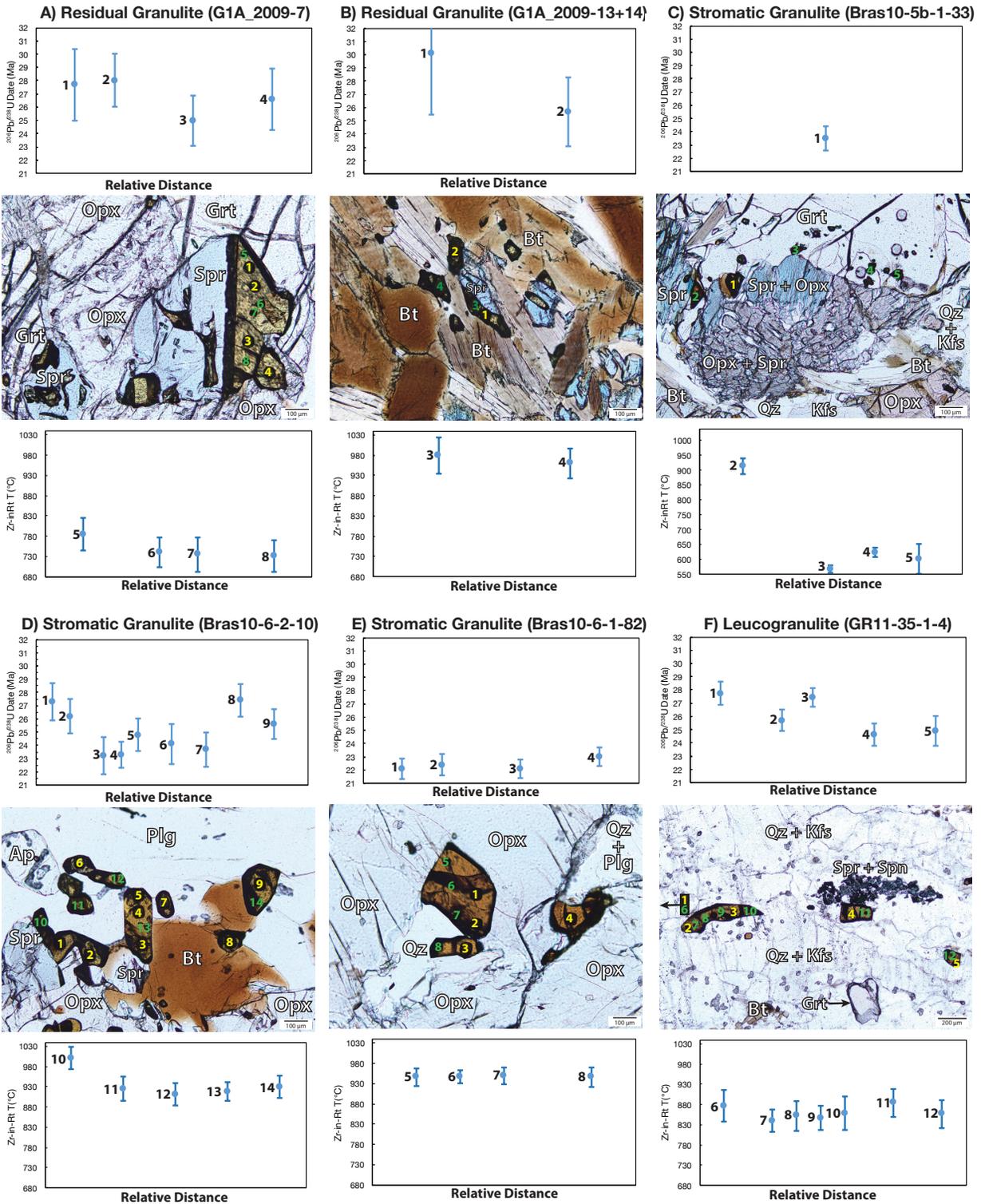


Figure 2.7

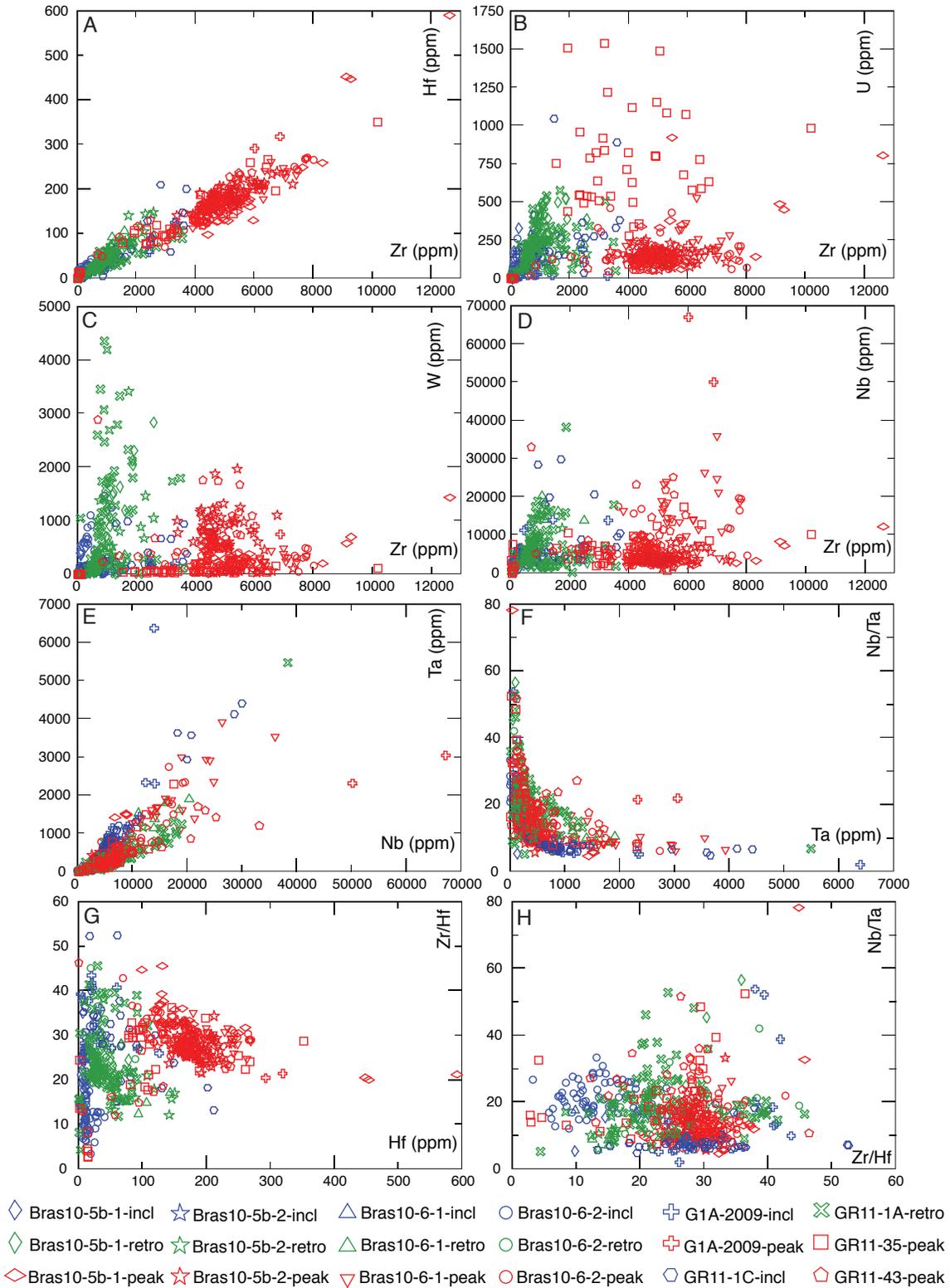


Figure 2.8

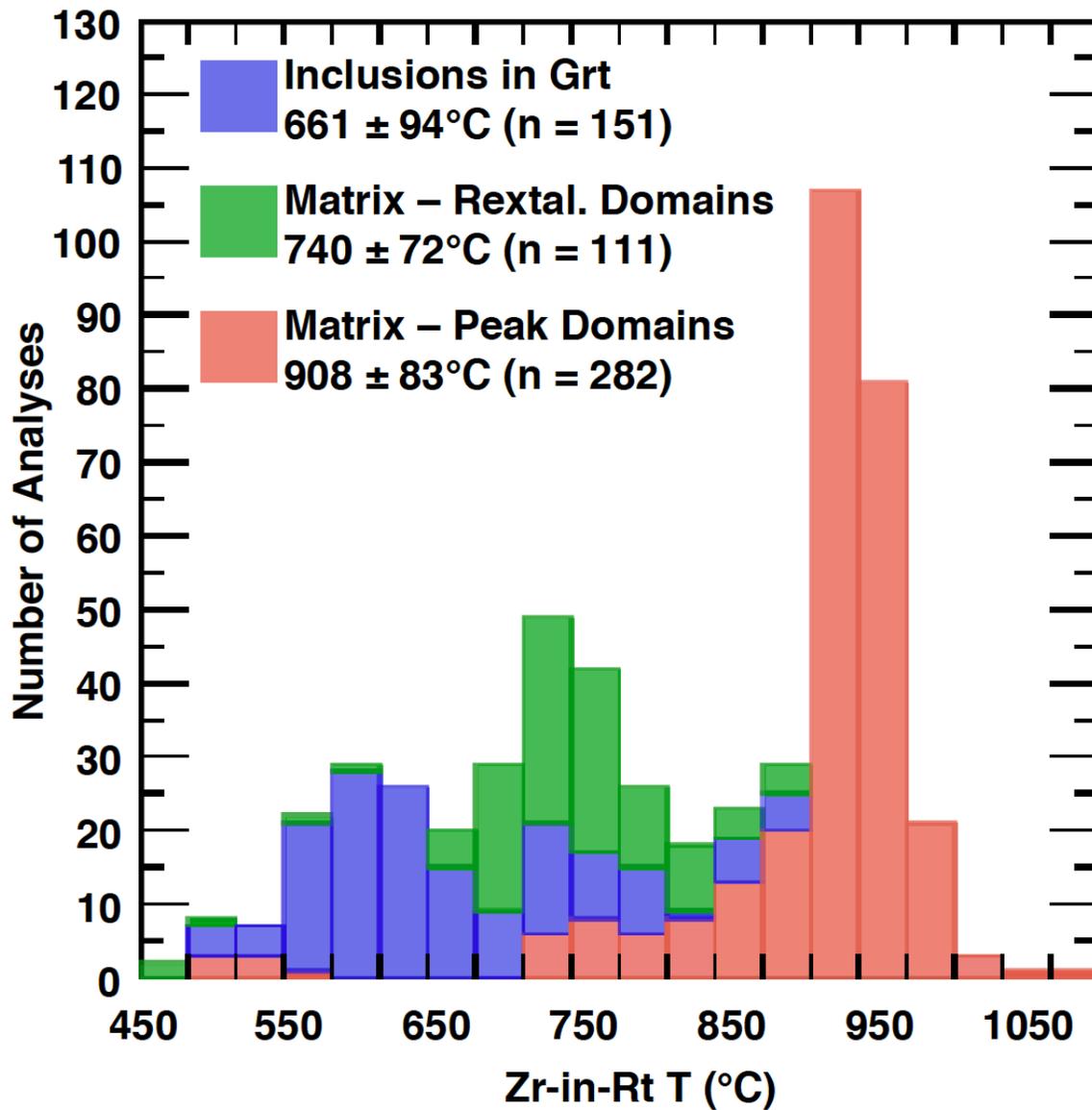


Figure 2.9

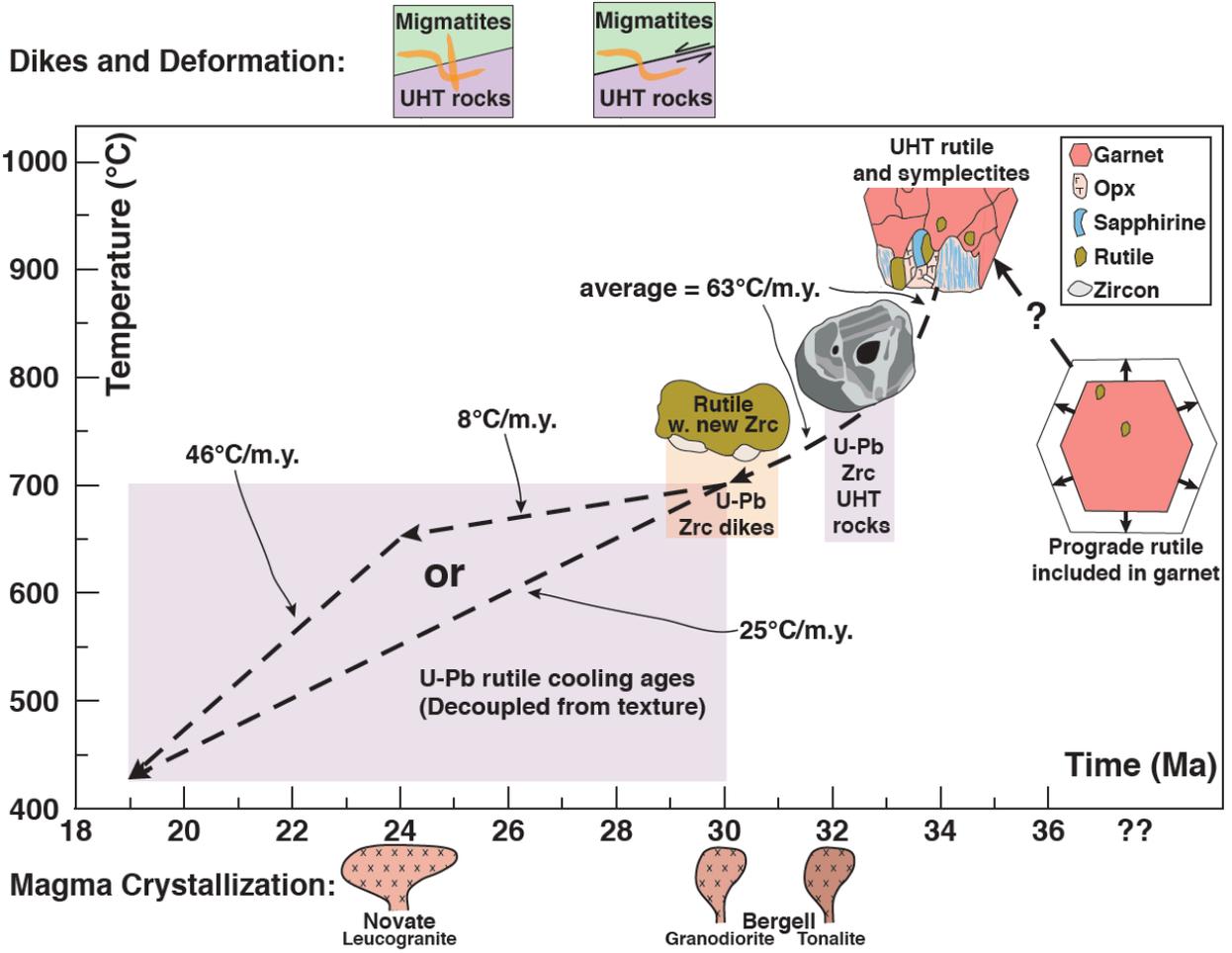


Figure 2.10

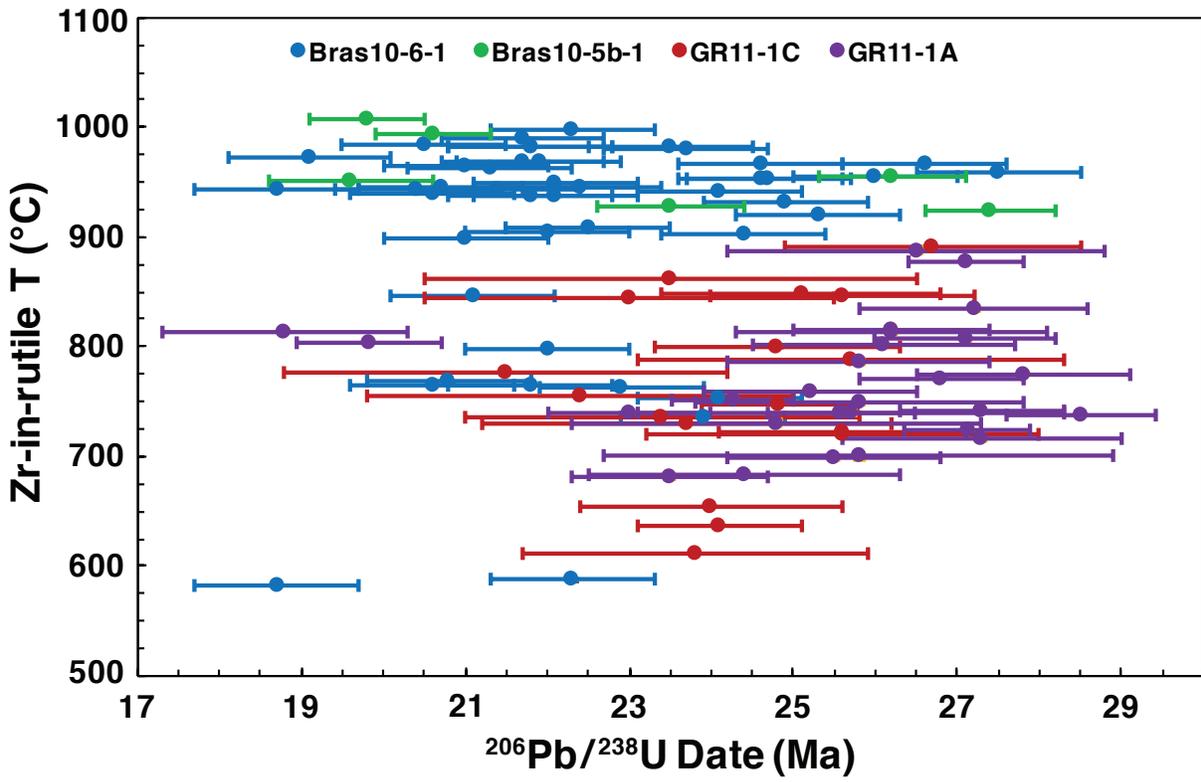


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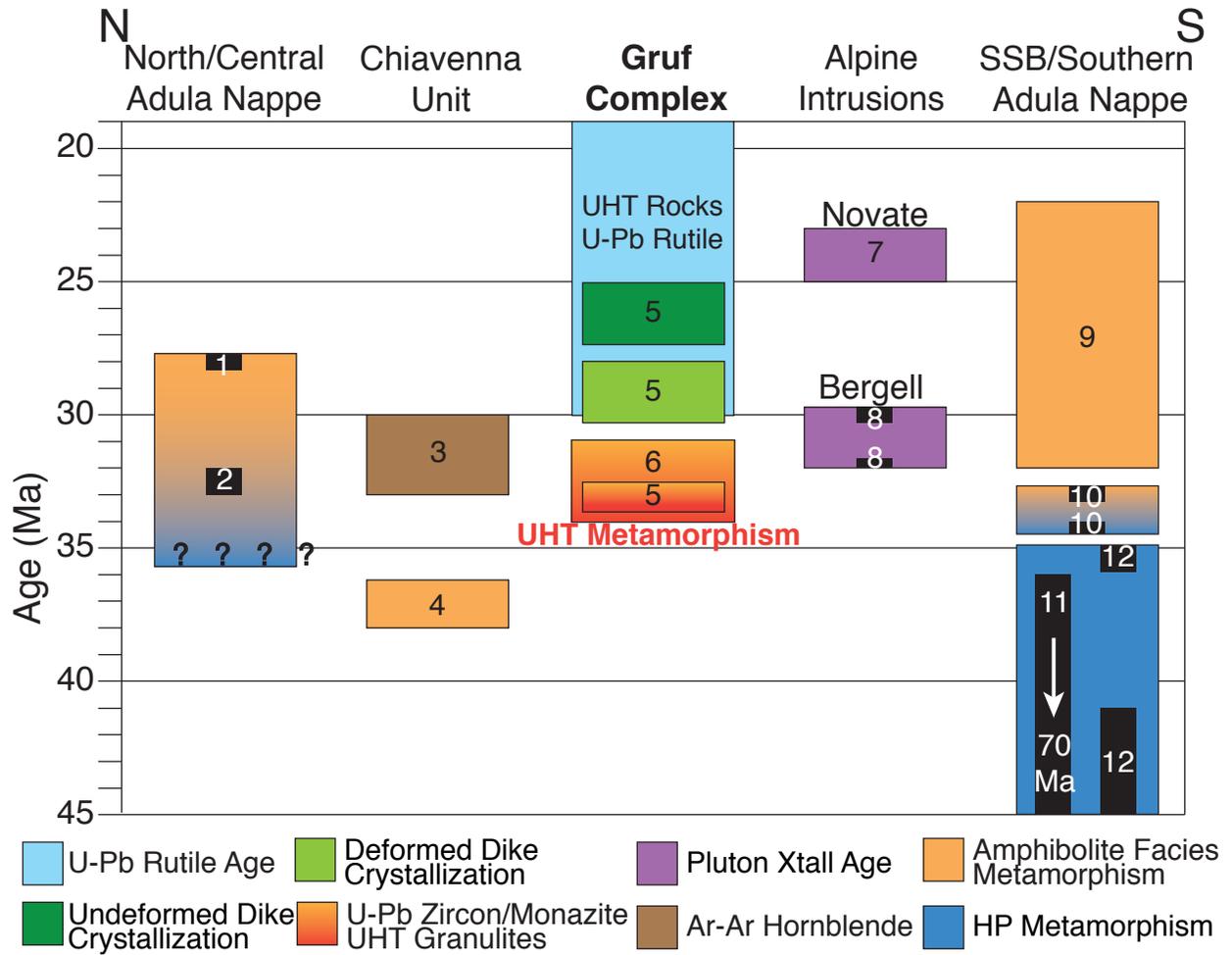


Figure 2.12

Table 2.1. LA-ICP-MS Analytical Methods

Laboratory & Sample Preparation	
Laboratory name	The University of Kansas, Dept. of Geology, Isotope Geochemistry Lab
Sample type/mineral	Rutile
Sample preparation	40 μm -thick, polished sections
Imaging	Petrographic microscope
Laser ablation system	
Make, model & type	ATL Arf excimer laser (193 nm), Photon Machines Analyte
Ablation cell & volume	Large volume 'Frames Cell' with teardrop-shaped insert
Laser wavelength	193 nm
Pulse width (ns)	5 ns
Fluence	3.7 J/cm ²
Repetition rate	10 Hz
Spot size (μm)	Trace elements: 22 μm ; U-Pb: 36 μm
Sampling mode / pattern	Single spots
Carrier gas	He, 0.4 l/min, Ar, 1.2 l/min
Ablation duration	30 s
Cell carrier gas flow	He, 0.4 l/min
ICP-MS Instrument	
Make, Model & type	Thermo Element2 magnetic sector field ICP-MS (single collector)
Sample introduction	Aerosol with sample + He was mixed with Ar using a Y-connector 15 cm upstream from torch. For some samples, a cyclonic spray chamber was attached at the inlet of the torch (downstream from the T-connector to improve stability of the aerosol introduction.
RF power	1100–1350 W
Make-up gas flow	Ar, 1.2 l/min
Sampling depth	20 μm
Detection system	single detector (SEM), counting & analog modes
Elements/ isotopes analyzed	Trace elements: 29Si, 49Ti, 51V, 53Cr, 90Zr, 93Nb, 95Mo, 118Sn, 120Sn, 121Sb, 178Hf, 181Ta, 184W, 206Pb, 208Pb, 232Th, 238U; U-Pb: 202Hg, 204Pb, 206Pb, 207Pb, 208Pb, 232Th, 235U, 238U
Integration time per peak (Sample Time)	Trace elements: 5 ms for all isotopes; U-Pb: 202=1 ms, 204=1 ms, 206=4 ms, 207=4 ms, 208=2 ms, 232=2 ms, 235=3 ms, 238=1 ms
Total integration time (Segment Duration)	Trace elements: 4 ms for all isotopes; U-Pb: 202Hg=4 ms, 204Hg/Pb=4 ms, 206Pb=16 ms, 207Pb=16 ms, 208Pb=8 ms, 232Th=8 ms, 235U=12 ms, 238U=4 ms

Table 2.1

Total method time	Trace elements: 1:31 (300 runs, 1 pass); U-Pb: 1:33 (230 runs, 5 passes)
Sensitivity (cps/ppm)	Si=464cps/ppm, Ti=1151cps/ppm, V=1390cps/ppm, Cr=bdl, Zr=1050cps/ppm, Nb=1391cps/ppm, Mo=2186cps/ppm, Sn=6048cps/ppm, Sb=2228cps/ppm, Hf= 2063cps/ppm, Ta=1743cps/ppm, W=2515cps/ppm, Pb=4228cps/ppm, Th=2278cps/ppm, 238U=5136cps/ppm
ICP Dead time	22 ns
UO+/U+	≤0.5%
238U+/232Th+	≤1.7
Data Processing	
Gas blank	Trace elements: 30 s, U-Pb: 10 s
Calibration strategy	Trace elements: standard-sampling bracketing + internal standardization assuming 100% TiO ₂ ; U-Pb: standard-sampling bracketing
Reference material info	Trace elements: GSD-1G (Guillong et al., 2005; Jochum et al., 2005); U-Pb: R10 (Luvizotto et al., 2009)
Internal std for trace elements	49Ti
Data processing package used / Correction for LIEF	IGOR PRO, Iolite 2.5 (Patton et al. 2011): Trace elements: TraceElementsIS data reduction scheme; U-Pb: U_Pb_Geochronology (Paton, et al. 2010) and VizualAge data reduction schemes (Petrus & Kamber, 2011), exponential LIEF correction for U-Pb ratios.
Common-Pb correction, composition and uncertainty	Not performed because of high Hg backgrounds
Uncertainty level & propagation	Trace elements: 2se internal uncertainty; U:Pb: 2se propogated uncertainty from VizualAge data reduction scheme. Concordia diagrams were plotted using ET_Redux (McLean et al. 2016) with 2s uncertainty ellipses.
Reproducibility	Trace elements: 5–7%. U-Pb ratios: 2–4%
Quality control / Validation	Trace elements: R10 and R19 (Luvizotto et al., 2009); U-Pb: R19

Table 2.2. Rutile U-Pb Data

Sample and Analysis Number	Isotopic Ratios				Dates (Ma)				Compositions										
	207Pb/ 235U	±2σ %	206Pb/ 238U	±2σ %	Rho 6/38-7/35	207Pb/ 206Pb	±2σ %	206Pb/ 238U	±2σ abs	207Pb/ 235U	±2σ abs	207Pb/ 206Pb	±2σ abs	% disc a 6/38-7/35	% disc b 6/38-7/6	U (ppm)	Th (ppm)	Th/U	
Retrogressed Granulite (G1A_2009)																			
G1A_2009-1a.FIN2	0.043	10.5	0.00379	7.4	0.55	0.0848	9.4	24.4	1.8	42.6	4.4	1300	120	42.7	98.1	50.9	0.000	6.1E-06	
G1A_2009-1b.FIN2	0.0481	14.6	0.00387	6.5	0.27	0.09	12.2	24.9	1.6	47.4	6.7	1350	160	47.5	98.2	52.8	0.001	6.1E-06	
G1A_2009-1c.FIN2	0.0387	10.6	0.00372	7.3	0.55	0.0746	7.5	24	1.8	38.5	3.9	1049	97	37.7	97.7	65.3	0.000	9.3E-06	
G1A_2009-1d.FIN2	0.0383	11.2	0.00372	5.9	0.64	0.0728	9.1	23.9	1.4	38	4.2	1030	110	37.1	97.7	64.3	0.000	3.0E-06	
G1A_2009-2.FIN2	0.0867	8.2	0.0045	5.1	0.30	0.141	7.8	28.9	1.5	84.1	6.6	2260	110	65.6	98.7	58.9	0.029	5.0E-04	
G1A_2009-3.FIN2	0.155	12.9	0.00489	9.0	0.75	0.222	9.0	31.5	2.8	144	18	3010	110	78.1	99.0	17.2	0.006	3.6E-04	
G1A_2009-4a.FIN2	0.091	13.2	0.00458	6.6	0.68	0.148	10.1	29.5	1.9	88	11	2340	150	66.5	98.7	18.9	0.068	3.6E-03	
G1A_2009-4b.FIN2	0.078	14.1	0.00416	5.8	0.72	0.131	13.0	26.8	1.5	76	11	2060	160	64.7	98.7	18.9	0.020	1.0E-03	
G1A_2009-4c.FIN2	0.087	14.9	0.00414	6.8	0.64	0.146	11.6	26.6	1.8	83	12	2280	130	68.0	98.8	19.1	0.009	4.8E-04	
G1A_2009-4d.FIN2	0.089	13.5	0.00437	7.6	0.36	0.146	13.0	28.1	2.1	86	11	2370	150	67.3	98.8	18.6	0.000	5.9E-06	
G1A_2009-5.FIN2	0.096	14.6	0.00416	7.5	0.61	0.157	12.1	26.8	2	92	13	2330	130	70.9	98.8	20.5	0.000	2.1E-05	
G1A_2009-6.FIN2	0.087	19.5	0.0042	11.0	0.65	0.143	14.7	27	3	83	16	2350	200	67.5	98.9	22.5	0.001	3.6E-05	
G1A_2009-7a.FIN2	0.076	14.5	0.00431	9.7	0.34	0.132	12.1	27.7	2.7	73	10	2030	150	62.1	98.6	27.5	0.003	1.2E-04	
G1A_2009-7b.FIN2	0.0631	12.5	0.00436	7.3	0.49	0.114	11.4	28	2	61.7	7.4	1790	150	54.6	98.4	29.4	0.003	9.7E-05	
G1A_2009-7c.FIN2	0.087	14.9	0.00389	7.7	0.10	0.168	19.0	25	1.9	84	12	2380	170	70.2	98.9	28.4	bdl	-	
G1A_2009-7d.FIN2	0.0642	12.3	0.00413	8.7	0.24	0.127	12.6	26.6	2.3	62.8	7.4	2070	190	67.2	98.7	32.6	0.001	3.6E-05	
G1A_2009-8.FIN2	0.078	15.4	0.00406	7.6	0.46	0.13	14.6	26.1	2	75	11	2190	190	65.2	98.8	27.4	0.003	1.2E-04	
G1A_2009-9a.FIN2	0.0407	12.8	0.00428	9.6	0.78	0.0675	7.4	27.5	2.6	40.4	5.1	909	86	31.9	97.0	76.9	bdl	-	
G1A_2009-9b.FIN2	0.0372	11.8	0.00386	6.7	0.55	0.0675	9.2	24.8	1.6	36.9	4.3	883	75	32.8	97.2	71.1	bdl	-	
G1A_2009-10.FIN2	0.0642	14.2	0.00425	7.8	0.65	0.102	10.8	27.3	2.1	62.6	8.5	1670	120	56.4	98.4	73.9	bdl	-	
G1A_2009-11.FIN2	0.041	11.0	0.00394	7.4	0.59	0.0737	7.1	25.4	1.8	40.6	4.3	1016	92	37.4	97.5	132.3	5.080	3.8E-02	
G1A_2009-12.FIN2	0.0349	10.6	0.00399	7.5	0.75	0.0695	6.6	25.7	1.9	34.7	3.6	722	96	25.9	96.4	144.0	1.283	8.9E-03	
G1A_2009-13.FIN2	0.199	19.1	0.00468	15.2	0.82	0.286	11.5	30.1	4.6	181	32	3382	77	83.4	99.1	308.0	0.026	8.4E-05	
G1A_2009-14.FIN2	0.111	13.5	0.004	10.3	0.86	0.194	7.2	25.7	2.6	106	14	2758	99	75.8	99.1	408.0	0.009	2.2E-05	
G1A_2009-15a.FIN2	0.0211	10.0	0.00314	7.3	0.69	0.0484	7.9	20.2	1.5	21.2	2.1	380	70	4.7	94.7	232.0	0.517	2.2E-03	
G1A_2009-15b.FIN2	0.0238	14.3	0.00344	9.6	0.85	0.0471	7.6	22.1	2.1	23.8	3.3	298	71	7.1	92.6	231.0	0.098	4.2E-04	
G1A_2009-16.FIN2	0.14	17.9	0.00461	9.8	0.86	0.196	10.2	29.6	2.9	133	22	2840	160	77.7	99.0	140.0	10.400	7.4E-02	
G1A_2009-17.FIN2	0.089	13.5	0.00369	10.3	0.89	0.184	7.1	23.7	2.4	86	11	2673	92	72.4	99.1	316.0	38.900	1.2E-01	
Retrogressed Granulite (GR11-1A)																			
GR111A-1.FIN2	0.0316	5.1	0.00377	2.9	0.57	0.0607	4.3	24.27	0.74	31.6	1.6	583	60	23.2	95.8	573.0	1.890	3.3E-03	
GR111A-2.FIN2	0.0309	3.2	0.00421	2.6	0.34	0.0536	3.4	27.1	0.7	30.9	1	344	50	12.3	92.1	646.0	0.780	1.2E-03	
GR111A-8.FIN2	0.0348	5.7	0.00443	3.2	0.66	0.0566	4.6	28.52	0.91	34.7	2	456	55	17.8	93.7	291.0	1.170	4.0E-03	
GR111A-17.FIN2	0.0239	6.3	0.00365	5.2	0.79	0.0468	3.8	23.5	1.2	24	1.5	114	35	2.1	79.4	1000.0	bdl	-	
GR111A-18.FIN2	0.029	3.8	0.00425	3.8	0.50	0.049	4.1	27.3	1	29.2	1.1	238	54	6.5	88.5	198.0	2.010	1.0E-02	
GR111A-22.FIN2	0.0424	13.2	0.00379	7.7	0.41	0.084	14.3	24.4	1.9	42.1	5.5	1260	270	42.0	98.1	126.1	0.003	2.8E-05	
GR111A-23a.FIN2	0.0274	5.1	0.00416	3.8	0.71	0.0482	3.5	26.8	1	27.4	1.4	133	27	2.2	79.8	335.0	bdl	-	
GR111A-23b.FIN2	0.0232	6.5	0.00357	4.5	0.76	0.0471	4.9	23	1	23.3	1.5	140	43	1.3	83.6	486.0	bdl	-	
GR111A-24.FIN2	0.0312	8.7	0.00425	6.4	0.62	0.0512	6.8	27.3	1.7	31.2	2.7	342	88	12.5	92.0	186.4	bdl	-	
GR111A-25.FIN2	0.0247	6.1	0.00391	5.1	0.83	0.0463	3.0	25.2	1.3	24.8	1.4	140	33	-1.6	82.0	654.0	bdl	-	
GR111A-27.FIN2	0.0288	4.9	0.00421	4.0	0.73	0.0503	3.6	27.1	1.1	28.8	1.4	258	49	5.9	89.5	299.0	bdl	-	
GR111A-28.FIN2	0.0262	5.7	0.00396	5.1	0.61	0.0485	4.1	28.5	1.3	26.3	1.4	229	44	3.0	88.9	256.0	bdl	-	
GR111A-28.FIN2	0.0283	7.8	0.00442	5.2	0.42	0.0468	7.5	23.5	1.5	28.3	2.2	175	52	-0.7	83.7	365.0	bdl	-	
GR111A-28.FIN2	0.0226	17.3	0.00291	7.9	0.76	0.0538	12.7	18.8	1.5	22.7	3.9	440	110	17.2	95.7	128.0	bdl	-	
GR111A-32.FIN2	0.0292	9.6	0.00407	7.4	0.57	0.0527	8.2	26.2	1.9	29.2	2.8	374	93	10.3	93.0	231.0	bdl	-	
GR111A-35.FIN2	0.0477	8.8	0.00424	5.2	0.41	0.085	8.5	27.2	1.4	47.2	4.1	1190	170	42.4	97.7	392.0	bdl	-	
GR111A-37a.FIN2	0.0311	6.8	0.00398	3.5	0.41	0.0571	5.6	25.58	0.9	31.1	2	496	70	17.7	94.8	244.8	bdl	-	
GR111A-37b.FIN2	0.159	11.3	0.00551	4.0	0.78	0.206	8.7	35.4	1.4	148	15	2870	130	76.1	98.8	226.0	bdl	-	

Table 2.2

Sample and Analysis Number	Isotopic Ratios										Dates (Ma)										Compositions				
	207Pb/235U	±2σ %	206Pb/238U	±2σ %	Rb0/638-735	±2σ %	207Pb/206Pb	±2σ %	207Pb/235U	±2σ abs	206Pb/238U	±2σ abs	207Pb/206Pb	±2σ abs	207Pb/206Pb	±2σ abs	% disc a/638-735	% disc b/638-735	U (ppm)	Th (ppm)	Th/U				
GR111A-38.FIN2	0.045	35.6	0.00461	4.8	0.87	0.0661	21.3	29.7	1.4	44	14	480	180	110	32.5	93.8	226.0	bdl	-	-	-				
GR111A-39.FIN2	0.034	13.8	0.00412	8.7	0.68	0.0583	8.1	26.5	2.3	33.9	4.6	620	110	110	21.8	95.7	303.0	bdl	-	-					
GR111A-40.FIN2	0.0263	4.9	0.00433	4.8	0.65	0.0449	4.5	27.8	1.3	26.3	1.3	123	35	35	-5.7	77.4	303.0	bdl	-	-					
GR111A-41.FIN2	0.0328	18.6	0.00386	10.1	0.84	0.0597	10.2	24.8	2.5	32.7	5.9	600	180	180	24.2	95.9	198.0	bdl	-	-					
GR111A-42.FIN2	0.214	7.9	0.00536	5.6	0.69	0.302	6.0	34.5	1.9	198	15	3465	83	82.6	99.0	231.0	bdl	-	-	-					
GR111A-43a.FIN2	0.0261	6.5	0.00406	6.4	0.94	0.04695	2.0	26.1	1.6	26.1	1.7	84	18	18	0.0	68.9	691.0	bdl	-	-					
GR111A-50.FIN2	0.0282	18.8	0.00401	12.2	0.69	0.0513	13.8	25.8	3.1	28.2	5.2	320	140	140	8.5	91.9	237.0	bdl	-	-					
GR111A-52.FIN2	0.0286	9.1	0.00401	6.0	0.67	0.0533	6.9	25.8	1.6	28.5	2.5	402	82	82	9.5	93.6	165.0	bdl	-	-					
GR111A-57.FIN2	0.0265	6.4	0.00397	5.0	0.70	0.0486	4.1	25.5	1.3	26.5	1.7	207	42	42	3.8	87.7	253.0	bdl	-	-					
GR111A-59a.FIN2	0.0358	10.6	0.00422	2.8	0.22	0.0615	10.6	27.12	0.76	35.7	3.7	740	210	210	24.0	96.3	367.0	bdl	-	-					
GR111A-59b.FIN2	0.0253	16.2	0.00399	10.3	0.84	0.0457	9.8	25.7	2.6	25.4	4	159	74	74	-1.2	83.8	580.0	bdl	-	-					
GR111A-59c.FIN2	0.0258	8.9	0.00401	7.7	0.81	0.0484	4.8	25.8	2	25.9	2.3	167	31	31	0.4	84.6	505.0	bdl	-	-					
GR111A-60.FIN2	0.076	17.1	0.00442	7.9	0.61	0.124	12.9	28.4	2.2	73	12	1960	210	61.1	98.6	106.6	bdl	-	-	-					
GR111A-79b.FIN2	0.043	9.1	0.00407	4.4	-0.08	0.0792	11.1	26.2	1.2	42.7	3.7	1140	200	38.6	97.7	129.0	bdl	-	-	-					
GR111A-81.FIN2	0.0287	3.8	0.00456	4.2	0.39	0.0467	4.3	29.4	1.2	28.8	1.1	126	53	53	-2.1	76.7	361.0	bdl	-	-					
GR111A-81.FIN2	0.026	5.4	0.004	5.0	0.73	0.0468	3.6	25.7	1.3	26.1	1.4	179	42	42	1.5	85.6	266.0	bdl	-	-					
GR111A-81.FIN2	0.021	10.5	0.00308	4.5	0.44	0.0517	10.3	19.83	0.88	21.1	2.2	330	110	110	6.0	94.0	164.3	bdl	-	-					
Retrospected Granulite (GR11-1C)																									
GR11-1C36-1.FIN2	0.0235	16.2	0.00365	12.9	0.86	0.0469	8.3	23.5	3	23.5	3.8	293	64	64	0.0	92.0	435.0	0.000	8.3E-07	-					
GR11-1C36-3.FIN2	0.0234	5.9	0.00386	3.9	0.29	0.0481	4.8	24.83	0.94	25.4	1.4	210	46	46	2.2	88.2	247.0	0.015	6.2E-05	-					
GR11-1C36-6a.FIN2	0.0248	16.1	0.00357	10.9	0.78	0.0478	11.7	23	2.5	24.8	4	411	79	73	7.3	94.4	373.0	0.029	7.8E-05	-					
GR11-1C36-9.FIN2	0.0419	10.0	0.00398	5.8	0.72	0.0739	7.0	25.6	1.5	41.5	4.1	1030	120	38.3	97.5	193.0	0.334	1.7E-03	-	-					
GR11-1C36-11.FIN2	0.0433	9.9	0.00375	4.3	0.54	0.0828	8.2	24.1	1	42.9	4.1	1220	110	43.8	98.0	253.0	0.015	6.0E-05	-	-					
GR11-1C36-12.FIN2	0.0398	16.6	0.00398	6.3	0.65	0.07	15.7	25.6	1.6	39.5	6.4	1020	160	35.2	97.5	55.3	0.449	8.1E-03	-	-					
GR11-1C36-13.FIN2	0.184	12.5	0.0052	6.7	0.78	0.248	8.9	33.4	2.2	169	19	3230	100	80.2	99.0	174.0	1.570	9.0E-03	-	-					
GR11-1C36-15.FIN2	0.0353	10.2	0.00373	6.7	0.45	0.0687	9.9	24	1.6	35.1	3.5	950	180	31.6	97.5	166.0	0.090	5.4E-04	-	-					
GR11-1C36-16a.FIN2	0.039	16.7	0.00386	6.2	0.40	0.073	15.1	24.8	1.5	38.6	6.2	1080	260	35.8	97.7	227.0	0.003	1.5E-05	-	-					
GR11-1C36-17a.FIN2	0.0237	14.3	0.00348	11.8	0.76	0.0625	10.3	22.4	2.6	23.7	3.4	352	78	5.5	93.6	313.0	0.082	2.6E-04	-	-					
GR11-1C36-17b.FIN2	0.0353	16.1	0.0037	8.6	0.66	0.0585	10.1	23.8	2.1	35	5.5	920	160	32.0	97.4	176.0	0.218	1.2E-03	-	-					
GR11-1C36-18a.FIN2	0.0345	20.3	0.00335	12.2	0.56	0.0619	12.8	21.5	2.7	34.2	6.8	950	260	37.1	97.7	370.0	0.070	1.9E-04	-	-					
GR11-1C36-22.FIN2	0.026	10.8	0.004	10.0	0.80	0.048	6.3	25.7	2.6	26.1	2.8	236	44	44	1.5	89.1	1434.0	0.000	1.5E-07	-					
GR11-1C36-23.FIN2	0.0231	12.6	0.00364	10.2	0.83	0.0487	7.2	23.4	2.4	23.1	2.9	310	51	51	-1.3	92.5	373.0	0.089	2.4E-04	-					
GR11-1C36-24.FIN2	0.0276	14.1	0.00398	9.3	0.89	0.0476	7.1	25.6	2.4	27.6	3.8	282	52	7.2	90.9	413.0	0.129	3.1E-04	-	-					
GR11-1C36-25.FIN2	0.0254	14.6	0.00368	10.6	0.51	0.0531	13.2	23.7	2.5	25.4	3.7	410	110	6.7	94.2	638.0	0.014	2.1E-05	-	-					
GR11-1C36-26.FIN2	0.04	9.0	0.0039	6.7	0.51	0.0735	7.9	25.1	1.7	39.7	3.5	1055	60	36.8	97.6	190.0	0.672	3.5E-03	-	-					
GR11-1C36-27.FIN2	0.0344	10.5	0.00415	6.7	0.64	0.0568	6.3	26.7	1.8	34.2	3.6	514	75	21.9	94.8	1103.0	0.028	2.5E-05	-	-					
Leucogranulite (GR11-3S)																									
GR11-35-1.FIN2	0.0489	6.3	0.00431	3.2	0.73	0.0801	3.9	27.75	0.88	47.7	2.7	1188	63	41.8	97.7	1063.0	0.367	3.5E-04	-	-					
GR11-35-2A.FIN2	0.0266	4.1	0.00399	3.0	0.75	0.0482	2.7	25.7	0.79	26.7	1.1	139	27	3.7	81.5	740.0	0.000	6.5E-08	-	-					
GR11-35-2B.FIN2	0.11	10.9	0.00513	4.1	0.85	0.152	7.9	33	1.4	105	11	2290	130	68.6	98.6	745.0	0.434	5.8E-04	-	-					
GR11-35-2C.FIN2	0.0268	3.4	0.00427	2.6	0.73	0.0464	2.4	27.45	0.7	26.85	0.89	106	26	-2.2	74.1	745.0	0.001	1.9E-06	-	-					
GR11-35-3.FIN2	0.0247	6.1	0.00383	3.4	0.81	0.0465	3.7	24.62	0.86	24.7	1.5	137	24	0.3	82.0	1022.0	0.000	2.5E-07	-	-					
GR11-35-4.FIN2	0.0257	6.2	0.00387	4.4	0.83	0.0466	3.0	24.9	1.1	25.7	1.6	113	20	3.1	78.0	1269.0	0.000	8.0E-08	-	-					
GR11-35-6.FIN2	0.0502	9.2	0.00415	4.6	0.42	0.0872	7.8	26.7	1.2	49.6	4.4	1390	140	46.2	98.1	1007.0	1.080	1.1E-03	-	-					
GR11-35-15.FIN2	0.0392	7.4	0.0041	4.9	0.57	0.0679	5.7	26.4	1.3	39	2.8	804	80	32.3	96.7	1181.0	bdl	-	-						
GR11-35-16.FIN2	0.096	21.9	0.00501	4.4	0.67	0.127	18.9	32.2	1.4	91	18	1880	300	64.6	98.3	1415.0	bdl	-	-						
GR11-35-18.FIN2	0.0291	4.5	0.0034	6.5	0.71	0.0633	4.1	21.9	1.4	29.1	1.3	744	49	24.7	97.1	1164.0	1.090	9.4E-04	-	-					
GR11-35-22.FIN2	0.238	17.6	0.00565	6.7	0.89	0.302	12.9	36.3	2.4	212	34	3420	210	82.9	97.0	35.400	bdl	-	-						
GR11-35-25.FIN2	0.0625	5.1	0.00439	4.1	0.69	0.1039	4.0	28.3	1.2	62	3.2	1709	60	54.4	98.3	1462.0	bdl	-	-						

Sample and Analysis Number	Isotopic Ratios										Dates (Ma)										Compositions																
	207Pb/235U		206Pb/238U		Rb0/638-735		207Pb/206Pb		±2σ		206Pb/±2σ		207Pb/±2σ		206Pb/±2σ		207Pb/±2σ		±2σ		207Pb/±2σ		±2σ		U		Th		Th/U								
	235U	±2σ%	238U	±2σ%	638-735	±2σ%	206Pb/206Pb	±2σ%	206Pb/238U	±2σ	207Pb/206Pb	±2σ	238U	±2σ	235U	±2σ	206Pb/206Pb	±2σ	207Pb/206Pb	±2σ	abs	abs	abs	abs	ppm	ppm	ppm	ppm	Th/U	Th/U							
GR11-35-38.FIN2	0.18	25.6	0.00523	7.3	0.81	0.236	20.8	33.6	2.4	168	40	2740	420	80.0	98.8	1116.0	10.700	9.6E-03																			
GR11-35-40.FIN2	0.0293	9.9	0.00441	5.9	0.86	0.047	6.0	28.4	1.7	293	2.9	162	44	3.1	82.5	1169.0	0.000	2.5E-07																			
GR11-35-41.FIN2	0.077	14.3	0.00428	4.2	0.69	0.137	12.4	27.5	1.1	75	10	1970	230	63.3	98.6	2145.0	2.820	1.3E-03																			
GR11-35-50.FIN2	0.091	17.6	0.00429	11.9	0.81	0.157	12.1	27.6	3.2	88	15	2410	130	68.6	98.9	1380.0	1.960	1.4E-03																			
Charnockite (GR11-45)																																					
GR11-43-1	0.113	11.5	0.0045	5.1	0.76	0.182	8.8	28.9	1.5	110	12	2690	120	73.7	98.9	193.0	0.010	5.0E-05																			
GR11-43-8	0.0279	12.2	0.00364	10.4	0.85	0.0558	7.0	23.4	2.4	27.9	3.4	545	89	16.1	95.7	166.0	0.000	-4.0E-08																			
GR11-43-12	0.0391	12.5	0.00373	9.1	0.57	0.0789	10.3	24	2.2	38.8	4.7	1213	96	38.1	98.0	268.0	0.000	2.6E-07																			
GR11-43-29	0.0393	4.8	0.00379	3.2	0.49	0.0762	4.7	24.38	0.79	39.1	1.9	1096	68	37.6	97.8	172.1	0.000	3.4E-07																			
GR11-43-50	0.0404	11.6	0.0038	7.4	0.66	0.0729	8.6	24.5	1.8	40.1	4.6	970	110	38.9	97.5	175.1	0.000	1.4E-07																			
GR11-43-31	0.54	7.4	0.00817	4.7	0.74	0.484	5.0	52.5	2.4	438	27	4233	35	88.0	98.8	174.9	bdl	-																			
GR11-43-34a	0.0255	11.8	0.00345	10.1	0.73	0.0548	7.7	22.2	2.2	25.5	3	473	51	12.9	95.3	183.0	0.000	1.0E-08																			
GR11-43-36	0.0263	14.1	0.0038	6.3	0.63	0.0536	12.5	24.4	1.5	26.3	3.7	400	100	7.2	93.9	105.2	0.001	4.9E-06																			
GR11-43-40	0.0467	9.2	0.00376	6.6	0.56	0.0963	8.9	24.2	1.6	46.2	4.2	1530	120	47.6	98.4	196.3	0.000	5.0E-07																			
GR11-43-41b	0.167	28.1	0.00495	10.9	0.89	0.204	18.6	31.8	3.5	139	34	2600	300	77.1	98.8	200.0	0.000	5.0E-07																			
GR11-43-49	0.363	12.1	0.00697	8.3	0.77	0.357	8.4	44.8	3.7	307	33	3792	82	83.4	98.8	179.0	0.004	2.2E-05																			
GR11-43-50	0.0288	26.4	0.00403	12.2	0.89	0.0469	15.6	25.9	3.1	28.7	7.5	290	110	9.8	91.1	131.0	bdl	-																			
Bras10-5a-1 (Stromatic Granulite)																																					
Bras105a1-1	0.06039	7.7	0.00366	5.1	0.66	0.1966	5.8	23.6	1.2	59.6	4.5	1951.1	107.4	60.4	98.8	121.7	nd	-																			
Bras105a1-2	0.06793	10.3	0.00324	4.1	0.40	0.13207	9.5	20.8	0.9	66.7	6.7	2369.4	171.9	68.8	99.1	135.7	nd	-																			
Bras105a1-3	0.02716	8.9	0.00308	4.4	0.49	0.06396	7.7	20.3	0.9	27.9	2.5	740.2	171.9	27.2	97.3	148.9	nd	-																			
Bras105a1-4	0.07868	15.0	0.00258	8.7	0.58	0.22117	12.2	16.6	1.4	76.6	11.1	2989.3	211.3	78.3	99.4	120.4	nd	-																			
Bras105a1-5	0.05246	13.9	0.00316	4.3	0.31	0.1204	13.2	19.8	0.9	50.7	6.9	1962.2	256.5	60.9	99.0	126.3	nd	-																			
Bras105a1-6	0.09454	7.0	0.00462	4.5	0.64	0.14841	5.4	34.2	1.5	104.8	7	2327.7	95.6	67.4	98.5	112.9	nd	-																			
Bras105a1-7	0.09891	6.5	0.00375	3.7	0.57	0.1913	5.3	23.8	0.9	94.6	5.9	2753.4	89.8	74.8	99.1	159.7	nd	-																			
Bras105a1-8	0.10007	6.5	0.00441	4.2	0.65	0.16457	5.0	28.9	1.2	98.7	6.1	2503.2	86.7	70.7	98.8	114.9	nd	-																			
Bras105a1-9	0.02344	7.4	0.00321	3.4	0.46	0.05296	6.6	20.5	0.7	23.3	1.7	327	157.1	12.0	204.0	nd	-																				
Bras105a1-10	0.02825	8.4	0.00349	3.6	0.43	0.05871	7.6	23.2	0.8	29.2	2.4	556.3	175	20.5	95.8	174.3	nd	-																			
Bras105a1-11	0.05122	5.8	0.00388	2.8	0.48	0.09575	5.1	24.6	0.7	50.1	2.8	1542.9	99	50.9	98.4	277.1	nd	-																			
Bras105a1-12	0.01955	7.9	0.00298	4.0	0.51	0.04759	6.8	19.9	0.8	20.3	1.6	78.9	78.8	2.0	74.8	196.2	nd	-																			
Bras105a1-13	0.04667	7.8	0.00409	3.2	0.41	0.08276	7.1	26.1	0.8	46	3.3	1263.7	145.4	43.3	97.9	157.1	nd	-																			
Bras105a1-14	0.08857	26.0	0.00411	4.9	0.19	0.15629	25.5	26.4	1.3	86.2	21.7	2416	513.7	69.4	98.9	231.8	nd	-																			
Bras105a1-15	0.03043	6.7	0.00334	2.9	0.43	0.06607	6.0	21.8	0.6	30.9	2	808.7	130.8	29.4	97.3	191.7	nd	-																			
Bras105a1-16	0.07162	7.1	0.00392	3.8	0.54	0.13251	6.0	25.7	1	71.6	4.9	2131.5	108.9	64.1	98.8	195.9	nd	-																			
Bras105a1-17	0.15536	6.2	0.00459	4.8	0.77	0.24549	4.0	30.3	1.4	150.1	8.7	3156	65	79.8	99.0	161.8	nd	-																			
Bras105a1-18	0.06733	7.3	0.00351	4.0	0.55	0.13913	6.1	22.4	0.9	65.5	4.6	2216.4	109.8	63.8	99.0	217.7	nd	-																			
Bras105a1-19	0.00959	11.9	0.00309	3.8	0.32	0.02251	11.3	20.1	0.8	9.8	1.2	0.1	0	-105.1	-20000.0	231.2	nd	-																			
Bras105a1-20	0.01447	9.8	0.00334	4.7	0.48	0.03086	8.6	21.2	1	14.1	1.4	0.1	0	-50.4	-21100.0	255.1	nd	-																			
Bras10-5b-1 (Stromatic Granulite)																																					
009_Bras10-5b_4b.txt	0.02979	8.5	0.00367	3.7	0.44	0.05888	7.6	23.5	0.9	29.7	2.5	562.6	174.8	20.9	95.8	136.9	nd	-																			
010_Bras10-5b_5.txt	0.02262	9.3	0.00322	4.5	0.48	0.05095	8.1	20.6	0.9	22.6	2.1	238.5	198.4	8.8	91.4	115.1	nd	-																			
012_Bras10-5b_6.txt	0.11672	6.5	0.0043	3.0	0.46	0.19686	5.8	27.4	0.8	111.2	6.9	2800.4	98.2	75.4	99.0	182.1	nd	-																			
013_Bras10-5b_7.txt	0.02039	6.8	0.00308	3.6	0.53	0.04802	5.8	20.1	0.7	20.8	1.4	100.3	100.2																								

Sample and Analysis Number	Isotopic Ratios										Dates (Ma)										Compositions				
	207Pb/235U	±2σ %	206Pb/238U	±2σ %	Rb/Sr	638-735	207Pb/206Pb	±2σ %	207Pb/235U	±2σ	207Pb/235U	±2σ	207Pb/206Pb	±2σ	207Pb/206Pb	±2σ	% disc a	638-735	% disc b	638-735	U	Th	Th/U		
	abs	abs	abs	abs	abs	abs	abs	abs	abs	abs	abs	abs	abs	abs	abs	abs	abs	abs	abs	ppm	ppm	ppm			
024_Bras10-5b-11.txt	0.06551	11.4	0.00398	3.4	0.30	0.11937	7.6	26.2	0.9	65.9	7.3	1946.9	208.9	60.2	98.7	80.0	nd	nd	nd	80.0	nd	nd	-		
025_Bras10-5b-12.txt	0.01734	8.8	0.00327	4.4	0.50	0.03845	10.9	20.9	0.9	17.3	1.5	0.1	0	-20.8	-208000.0	120.3	nd	nd	nd	120.3	nd	-			
027_Bras10-5b-13.txt	0.02168	7.3	0.00322	3.5	0.48	0.04883	6.4	20.7	0.7	21.8	1.6	139.7	139.6	5.0	85.2	182.3	nd	nd	nd	182.3	nd	-			
028_Bras10-5b-14.txt	0.02023	6.3	0.0031	3.6	0.57	0.04734	5.2	19.8	0.7	20.1	1.3	66.3	66.2	1.5	70.1	201.9	nd	nd	nd	201.9	nd	-			
030_Bras10-5b-15.txt	0.02143	8.0	0.00332	3.6	0.45	0.04682	7.2	20.7	0.7	20.9	1.7	40	39.9	1.0	48.3	149.9	nd	nd	nd	149.9	nd	-			
031_Bras10-5b-16.txt	0.02092	10.3	0.00319	3.6	0.35	0.04757	9.7	20.6	0.7	21.1	2.1	78.2	78.1	2.4	73.7	135.6	nd	nd	nd	135.6	nd	-			
033_Bras10-5b-17.txt	0.18932	7.0	0.00534	2.8	0.40	0.25713	6.4	34.6	1	177.2	11.4	3229.2	104.7	80.5	98.9	144.9	nd	nd	nd	144.9	nd	-			
Bras10-5b-2 (Stromatic Granulite)																									
Bras105b2-1	0.0328	6.4	0.00501	4.6	0.22	0.0475	7.2	32.2	1.5	32.8	2	194	66	1.8	83.4	280.1	0.000	0.000	0.000	280.1	0.000	1.5E-06			
Bras105b2-2	0.0299	4.0	0.00463	2.2	0.25	0.0465	4.1	29.75	0.65	29.9	1.2	197	42	0.5	84.9	293.2	0.000	0.000	0.000	293.2	0.000	5.1E-07			
Bras105b2-3	0.0353	4.2	0.004265	1.9	0.37	0.0606	4.0	27.43	0.54	35.2	1.3	605	56	22.1	95.5	232.7	0.000	0.000	0.000	232.7	0.000	5.2E-07			
Bras105b2-4	0.029	5.5	0.004281	2.2	-0.03	0.0475	5.9	27.54	0.6	29	1.6	438	68	5.0	93.7	138.9	0.000	0.000	0.000	138.9	0.000	2.2E-07			
Bras105b2-5	0.0221	10.0	0.00608	3.5	0.78	0.256	7.8	39	1.3	200	18	3130	140	80.5	98.8	140.6	0.001	0.001	0.001	140.6	0.001	4.0E-06			
Bras105b2-6	0.0254	5.9	0.003869	2.5	0.07	0.0489	7.0	24.89	0.62	25.4	1.4	331	68	2.0	92.5	181.8	0.000	0.000	0.000	181.8	0.000	6.1E-07			
Bras105b2-7	0.0252	7.5	0.0038	3.2	-0.22	0.0469	8.3	24.46	0.77	25.2	1.9	265	57	2.9	90.8	222.3	0.000	0.000	0.000	222.3	0.000	2.2E-06			
Bras105b2-8	0.089	11.2	0.00452	3.5	0.70	0.138	8.7	29.1	1	86.1	9.3	2120	150	66.2	96.6	191.4	0.004	0.004	0.004	191.4	0.004	2.0E-05			
Bras105b2-9	0.0254	4.3	0.003852	2.4	0.20	0.0478	4.8	24.78	0.59	25.5	1.1	195	43	2.8	87.3	150.7	0.001	0.001	0.001	150.7	0.001	5.8E-06			
Bras105b2-10	0.0228	4.8	0.003543	2.3	0.16	0.0461	5.2	22.8	0.52	22.8	1.1	222	53	0.0	89.7	195.2	0.001	0.001	0.001	195.2	0.001	3.8E-06			
Bras105b2-11	0.0221	5.9	0.003306	2.6	0.12	0.0498	6.0	21.27	0.55	22.1	1.3	354	60	3.8	94.0	211.8	0.000	0.000	0.000	211.8	0.000	1.7E-06			
Bras105b2-12	0.0485	10.1	0.00382	3.1	0.44	0.0922	9.0	24.56	0.76	47.9	4.7	1450	160	48.7	98.3	189.6	0.001	0.001	0.001	189.6	0.001	4.7E-06			
Bras105b2-13	0.022	5.9	0.003469	2.5	0.08	0.0468	6.2	22.32	0.55	22.1	1.3	219	54	-1.0	89.8	285.1	0.002	0.002	0.002	285.1	0.002	7.7E-06			
Bras105b2-14	0.02367	4.0	0.003531	2.2	0.24	0.0491	3.9	22.72	0.49	23.75	0.93	234	39	4.3	90.3	353.7	0.003	0.003	0.003	353.7	0.003	9.0E-06			
Bras105b2-15	0.0275	5.8	0.00399	3.0	0.08	0.0504	6.2	25.64	0.78	27.6	1.6	362	73	7.1	92.9	150.8	0.004	0.004	0.004	150.8	0.004	2.9E-05			
Bras105b2-16	0.0389	5.9	0.00376	2.9	0.16	0.0737	5.6	24.16	0.69	39	2.2	1048	62	38.1	97.7	171.9	0.002	0.002	0.002	171.9	0.002	1.2E-05			
Bras105b2-17	0.0223	5.8	0.00323	2.8	0.12	0.049	5.9	21.39	0.59	22.4	1.3	368	63	4.5	94.2	182.0	0.004	0.004	0.004	182.0	0.004	2.4E-05			
Bras105b2-18	0.0244	5.3	0.00339	3.8	0.31	0.0515	6.6	21.83	0.86	24.5	1.3	311	73	10.9	93.0	322.2	0.001	0.001	0.001	322.2	0.001	3.4E-06			
Bras105b2-19	0.0231	4.8	0.003522	2.4	0.11	0.0478	5.0	22.66	0.55	23.2	1.1	292	63	2.3	92.2	310.0	0.005	0.005	0.005	310.0	0.005	1.7E-05			
Bras105b2-20	0.0247	5.3	0.00375	2.7	-0.04	0.0486	5.3	24.82	0.66	24.8	1.2	274	71	2.7	91.2	332.7	0.004	0.004	0.004	332.7	0.004	1.1E-05			
Bras105b2-21	0.0342	3.8	0.00386	3.1	-0.27	0.0638	6.4	24.82	0.78	34.1	1.3	769	91	27.2	96.8	254.9	bdl	bdl	bdl	254.9	bdl	-			
Bras105b2-22	0.0267	4.1	0.00415	2.1	-0.04	0.0467	4.3	26.47	0.56	26.8	1.1	162	36	1.2	83.7	301.2	bdl	bdl	bdl	301.2	bdl	-			
Bras105b2-23	0.0377	6.1	0.00415	2.9	-0.02	0.0628	6.8	26.72	0.79	37.5	2.3	788	69	28.7	96.6	232.2	0.025	0.025	0.025	232.2	0.025	1.1E-04			
Bras105b2-24	0.0299	5.4	0.00395	3.3	0.18	0.0566	6.2	25.39	0.82	29.9	1.3	490	81	15.1	94.8	305.8	0.001	0.001	0.001	305.8	0.001	2.9E-06			
Bras105b2-25	0.0272	7.7	0.004	3.5	0.33	0.0493	7.3	25.73	0.93	27.2	2.1	202	88	5.4	87.3	505.0	bdl	bdl	bdl	505.0	bdl	-			
Bras10-6-1 (Stromatic Granulite)																									
003Br_1c.txt	0.01877	7.5	0.00291	3.8	0.51	0.04678	6.5	18.7	0.7	18.8	1.4	37.8	37.7	0.5	50.5	263.4	nd	nd	nd	263.4	nd	-			
004Br_1i.txt	0.0217	5.8	0.00314	3.3	0.57	0.05013	4.8	20.6	0.7	22.2	1.3	201.1	115.4	7.2	89.8	262.2	nd	nd	nd	262.2	nd	-			
006Br_1r.txt	0.02003	6.5	0.00313	3.7	0.57	0.04641	5.3	20.7	0.8	20.7	1.3	18.8	18.7	0.0	-10.1	261.5	nd	nd	nd	261.5	nd	-			
007Br_2c.txt	0.02615	5.7	0.00345	3.1	0.54	0.05497	4.8	22	0.7	26	1.5	411	111.1	15.4	94.6	266.1	nd	nd	nd	266.1	nd	-			
009Br_2r.txt	0.02678	5.7	0.00339	3.1	0.54	0.05729	4.8	22.4	0.7	27.6	1.6	502.9	109.3	18.8	95.5	256.3	nd	nd	nd	256.3	nd	-			
010Br_2b.txt	0.02784	4.6	0.00342	2.9	0.63	0.05903	3.6	22.1	0.6	28	1.3	568.1	80.3	21.1	96.1	281.4	nd	nd	nd	281.4	nd	-			
012Br_3a.txt	0.01981	6.5	0.00299	3.4	0.52	0.04805	5.5	19.1	0.6	19.8	1.3	101.9	101.8	3.5	81.3	245.2	nd	nd	nd	245.2	nd	-			
013Br_3b.txt	0.01915	6.6	0.00292	3.4	0.52	0.04598	5.7	20.1	0.7	19.9	1.3	0.1	0	-1.0	-20000.0	217.1	nd	nd	nd	217.1	nd	-			
015Br_4c.txt	0.021	5.0	0.00326	2.9	0.58	0.04672	4.1	21.3	0.6	21.4	1.1	35	34.9	0.5	39.1	335.2	nd	nd	nd	335.2	nd	-			
016Br_4r.txt	0.02055	5.0	0.00307	2.8	0.56	0.04855	4.1	20.4	0.6	21.4	1.1	126.4	99.5	4.7	83.9	428.7	nd	nd	nd	428.7	nd	-			
018Br_4r2.txt	0.02063	5.2	0.00322	2.7	0.52	0.04646	4.5	20.4	0.5	20.4	1.1	21.7	21.6	0.0	6.0	396.0	nd	nd	nd	396.0	nd	-			
019Br_5a.txt	0.0206	4.5	0.00318	2.6	0.58	0.04699	3.7	20.8	0.5	21.1	0.9	48.9	48.8	1.4	57.5	468.7	nd	nd	nd	468.7	nd	-			
021Br_5b.txt	0.02373	4.7	0.00341	2.6	0.55	0.05048	3.9	22	0.6	23.8	1.1	217.2	92.9	7.6	89.9	439.8	nd	nd	nd	439.8	nd	-			
022Br_5c.txt	0.02095	4.5	0.00318	2.6	0.58	0.04779	3.7	20.6	0.5	21.2	0.9	88.7	88.6	2.8	76.8	521.9	nd	nd	nd	521.9	nd	-			
024Br_6.txt	0.0292	4.5	0.00411	3.0	0.67	0.05152	3.3	23.9	0.7	26.4	1.2	264.3	77.5	9.5	91.0	442.1	nd	nd	nd	442.1	nd	-			

Sample and Analysis Number	Isotopic Ratios					Dates (Ma)					Compositions							
	207Pb/ 235U	±2σ %	206Pb/ 238U	±2σ %	Rb0 638-735	207Pb/ 206Pb	±2σ %	206Pb/ 238U	±2σ abs	207Pb/ 235U	±2σ abs	207Pb/ 206Pb	±2σ abs	% disc a 638-735	% disc b 638-735	U (ppm)	Th (ppm)	Th/U
025Br_7.txt	0.02566	3.8	0.00381	2.3	0.61	0.04884	3.0	24.6	0.6	25.8	1	140.1	72	4.7	82.4	637.3	nd	-
027Br_8.txt	0.02087	4.3	0.00323	2.4	0.56	0.04687	3.6	20.6	0.5	20.8	0.9	42.4	42.3	1.0	51.4	525.5	nd	-
028Br_9.txt	0.02161	4.5	0.00322	2.4	0.53	0.04868	3.8	20.8	0.5	21.8	1	132.5	91.9	4.6	84.3	439.0	nd	-
030Br_10.txt	0.02339	4.8	0.00338	2.7	0.56	0.0502	4.0	21.9	0.6	23.6	1.1	204.4	95.5	7.2	89.3	437.9	nd	-
031Br_11.txt	0.05437	8.7	0.00396	2.7	0.31	0.09398	8.3	26.4	0.7	35.7	4.7	1616.2	163.1	52.6	98.4	243.3	nd	-
033Br_13a.txt	0.03577	5.8	0.00392	3.6	0.64	0.06618	4.3	25.3	0.9	35.8	2	812.2	92.6	29.3	96.9	208.1	nd	-
034Br_13b.txt	0.02719	5.6	0.0033	2.3	0.40	0.05975	5.3	21.2	0.5	27.2	1.6	594.6	119.2	22.1	96.4	427.4	nd	-
037Br_14.txt	0.02235	7.1	0.00342	3.4	0.48	0.0474	6.2	22.9	0.8	23.3	1.6	69.5	69.4	1.7	67.1	257.0	nd	-
039Br_15.txt	0.02878	6.5	0.00338	2.7	0.42	0.06176	5.9	21.7	0.6	28.7	1.8	665.7	131.6	24.4	96.7	443.8	nd	-
040Br_16.txt	0.0204	5.3	0.00313	3.3	0.62	0.04728	4.2	20.1	0.7	20.4	1.1	63.5	63.4	1.5	68.3	345.9	nd	-
039004	0.03004	9.6	0.00418	4.0	0.42	0.05212	8.7	26.6	1.1	29.7	2.8	290.8	212	10.4	90.9	142.4	nd	-
043Br_18a.txt	0.02585	8.4	0.00395	3.8	0.45	0.04746	7.5	25.7	1.1	26.2	2.2	72.7	72.6	1.9	64.6	147.3	nd	-
044Br_18b.txt	0.02818	7.3	0.00385	3.7	0.51	0.05308	6.3	24.7	0.9	28.1	2	332.5	149.6	12.1	92.6	150.9	nd	-
046Br_19.txt	0.02314	4.4	0.0036	2.3	0.52	0.04661	3.7	23.1	0.5	23.2	1	29.5	29.4	0.4	21.7	541.9	nd	-
047Br_20a.txt	0.02209	6.3	0.00329	3.0	0.48	0.0487	5.5	21.3	0.6	22.3	1.4	133.3	133.2	4.5	84.0	285.4	nd	-
048Br_20b.txt	0.02232	6.1	0.00326	2.5	0.41	0.04966	5.6	24.1	0.6	25.7	1.5	179.1	136	6.2	86.5	293.5	nd	-
050Br_21.txt	0.02281	6.3	0.00348	2.4	0.38	0.04753	5.8	22.3	0.5	22.8	1.4	75.8	75.7	2.2	70.6	261.8	nd	-
051Br_22a.txt	0.02451	6.4	0.00333	2.5	0.39	0.05338	5.9	21.1	0.5	24.3	1.5	344.9	139.3	13.2	93.9	325.4	nd	-
052Br_23.txt	0.02148	5.0	0.00311	2.5	0.50	0.0501	4.3	20.5	0.5	22.1	1.1	199.8	103	7.2	89.7	353.5	nd	-
054Br_24.txt	0.02095	4.3	0.00329	2.4	0.56	0.04619	3.6	21.3	0.5	21.1	0.9	7.4	7.3	-0.9	-187.8	391.1	nd	-
055Br_25.txt	0.02158	6.4	0.00324	3.2	0.50	0.04687	3.5	21.6	0.7	21.8	1.4	42.6	42.5	0.9	49.3	272.5	nd	-
056Br_26.txt	0.02172	4.2	0.00327	2.1	0.50	0.04818	3.6	21	0.4	21.8	0.9	108.2	87.3	3.7	80.6	357.9	nd	-
058Br_27a.txt	0.02416	4.9	0.00337	2.6	0.53	0.05199	4.1	21.8	0.6	24.4	1.2	285	96.6	10.7	92.4	354.8	nd	-
059Br_27b.txt	0.02958	5.4	0.00447	2.6	0.48	0.04799	4.7	28.2	0.7	29	1.5	98.9	98.8	2.8	71.5	310.8	nd	-
061Br_28.txt	0.02324	5.3	0.00332	2.2	0.42	0.05077	4.8	21.8	0.5	23.8	1.2	230.4	114.9	8.4	90.5	431.7	nd	-
062Br_29a.txt	0.02725	6.4	0.00321	3.1	0.48	0.06156	5.6	21	0.6	27.7	1.7	638.9	124.9	24.2	96.8	236.5	nd	-
063Br_29b.txt	0.03199	5.7	0.00376	3.0	0.53	0.06171	4.8	24.1	0.7	31.9	1.8	664.1	106.3	24.5	96.4	300.9	nd	-
065Br_30.txt	0.023	6.5	0.00335	3.0	0.46	0.0498	5.8	22.3	0.7	23.9	1.5	185.6	140.9	6.7	88.0	243.9	nd	-
066Br_31.txt	0.02291	7.2	0.00338	3.0	0.42	0.04915	6.5	21.7	0.7	23	1.6	154.9	154.8	5.7	86.0	247.4	nd	-
067Br_32.txt	0.02219	6.1	0.00332	3.2	0.52	0.04848	5.2	21	0.7	21.9	1.3	122.6	122.5	4.1	82.9	215.5	nd	-
069Br_33.txt	0.04121	5.2	0.00351	3.0	0.58	0.08515	4.2	21.7	0.6	39.4	2	1319	83.7	44.9	98.4	203.4	nd	-
070Br_34.txt	0.02112	5.8	0.00335	3.2	0.55	0.04572	4.8	21.4	0.7	21	1.2	0.1	0	-1.9	-21300.0	220.3	nd	-
071Br_34b.txt	0.02225	7.4	0.00331	2.9	0.39	0.04876	6.8	21.8	0.6	22.9	1.7	136.3	136.2	4.8	84.0	239.2	nd	-
073Br_35.txt	0.02258	7.1	0.00331	3.6	0.51	0.04947	6.1	21.6	0.8	23	1.6	170.3	149.1	6.1	87.3	152.0	nd	-
074Br_35c.txt	0.02141	7.1	0.00339	3.6	0.51	0.04581	6.1	21.8	0.8	21.5	1.5	0.1	0	-1.4	-21700.0	152.0	nd	-
075Br_36.txt	0.02555	5.8	0.00367	3.0	0.52	0.0505	5.0	23.5	0.7	25.5	1.5	218.3	120	7.8	89.2	191.2	nd	-
077Br_37.txt	0.02083	8.1	0.00295	3.4	0.42	0.05121	7.3	19.6	0.7	21.6	1.7	250.2	177.3	9.3	92.2	185.8	nd	-
078Br_38a.txt	0.01842	7.3	0.00326	3.4	0.47	0.04097	6.5	20.5	0.7	18.1	1.3	0.1	0	-13.3	-20400.0	183.2	nd	-
079Br_38b.txt	0.02006	6.4	0.00318	2.6	0.41	0.04576	5.9	20.6	0.5	20.3	1.3	0.1	0	-1.5	-20500.0	246.0	nd	-
081Br_39a.txt	0.02762	6.8	0.00344	3.8	0.56	0.05823	5.6	22.8	0.9	28.5	1.9	538.6	127.5	20.0	95.8	167.8	nd	-
082Br_39b.txt	0.02121	8.4	0.00299	4.4	0.52	0.05146	7.1	19.7	0.9	21.8	1.8	261.5	171.8	9.6	92.5	160.9	nd	-
083Br_40.txt	0.02143	6.5	0.0032	3.2	0.49	0.04856	5.7	20.3	0.6	21.2	1.4	126.8	126.7	4.2	84.0	224.9	nd	-
083Br_43a.txt	0.02514	8.5	0.0034	4.7	0.55	0.05363	7.1	22	1	25.4	2.1	355.4	168.9	13.4	93.8	170.1	nd	-
086Br_43b.txt	0.0214	6.3	0.00318	3.6	0.57	0.0488	5.2	20.4	0.7	21.5	1.3	138.5	126.9	5.1	85.3	233.9	nd	-
087Br_44a.txt	0.02479	6.5	0.00378	2.8	0.43	0.04756	5.9	24.1	0.7	24.7	1.6	77.6	77.5	2.4	68.9	239.6	nd	-
089Br_44b.txt	0.02135	5.5	0.00323	2.9	0.53	0.04793	4.7	20.8	0.6	21.5	1.2	96.1	96	3.3	78.4	311.8	nd	-
090Br_45.txt	0.02179	5.0	0.00337	2.8	0.56	0.0469	4.2	20.7	0.6	21.9	1.1	44.3	44.2	0.9	51.0	400.3	nd	-
091Br_46.txt	0.02809	3.9	0.00395	2.3	0.59	0.05158	3.1	25.3	0.6	28	1.1	267	72.7	9.6	90.5	891.9	nd	-
093Br_47.txt	0.02946	5.3	0.00386	2.3	0.43	0.04536	4.8	25.1	0.6	29.8	1.6	426.9	110.7	15.8	94.1	652.1	nd	-

Sample and Analysis Number	Isotopic Ratios					Dates (Ma)					Compositions							
	207Pb/ 235U	±2σ %	206Pb/ 238U	±2σ %	Rb/Sr 638-735	207Pb/ 206Pb	±2σ %	207Pb/ 238U	±2σ abs	207Pb/ 235U	±2σ abs	207Pb/ 206Pb	±2σ abs	% disc a 638-735	% disc b 638-735	U (ppm)	Th (ppm)	Th/U
094Br_48a.txt	0.0403	4.5	0.00421	2.1	0.47	0.06943	4.0	27.5	0.6	40.8	1.8	911.7	84.6	32.6	97.0	489.1	nd	-
095Br_48b.txt	0.0388	6.7	0.00418	3.0	0.45	0.06733	3.0	26.3	0.8	37.8	2.5	848.1	130	30.4	96.9	470.8	nd	-
097Br_48c.txt	0.03421	6.5	0.00384	2.7	0.42	0.06462	5.9	25	0.7	34.5	2.2	761.9	129.6	27.5	96.7	538.0	nd	-
098Br_48d.txt	0.0332	5.9	0.00403	3.0	0.51	0.06575	5.1	26	0.8	33.3	2.1	594.8	114.6	21.9	95.6	503.3	nd	-
099Br_49.txt	0.04877	4.5	0.00401	2.4	0.53	0.08821	3.8	26.1	0.6	48.9	2.1	1387.1	74.8	46.6	98.1	511.5	nd	-
101Br_50a.txt	0.03619	4.5	0.00413	2.5	0.56	0.06356	3.7	27.1	0.7	36.8	1.6	726.9	85.4	26.4	96.3	589.2	nd	-
102Br_50b.txt	0.03499	4.9	0.00417	2.3	0.47	0.06085	4.3	27.1	0.6	35.3	1.7	634	95.4	23.2	95.7	586.1	nd	-
103Br_50c.txt	0.04968	14.3	0.00377	3.0	0.21	0.09557	14.0	34	0.7	48.8	6.8	1539.4	289.3	50.8	98.4	189.3	nd	-
104Br_51.txt	0.06877	14.0	0.00461	3.2	0.23	0.10819	13.6	34	1.1	77	10.4	1769.2	271.6	55.8	98.1	190.3	nd	-
106Br_52.txt	0.03601	7.9	0.00364	2.8	0.35	0.07175	7.4	24.6	0.7	37.7	2.9	978.8	158.5	34.7	97.5	433.5	nd	-
107Br_54.txt	0.03299	6.4	0.00356	2.5	0.39	0.06721	5.9	22.7	0.6	32.7	2.1	844.2	127.8	30.6	97.3	255.2	nd	-
108Br_55a.txt	0.02678	5.3	0.00337	2.9	0.55	0.05764	4.4	21.8	0.6	27	1.4	516	99.7	19.3	95.8	375.5	nd	-
109Br_55b.txt	0.02603	6.2	0.00343	2.6	0.42	0.05505	5.6	22.1	0.6	26.2	1.6	414.3	130.3	15.6	94.7	341.3	nd	-
111Br_56.txt	0.02749	3.8	0.00393	2.4	0.63	0.05073	3.0	24.9	0.6	27.1	1	228.7	70.9	8.1	89.1	784.7	nd	-
112Br_57a.txt	0.02456	8.4	0.00357	2.4	0.29	0.0499	8.1	22.5	0.5	24.2	2	190.3	190.2	7.0	88.2	273.0	nd	-
113Br_57b.txt	0.02571	6.6	0.00365	2.8	0.42	0.05108	6.0	23.7	0.7	26	1.7	244.6	144.4	8.8	90.3	254.3	nd	-
114Br_58.txt	0.02156	5.8	0.00328	3.1	0.53	0.04768	4.9	21.5	0.7	22.1	1.3	83.7	83.6	2.7	74.3	301.9	nd	-
116Br_59a.txt	0.02304	7.2	0.00338	3.4	0.47	0.04944	6.4	20.6	0.7	21.9	1.6	168.8	156.7	5.9	87.8	197.6	nd	-
117Br_59c.txt	0.07117	12.8	0.0225	6.4	0.50	0.22941	11.1	15.1	1	72.8	9	3048	189.9	79.3	99.5	176.1	nd	-
118Br_59i.txt	0.04414	6.0	0.00338	3.4	0.57	0.09472	4.9	22.6	0.8	45.5	2.7	1522.6	95.3	50.3	98.5	189.9	nd	-
119Br_60.txt	0.02307	5.6	0.00352	2.8	0.50	0.04753	4.9	22.4	0.6	22.9	1.3	75.8	75.7	2.2	70.4	283.2	nd	-
121Br_61.txt	0.02483	6.0	0.00342	3.3	0.55	0.05266	5.0	21.7	0.7	24.5	1.5	314.2	117.9	11.4	93.1	222.1	nd	-
122Br_62a.txt	0.02597	4.5	0.00394	1.9	0.42	0.04781	4.1	25.4	0.5	26.1	1.2	89.8	89.7	2.7	71.7	558.4	nd	-
123Br_62b.txt	0.02491	4.3	0.00384	2.8	0.65	0.04705	3.3	25.3	0.7	25.5	1.1	51.7	51.6	0.8	51.1	498.5	nd	-
124Br_63.txt	0.0256	5.1	0.00357	2.7	0.53	0.05201	4.3	23.2	0.6	25.9	1.3	285.9	101.4	10.4	91.9	409.4	nd	-
126Br_64a.txt	0.02356	5.8	0.00328	3.1	0.53	0.05209	4.9	21.8	0.7	24.4	1.4	289.5	116.1	10.7	92.5	372.2	nd	-
127Br_64b.txt	0.00246	30.2	0.00329	2.8	0.09	0.00542	30.1	21.4	0.6	2.5	0.8	0.1	0	-756.0	-21300.0	376.7	nd	-
128Br_65.txt	0.02402	3.4	0.00382	2.2	0.65	0.04561	2.6	24.6	0.5	24.2	0.8	0.1	0	-1.7	-24500.0	931.5	nd	-
129Br_66.txt	0.02517	7.5	0.00356	3.1	0.41	0.05128	6.8	22.8	0.7	25.2	1.9	253.5	164.4	9.5	91.0	171.0	nd	-
131Br_67a.txt	0.02955	7.2	0.00386	2.1	0.29	0.05552	6.9	24.2	0.5	28.9	2.1	433.3	161.5	16.3	94.4	546.2	nd	-
132Br_67b.txt	0.13689	7.1	0.00406	4.2	0.59	0.24453	5.7	24.5	1	122.6	8.2	3149.7	93.5	80.0	99.2	248.5	nd	-
133Br_67c.txt	0.22182	5.3	0.00457	3.3	0.62	0.35204	4.2	29.8	1	206	9.9	3715.9	65.5	83.5	99.2	403.1	nd	-
134Br_68a.txt	0.08344	13.8	0.00467	5.7	0.41	0.12958	12.6	30.9	1.8	83.6	11.1	2092.4	240	63.0	98.5	38.9	nd	-
136Br_68b.txt	0.04739	16.4	0.00313	8.5	0.52	0.10982	14.0	20.5	1.7	47.9	7.7	1796.4	279.4	57.2	98.9	50.1	nd	-
137Br_68c.txt	0.04325	13.3	0.0041	5.3	0.40	0.0765	12.2	26	1.4	42.3	5.5	1108.2	265.2	38.5	97.7	43.5	nd	-
138Br_68d.txt	0.41768	14.2	0.00686	12.4	0.87	0.41359	7.0	44.3	5.5	356	43.6	4057.2	108.5	87.6	98.9	41.3	nd	-
146Br_69i.txt	0.04718	12.0	0.00305	6.7	0.56	0.11218	12.0	19.4	1.3	46.2	5.4	1835.1	193.1	58.0	98.9	74.1	nd	-
147Br_70a.txt	0.0229	4.5	0.00351	2.3	0.51	0.04731	4.5	23.2	0.5	23.6	1	64.8	64.7	1.7	64.2	499.1	nd	-
148Br_70b.txt	0.02723	7.0	0.00356	5.0	0.71	0.05548	7.0	23.6	1.2	28.1	1.9	431.7	113.1	16.0	94.5	266.0	nd	-
158Br_71.txt	0.09112	10.8	0.00391	6.6	0.61	0.16902	10.8	25.1	1.7	88.4	9.2	2548	149.9	71.6	99.0	32.3	nd	-
166Br_72f.txt	0.01925	11.1	0.00316	5.2	0.47	0.04418	11.1	18.7	1	17.8	2	0.1	0	-5.1	-18600.0	156.0	nd	-
167Br_72g.txt	0.02516	13.3	0.00341	6.1	0.46	0.05351	13.3	22.3	1.4	25.7	3.4	350.7	291.3	13.2	93.6	72.9	nd	-
178Br_73a.txt	0.03889	33.7	0.00385	8.7	0.26	0.07326	33.7	24.3	2.1	38.1	12.7	1021.2	855.4	36.2	97.6	40.9	nd	-
179Br_73b.txt	0.0249	9.1	0.00348	4.3	0.47	0.0519	9.1	22.6	1	25.2	2.3	281.3	194.2	10.3	92.0	103.0	nd	-
181Br_74a.txt	0.02382	13.2	0.00325	5.1	0.39	0.05316	13.2	20.5	1	23.4	3.1	335.7	303.1	12.4	93.9	80.9	nd	-
182Br_74b.txt	0.03227	13.4	0.00437	5.1	0.38	0.05355	13.4	27	1.4	30.9	4.1	332.3	307.6	12.6	92.3	88.8	nd	-
183Br_75.txt	0.02132	5.6	0.00357	3.0	0.54	0.04603	5.6	21.8	0.7	21.6	1.2	0.1	0	-0.9	-21700.0	333.4	nd	-
184Br_76.txt	0.0223	5.4	0.00353	3.2	0.59	0.04581	5.4	26.7	0.9	26.2	1.4	0.1	0	-1.9	-26600.0	446.0	nd	-
186Br_77.txt	0.02101	6.3	0.00333	3.0	0.48	0.04575	6.3	21.9	0.7	21.6	1.3	0.1	0	-1.4	-21800.0	254.0	nd	-

Sample and Analysis Number	Isotopic Ratios				Dates (Ma)				Compositions								
	207Pb/ 235U	206Pb/ 238U	Rb/Sr 638-735	±2σ%	207Pb/ 206Pb	±2σ%	207Pb/ 238U	±2σ	207Pb/ 235U	±2σ	207Pb/ 206Pb	±2σ	U (ppm)	Th (ppm)	Th/U		
188Br_78a.txt	0.02373	6.4	0.00348	3.2	0.50	0.04945	6.4	22.2	0.7	23.6	1.5	169.3	136.3	5.9	86.9	240.4	nd
188Br_79a.txt	0.02424	6.9	0.00365	3.3	0.48	0.04816	6.9	23.2	0.8	24	1.6	107.1	107	3.3	78.3	199.8	nd
191Br_79c.txt	0.02578	6.6	0.00373	3.4	0.52	0.05013	6.6	23.3	0.8	25.1	1.6	201.3	138	7.2	88.4	568.6	nd
192Br_80a.txt	0.02191	7.1	0.00333	3.4	0.48	0.04816	7.1	21.3	0.7	22	1.5	107.2	107.1	3.2	80.1	224.4	nd
193Br_80c.txt	0.02495	7.4	0.00352	4.0	0.54	0.0514	7.4	22	0.9	24.3	1.8	258.9	149.1	9.5	91.5	214.4	nd
194Br_81a.txt	0.02317	5.2	0.00361	3.0	0.58	0.04655	5.2	23.1	0.7	23.2	1.2	26	25.9	0.4	11.2	334.0	nd
196Br_82a.txt	0.02271	6.6	0.00337	3.7	0.56	0.04888	6.6	22.1	0.8	23.2	1.5	142.3	134.4	4.7	84.5	254.6	nd
197Br_82b.txt	0.02208	6.3	0.00349	3.5	0.56	0.04589	6.3	22.4	0.8	22.2	1.4	0.1	0	-0.9	-22300.0	325.3	nd
198Br_82c.txt	0.02279	5.9	0.00339	3.2	0.54	0.04876	5.9	22.1	0.7	23.2	1.4	136.6	119.4	4.7	83.8	388.5	nd
199Br_83a.txt	0.02316	5.9	0.00358	3.1	0.53	0.04691	5.9	23	0.7	23.2	1.4	44.6	44.5	0.9	48.4	463.3	nd
201Br_84a.txt	0.02439	6.0	0.00375	3.1	0.52	0.04717	6.0	24.4	0.8	24.8	1.5	57.9	57.8	1.6	57.9	302.4	nd
202Br_84b.txt	0.02128	6.6	0.00347	4.0	0.61	0.04448	6.6	22	0.9	21	1.4	0.1	0	-4.8	-21900.0	295.3	nd
203Br_86a.txt	0.02888	6.4	0.00376	3.9	0.61	0.0557	6.4	24.9	1	29.7	1.9	440.4	117.7	16.2	94.3	307.4	nd
204Br_87a.txt	0.02251	6.4	0.00354	3.8	0.59	0.04611	6.4	22.5	0.9	22.3	1.4	3.3	3.2	-0.9	-581.8	226.0	nd
206Br_87b.txt	0.02166	4.9	0.00335	2.7	0.55	0.0469	4.9	21.7	0.6	21.9	1.1	44.3	44.2	0.9	51.0	405.0	nd
207Br_87c.txt	0.02504	7.7	0.00311	4.6	0.60	0.05839	7.7	20.4	0.9	25.5	1.9	544.4	141.5	20.0	96.3	223.3	nd
208Br_87d.txt	0.02168	6.5	0.00329	3.3	0.51	0.0478	6.5	21.1	0.7	21.7	1.4	89.2	89.1	2.8	76.3	326.2	nd
209Br_87e.txt	0.02081	5.1	0.0033	2.9	0.57	0.04573	5.1	21.1	0.6	20.7	1	0.1	0	-1.9	-21000.0	453.4	nd
Brast 10-6-2 (Stromatolite)																	
003Br_1.txt	0.02589	9.2	0.00392	4.5	0.49	0.0479	8.0	25.4	1.1	26.2	2.4	94.3	94.2	3.1	73.1	81.8	nd
004Br_1b.txt	0.02281	8.1	0.00339	4.8	0.59	0.0488	6.5	22.2	1.1	23.3	1.9	138.2	138.1	4.7	83.9	78.4	nd
006Br_2a.txt	0.02474	8.0	0.00382	4.4	0.55	0.04697	6.7	24.2	1.1	24.4	1.9	47.8	47.7	0.8	49.4	104.7	nd
007Br_2b.txt	0.02768	7.4	0.00396	3.6	0.49	0.05069	6.5	25.4	0.9	27.6	2	226.7	157.6	8.0	88.8	117.2	nd
009Br_2c.txt	0.02388	7.6	0.00352	3.9	0.51	0.04714	6.5	23.1	0.9	23.4	1.8	56.4	56.3	1.3	59.0	112.9	nd
010Br_2d.txt	0.02576	7.7	0.00398	3.9	0.51	0.04694	6.6	25.8	1	26	2	46.4	46.3	0.8	44.4	112.9	nd
012Br_2e.txt	0.02834	6.9	0.00418	3.8	0.55	0.04918	5.7	26.9	1	28.4	1.9	156.4	139.2	5.3	82.8	114.5	nd
013Br_2f.txt	0.02663	6.8	0.004	3.5	0.51	0.04828	5.8	25.1	0.9	26.1	1.8	113	112.9	3.8	77.8	127.1	nd
015Br_2g.txt	0.02469	6.2	0.00388	3.6	0.58	0.04616	5.1	24.6	0.9	24.4	1.5	5.9	5.8	-0.8	-316.9	133.7	nd
016Br_2h.txt	0.03364	8.9	0.00369	3.7	0.42	0.06611	8.1	23.6	0.9	33.4	2.9	810	179.3	29.3	97.1	108.5	nd
022Br_2i.txt	0.02612	6.7	0.00367	3.5	0.52	0.05161	5.7	23.6	0.8	26.2	1.7	268.4	136.3	9.9	91.2	116.7	nd
028Br_2p.txt	0.03417	6.3	0.0041	3.4	0.54	0.06045	5.3	26.6	0.9	34.5	2.1	619.8	118.7	22.9	95.7	139.7	nd
030Br_2q.txt	0.03094	7.2	0.00393	3.4	0.47	0.0571	6.4	24.8	0.8	30.4	2.2	495.5	147.6	18.4	95.0	122.8	nd
031Br_2r.txt	0.02367	9.2	0.00362	4.3	0.47	0.04742	8.1	22.2	1	22.6	2.1	70.5	70.4	1.8	68.5	128.1	nd
033Br_2s.txt	0.12369	16.0	0.00231	14.6	0.91	0.38835	6.5	10.6	1.5	85.5	13.2	3864.6	101.6	87.6	99.7	85.7	nd
034Br_2t.txt	0.03122	23.8	0.00407	12.0	0.50	0.05564	20.6	26.2	3.1	31.2	7.3	438.2	438.1	16.0	94.0	83.7	nd
036Br_3a.txt	0.05062	17.3	0.00328	11.6	0.67	0.11193	12.9	20.8	2.4	49.4	8.4	1831	254.3	57.9	98.9	16.8	nd
037Br_3b.txt	0.02926	9.6	0.00372	3.6	0.38	0.05705	8.9	23	0.8	28.2	2.7	493.4	209.2	18.4	95.3	80.8	nd
039Br_3d.txt	0.0302	8.8	0.00452	3.7	0.42	0.04846	8.0	28.3	1	29.5	2.6	121.9	121.8	4.1	76.8	68.7	nd
040Br_3e.txt	0.02675	8.0	0.00375	3.6	0.45	0.05173	7.2	24.4	0.9	27.1	2.1	273.7	174	10.0	91.1	79.8	nd
042Br_3f.txt	0.0264	9.2	0.00398	4.2	0.46	0.04811	8.2	25.5	1.1	26.3	2.4	104.6	104.5	3.0	75.6	87.5	nd
043Br_3g.txt	0.05408	10.4	0.00383	3.6	0.35	0.1024	9.8	24.1	0.9	52.4	5.3	1668.1	193.1	54.0	98.6	99.4	nd
044Br_3h.txt	0.06342	14.8	0.00408	9.9	0.67	0.11273	11.0	26.4	2.6	62.8	9.1	1843.9	213.7	58.0	96.6	94.5	nd
046Br_3i.txt	0.02814	10.3	0.00373	4.6	0.45	0.05472	9.2	24.6	1.1	28.9	2.9	400.9	220.5	14.9	93.9	47.0	nd
047Br_3j.txt	0.0272	11.9	0.00376	5.3	0.45	0.05247	10.6	24.1	1.3	27.2	3.2	305.9	261.4	11.4	92.1	51.6	nd
048Br_4a.txt	0.02703	7.4	0.00401	4.0	0.54	0.04889	6.2	30.2	1.2	31.6	2.3	142.7	142.6	4.4	78.8	114.6	nd
050Br_4b.txt	0.02514	9.5	0.0036	4.8	0.51	0.05064	8.2	23.5	1.1	25.6	2.4	224.6	201.5	8.2	89.5	193.7	nd
055Br_4f.txt	0.02562	15.3	0.00381	7.1	0.46	0.04877	13.5	24.8	1.8	26	5.9	136.7	136.6	4.6	81.9	53.3	nd
058Br_5.txt	0.04828	6.2	0.00427	3.2	0.52	0.08201	5.3	27.5	0.9	47.9	2.9	1245.7	107.4	42.6	97.8	170.2	nd
059Br_6a.txt	0.027	8.5	0.00408	4.1	0.48	0.04799	7.5	26.2	1.1	27	2.3	98.7	98.6	3.0	73.5	63.5	nd

Sample and Analysis Number	Isotopic Ratios					Dates (Ma)					Compositions							
	207Pb/ 235U	±2σ %	206Pb/ 238U	±2σ %	Rho 638-735	207Pb/ 206Pb	±2σ %	206Pb/ 238U	±2σ	207Pb/ 235U	±2σ	207Pb/ 206Pb	±2σ	% disc a 638-735	% disc b 638-735	U (ppm)	Th (ppm)	Th/U
060Br_6b.kst	0.29301	8.1	0.00613	5.9	0.73	0.34667	5.6	38.9	2.3	25.7	18.6	3692.5	88.2	84.9	98.9	29.6	nd	-
062Br_6c.kst	0.0274	9.6	0.00381	4.3	0.45	0.05216	8.6	25.7	1.1	28.8	2.7	292.5	209.3	10.8	91.2	60.1	nd	-
063Br_6d.kst	0.02533	6.1	0.00396	3.3	0.54	0.04639	5.1	25.6	0.8	25.5	1.5	17.9	17.8	-0.4	-43.0	153.6	nd	-
064Br_6e.kst	0.03138	9.0	0.00385	4.4	0.49	0.03911	7.8	25.5	1.1	32.3	2.9	571.2	179.4	21.1	95.5	77.5	nd	-
066Br_6f.kst	0.03194	7.4	0.00365	4.8	0.65	0.06346	5.6	24.4	1.2	33.1	2.4	723.8	123.5	26.3	96.6	89.9	nd	-
067Br_6g.kst	0.10682	6.9	0.00538	4.1	0.59	0.144	5.5	34.2	1.4	102.1	6.7	2276	98	66.5	98.5	74.7	nd	-
068Br_6h.kst	0.03871	4.8	0.00457	3.1	0.65	0.06144	3.6	29.3	0.9	38.5	1.8	654.6	79.2	23.9	95.5	248.8	nd	-
070Br_7a.kst	0.02674	7.1	0.00396	3.7	0.52	0.04897	6.1	25.1	0.9	26.4	1.9	146.4	4.9	82.9	105.0	nd	nd	-
071Br_7b.kst	0.06192	13.0	0.00442	6.2	0.48	0.10161	11.4	28.7	1.8	61.6	7.8	1653.8	227.6	53.4	98.3	29.8	nd	-
072Br_7c.kst	0.0768	15.1	0.00391	7.1	0.47	0.14245	13.3	21.9	1.6	65.8	9.7	2257.3	249.8	66.7	99.0	50.5	nd	-
074Br_7d.kst	0.66327	8.5	0.01014	7.8	0.92	0.47441	3.5	67.5	5.2	531.6	35.8	4163.6	52.8	87.3	98.4	34.0	nd	-
075Br_9.kst	0.02569	9.0	0.00398	4.3	0.48	0.04681	7.9	25.6	1.1	25.7	2.3	39.6	39.5	0.4	35.4	62.4	nd	-
076Br_10a.kst	0.02266	10.0	0.00376	4.1	0.41	0.0437	9.1	23.3	1	21.9	2.2	0.1	0	-6.4	-23200.0	69.3	nd	-
078Br_10b.kst	0.02496	9.5	0.00379	4.9	0.52	0.04776	8.1	24.8	1.2	25.5	2.4	87.5	87.4	2.7	71.7	66.8	nd	-
079Br_10c.kst	0.05472	12.8	0.0036	6.1	0.48	0.11024	11.2	23.2	1.4	54.1	6.8	1803.5	219	57.1	98.7	72.5	nd	-
080Br_10d.kst	0.02809	9.9	0.00409	5.0	0.51	0.04982	8.5	26.2	1.3	28	2.7	186.5	186.4	6.4	86.0	73.8	nd	-
082Br_10e.kst	0.03067	9.1	0.0042	5.0	0.55	0.05297	7.6	27.3	1.4	30.9	2.8	327.7	182.4	11.7	91.7	71.4	nd	-
083Br_10f.kst	0.02292	15.0	0.00357	6.1	0.41	0.04657	13.7	24.1	1.5	24.1	3.6	27.1	27	0.0	11.1	78.5	nd	-
084Br_10g.kst	0.02666	12.7	0.00373	5.4	0.43	0.05183	11.5	23.7	1.3	26.4	3.3	277.8	277.7	10.2	91.5	62.7	nd	-
086Br_10h.kst	0.02752	8.9	0.00427	4.2	0.47	0.04675	7.9	27.4	1.2	27.5	2.4	36.3	36.2	0.4	24.5	73.9	nd	-
087Br_10i.kst	0.02685	8.5	0.00413	4.6	0.54	0.04715	7.1	26.7	1.2	27	2.3	57	56.9	1.1	53.2	75.5	nd	-
088Br_10j.kst	0.0325	8.5	0.00394	4.4	0.52	0.05982	7.3	26.2	1.2	33.6	2.8	597.1	166.6	22.0	95.6	76.5	nd	-
090Br_10k.kst	0.03977	8.7	0.00505	4.3	0.49	0.05712	7.6	33.4	1.4	40.7	3.5	496.2	176.8	17.9	93.3	79.3	nd	-
091Br_10l.kst	0.021	15.1	0.00326	7.3	0.48	0.04671	13.2	21.2	1.5	21.3	3.2	34.4	34.3	0.5	38.4	86.9	nd	-
092Br_10m.kst	0.02852	8.2	0.00407	4.3	0.52	0.05083	7.0	26.1	1.1	28.5	2.3	233.3	170.1	8.4	88.8	98.6	nd	-
094Br_10n.kst	0.02828	8.2	0.00401	4.5	0.55	0.05115	6.9	25.2	1.1	27.7	2.2	247.7	167.2	9.0	89.8	86.0	nd	-
095Br_10o.kst	0.03161	8.7	0.00441	4.7	0.54	0.05198	7.3	28.4	1.3	31.7	2.7	284.7	176.2	10.4	90.0	83.0	nd	-
096Br_10p.kst	0.13077	8.8	0.00549	4.0	0.45	0.17276	7.8	35.4	1.4	125.2	10.4	2584.6	136.5	71.7	98.6	93.7	nd	-

Notes:

Rho = correlation coefficient between 206Pb/238U and 207Pb/235U ratios

a % discordance = 100 * (206Pb/238U date) / (207Pb/235U date)

b % discordance = 100 * (100 * (206Pb/238U date) / (207Pb/206Pb date))

bdl = below detection limit

nd = not determined

Sample and Analysis Number	Texture	Ti	V	Zr	Zr + 2SE	Zr-in-RT	2SE	Nb	Mo	Sn	Sb	HF	Ta	W	200Pb	208Pb	Th	U	2SE														
																				(ppm)													
BR11-33-37d	Peak	3.0E+07	2.8E+06	6770	530	65.4	90	538	48	656	23	4.4	1.4	nd	-	1.2	6.4	1.1	120.8	4.8	10.2	5.3	6.50	2.60	3.3	0.8	3.8	0.74					
BR11-33-37d	Gr	4.1E+07	2.6E+06	881	33	74.1	3.9	5.3	549	29	1567	60	0.68	0.28	nd	-	1.188	0.082	3.49	0.47	94.5	2.9	394	24	0.4	0.2	0.22	0.14	4.2	1.3	14.3	1.3	
BR11-33-37d	Gr	3.4E+07	2.6E+06	1275	37	98.1	2.5	2.5	566	14	1452	48	2.07	0.28	nd	-	1.22	0.16	0.76	0.3	93.7	1.4	596	17	0.3	0.1	0.05	0.01	0.0	0.0	11.2	0.46	
BR11-33-37d	Gr	3.5E+07	2.6E+06	116	16	10.6	0.6	0.6	57	1	10	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02	0.02
BR11-33-37b	Gr	3.1E+07	2.3E+06	741	10	99.7	1.7	1.7	957	10	1477	35	0.56	0.11	nd	-	1.99	0.21	4.15	0.32	2.4	355	14	2.7	0.3	0.22	0.07	0.6	0.3	106.6	2.7		
BR11-33-37a	Gr	2.9E+07	4.1E+06	1120	24	99.9	2.4	2.4	568	14	1395	48	2.54	0.49	nd	-	1.37	0.21	4.28	0.31	82.6	2.1	514	15	0.8	0.18	0.11	0.0	0.0	22.23	0.75		
BR11-33-37c	Gr	3.3E+07	4.0E+06	995	29	103.7	2.8	2.7	570	15	1857	39	2.59	0.33	nd	-	2.14	0.17	4.16	0.39	93.5	2.3	392	13	3.9	0.3	0.02	0.02	0.4	0.1	181.5	4.4	
BR11-33-37c	Gr	3.1E+07	3.5E+06	894	21	107.9	2.2	2.0	579	12	1853	11	0.92	0.13	nd	-	2.39	0.21	5.24	0.31	98.5	2.2	447	12	2.6	0.1	0.03	0.02	0.0	0.0	130.7	4.1	
BR11-33-37c	Gr	3.2E+07	3.5E+06	894	21	107.9	2.2	2.0	579	12	1853	11	0.92	0.13	nd	-	2.39	0.21	5.24	0.31	98.5	2.2	447	12	2.6	0.1	0.03	0.02	0.0	0.0	130.7	4.1	
BR11-33-37c	Gr	4.2E+07	2.4E+06	1191	72	169	14	8.3	603	50	1386	45	2.24	0.16	nd	-	3.113	0.092	6.71	0.71	78.7	3.4	531.4	3.5	1.4	0.4	0.95	0.01	0.9	0.4	19.41	0.66	
BR11-33-34d	Gr	3.1E+07	2.2E+06	1112	20	218.1	5.8	2.7	622	17	1481	46	2.77	0.31	nd	-	0.559	0.051	6.21	0.39	88.6	3.4	555	2.0	2.0	0.2	0.02	0.02	0.0	0.0	96.7	3.6	
BR11-33-34b	Gr	3.4E+07	3.9E+06	800	21	226.6	5.4	2.4	625	15	1464	43	2.05	0.31	nd	-	0.846	0.088	5.87	0.35	81.8	2.7	565	1.9	1.8	0.2	0.05	0.03	0.0	0.0	82.1	2.8	
BR11-33-34b	Gr	3.4E+07	3.9E+06	800	21	226.6	5.4	2.4	625	15	1464	43	2.05	0.31	nd	-	0.846	0.088	5.87	0.35	81.8	2.7	565	1.9	1.8	0.2	0.05	0.03	0.0	0.0	82.1	2.8	
BR11-33-34b	Gr	4.2E+07	2.5E+06	2230	150	778	31	4.0	729	29	1850	66	20.5	3.2	nd	-	0.083	0.001	25.6	2.4	84.7	3.1	503	8.0	1.7	0.52	0.24	0.0	0.0	41.0	4.4		
BR11-33-26b	Rehtal	2.4E+07	1.5E+06	1359	29	95.1	3.3	3.5	746	26	2346	55	26.4	1.1	nd	-	0.164	0.076	31.3	1.2	137.9	5.7	1712	49	9.3	0.7	0.45	0.14	0.0	0.0	382.6	9.9	
BR11-33-24a	Rehtal	3.6E+07	3.3E+06	1889	37	95.1	1.8	1.9	746	14	2011	37	23.7	1.3	nd	-	0.214	0.037	33.02	0.77	77.3	2.2	703	15	10.8	0.4	1.28	0.20	0.0	0.0	463	13	
BR11-33-26c	Rehtal	3.0E+07	3.1E+06	2461	47	95.3	3.4	3.6	746	27	1951	38	33.36	0.89	nd	-	0.188	0.066	35.4	1.2	96.5	2.1	1008	28	9.4	0.5	0.13	0.04	0.0	0.0	418	11	
BR11-33-26c	Rehtal	3.0E+07	3.1E+06	2461	47	95.3	3.4	3.6	746	27	1951	38	33.36	0.89	nd	-	0.188	0.066	35.4	1.2	96.5	2.1	1008	28	9.4	0.5	0.13	0.04	0.0	0.0	418	11	
BR11-33-26b	Rehtal	2.7E+07	3.8E+06	5772	79	1020	32	5.1	753	24	1801	49	32.3	1.4	nd	-	0.137	0.18	33.15	0.84	87.4	2.2	680	22	12.1	0.6	0.37	0.09	0.0	0.0	488	16	
BR11-33-24b	Rehtal	3.6E+07	3.7E+06	2055	41	1085	42	3.9	759	29	2186	62	24.3	1.2	nd	-	0.213	0.056	37	1.2	93.9	2.7	907	32	11.3	0.5	1.52	0.15	0.0	0.0	474	15	
BR11-33-24c	Rehtal	3.7E+07	3.0E+06	1596	42	1090	27	2.5	759	19	2128	41	23.78	0.69	nd	-	0.176	0.048	38.4	1.4	106.6	1.8	1053	37	11.4	0.4	0.98	0.13	0.0	0.0	485	15	
BR11-33-24c	Rehtal	3.7E+07	3.0E+06	1596	42	1090	27	2.5	759	19	2128	41	23.78	0.69	nd	-	0.176	0.048	38.4	1.4	106.6	1.8	1053	37	11.4	0.4	0.98	0.13	0.0	0.0	485	15	
BR11-33-24a	Rehtal	3.8E+07	4.5E+06	2002	57	1178	34	2.9	762	22	2190	59	24.4	1.3	nd	-	0.113	0.063	41.7	1.6	123.5	3.1	1462	44	13.6	0.4	0.91	0.15	0.0	0.0	519	14	
BR11-33-28c	Rehtal	2.8E+07	3.5E+06	1993	48	1205	31	2.6	769	20	2240	44	24.7	1.1	nd	-	0.253	0.054	44.7	1.2	129.9	3.7	1531	41	11.3	0.5	0.80	0.11	0.0	0.0	510	13	
BR11-33-28c	Rehtal	2.8E+07	3.5E+06	1993	48	1205	31	2.6	769	20	2240	44	24.7	1.1	nd	-	0.253	0.054	44.7	1.2	129.9	3.7	1531	41	11.3	0.5	0.80	0.11	0.0	0.0	510	13	
BR11-33-2	Rehtal	1.6E+07	3.1E+06	938	23	1277	67	5.2	775	41	4810	140	154.5	6	nd	-	0.54	0.13	35.7	2.4	84.8	3.5	1358	0.74	0.6	0.10	0.05	0.0	0.0	293	14		
BR11-33-27	Rehtal	3.1E+07	4.6E+06	1461	43	1588	41	2.6	798	21	2230	70	18.6	0.59	nd	-	0.178	0.055	71.9	3.4	161.1	7	1202	54	11.8	0.5	1.50	0.25	0.0	0.0	501	20	
BR11-33-25f	Rehtal	3.1E+07	4.1E+06	2693	58	1857	65	3.5	814	29	2648	67	30.4	1.6	nd	-	0.157	0.092	66	3	161.4	6.4	2130	100	11.3	0.5	0.21	0.04	0.0	0.0	526	14	
BR11-33-25c	Rehtal	3.0E+07	3.6E+06	2652	63	1859	49	2.6	815	21	2648	65	31	1.6	nd	-	0.157	0.092	66	3	161.4	6.4	2130	100	11.3	0.5	0.21	0.04	0.0	0.0	526	14	
BR11-33-29b	Rehtal	2.6E+07	3.6E+06	914	25	2668	78	3.0	851	26	2901	81	39	1.7	nd	-	0.164	0.097	94.8	3.1	115.6	4.1	2840	110	8.7	0.4	0.75	0.10	0.0	0.0	390	12	
BR11-33-29b	Rehtal	2.6E+07	3.6E+06	914	25	2668	78	3.0	851	26	2901	81	39	1.7	nd	-	0.164	0.097	94.8	3.1	115.6	4.1	2840	110	8.7	0.4	0.75	0.10	0.0	0.0	390	12	
BR11-33-33c	Peak	2.9E+07	2.7E+06	1216	23	4270	120	2.8	912	17	8770	170	30.6	6.2	nd	-	0.046	0.036	127.9	3	152.1	52	524	14	1.8	0.2	0.02	0.01	0.0	0.0	89.6	2.4	
BR11-33-33c	Peak	2.9E+07	2.7E+06	1216	23	4270	120	2.8	912	17	8770	170	30.6	6.2	nd	-	0.046	0.036	127.9	3	152.1	52	524	14	1.8	0.2	0.02	0.01	0.0	0.0	89.6	2.4	
BR11-33-32c	Peak	3.7E+07	3.2E+06	3193	44	4900	110	2.6	914	23	8800	230	11.3	1.4	nd	-	0.079	0.045	30.9	4.1	1499	35	397	25	1.9	0.1	0.00	0.01	0.0	0.0	94.7	1.5	
BR11-33-21c	Peak	3.1E+07	4.0E+06	765	20	4412	98	2.2	918	20	670	53	228.4	6.7	nd	-	0.396	0.044	96.6	2.2	34.1	1.1	3	0.24	7.7	0.3	0.40	0.0	0.0	342	10		
BR11-33-21c	Peak	3.1E+07	4.0E+06	765	20	4412	98	2.2	918	20	670	53	228.4	6.7	nd	-	0.396	0.044	96.6	2.2	34.1	1.1	3	0.24	7.7	0.3	0.40	0.0	0.0	342	10		
BR11-33-39c	Peak	1.8E+07	1.3E+06	2620	97	4720	160	3.4	924	31	2172	70	31.4	1.6	nd	-	0.235	0.071	126.9	6.5	107.6	7	235	11	3.1	0.5	0.05	0.03	0.0	0.0	109.7	6.2	
BR11-33-39c	Peak	1.8E+07	1.3E+06	2620	97	4720	160	3.4	924	31</																							

Table 2.4. Zr-in-rutile temperature and U-Pb dates for selected analyses.

Sample and Analysis Number	Texture	Zr-in-rutile T (°C)	206Pb/238U Date (Ma)	± 2SE (Ma)
Retrogressed Granulite (GR11-1A)				
GR11-1A-17.FIN2	Rextal	681	23.5	1.2
GR11-1A-22.FIN2	Rextal	683	24.4	1.9
GR11-1A-57.FIN2	Rextal	699	25.5	1.3
GR11-1A-50.FIN2	Rextal	700	25.8	3.1
GR11-1A-24.FIN2	Rextal	715	27.3	1.7
GR11-1A-59a.FIN2	Rextal	725	27.1	0.8
GR11-1A-41.FIN2	Rextal	731	24.8	2.5
GR11-1A-8.FIN2	Rextal	738	28.5	0.9
GR11-1A-37a.FIN2	Rextal	740	25.6	0.9
GR11-1A-23b.FIN2	Rextal	740	23.0	1.0
GR11-1A-59b.FIN2	Rextal	740	25.7	2.6
GR11-1A-18.FIN2	Rextal	742	27.3	1.0
GR11-1A-59c.FIN2	Rextal	750	25.8	2.0
GR11-1A-1.FIN2	Rextal	751	24.3	0.7
GR11-1A-25.FIN2	Rextal	760	25.2	1.3
GR11-1A-23a.FIN2	Rextal	771	26.8	1.0
GR11-1A-40.FIN2	Rextal	775	27.8	1.3
GR11-1A-52a.FIN2	Rextal	786	25.8	1.6
GR11-1A-43a.FIN2	Rextal	802	26.1	1.6
GR11-1A-81.FIN2	Rextal	805	19.8	0.9
GR11-1A-27.FIN2	Rextal	807	27.1	1.1
GR11-1A-32.FIN2	Rextal	814	26.2	1.9
GR11-1A-28.FIN2	Rextal	814	18.8	1.5
GR11-1A-79b.FIN2	Rextal	814	26.2	1.2
GR11-1A-35.FIN2	Rextal	834	27.2	1.4
GR11-1A-2.FIN2	Rextal	877	27.1	0.7
GR11-1A-39.FIN2	Rextal	888	26.5	2.3
Retrogressed Granulite (GR11-1C)				
GR11-1C-17b.FIN2	Grt	612	23.8	2.1
GR11-1C-11.FIN2	Grt	637	24.1	1.0
GR11-1C-15.FIN2	Grt	653	24.0	1.6
GR11-1C-24.FIN2	Grt	721	25.6	2.4
GR11-1C-9.FIN2	Grt	722	25.6	1.5
GR11-1C-25.FIN2	Grt	729	23.7	2.5
GR11-1C-23.FIN2	Grt	735	23.4	2.4
GR11-1C-3.FIN2	Grt	748	24.8	0.9
GR11-1C-17a.FIN2	Grt	754	22.4	2.6
GR11-1C-18a.FIN2	Grt	776	21.5	2.7
GR11-1C-22.FIN2	Grt	789	25.7	2.6
GR11-1C-16a.FIN2	Grt	799	24.8	1.5
GR11-1C-7.FIN2	Grt	844	23.0	2.5
GR11-1C-12.FIN2	Grt	847	25.6	1.6
GR11-1C-26.FIN2	Grt	849	25.1	1.7
GR11-1C-1.FIN2	Grt	862	23.5	3.0
GR11-1C-27.FIN2	Grt	892	26.7	1.8
Stromatic Granulite (Bras10-5b-1)				
Bras105b-6	Peak	924	27.4	0.8
Bras105b-4	Peak	928	23.5	0.9
Bras105b-9ta	Peak	951	19.6	1.0
Bras105b-11ta	Peak	954	26.2	0.9
Bras105b-17t	Peak	994	20.6	0.7

Table 2.4

Sample and Analysis		Zr-in-rutile T	206Pb/238U	± 2SE
Number	Texture	(°C)	Date (Ma)	(Ma)
Bras105b-14t	Peak	1007	19.8	0.7
Stromatic Granulite (Bras10-6-1)				
72ta	Grt	582	18.7	1.0
72tb	Grt	587	22.3	1.4
6	Rextal	735	23.9	0.7
20t	Rextal	754	24.1	0.6
14	Rextal	762	22.9	0.8
27t	Rextal	764	21.8	0.6
5d	Rextal	765	20.6	0.5
5a	Rextal	768	20.8	0.5
5b	Rextal	798	22.0	0.6
22	Rextal	847	21.1	0.5
32	Peak	898	21.0	0.7
84	Peak	902	24.4	0.8
80	Peak	904	22.0	0.9
87a	Peak	908	22.5	0.9
13a	Peak	921	25.3	0.9
86	Peak	932	24.9	1.0
35	Peak	937	21.8	0.8
2	Peak	938	22.1	0.6
59	Peak	938	20.6	0.7
44a	Peak	941	24.1	0.7
34	Peak	942	21.4	0.7
1c	Peak	942	18.7	0.7
43b	Peak	942	20.4	0.7
1r	Peak	946	20.7	0.8
82a	Peak	946	22.1	0.8
82b	Peak	946	22.4	0.8
82c	Peak	949	22.1	0.7
18b	Peak	953	24.7	0.9
65t	Peak	953	24.6	0.5
48dt	Peak	955	26.0	0.8
48a	Peak	958	27.5	0.6
4	Peak	963	21.3	0.6
29a	Peak	964	21.0	0.6
7	Peak	966	24.6	0.6
17t	Peak	967	26.6	1.1
45	Peak	968	21.7	0.6
77	Peak	969	21.9	0.7
3a	Peak	972	19.1	0.6
57b	Peak	981	23.7	0.7
34b	Peak	981	21.8	0.6
36	Peak	982	23.5	0.7
75	Peak	983	21.8	0.7
23	Peak	984	20.5	0.5
15t	Peak	990	21.7	0.6
30	Peak	997	22.3	0.7

Notes:

Textures: Grt = rutile included in garnet; Peak = rutile in peak matrix assemblage; Rextal = rutile in recrystallized matrix

Zr-in-rutile T(°C): temperature calculated from Zr concentration using the equation of Tomkins et al. (2007) at 9 kbar

CHAPTER 3

Constraining the timing of UHT metamorphism, melt crystallization, and deformation with zircon geochronology and trace element geochemistry: The Gruf Complex, Central Alps

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Abstract

The results of U-Pb geochronology and trace element geochemistry of zircon are used to constrain the timing of ultra-high temperature (UHT) metamorphism of granulites and charnockites, ductile deformation, and exhumation of the UHT rocks into the migmatitic middle crust. Titanium-in-zircon thermometry and rare earth element patterns of zircon reveal that zircon in the granulites was resorbed at UHT conditions, and featureless rims, which overgrew resorbed Permian–Cretaceous oscillatory-zoned cores, formed shortly after UHT garnet breakdown at 32.7 ± 0.7 Ma. The UHT rocks were juxtaposed against upper amphibolite facies migmatites along mylonitic shear zones between 30 and 27 Ma. Ductile deformation ended by ca. 26 Ma, but felsic magmatism continued until ca. 24 Ma, culminating in the intrusion of the Novate leucogranite.

The most plausible geodynamic scenario is that UHT metamorphism occurred as a result of advection of asthenospheric material or magmas to the base of the crust at ca. 34 Ma after slab breakoff and partial exhumation of the (ultra)-high pressure nappes in the Central Alps. The lower crustal rocks were then exhumed to the middle crust, likely aided by the buoyant, intruding magmas of the Bergell intrusion.

Introduction

Ultra-high temperature (UHT) metamorphism, which is defined as temperature in excess of 900°C within the stability field of sillimanite (Harley, 1998; Brown, 2007; Harley, 2008), requires an elevated geothermal gradient (Kelsey, 2008) or advection of an external heat source (Collins, 2002; Sizova et al., 2010; Clark et al., 2011). Therefore, by understanding the timing of UHT metamorphism geodynamic processes and their rates and temperatures can be better constrained.

With applications in studies of igneous, metamorphic, and sedimentary systems, zircon is the most widely used U-Pb geochronometer (e.g. Hanchar and Hoskin, 2003; Harley and Kelly, 2007; Schoene, 2014). However, relating zircon dates to geological processes can be challenging, especially when dealing with rocks that have experienced multiple high temperature metamorphic or magmatic events (e.g. Möller et al., 2002; Harley et al., 2007). Combining zircon dates with trace element compositions of zircon and other minerals and their textural relationships can be used to help determine the pressure-temperature (P-T) conditions of zircon growth in metamorphic rocks (e.g. Rubatto, 2002; Möller et al., 2003; Whitehouse and Platt, 2003; Kelly and Harley, 2005), and thus help unravel the geodynamic evolution of terranes.

The Gruf Complex of the Central Alps, which is characterized by UHT granulites and charnockites juxtaposed against upper amphibolite facies migmatites, is enigmatic in that it is the

only occurrence of UHT rocks within the Central Alps (Bousquet et al., 2012a). Therefore, knowing the age of UHT metamorphism is important for reconstructions of the tectonometamorphic evolution of the Central Alpine orogeny, but this has been a matter of debate. Recent works by Galli *et al.* (2011; 2012; 2013) have concluded that UHT metamorphism occurred during Permian rifting of the European margin, whereas previous studies (Liatì and Gebauer, 2003; Schmitz et al., 2009) interpreted the UHT sapphirine-forming reactions to occur during Alpine orogenesis.

In this study, we combine U-Pb zircon geochronology with Ti-in-zircon thermometry and zircon trace element geochemistry to constrain the timing of UHT metamorphism in the granulites of the Gruf Complex. Samples of different types of felsic dikes, which trend from less evolved and structurally earlier to more evolved and less deformed, were also collected for U-Pb zircon geochronology to establish the absolute timing of crystallization and deformation. We use these dike crystallization ages to bracket the timing of deformation and juxtaposition of the UHT rocks and migmatites, which lack evidence for ever experiencing UHT metamorphism.

Geologic Setting

The Central Alps (Figure 3.1) consist of several thrust-bound terranes that were amalgamated during the Alpine orogeny (e.g. Stampfli et al., 1998; Bousquet et al., 2012b). The nappes were originally part of either the pre-Alpine European margin or the Penninic units, which consist of the Piemonte-Liguria and Valais oceans and the Briançonnais microcontinent (Figure 3.1). Many of the units in the Central Alps experienced (ultra)-high pressure ((U)HP) metamorphism during the continental subduction phase of the Alpine orogeny (e.g. Becker, 1993; Gebauer, 1996; Brouwer et al., 2005; Hermann et al., 2006). After amalgamation, a Barrovian metamorphic field gradient formed across a large part of the Central Alps (the Lepontine Dome),

resulting in greenschist facies metamorphism in the north and upper amphibolite facies migmatization in the south at the Southern Steep Belt (Niggli and Niggli, 1965; Trommsdorff, 1966; Todd and Engi, 1997; Frey and Ferreiro Mählmann, 1999; Burri et al., 2005; Berger et al., 2011). Migmatization in the Southern Steep Belt continued for ca. 10 m.y. between 32 and 22 Ma (Rubatto et al., 2009).

The Gruf Complex, which lies at the eastern extent of the Central Alps, is structurally overlain by the 32–30 Ma (von Blanckenburg, 1992; Samperton et al., 2015) Bergell intrusion (Berger et al., 1996; Davidson et al., 1996). To the north, the Gruf complex is bordered by the Chiavenna ophiolite and Tambo nappe, which are Penninic units. To the west, across Valle della Mera, is the Adula nappe, which experienced UHP metamorphism (Nimis and Trommsdorff, 2001). Originally the Gruf Complex was thought to be part of the Adula nappe, but more recent studies (Ciancaleoni and Marquer, 2006; Galli et al., 2013) have concluded that the normal sense Forcula fault separates the two units (Figure 3.1), which experienced different pressure-temperature histories. The 24 ± 1 Ma (Liati et al., 2000) Novate leucogranite intruded the southwest part of the Gruf Complex.

Based on the detailed map presented by Galli et al. (2013), the Gruf Complex primarily comprises upper amphibolite facies migmatitic ortho- and paragneisses (Figure 3.2) that underwent peak metamorphism at 700–750°C and ca. 6.5–7.5 kbar (Galli et al., 2013; Oalman et al., in preparation-a). However, several charnockite units ranging from meter- to kilometer-wide layers are also present. The largest of these charnockite layers occur in the center of the Gruf Complex at the highest elevations (Figure 3.2). Scarce sapphirine-bearing granulites (Figure 3.3) occur within the charnockite bodies as meter-scale rafts and pods and finer-scale melanosomes. The granulites and charnockites underwent UHT metamorphism as garnet was breaking down

along a decompression P-T path at 900–960°C and 8–10 kbar (Guevara and Caddick, 2016; Oalman et al., in preparation-a). Mylonitic shear zones (up to 50 cm wide) separate the charnockite/granulite units from the migmatite units (Figure 3.3A).

Felsic dikes are common throughout the Gruf Complex and range from biotite granite to muscovite ± garnet ± tourmaline ± beryl-bearing pegmatites. The dikes show different degrees of deformation (Figure 3.4). The most mineralogically primitive (i.e. muscovite-free and muscovite-poor) dikes are ductily deformed or deformed within the mylonitic shear zones (Figure 3.4A) that separate the UHT rocks from the migmatites. The most mineralogically evolved (i.e. muscovite-rich) dikes are not deformed within the mylonitic shear zones (Figure 3.4B) and crosscut some of the less evolved dikes (Figure 3.4D, E) and foliations in the migmatites and UHT rocks. This observed correlation provides a relative time–composition framework of melt crystallization within the Gruf complex that will be explored using U-Pb zircon dating.

Sample Descriptions

Leucogranulite

Sample GR11-37 is a leucogranulite (Figure 3.3B) consisting mainly of quartz + K-feldspar + plagioclase leucosome. Minor mafic minerals (biotite + orthopyroxene + garnet + sapphirine + spinel) and very scarce, elongate sillimanite laths (Figure 3.5A) typically occur in thin layers (Figure 3.A1A). Biotite commonly fringes orthopyroxene, and thus likely grew during retrograde rehydration. Garnet prophyroblasts are subhedral to anhedral and commonly have embayments filled with orthopyroxene, biotite, plagioclase, sapphirine, spinel, and/or oxide minerals. These garnet-breakdown textures are interpreted to have formed during UHT decompression (Guevara and Caddick, 2016; Oalman et al., in preparation-a).

Zircon occurs mainly in the leucosome of the leucogranulite sample (Figure 3.5A, B), but some grains occur within garnet breakdown textures (Figure 3.A1A). The grains typically have rounded to square shapes and reach 300 μm length (Figure 3.5A, B).

Residual Granulite

The residual granulite sample (GR11-12b) is characterized by large garnet porphyroblasts (up to 1 cm) and abundant orthopyroxene and sapphirine porphyroblasts (Figure 3.3C). Garnet, orthopyroxene, and sapphirine are commonly separated from one another by cordierite moats. Garnet is fringed by orthopyroxene + sapphirine \pm cordierite intergrowths or symplectites of cordierite + orthopyroxene (Figure 3.5C). Elongate sillimanite porphyroblasts are commonly fringed by sapphirine \pm spinel symplectites, which are surrounded by cordierite moats in some cases. Elongate bands of subrounded quartz grains occur mostly in close proximity to the sillimanite porphyroblasts. Quartz and sapphirine that are fringing sillimanite, are in contact in some rare cases. Biotite occurs as aggregates (up to 3 mm) that are commonly in contact with the other mafic minerals. Leucosomes are present as pods or bands of quartz + K-feldspar + plagioclase.

Zircon grains range from equant (square or round) to elongate (up to 300 μm long) and occur mostly as inclusions in biotite or cordierite or in the leucosomes (Figure 3.A1B). One large, elongate grain (grain 6) emanates from the largest garnet porphyroblast into an orthopyroxene + cordierite symplectite (Figure 3.5C, A1B).

Stromatic Granulite

Stromatic granulite sample Bras10-6-2 is characterized by alternating leucosome and melanosome zones (Figure 3.3D). The thin section used in this study is from a melanosome that

contains two large, subhedral garnet porphyroblasts surrounded by orthopyroxene, biotite, and sapphirine plus leucosomes of quartz + K-feldspar + plagioclase (Figure 3.A1C). The garnets have inclusion rich cores and inclusion poor rims, indicating resorption of the cores. The center of one garnet surrounds an orthopyroxene + sapphirine intergrowth (Figure 3.A1C), which is interpreted to have formed at the expense of garnet.

Of all studied granulite samples, this thin section contains the fewest zircon grains, which are typically anhedral. Zircon occurs as inclusions in garnet (Figure 3.5D) and within the leucosome (Figure 3.A1C).

Leucosomes and Dikes

The most deformed dike/leucosome sample (GR11-55) is a biotite-granite leucosome that forms the matrix of a magmatic breccia containing clasts of altered mafic and ultramafic rock (Figure 3.4C). The overall fabric within the breccia and the orientation of the breccia itself are subparallel to the foliation in the migmatitic orthogneiss that hosts it (Figure 3.2; 4C).

Sample GR11-54 is a biotite-bearing granitic pegmatite that is folded and boudinaged within a migmatitic biotite orthogneiss (Figure 3.4D). The width of the pegmatite changes in a wave-like pattern along strike, which is parallel to the foliation in the host migmatites. We interpret this geometry to indicate that the pegmatite intruded the migmatite while the migmatite was undergoing ductile deformation and possibly still contained melt.

Sample GR11-27 is a biotite granite pod that crosscuts the foliation in the charnockite it intruded (Figure 3.4E). This relationship indicates that the charnockite was not ductily deforming when the granitic mass intruded and crystallized. Two generations of aplitic dikes crosscut the contact between the granite and the charnockite.

Sample GR11-32 is a medium-grained biotite-granite dike that contains <1 volume percent muscovite and crosscuts the foliation in the charnockite (Figure 3.4F). This dike is mineralogically similar to the largest deformed muscovite-poor dike in the mylonite zone shown in Figure 3.4A, indicating that it intruded prior to or coeval with the juxtaposition of the charnockite/granulite and migmatitic gneiss units. This dike is crosscut by muscovite-rich pegmatitic dikelets (Figure 3.4F).

Sample GR11-34 is a garnet-biotite-muscovite granitic pegmatite dike that crosscuts the foliation in ultramafic rocks included in the charnockite (Figure 3.4G). This dike is mineralogically similar to the dikelets that crosscut the biotite-granite dike sample GR11-32 (Figure 3.4F) and the dike that is not deformed within the mylonite zone (Figure 3.4B). Therefore, we interpret that this muscovite-rich dike intruded later than the muscovite-poor dike (sample GR11-32) and subsequent to the juxtaposition of the charnockite/granulite and migmatitic gneiss units.

Sample VER11-1 is a pegmatitic two mica dike that crosscuts the foliation in migmatitic orthogneisses (Figure 3.4H). The sample was collected south of the Novate intrusion, but this dike is similar in mineralogy to dikes that crosscut the migmatites within the Gruf Complex proper. Therefore, we interpret that this dike intruded at approximately the same time as the most evolved pegmatites, which crosscut the Gruf charnockites and migmatites (e.g. sample GR11-34).

Two samples of the Novate leucogranite were collected from a quarry ca. 1 km north of the town of Novate Mezzola along highway SS36 (Via Spluga). The samples (NOV11-1A and -1B) look similar in hand sample, with sample A containing a slightly higher proportion of biotite than sample B. Both samples are leucocratic and also contain muscovite, biotite, and garnet.

Analytical Methods

For the UHT granulites, zircon grains were analyzed directly within thin sections to preserve textural context (*in situ*). Color cathodoluminescence (CL) images were collected using

a microscope CL system in the Geology Department at the University of Kansas. Black and white CL images were collected using a FEI Versa 3D dual beam scanning electron microscope (SEM) coupled to a Gatan CL detector at the University of Kansas Microscopy and Analytical Imaging (MAI) Laboratory. To acquire the images, an accelerating voltage of 15kV, a beam current of 45 nA, and a working distance of 6 mm were used in low vacuum mode.

Leucosome and dike samples underwent standard mineral separation procedures to extract zircon grains. For most samples, the grains were then picked and mounted in epoxy. CL imaging was performed at the University of Kansas MAI Laboratory using a Zeiss LEO SEM with a Gatan CL detector. Core and rim domains were targeted using the LA-ICP-MS system at the University of Kansas Isotope Geochemistry Laboratory to assess both inheritance and last crystallization of the zircon grains. For two samples, grains were simply mounted on double sided tape mounts and ablated from the outside of the grain in order to target the youngest overgrowths by depth-profiling.

For the granulites, the U-Pb and trace element data were collected simultaneously by laser ablation (LA) coupled with a Thermo-Scientific Element 2 (single collector) sector field inductively coupled plasma mass spectrometer (ICP-MS) to ensure that the U-Pb isotopic and compositional data came from the same volume of zircon. This resulted in shorter counting times on U and Pb isotopes than standard U-Pb analyses, increasing the reproducibility uncertainty of U-Pb ratios from 1–2% to 3–4%. These analyses were carried out at the University of Kansas Isotope Geochemistry Laboratory. The U-Pb data for the granulite and dike samples were reduced using the VizualAge data reduction scheme (Petrus and Kamber, 2011) for Iolite (Paton et al., 2011). The software ET_Redux (McLean et al., 2016) was used to construct concordia plots and calculate ages. Trace element data from the granulite zircons were reduced using the TraceElementsIS data reduction scheme of Iolite. GJ-1 zircon (Jackson et al., 2004) was the

calibration standard used for U-Pb data reduction, and NIST612 glass (Jochum et al., 2005) was used as the calibration standard for trace element data reduction. Plešovice (Sláma et al., 2008) and GJ-1 zircon were the validation reference materials for trace element compositions, and only Plešovice zircon was used as a validation reference material for the dating. The specific analytical parameters used for the granulite and dike samples are shown in Tables 1 and 2, respectively. The Ti-in-zircon temperatures were calculated using the calibration of Watson et al. (2006).

Whole-rock major element compositions of leucosome and dike samples were obtained by x-ray fluorescence (XRF) with a Thermo-ARL automated XRF spectrometer at the Peter Hooper GeoAnalytical Lab at Washington State University. These data were used for calculating zircon saturation temperatures with the equation of Watson and Harrison (1983). The compositional data and calculated temperatures are shown in Table 3.

Zircon CL Textures

Granulites

Zircon grains in the granulites range in morphology from equant to prismatic, and show a wide variety of complex zoning patterns visible in the CL images (Figure 3.6). Zircon crystals from leucogranulite GR11-37 and stromatic granulite Bras10-6-2 (Figure 3.6A, C) have one or more bluish or CL-dark, featureless, patchy zoned, or more rarely strong to weak oscillatory-zoned interior domains and at least one yellowish, featureless overgrowth domain. Most of the grains from the residual granulite sample GR11-12B (Figure 3.6B) have at least one interior domain that exhibits strong oscillatory zoning and featureless exterior domains. The largest grains have up to 6 distinct domains based on their CL characteristics (e.g. the two largest grains from sample GR11-12B, Figure 3.6B grains 6 and 24), but only 2 or 3 domains are visible in most grains (Figure 3.6). The majority of the outermost domains have euhedral boundaries with the matrix minerals,

whereas internal domains are typically irregular, indicating some resorption. However, most of the grains from the stromatic granulite sample Bras10-6-2 have irregular outer grain boundaries (Figure 3.6C). Some grains have bright yellow curvilinear features, which crosscut domain boundaries (e.g. Figure 3.6A grain D2-3), that are interpreted as the result of recrystallization along fractures or fracture fills.

Leucosomes and Dikes

Zircon separated from the leucosome of the magmatic breccia (sample GR11-55) comprises mostly CL-dark or -bright, equant (soccer ball) grains (Figure 3.7A). However, some grains are characterized by a CL-bright or -dark, weakly to strongly oscillatory-zoned interior domain surrounded by a CL-dark, featureless domain with a very thin (<5 μm), CL-bright rim (Figure 3.7A).

The biotite leucogranite dike sample GR11-32 contains zircons dominated by oscillatory zoning, but sector-zoned and featureless domains are also observed (Figure 3.7B). Most grains exhibit two or three domains with distinct zoning patterns and CL brightness. The zoning patterns of interior domains are commonly truncated by the zoning patterns of the outer domains, and in some cases, interior domains have a mottled appearance (e.g. grain 7e).

Zircon grains from the garnet-biotite-muscovite pegmatite sample GR11-34 commonly consist of oscillatory-zoned or mottled, CL-dark interior domains with featureless, CL-bright exteriors (Figure 3.7C). Zoning patterns of the interior domains are commonly truncated by the exterior domains. Some of the featureless outer domains have irregular (anhedral) outer grain boundaries (grains 4c, 10b). Scarce prismatic, featureless or sector-zoned grains are also observed.

The biotite-muscovite pegmatite sample VER11-1 yielded mostly euhedral (prismatic), oscillatory- or sector-zoned zircon grains (Figure 3.7D). A variety of core-rim relationships are

observed, including CL-dark interiors with CL-bright exteriors and vice versa, as well as grains with homogenous CL (Figure 3.7D).

The zircons from the Novate leucogranite sample NOV11-1B almost invariably exhibit oscillatory zoning patterns (Figure 3.7E). The zoning patterns of some interior domains are truncated by the zoning patterns of exterior domains. Scarce mottled domains are also present.

U-Pb Ages

Granulites

Zircon domains from the granulites fall into three main age populations: (1) Oscillatory-zoned cores and intermediate zones, which have the largest range of dates spreading along concordia from 200 to 350 Ma (Figure 3.6, 8A-C; Table 4). (2) Oscillatory-zoned and featureless domains ranging from 75–175 Ma (Figure 3.6, 8A-C). (3) Featureless overgrowths and scarce whole grains with dates from 29–35 Ma (Figure 3.6, 8D-E).

Leucogranulite sample GR11-37 (47 analyses from 18 grains) yielded one concordant analysis at ca. 350 Ma and a tight cluster of concordant analyses between 250 and 300 Ma (Figure 3.8A). Three concordant and two discordant analyses have dates between 200 and 250 Ma (Figure 3.8A). Three analyses are concordant or slightly discordant at ca. 130 Ma, and there is one discordant analysis at 75 Ma (Figure 3.8A). The largest number of grains with young (<40 Ma) overgrowths (Figure 3.6A, 8A, D; Table 4) occur in this sample. Most of the overgrowths have dates that overlap with one another (Figure 3.8D). The weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of the analyses that are <10% discordant with $^{206}\text{Pb}/^{238}\text{U}$ dates between 31.0 and 33.5 (Figure 3.8D; Table 5) is 32.7 ± 0.7 Ma (MSWD = 0.65, n = 8). One grain (Figure 3.5B, 6A grain D3-11) has a CL-dark, featureless core with a $^{206}\text{Pb}/^{238}\text{U}$ date of 32.5 ± 1.5 Ma. The overgrowth, characterized by alternating dark and bright-yellow CL bands, has $^{206}\text{Pb}/^{238}\text{U}$ dates of 29.1 ± 0.6 and 29.1 ± 0.9 Ma

with a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 29.1 ± 0.5 (MSWD = 0.001, n = 2). One bright yellow CL rim has an intermediate $^{206}\text{Pb}/^{238}\text{U}$ date of 30.9 ± 1.1 Ma but is more than 10% discordant and not used for further interpretations.

From oldest to youngest, the results from residual granulite sample GR11-12B (48 analyses from 16 grains) include analyses that spread along concordia from 220–330 Ma (Figure 3.8B). Three analyses are concordant at 150 Ma, and four discordant analyses define a discordia that intersects these concordant analyses (Figure 3.8B). One domain is concordant at 75 Ma. Several grains with featureless (CL dark or bright) overgrowths (Figure 3.6B) occur in this sample. However, only one of these grains (Figure 3.6B grain 14, 8D) yielded Oligocene dates. The core and rim domains from this young grain gave overlapping dates (Figure 3.6B grain 14) with a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 34.8 ± 1.1 Ma (MSWD = 0.16, n = 5) (Figure 3.8D).

The zircon analyses from a thin section of a melanosome in the stromatic granulite sample Bras10-6-2 (22 analyses from 7 grains) yielded clusters of concordant analyses at 310 Ma, 290 Ma, and 255 Ma. One discordant analysis occurs at 200 Ma, and one analysis is concordant at 160 Ma, albeit with large uncertainty. This sample did not yield any grains with overgrowths younger than ca. 45 Ma (Figure 3.6C, 8C; Table 4). Based on their discordance, two analyses that are younger than 100 Ma (Figure 3.8C) probably represent mixing, due to analyzing different age zones, of older and younger domains. One grain with a 65.5 ± 2.6 Ma exterior domain (Figure 3.5D, 6C grain 2) is included in a garnet core (analysis 02-3r in Table 4; grain 2 in Figure 3.A1C). The zircon grains from this sample have the most irregular grain boundaries (Figure 3.6C), which may be interpreted as indicating late resorption.

Leucosomes and Dikes

All of the dike and leucosome samples contain inherited zircon (Figure 3.7, 9). Most samples do not contain grains with domains older than 300 Ma, however, some samples contain a few grains (up 17% of the analyzed population) as old as 1800 Ma (Figure 3.7; Table 5), with the most compositionally evolved dikes containing the most >300 Ma inherited domains. The concordia plots in Figure 3.9 only show results up to 400 Ma for clarity.

The leucosome of the magmatic breccia (sample GR11-55; 85 analyses from 29 grains) contains inherited zircon cores and scarce whole grains with concordant dates of 300–260 Ma, ca. 200 Ma, and ca. 150 Ma (Figure 3.9A). However, the vast majority of the grains (68 of 85) from this sample form a well-defined, concordant population with a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 32.4 ± 0.1 Ma (MSWD = 0.91, n = 68) (Figure 3.9I; Table 5). Most of these young grains have equant (soccer ball) shapes (Figure 3.7A). However, some CL-dark interior domains also yielded <35 Ma dates (Figure 3.7A, grain 7e).

The boudinaged biotite pegmatite dike sample GR11-54 (42 analyses from 42 grains) yielded two analyses with nearly concordant dates of ca. 1800 Ma (Table 5) and one nearly concordant date of ca. 600 Ma. Several analyses from interior and exterior domains spread along concordia from 250 to 350 Ma (Figure 3.9B). The youngest population has a lower intercept age of $29.2 +1.0/-1.1$ Ma (MSWD = 1.2, n = 22) with a common-Pb upper intercept (Figure 3.9J), excluding four nearly concordant analyses that have $^{206}\text{Pb}/^{238}\text{U}$ dates >33.5 Ma and one young discordant analysis that lie off the discordia.

Sample GR11-27 (66 analyses from 66 grains), the biotite granite pod, contains one inherited zircon with a concordant date of ca. 450 Ma (Table 5; not shown in the concordia diagram). Several concordant analyses have dates between 240 and 310 Ma, and several analyses form a discordia between ca. 250 Ma and the youngest population (Figure 3.9C). The youngest,

concordant (<10% discordance; Table 5) analyses with $^{206}\text{Pb}/^{238}\text{U}$ dates <31 Ma give a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 29.8 ± 0.4 Ma (MSWD = 0.53, n = 14) (Figure 3.9K).

Sample GR11-32 (36 analyses from 36 grains) is a biotite-bearing leucogranite dike that crosscuts the main foliation in the charnockites it intrudes. Excluding two analyses, which are nearly concordant at ca. 75 Ma and ca. 310 Ma (Figure 3.7B, 9D), most of the analyses form a discordia with a lower intercept of $27.2 +0.5/-0.6$ Ma (MSWD = 0.93, n = 31) and upper intercept of ca. 1200 Ma. This young population consists of analyses of oscillatory-zoned overgrowths and homogenous grains (Figure 3.7B). Three analyses that are >15% discordant and do not plot with the main discordia trend (Figure 3.9L; Table 5) were not included in the age calculation. We prefer the lower intercept age to the weighted mean age for this sample because of the well-defined discordia formed by the analyses.

Garnet-biotite-muscovite pegmatite sample GR11-34 (33 analyses from 32 grains) contains inherited zircon spreading along concordia from ca. 230 to 280 Ma, two analyses at ca. 180 Ma, and a few younger discordant analyses at ca. 70 Ma and 43 Ma (Figure 3.9E). Most of these older dates come from oscillatory-zoned interior domains, but some of the featureless overgrowths have dates as old as 276 ± 10 Ma (Figure 3.7C grain 4b). The youngest zircon population, which exclusively comprises analyses from featureless overgrowths and prismatic grains (Figure 3.7C), has a lower intercept age of $29.6 +0.3/-0.3$ Ma (MSWD = 1.2, n = 18) with a common-Pb upper intercept (Figure 3.9M). Based on this dike's crosscutting nature and highly evolved composition, however, we do not interpret this to be the crystallization age of the pegmatite itself (see '*Zircon saturation in the dike samples*' below).

Inherited zircon, consisting of oscillatory-zoned interior and scarce exterior domains (Figure 3.7D), is abundant in the biotite-muscovite pegmatite sample (VER11-1; 50 analyses from

36 grains). Clusters of concordant analyses occur at 600 Ma, 500 Ma, 250–300 Ma, and ca. 175–200 Ma (Table 5, Figure 3.9F). One analysis is concordant, albeit with large uncertainty at ca. 34 Ma, and there is one near-concordant grain at ca. 27 Ma. The weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of the youngest concordant population ($^{206}\text{Pb}/^{238}\text{U}$ dates between 24 and 27 Ma and <10% discordance) is 25.6 ± 0.3 Ma (MSWD = 0.54, n = 19) (Figure 3.9N). These analyses are from oscillatory-zoned, sector-zoned, and featureless whole grains and exterior domains (Figure 3.7D).

Although the two samples of the Novate leucogranite (NOV11-1A and B) look very similar in hand sample, they have distinct zircon populations (Figure 3.9G-H, O). The youngest zircon grains found in sample A (131 analyses from 60 grains) are >250 Ma (Figure 3.9G), whereas abundant Mesozoic and Cenozoic and some Proterozoic grains were found in sample B (170 analyses from 64 grains) (Figure 3.9H). Both samples yielded concordant dates between 370 and 600 Ma, and >90% of analyses from sample A and <10% of analyses from sample B have concordant dates between 250 and 300 Ma (Figure 3.9G-H). Sample B also contains zircons with concordant dates between 200 and 250 Ma, and there are Jurassic and Cretaceous populations with lower intercept ages of ca. 150 Ma and ca. 75 Ma, respectively (Figure 3.9H). One grain (Figure 3.7E) consists of a featureless, dark-CL core with a $^{206}\text{Pb}/^{238}\text{U}$ date of 33.5 ± 2.1 Ma and an oscillatory-zoned rim with a $^{206}\text{Pb}/^{238}\text{U}$ date of 27.4 ± 1.0 Ma, both of which are concordant. Of the 170 spots analyzed in sample B, 111 analyses have $^{206}\text{Pb}/^{238}\text{U}$ dates between 21.6 and 27.4 Ma (Table 5; Figure 3.9O, P). There are two separate discordia trends with lower intercepts between 26 and 22 Ma (Figure 3.9O). The youngest of these trends has a lower intercept age of $23.8 \pm 0.1/-0.1$ Ma (MSWD = 1.2, n = 93) with a common-Pb upper intercept. This age excludes the data that define the slightly older lower intercept age and two analyses, one of which is concordant, that define a discordia with a lower intercept of 21.5 Ma.

Granulite Zircon Trace Element Compositions

Zircon Rare Earth Element Patterns

Figure 3.10 shows chondrite-normalized rare earth element (REE) plots colored by $^{206}\text{Pb}/^{238}\text{U}$ date and grouped by sample. Most analyses have similar REE patterns with enrichment of heavy REE (HREE) relative to light REE (LREE), positive Ce anomalies, and negative Eu anomalies. The slopes of the HREE patterns (shown also as chondrite-normalized Yb/Dy ratios - YbN/DyN - in Figure 3.11A-C) vary with date and from sample to sample (Figure 3.10 and 11A-C). This will be discussed in more detail below.

In leucogranulite sample GR11-37, the youngest zircon domains are depleted in LREE and middle REE (MREE) compared to the older domains (Figure 3.10A), and the HREE patterns of the youngest zircon domains have steeper positive slopes than the older domains (Figure 3.10A and 11A). The normalized Yb/Dy ratios (YbN/DyN) of the young (29.1–43.5 Ma) domains range from 9.1–32.2, and domains between 45.3 and 135 Ma have YbN/DyN values between 9.3 and 15.1 (Table 4; Figure 3.11A). Domains older than 194 Ma exhibit two distinct HREE slopes (Figure 3.10A, 11A): moderately positive (YbN/Dy = 3.1–13.5) and nearly flat (YbN/DyN = 1.0–2.1) (Table 4; Figure 3.10A, 11A). All except 2 of these 13 domains with YbN/DyN = 3.1–13.5 are older than 240 Ma (Table 4; Figure 3.10A, 11A).

The younger domains in the residual granulite sample (GR11-12b) are enriched in LREE and MREE compared to most of the older domains, and the younger domains have shallower (negative in one case) HREE slopes than the older domains (Table 4; Figure 3.10B, 11B). The YbN/DyN values for the young (34.5–43.9 Ma) domains range from 0.3–4.0, whereas YbN/DyN values for >135 Ma domains range from 4.8–12.1 (Table 4; Figure 3.10B, 11B). One Cretaceous

zircon domain ($^{206}\text{Pb}/^{238}\text{U}$ date = 75.2 ± 5.9 Ma) has a YbN/DyN value of 4.9 (Figure 3.11B), which is intermediate between that of most of the older grains and the younger domains.

All domains from the stromatic granulite sample Bras10-6-2 have very small Eu anomalies ($\text{Eu}/\text{Eu}^* = 0.28\text{--}0.68$) and YbN/DyN between 5.9 and 12.9 (except for one ca. 285 Ma domain with YbN/DyN = 32.9, which is also more depleted in MREE compared to all other domains) (Table 4; Figure 3.10C, 11C). The younger zircon domains in this sample are depleted in all REE compared to most of the older domains (Table 4; Figure 3.11C).

Ti-in-zircon Thermometry

Figure 3.11D-F shows that there is no simple relationship between $^{206}\text{Pb}/^{238}\text{U}$ dates and Ti-in-zircon temperatures (TiZrcT). However, there is a general increase in TiZrcT with decreasing date from 350–150 Ma and decrease in TiZrcT with decreasing date from 150–30 Ma. Most of the analyzed domains give nominal Ti-in-zircon temperatures between 542 and 870°C, with most of those again ranging from 650–750°C (Figure 3.11D–F). Several analyses (3 from the residual granulite sample GR11-12B and 1 from the stromatic granulite sample Bras10-6-2) have Ti-in-zircon temperatures between 1040°C and 1275°C (>140 ppm Ti; Table 4), but based on the sharp increases (spikiness) of Ti intensity in the time resolved signals, these results are likely the result of ablating sub-micron Ti-rich inclusions (Chapman et al., 2016). Two of these high Ti analyses are from the tips of zircon grain G12b-14 in the residual granulite sample (Table 4; Figure 3.6D grain 14). This grain is included in high-Ti biotite, and thus inclusions or intergrowths of biotite or rutile or ilmenite needles exsolved from the biotite may have been inadvertently ablated and analyzed. The apparent Ti concentration of one analysis from the residual granulite sample GR11-12B is 0.6 ± 1.6 ppm, corresponding to a Ti-in-zircon temperature of $542 \pm 299^\circ\text{C}$ (Table 4), which

is at the detection limit of Ti with this method. Because of the extremely high uncertainty, this analysis will not be considered further.

Th-U Ratios

There is no simple correlation between Th-U ratios and the calculated dates (Figure 3.11G–I; Table 4). Overall, the Th-U ratios range from 0.001 to 1.4, and the >200 Ma domains have the widest ranges. The Th-U ratios of domains older than 75 Ma from the leucogranulite sample GR11-37 range from 0.01 to 0.68, and the youngest domains (29–36 Ma) have the lowest Th-U ratios of all the analyses, ranging from 0.001–0.009 (Table 4; Figure 3.11F). In the residual granulite sample (GR11-12B), domains older than 135 Ma have Th-U ratios between 0.1 and 1.0, and younger domains have Th/U values ranging from 0.1 to 0.31 (Table 4; Figure 3.11E). The domains from the stromatic granulite sample Bras10-6-2 have the most diverse Th-U ratios; Th-U ratios range from 0.07–1.4 and 0.03–0.4, respectively for domains older and younger than 250 Ma (Table 4; Figure 3.11F).

Correlations Between REE and Ti

Figure 3.11J shows that there is a slight negative correlation between $TiZrcT$ and YbN/DyN . In general, domains with the highest YbN/DyN , which are also the youngest domains, have the lowest $TiZrcT$. However, there is no correlation between $TiZrcT$ and YbN/DyN in the residual and stromatic granulite samples (Figure 3.11K–L).

Discussion

Validity of Ti-in-zircon Temperature Estimates

Rutile is present in all the investigated samples, but it is unlikely that rutile was always present throughout the history of these rocks. Rutile, which records UHT conditions, is abundant in the matrices and leucosomes of the leucogranulite and stromatic granulite samples (Oalmann et al., submitted). However, it is very unlikely that rutile was present prior to metamorphism of these samples. Thus, it cannot be assumed that the activity of TiO_2 (a_{TiO_2}) was unity throughout the history of the granulites, especially at the time the inherited zircon domains first crystallized from their magmas, which rarely precipitate rutile. Therefore, it could be argued that the apparent Ti-in-zircon temperature estimates from zircon domains that crystallized from melts prior to UHT metamorphism may not represent true zircon crystallization temperatures (Watson and Harrison, 2005; Watson et al., 2006). For non-rutile-saturated systems, TiZrcT are minimum estimates of zircon crystallization temperature (Watson et al., 2006). Decreasing a_{TiO_2} to 0.6, which is a minimum estimate of a_{TiO_2} in basic rocks (Ghent and Stout, 1984), would increase the temperature estimates by ca. 40–50°C in most cases (Table 4), resulting in only one >900°C (930°C) temperature estimate from a 155 Ma domain in the stromatic granulite (Table 4). However, the granulites are Mg-Al-rich metapelites (Galli et al., 2011; Oalmann et al., in preparation-a), and a_{TiO_2} in most pelitic systems is between 0.9 and 1.0 (Ghent and Stout, 1984; Watson et al., 2006), indicating that it is unlikely that a_{TiO_2} was as low as 0.6 in these aluminous samples. Therefore, we interpret the TiZrcT estimates presented here to be robust indicators of zircon crystallization temperatures within analytical uncertainty, which ranges from $\pm 14^\circ\text{C}$ to $\pm 55^\circ\text{C}$ (2SE) and is shown as vertical error bars in Figure 3.11D–F.

Conditions of zircon growth and timing of UHT metamorphism in the Granulites

Although there is strong petrologic evidence that the Gruf granulites underwent metamorphism in excess of 900°C (Galli et al., 2011; Guevara and Caddick, 2016; Oalmann et al.,

in preparation-a; Oalman et al., submitted), none of the robust Ti-in-zircon temperature estimates are higher than 850°C (Figure 3.11G-I), excluding the four analyses that are considered to be erroneous measurements due to ablation of Ti-rich microinclusions (see *Ti-in-zircon Thermometry* section above). We therefore interpret the lack of >850°C zircon domains combined with the observed irregular grain and domain boundaries (Figure 3.6) to indicate that zircon was not (re)crystallizing but resorbing during UHT metamorphism. Domains ranging from 140–300 Ma commonly exhibit irregular grain boundaries (Figure 3.6), and in some cases, <34 Ma euhedral rims overgrow older domains along anhedral interfaces (e.g. Figure 3.6A). Based on the more euhedral nature of the younger domains (e.g. Figure 3.6A), it is likely that resorption of zircon and thus UHT metamorphism occurred sometime after ca. 140 Ma but prior to ca. 33 Ma. At UHT conditions, Zr may have been sequestered in rutile, melt, and garnet.

Petrographic evidence (e.g. orthopyroxene + sapphirine symplectites replacing garnet) and results from Zr-in-rutile thermometry (Oalman et al., submitted) and thermobarometric modeling of whole-rock samples and microdomains (Galli et al., 2011; Guevara and Caddick, 2016; Oalman et al., in preparation-a) indicate that UHT metamorphism of the Gruf granulites occurred during decompression, resulting in the formation of orthopyroxene, sapphirine, and cordierite at the expense of garnet. The steep HREE patterns ($YbN/DyN \geq 15$) of the 29.1–33.5 Ma zircon domains in leucogranulite sample GR11-37 (Figure 3.10A, 11A) indicate that these domains grew after or coeval with garnet breakdown (Rubatto, 2002) and thus after or coeval with UHT metamorphism. The shallower HREE slopes in the older zircon domains from this sample (Figure 3.10A, 11A) indicate that garnet was stable prior to ca. 33.5 Ma (date of oldest concordant zircon overgrowth with $YbN/DyN > 15$) (Rubatto, 2002). Some of the 210–280 Ma zircon domains from the leucogranulite sample GR11-39 have nearly flat HREE patterns ($YbN/DyN = 1.0–2.1$),

possibly indicating that at least some garnet was present during Permo-Triassic time (Rubatto, 2002). We interpret the extremely low Th/U values (<0.01) of the <34 Ma zircon domains from this sample to indicate that monazite, which formed at near-UHT, post-peak conditions (Oalman et al., in preparation-b), grew before or concurrently with these domains, outcompeting zircon for Th.

In contrast to the leucogranulite, the youngest zircon grain from residual granulite sample GR11-12B has YbN/DyN values of ≤ 4.0 (Table 4; Figure 3.10B, 11B). We interpret these very shallow HREE patterns to indicate that this grain crystallized while garnet was stable (Rubatto, 2002). Therefore, UHT garnet breakdown must have occurred after (re)crystallization of the youngest zircon in the residual granulite sample (weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age = 34.8 ± 1.1 Ma) and prior to (re)crystallization of the 32.7 ± 0.7 Ma zircon population in the leucogranulite sample. We interpret the relatively high YbN/DyN values of the >135 Ma zircon domains in the residual granulite sample (4.8–12.1 compared to <4.0 for the younger domains; Table 4, Figure 3.10B, 11B) to indicate that this sample did not experience significant garnet growth or high-grade metamorphism during the Permian through Cretaceous zircon-producing events.

Therefore, we interpret UHT metamorphism to have occurred between 34.8 ± 1.1 Ma (age of zircon in equilibrium with garnet from residual granulite) and ca. 32.7 ± 0.7 Ma (age of zircon that grew during/after garnet breakdown). The corroded nature of Permian through Mesozoic zircon domains (Figure 3.6) also supports the UHT event to be younger than Mesozoic. The presence of a zircon grain with a 65.5 ± 2.6 Ma (albeit discordant) exterior domain included in a garnet core, which was breaking down during UHT metamorphism (Oalman et al., in preparation-a), in the stromatic granulite sample Bras10-6-2 (analysis 02-3r in Table 4; grain 2 in Figure 3.A1C) provides further evidence for a young (<65 Ma) UHT event. We interpret the ca. 29 Ma

zircon rim surrounding a ca. 32.5 Ma core (Figure 3.6B grain D3-11) in the leucogranulite sample GR11-37 to indicate that there was melt infiltration into the granulites at ca. 29 Ma, coeval with emplacement of the deformed dikes (see '*Dating deformation using crystallization ages of leucosomes and felsic dikes*' below).

Independent evidence, including preservation of complex major element zoning profiles in garnet (Oalman et al., in preparation-a) and preservation of UHT Zr-in-rutile temperatures (Oalman et al., submitted), indicates that the UHT event was short-lived (≤ 1 m.y.). In a recent review, Harley (2016) noted that zircon only forms at UHT conditions in long-lived (30–100 m.y.) UHT terranes when trapped melt interacts with the surrounding rock, leading to zircon crystallization. The lack of zircon crystallizing at UHT conditions in the Gruf Complex is therefore consistent with the short-lived UHT event inferred from the zircon ages.

Our interpretation of an Alpine UHT event is in contrast to that of Galli et al. (2012) who concluded that UHT metamorphism of the Gruf Complex occurred during late Variscan (Permian) rifting of the European margin. These authors based their conclusion in part on the presence of scarce orthopyroxene inclusions (with up to 8 wt. % Al_2O_3) in Permian zircon domains from the charnockites. However, the images of the inclusion textures provided show clearly that the minerals interpreted as inclusions are aligned along cracks in the zircon grains, which show undated recrystallization textures along their edges. The zircon domains yielding the Permian ages are distant from the cracks and have oscillatory zoning patterns, placing doubt on the interpretation as true inclusions. Additionally, orthopyroxene porphyroblasts in the granulites have up to 12 weight percent Al_2O_3 (Oalman et al., in preparation-a). This may either indicate that the plutonic protoliths of the charnockites originally crystallized orthopyroxene, but at a lower temperature than UHT conditions, which are recorded by the sapphirine- and orthopyroxene-forming garnet-

breakdown reactions in the granulites, or it indicates that the orthopyroxene "inclusions" were incorporated during a later metamorphic stage. Galli et al. (2012) also used zircon saturation temperatures of the charnockites to interpret that the presence of abundant Permian zircon indicates the UHT event could not have occurred during Alpine orogenesis. However, a short-lived UHT event would not have resulted in complete dissolution of zircon in dry melts because of kinetic factors (Harrison and Watson, 1983; Watson and Harrison, 1983). (Ultra)-high temperature rocks commonly contain zircon inherited from lower temperature magmatic and metamorphic events (see review by Taylor et al., 2016).

All zircon grains with <34 Ma domains (e.g. Figure 3.6A) occur within the leucosome-rich leucogranulite sample GR11-37. We interpret this to indicate that most of the zircon produced during the Alpine event formed while melts crystallized. The lack of young, concordant zircon domains in the residual and stromatic granulite samples, with the exception of the single ca. 35 Ma grain in the residual granulite and undateable microzircon growing on rutile in the stromatic granulites (Oalmann et al., submitted), indicate that solid-state reactions did not play a major role in producing zircon in the Gruf granulites. This has also been observed in other UHT granulite terranes (e.g. Kelsey and Hand, 2015; Harley, 2016; Taylor et al., 2016). This is exemplified by the elongate, oscillatory zoned grain within the garnet breakdown texture in the residual granulite sample that lacks overgrowths younger than ca. 200 Ma (Figure 3.5C, grain 6 in Figure 3.6E, A1B). The shape and type of zoning leads us to interpret this to be an inherited (igneous) grain, indicating that it did not participate in the UHT garnet-breakdown reaction in any way. The lack of young zircons in the residual and stromatic granulite types also leads to the interpretation that these rocks lost melt prior to 33.5 Ma.

Zircon saturation in the leucosome and dike samples

The calculated zircon saturation temperatures (Table 3, Figure 3.12) are considered maximum melt temperatures because all of the dike samples contain inherited zircon (Miller et al., 2003). Therefore, it is likely that some of these evolved felsic magmas that intruded after the peak of Alpine metamorphism never reached their zircon saturation temperatures, and thus the zircon ages do not represent the timing of dike emplacement. For example, one of the Novate leucogranite samples (NOV11-1A) lacks zircon younger than ca. 250 Ma, but it is clear from the other sample (NOV11-1B) and from previous studies (Liati et al., 2000) that the intrusion crystallized at ca. 24 Ma. The lack of new zircon growth was likely the result of the very high calculated zircon saturation temperature (ca. 850°C) for sample NOV11-1A (Figure 3.9G), above the temperature of the Novate magma, which is constrained to be $\leq 700^\circ\text{C}$ (zircon saturation temperature calculated from sample NOV11-1B; Table 3).

The age obtained for the youngest zircon domains in the garnet-biotite-muscovite pegmatite dike sample GR11-34 (29.6 ± 0.3 Ma) overlaps with the ages of deformed dike samples (GR11-27: 29.8 ± 0.4 Ma, and GR11-54: $29.2 +1.0/-1.1$ Ma), but the pegmatite clearly crosscuts the foliation of its host rocks and has a more evolved mineralogy than the deformed dikes, which lack abundant muscovite and garnet. Therefore, we argue that this garnet-biotite-muscovite pegmatite did not crystallize new zircon at the time it was emplaced. The zircon saturation temperature calculated for this sample is ca. 730°C (Figure 3.12, Table 3). However, the mineralogy, which is similar to that of the Novate granite (zircon saturation temperature = 699°C), and pegmatitic texture likely indicate that it intruded at a lower temperature and was therefore unable to crystallize new zircon. We interpret the anhedral shape of some of the young (ca. 30 Ma) exterior zircon domains (Figure 3.7C), to have formed as a result of resorption. This may indicate that these grains were in contact with melt after ca. 30 Ma, and thus provides further evidence for

post 30 Ma emplacement of this pegmatite. Therefore, this evolved pegmatite was likely emplaced at approximately the same time as the biotite-muscovite pegmatite (VER11-1) and Novate leucogranite, between 26 and 24 Ma.

The most evolved pegmatite samples (GR11-34 and VER11-1) and the Novate leucogranite samples (NOV11-1A, -1B) contain 29–33 Ma zircon grains and domains (Figure 3.7, 9; Table 4). We interpret this zircon to have been inherited from the source of the melt that formed the dikes. Therefore, the source of the latest dikes must have been felsic material (e.g. migmatites, granulites, older dikes, and/or unexposed plutons) that crystallized zircon between 29 and 33 Ma.

Dating deformation using crystallization ages of leucosomes and felsic dikes

With the exception of the garnet-biotite-muscovite pegmatite (sample GR11-34) discussed above, there is a good correlation between degree of deformation and U-Pb zircon ages of the felsic dikes (Figure 3.4, 9) and a general decrease in age towards younger, more evolved compositions. Therefore, we can use the zircon crystallization ages of the dikes to constrain the timing of deformation and juxtaposition of the lower crustal UHT granulites and charnockites against the mid crustal migmatites. The biotite-granite dike sample GR11-32, which is mineralogically similar to a dike that is deformed within the mylonite zone shear zone separating the UHT rocks from the migmatitic gneisses (Figure 3.4A), has an age of $27.2 \pm 0.5/-0.6$ Ma, indicating that amalgamation of the Gruf Complex was not complete at this time. Undeformed muscovite-rich pegmatites (samples GR11-34 and VER11-1) and the Novate leucogranite, which are mineralogically similar to pegmatites that intruded a mylonite zone after mylonitization had ceased (Figure 3.4B) and crosscut the biotite-granite dike (Figure 3.4F), were emplaced between 24 and 26 Ma. This leads to the interpretation that juxtaposition of the lower crustal UHT rocks was completed by ca. 26 Ma. The overlapping ages of the biotite pegmatite dike (GR11-54: 29.2

+1.0/-1.1 Ma), which is deformed within the migmatites, and the biotite-granite pod (GR11-27: 29.8 ± 0.4 Ma), which crosscuts the foliation in the charnockite, indicates that internal ductile deformation was still occurring within the migmatite but had ceased within the charnockite/granulite unit by ca. 29 Ma.

The older (32.4–29.2 Ma) dike and leucosome samples completely lack muscovite, and the $27.2 \pm 0.5/-0.6$ Ma dike contains very little muscovite. These intrusions are deformed within the migmatites and mylonitic shear zones (Figure 3.4). In contrast, the younger (25.6–23.8 Ma) pegmatites and Novate leucogranite samples, which crosscut the mylonitic shear zones (Figure 3.4), contain abundant magmatic muscovite, and the biotite-muscovite pegmatite (VER11-1) and one of the Novate leucogranite samples (NOV11-1A) have the highest alumina saturation indices (ASI = molar Al/Na+K+Ca = 1.10 and 1.11, respectively; Table 3) of all the leucosome and dike samples, implying a predominately crustal source for these magmas (Chappell and White, 1974). This correlation between increasing crustal melt component (or decreasing mantle melt component) and decreasing contractional deformation suggests that the amalgamated Gruf Complex was undergoing extension and exhumation beginning at ca. 26 Ma. Further evidence for exhumation at this time includes extensional deformation along the normal-sense Forcola fault (Figure 3.1), which was coeval with emplacement of the Novate leucogranite (Ciancaleoni and Marquer, 2006) at 23.8 ± 0.1 Ma.

Temperature-deformation-time Implications

The results presented and discussed here provide constraints on several aspects of the temperature-deformation-time evolution of the Gruf Complex that are shown in Figure 3.13. Prior to UHT metamorphism there was minor zircon growth at 34.8 ± 1.1 Ma, in equilibrium with garnet during prograde metamorphism in the residual granulite sample. Based on the Ti-in-zircon

temperatures of the youngest rims in the leucogranulite sample GR11-37, the granulite/charnockite until had cooled from UHT conditions to ca. 700–800°C by 32.7 ± 0.7 Ma. Therefore, the age of UHT metamorphism is bracketed to ca. 34 Ma. The age of the youngest zircon population in the leucosome of the magmatic breccia (32.4 ± 0.1 Ma), which is deformed within the migmatitic orthogneiss unit, indicates that the migmatites began to crystallize zircon at nearly the same time as the zircon rims formed in the leucogranulite sample. The most primitive magmas within immediate vicinity of the Gruf Complex (the Bergell tonalite and granodiorite) were emplaced in the middle crust at 32–30 Ma (von Blanckenburg, 1992; Samperton et al., 2015). Based on the zircon crystallization ages of deformed and crosscutting dike samples studied here, the UHT granulite/charnockite unit was exhumed into the middle crust and juxtaposed against the upper amphibolite facies migmatites between 30 and 27 Ma. Between 26 and 24 Ma, highly evolved pegmatitic dikes and the Novate leucogranite intruded the Gruf complex during extensional deformation. Based on U-Pb rutile ages presented by Oalman et al. (submitted), the Gruf Complex remained above ca. 400°C until about 19 Ma. This temperature-time scenario results in the following possible cooling rates: (a) ca. 60°C/m.y. between 34 and 30 Ma for the granulite/charnockite unit and either (b) 25°C/m.y. between 30 and 19 Ma or (c) ca. 8°C/m.y. between 30 and 24 Ma increasing to ca. 45°C/m.y. between 24 and 19 Ma for the amalgamated Gruf Complex (Figure 3.13).

Regional Geodynamic Implications

Although we cannot directly date the UHT metamorphism of the Gruf granulites because zircon did not crystallize during the short-lived UHT event, we can bracket the timing of UHT metamorphism to ca. 34 Ma (between the 34.8 ± 1.1 Ma age of zircon in equilibrium with garnet in residual granulite sample and the 32.7 ± 0.7 Ma age of post-UHT zircon in leucogranulite

sample). In order to constrain the geodynamic mechanism responsible for UHT metamorphism, we compare the ages obtained from the Gruf Complex to the ages of other metamorphic and plutonic events in the Central Alps (Figure 3.14). Several nappes within the Central Alps, including the Adula nappe, experienced UHP metamorphism between 43 ± 2 and 35.4 ± 0.5 Ma (e.g. Gebauer, 1996). These units were exhumed beginning at 34.2 ± 0.2 Ma (Hermann et al., 2006) and underwent a high temperature (Barrovian) overprint between 32 and 22 Ma (Rubatto et al., 2009). The coinciding time of UHT metamorphism of the Gruf lower crustal rocks and the switch from UHP to Barrovian metamorphism in the other units likely implies that exhumation of the UHP units and UHT metamorphism are genetically linked. A possible scenario we favor is that UHT metamorphism occurred in the lower crust after the UHP units had been exhumed to the middle crust.

Obtaining UHT conditions at ca. 8–10 kbar requires a very high geothermal gradient ($\gg 20^\circ\text{C}/\text{km}$ (Kelsey et al., 2008)) or advection of hot material to the base of the crust (Collins, 2002; Sizova et al., 2010; Clark et al., 2011). Because there is no strong evidence for an extreme geothermal gradient in the Central Alps, we conclude that heat for UHT metamorphism came from advection of hot material to the base of the crust. Many authors agree that the exhumation of Alpine (U)HP units occurred as a result of slab breakoff (e.g. Davies and von Blanckenburg, 1995; von Blanckenburg and Davies, 1995; Handy et al., 2010). Therefore, we postulate that slab breakoff (von Blanckenburg and Davies, 1995) resulted in asthenospheric upwelling, which then provided the necessary heat for UHT metamorphism of the Gruf charnockites and granulites. The lithospheric thinning model related to slab rollback (Beltrando et al., 2010) is not favored because it cannot accurately explain the timing and location of Bergell magmatism. Shortly after UHT metamorphism, the charnockites and granulites were exhumed into the middle crust. This

exhumation may have been aided by the Bergell magmas, which were intruding the middle crust at the same time.

Conclusions

The zircon grains within the Gruf Complex granulites have multiple textural and age domains, but we conclude that during brief attainment of UHT conditions zircon was being resorbed and not growing. Therefore, the UHT event cannot be directly dated by U-Pb zircon geochronology. Based on the low HREE abundances in the youngest zircon grain from the residual granulite sample, we interpret that there was minor zircon growth in equilibrium with garnet during Alpine prograde metamorphism at 34.8 ± 1.1 Ma. However, based on the lack of UHT Ti-in-zircon temperatures, and enriched HREE contents found in the youngest zircon rims in the leucogranulite, we conclude that most of the Alpine zircon growth resumed during the retrograde stages of the UHT event, concurrent with crystallization of partial melts, between 29.1 and 33.5 Ma. The location of these domains in the leucosomes of the granulites and their extremely low Th-U ratios indicate that they crystallized from partial melts and not from solid-state reactions, concurrently with monazite. In summary, we conclude that the UHT event occurred at ca. 34 Ma. Heat for UHT metamorphism possibly came from advection of asthenosphere or mantle derived melts resulting from slab breakoff (Davies and von Blanckenburg, 1995; von Blanckenburg and Davies, 1995; Handy et al., 2010) and after exhumation of the (U)HP nappes in the Central Alps.

The U-Pb zircon ages of felsic dikes analyzed in this study indicate that the UHT granulites/charnockites and upper amphibolite facies migmatites were amalgamated, and contractional ductile deformation ended between $27.2 +0.5/-0.6$ Ma and 25.6 ± 0.4 Ma. The lower crustal, UHT rocks and the Bergell magmas (von Blanckenburg, 1992; Samperton et al., 2015) were emplaced into the migmatitic middle crust at the same time, indicating a temporal and genetic

link between UHT metamorphism, large-scale plutonism, and exhumation in the Central Alps. Felsic magmatism resulting from crustal anatexis in and around the Gruf Complex continued until ca. 24 Ma. Future research on felsic dikes in the surrounding units (e.g. Adula and Chiavenna) could shed light on the timing of amalgamation of the Central Alps.

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Figure Captions

Figure 3.1: Tectonic map of the Central Alps color coded by paleogeographic affinity. Ages for the Bergell and Novate intrusions are from (von Blanckenburg, 1992) and Liati et al. (2000), respectively. Modified from Bousquet et al. (2012b).

Figure 3.2: Geologic map of the Gruf Complex and surrounding units with granulite and dike sample locations. Modified from (Galli., 2010).

Figure 3.3: Outcrop photos of the granulite types: (A) Mylonitic shear zones separating charnockite (bottom of photo) from migmatitic orthogneiss (top of photo). (B) Leucogranulite

sample GR11-37. (C) Residual granulite sample GR11-12b. (D) Stromatic granulite sample Bras10-6-2.

Figure 3.4: Photos of leucosome and dike samples: (A) Muscovite-poor biotite-granite dikes (mineralogically similar to sample GR11-32) that are deformed within a mylonitic shear zone, which separate charnockite from migmatitic orthogneiss. (B) Garnet-biotite-muscovite pegmatite (mineralogically similar to sample GR11-34) and quartzite that lack internal deformation within a mylonite zone, which separates charnockite from migmatitic orthogneiss (not shown in photo). (C) Magmatic breccia with biotite granite leucosome (sample GR11-55) and ultramafic and mafic clasts. (D) Biotite-bearing granite pegmatite dike (sample GR11-54) boudinaged within foliation of migmatitic host rocks. (E) Biotite-granite pod (sample GR11-27) crosscutting foliation in charnockite and crosscut by two small aplitic dikes. (F) Medium-grained biotite-granite dike (sample GR11-32) crosscutting foliation in charnockite and leucogranulite. (G) Garnet + muscovite + biotite granitic pegmatite dike (sample GR11-34) that crosscuts the foliation in the charnockite and entrains an ultramafic pod. (H) Muscovite + biotite granitic pegmatite dike (sample VER11-1) that crosscuts migmatitic orthogneiss.

Figure 3.5: Photomicrographs of selected zircon-bearing textures. (A) Zircon between sillimanite and leucosome in leucogranulite sample GR11-37. (B) Zircon in leucosome in leucogranulite sample GR11-37. (C) Zircon emanating from garnet into orthopyroxene + cordierite symplectite in residual granulite sample GR11-12B. (D) Zircon included in garnet in stromatic granulite sample Bras10-6-2.

Figure 3.6: CL images of representative zircon grains from each of the granulite samples showing grain/analyses numbers, location of laser ablation pits, and $^{206}\text{Pb}/^{238}\text{U}$ dates: (A) Leucogranulite sample GR11-37. (B) Residual granulite sample GR11-12b. (C) Stromatic granulite sample Bras10-6-2. All pit diameters are 15 μm . Italicized dates are >10% discordant and not used in the age calculations in Figure 3.8.

Figure 3.7: CL images of representative zircon grains from leucosome and dike samples showing grain/analysis number, location of laser ablation pits, and $^{206}\text{Pb}/^{238}\text{U}$ dates: (A) Biotite-bearing leucosome of magmatic breccia (sample GR11-55). (B) Medium-grained biotite-granite dike (sample GR11-32). (C) Garnet-biotite-muscovite granitic pegmatite dike (sample GR11-54). (D) Muscovite + biotite-bearing granitic pegmatite dike (sample VER11-1). (E) Garnet-biotite-muscovite leucogranite (sample NOV11-1B). There are no CL images of zircon from the other dike samples because these grains were analyzed on tape mounts. All pit diameters are 30 μm . Italicized dates are >10% discordant and not used in age calculations shown in Figure 3.9.

Figure 3.8: Wetherill concordia diagrams for granulite samples: (A-C) All analyses between 0 and 400 Ma for each of the samples. (D-E) Youngest populations from the leucogranulite GR11-37 and residual granulite GR11-12b samples with ages. Blue ellipses were used to calculate the ages.

Figure 3.9: Wetherill concordia diagrams for leucosome and dike samples: (A-H) All analyses between 0 and 400 Ma for all dike and leucosome samples. (I-O) Youngest populations with ages for all samples except NOV11-1A, which did not yield young zircon.

Figure 3.10: Chondrite-normalized REE plots for zircon from granulite samples color coded by $^{206}\text{Pb}/^{238}\text{U}$ date: (A) Leucogranulite (sample GR11-37). (B) Residual granulite sample GR11-12b. (C) Stromatic granulite sample Bras10-6-2. REE concentrations are normalized to the CI carbonaceous chondrite values of McDonough and Sun (1995).

Figure 3.11: Trace element plots for zircon from leucogranulite (GR11-37), residual granulite (GR11-12B), and stromatic granulite (Breas10-6-2) samples. (A–C) Chondrite normalized Yb-Dy ratio (YbN/DyN) versus $^{206}\text{Pb}/^{238}\text{U}$ date color coded by concordance. (D-F) Ti-in-zircon temperature versus $^{206}\text{Pb}/^{238}\text{U}$ date color coded by concordance. (G-I) Th/U versus $^{206}\text{Pb}/^{238}\text{U}$ date color coded by concordance. (M-O) Chondrite-normalized Yb-Dy ratio (YbN/DyN) versus Ti-in-zircon temperature color coded by $^{206}\text{Pb}/^{238}\text{U}$ date.

Figure 3.12: Zircon saturation temperature versus age plot for leucosome and dike samples.

Figure 3.13: Schematic temperature-time diagram showing the evolution of zircon, dike emplacement, deformation, and major mineral reactions in the Gruf granulites and dikes.

Figure 3.14: Age ranges of metamorphic and plutonic events in the Central Alps. The numbers refer to the following references: (1) U-Pb staurolite: ca. 28 Ma (Nagel, 2002), (2) U-Pb zircon (rims) 33–32 Ma (Liati et al., 2009), (3) Ar-Ar phengite: 33–30 Ma (Talerico, 2000), (4) U-Pb zircon: 37.1 ± 0.9 Ma (Liati et al., 2003), (5) U-Pb rutile: 30–19 Ma (Oalman et al., in preparation-b), (6) U-Pb zircon: 24.0 ± 1.2 Ma (Liati et al., 2000), which envelopes our age of 23.8 ± 0.1 Ma, (7) U-Pb zircon (tonalite): 31.88 ± 0.09 Ma, U-Pb allanite (granodiorite): 30.1 ± 0.25 Ma (von

Blanckenburg, 1992), (8) U-Pb zircon: 32–22 Ma (Rubatto et al., 2009), (9) U-Pb zircon: 34.2 ± 0.2 and 32.9 ± 0.3 Ma (Hermann et al., 2006), (10) Lu-Hf garnet: 70–36 Ma (Brouwer et al., 2005), (11) U-Pb zircon: 43 ± 2 and 35.4 ± 0.5 Ma (Gebauer, 1996).

Appendix 3.A. Thin Section Scans

Figure 3.A1: Scans of thin sections showing location of analyzed zircon grains in the granulite samples. A) Leucogranulite sample GR11-37. (B) Residual granulite sample GR11-12b. (C) Stromatic granulite sample Bras10-6-2.

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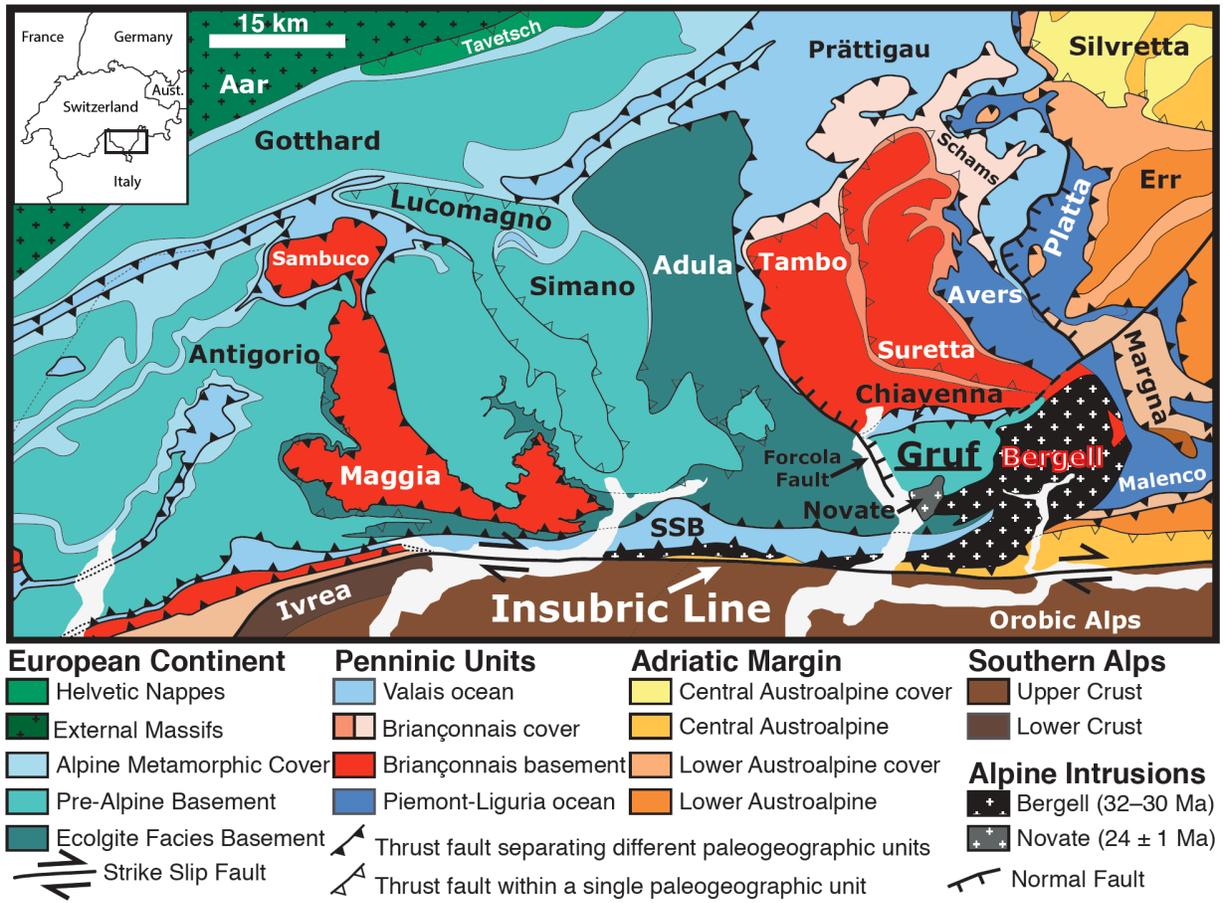


Figure 3.1

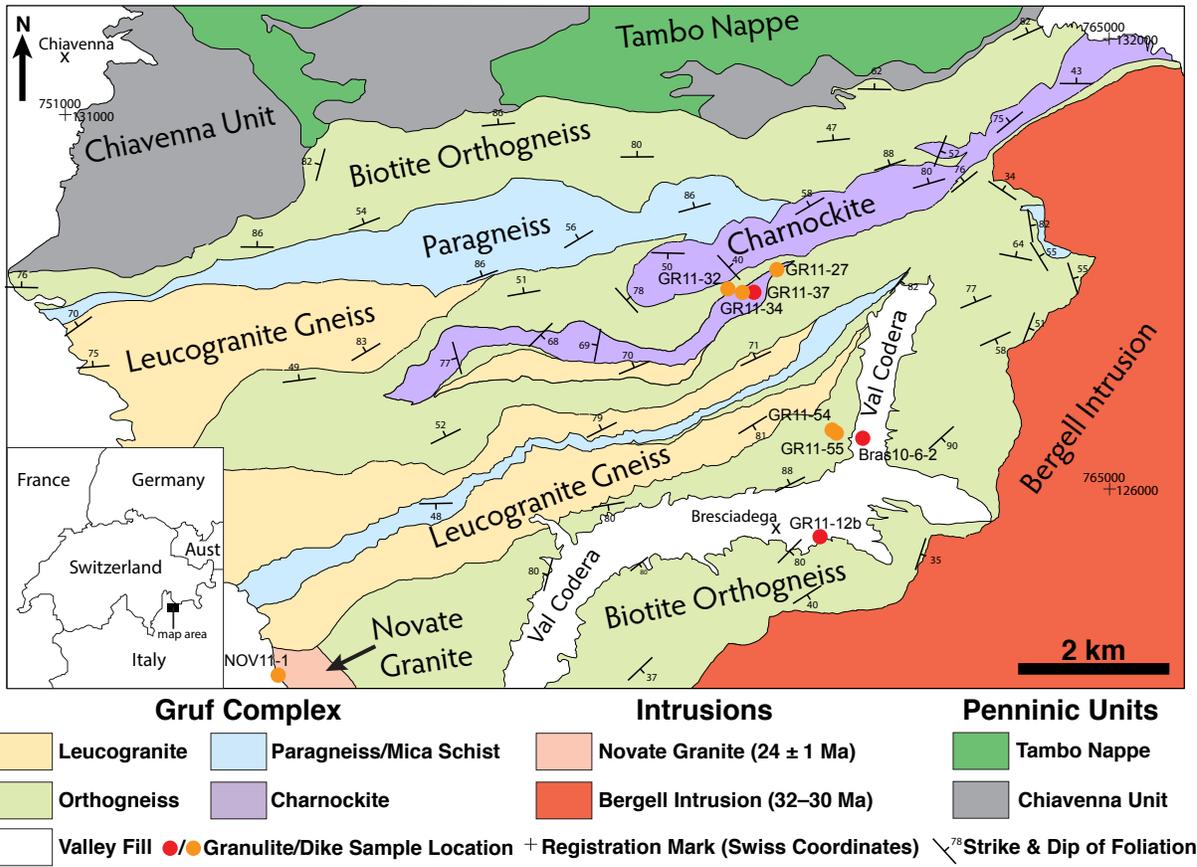


Figure 3.2

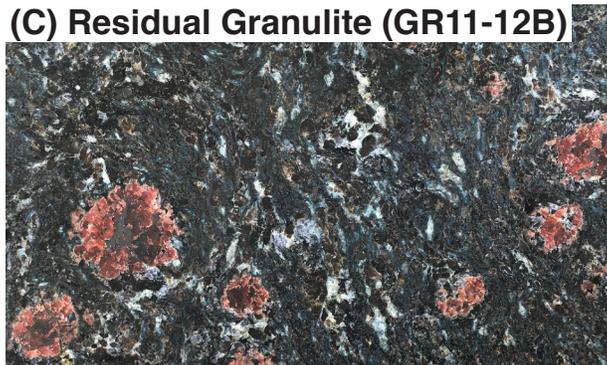
(A) Mylonitic Shear Zone



(B) Leucogranulite (GR11-37)



(C) Residual Granulite (GR11-12B)



(D) Stromatic Granulite (Bras10-6-2)



Figure 3.3

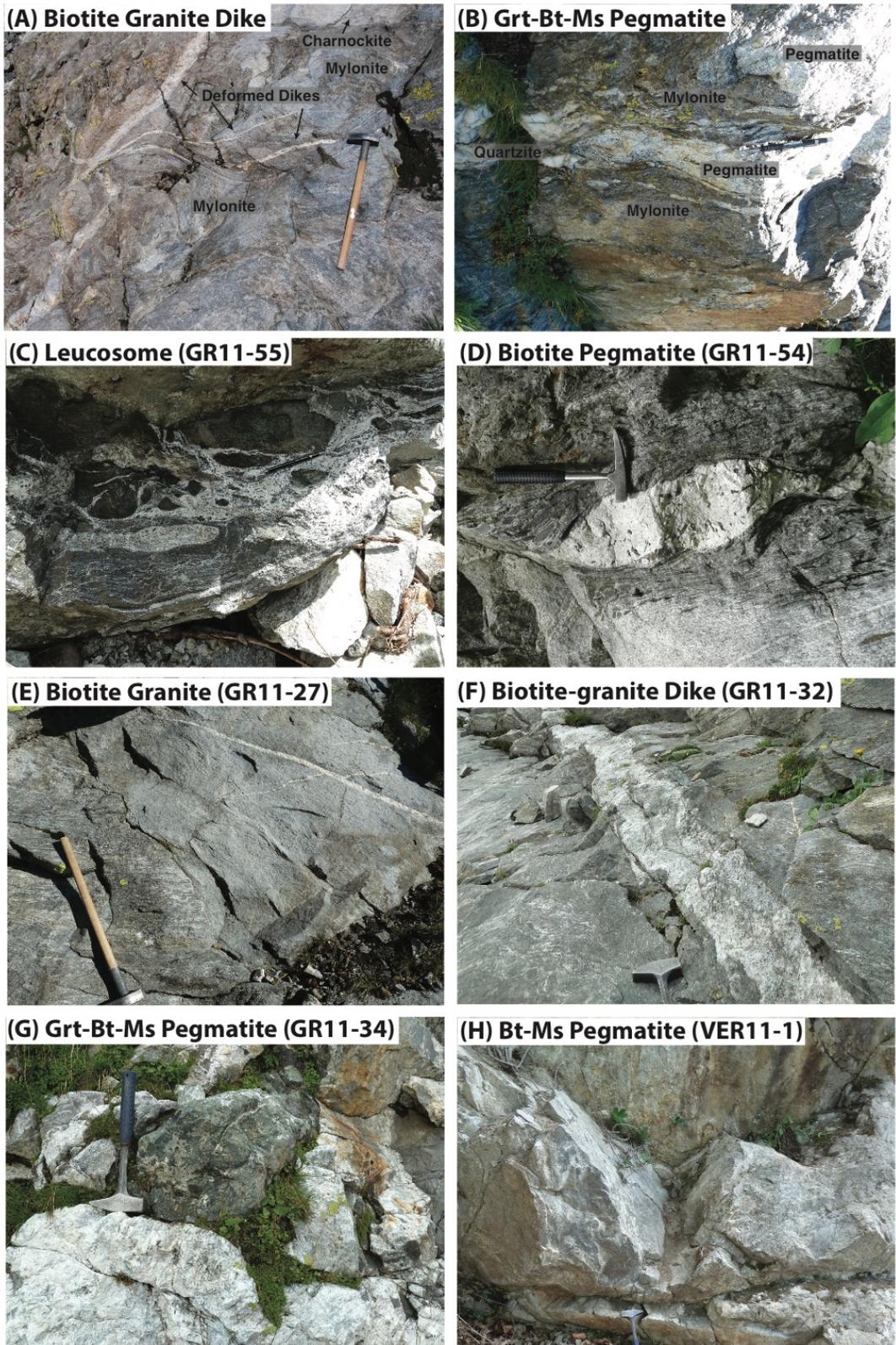


Figure 3.4

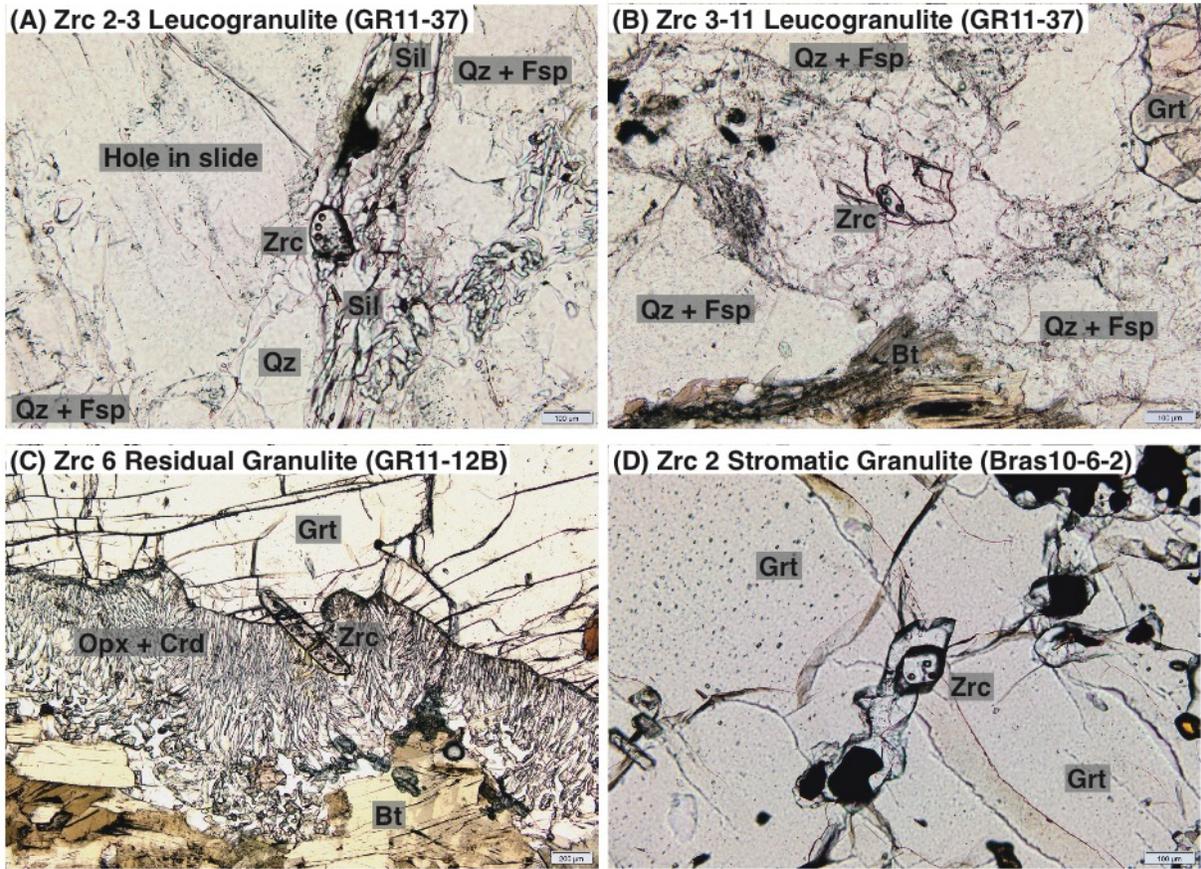


Figure 3.5

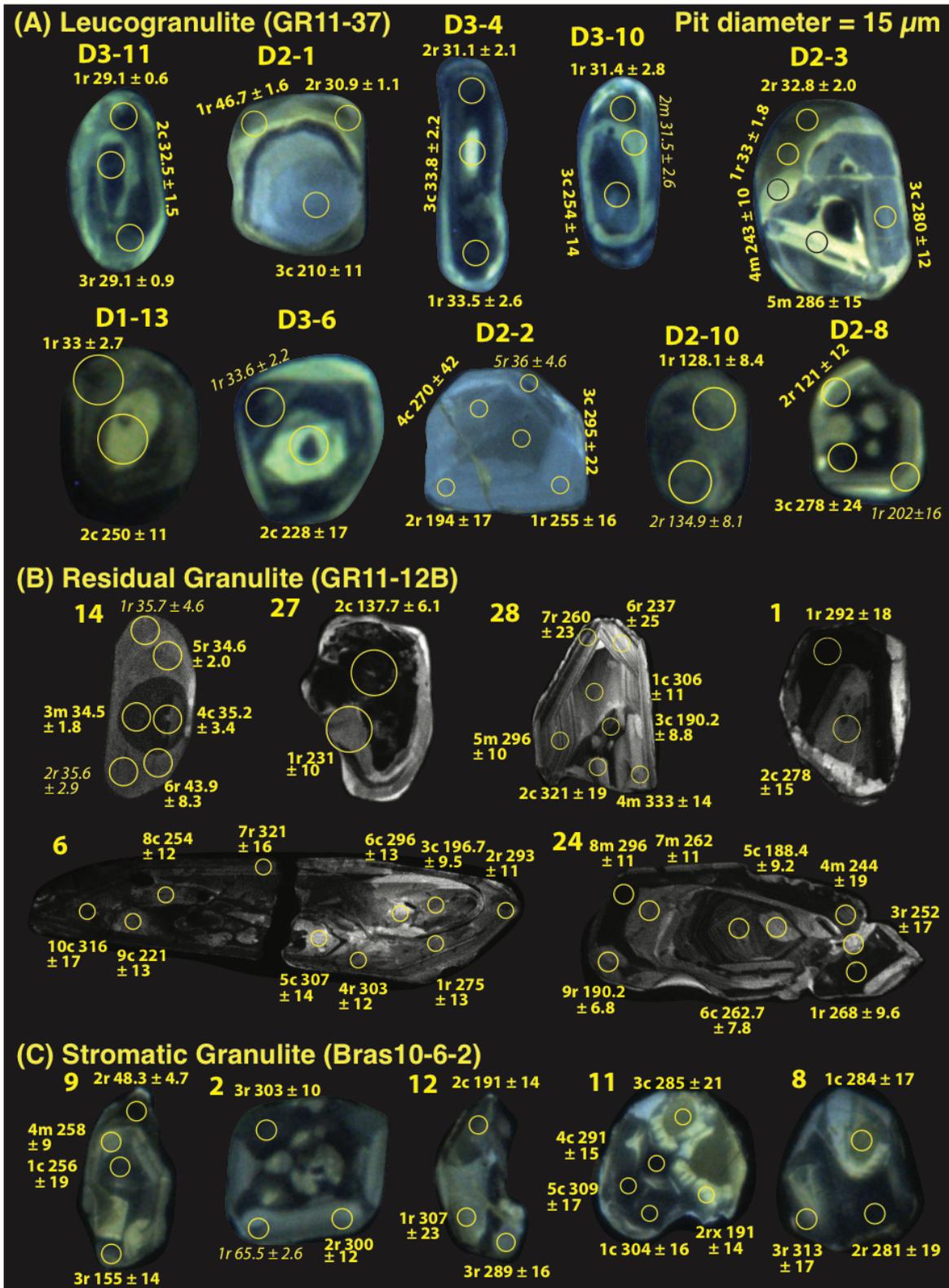


Figure 3.6

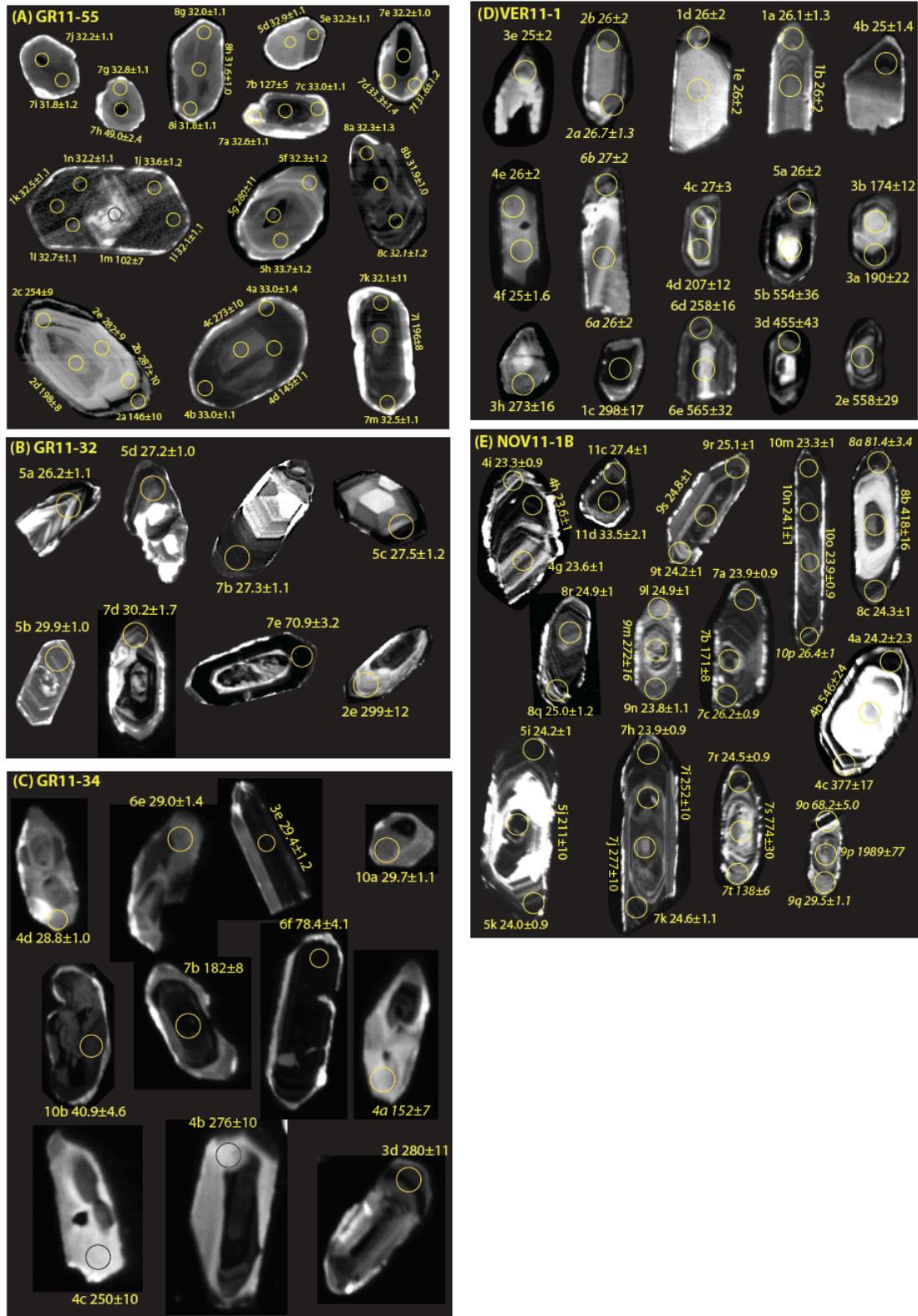


Figure 3.7

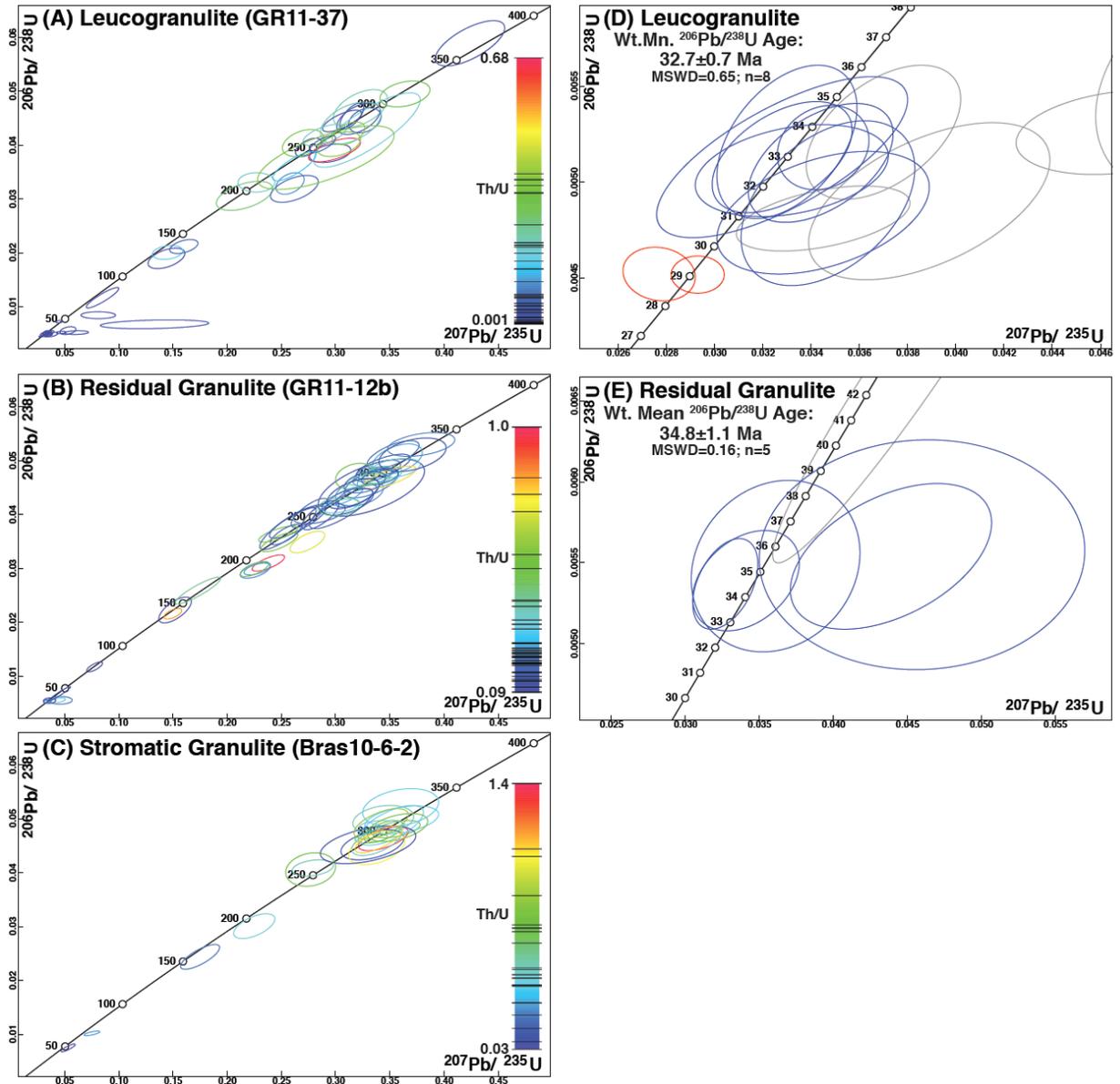


Figure 3.8

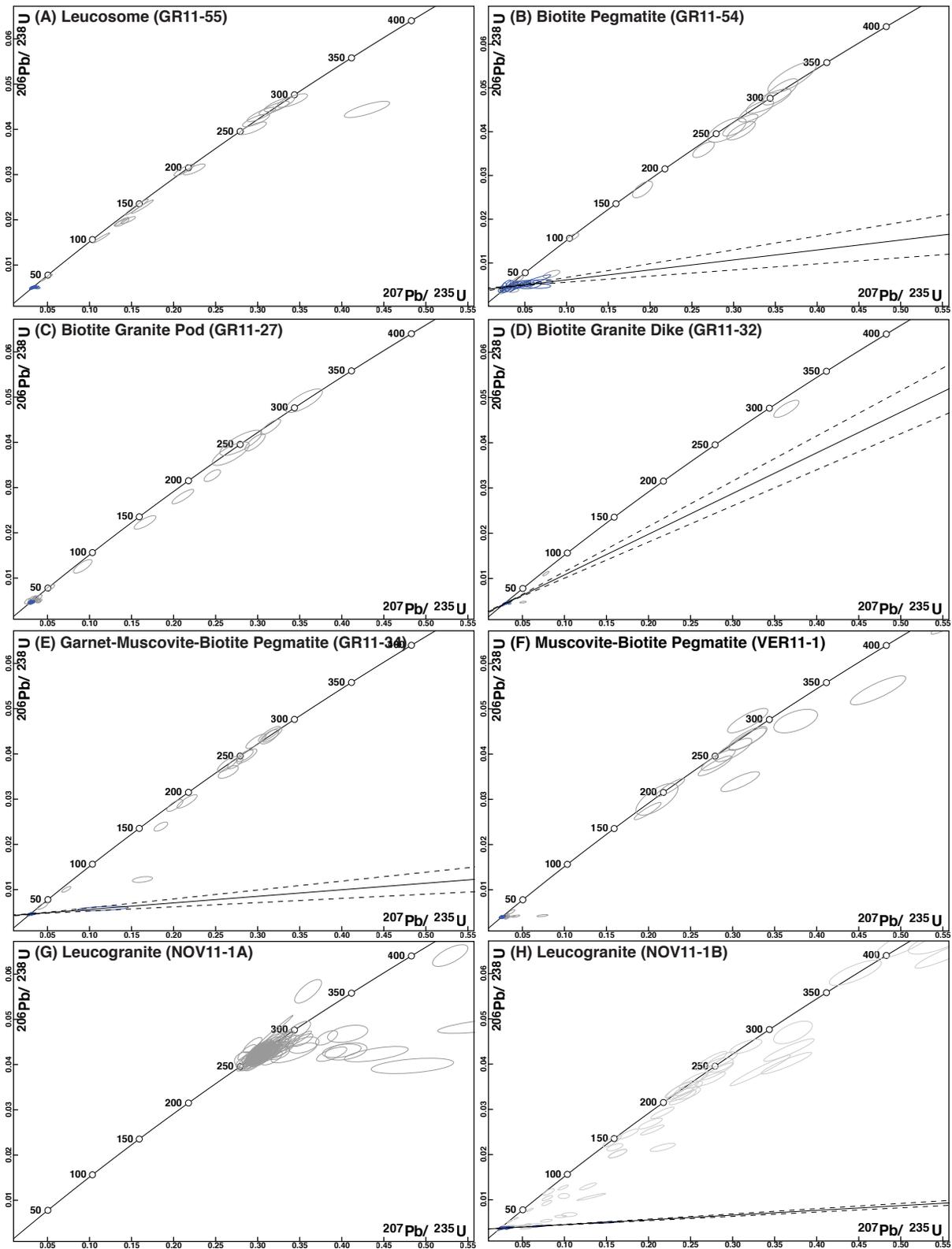


Figure 3.9

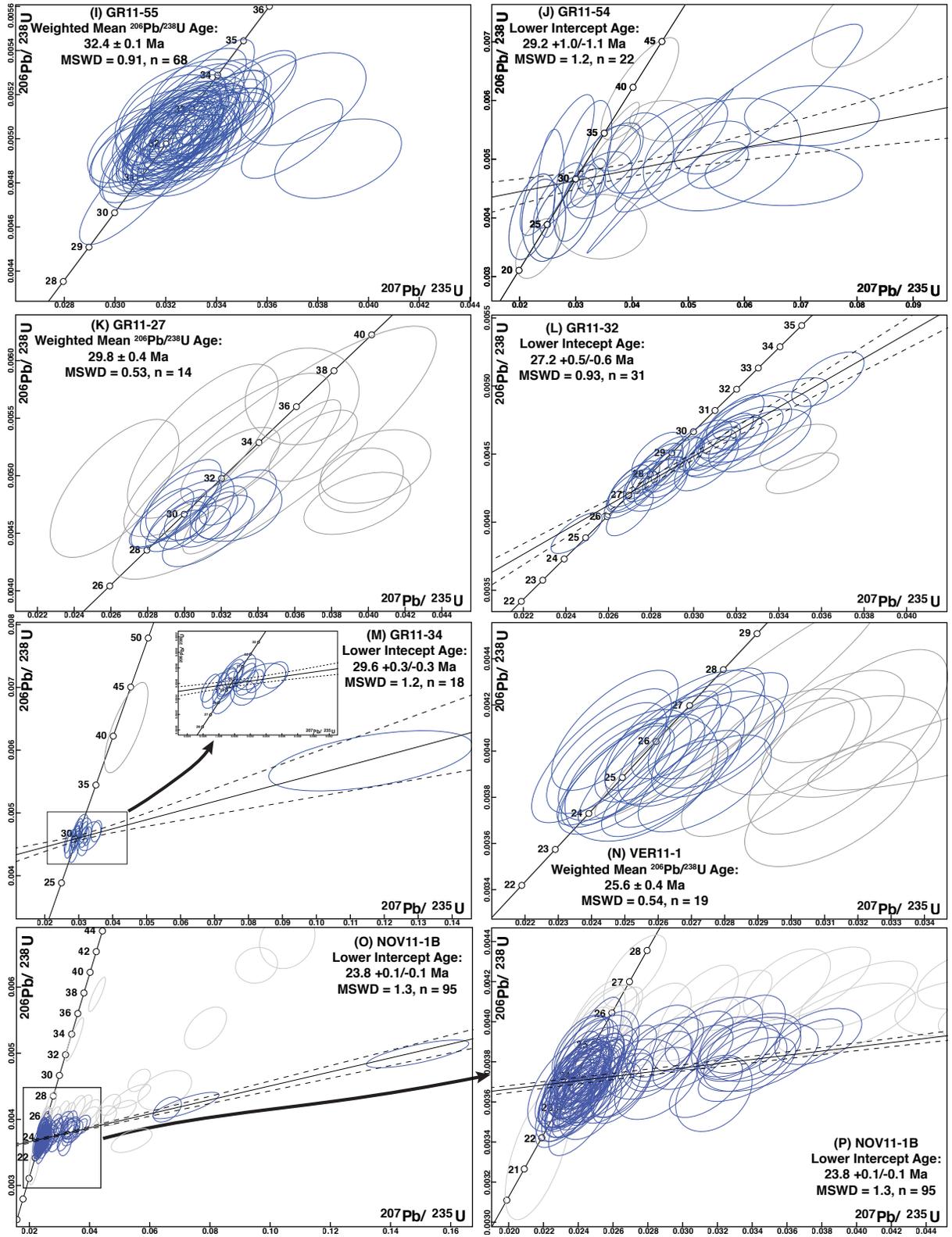


Figure 3.9 (continued)

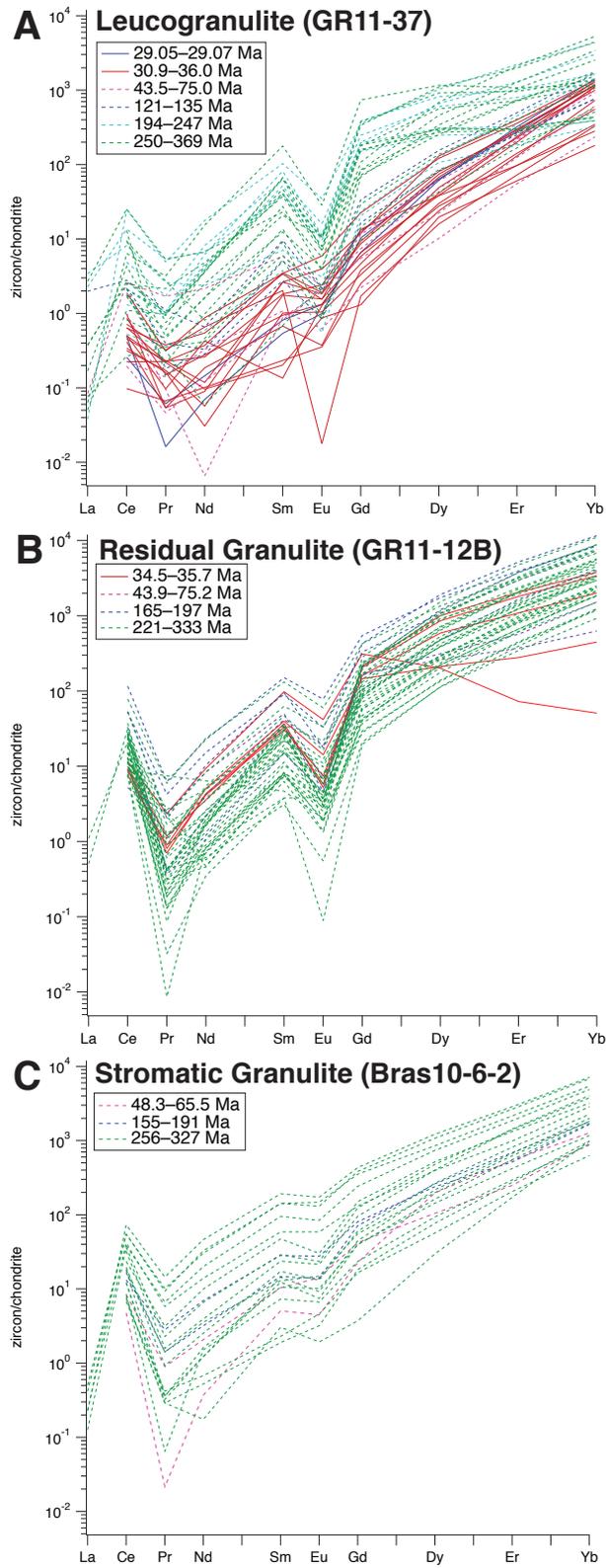


Figure 3.10

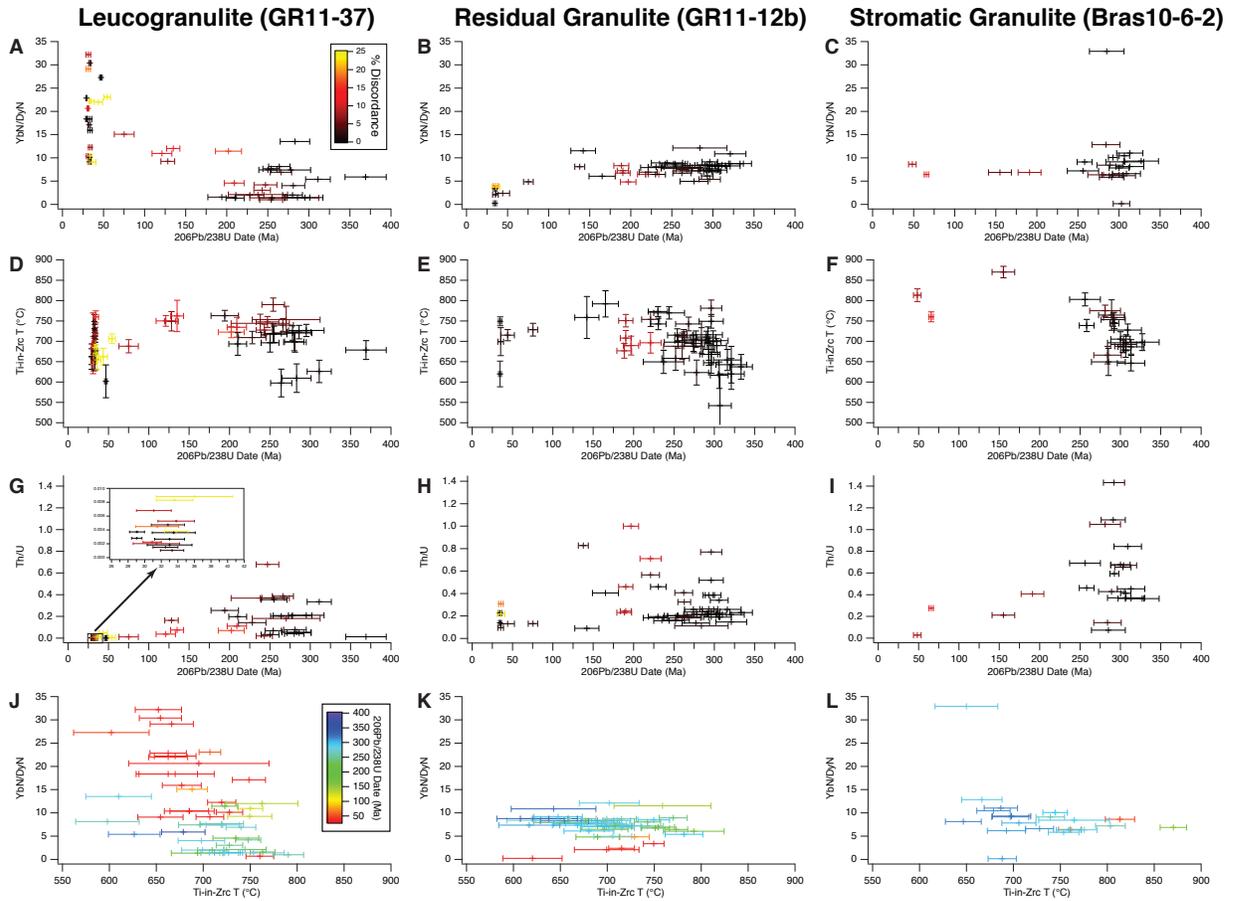


Figure 3.11

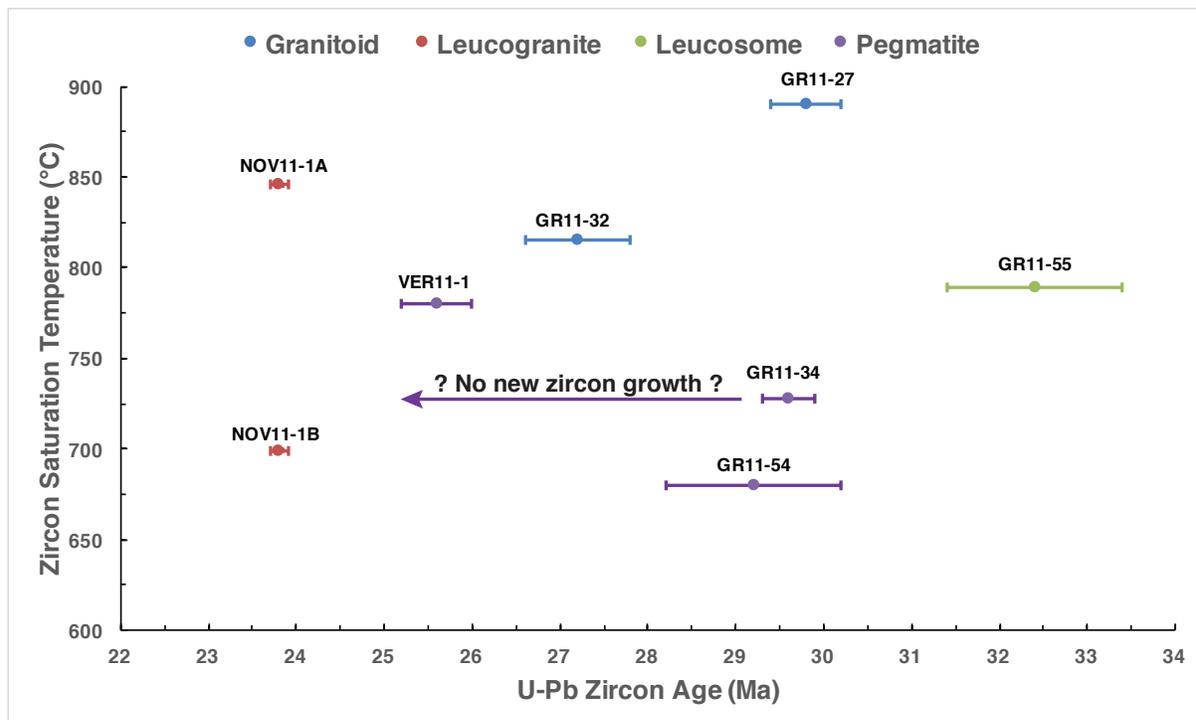


Figure 3.12

Major Mineral Reactions: Garnet Sapphirine
Opx Rutile

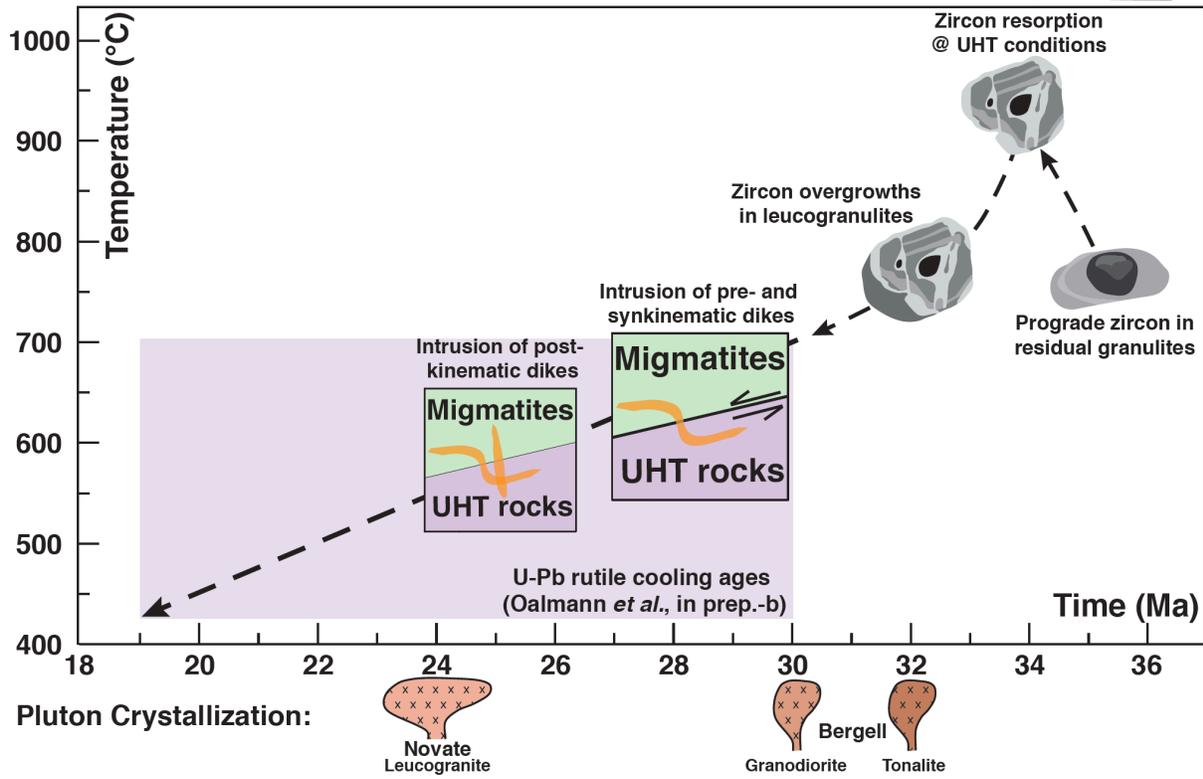
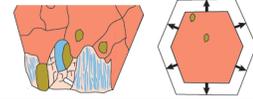


Figure 3.13

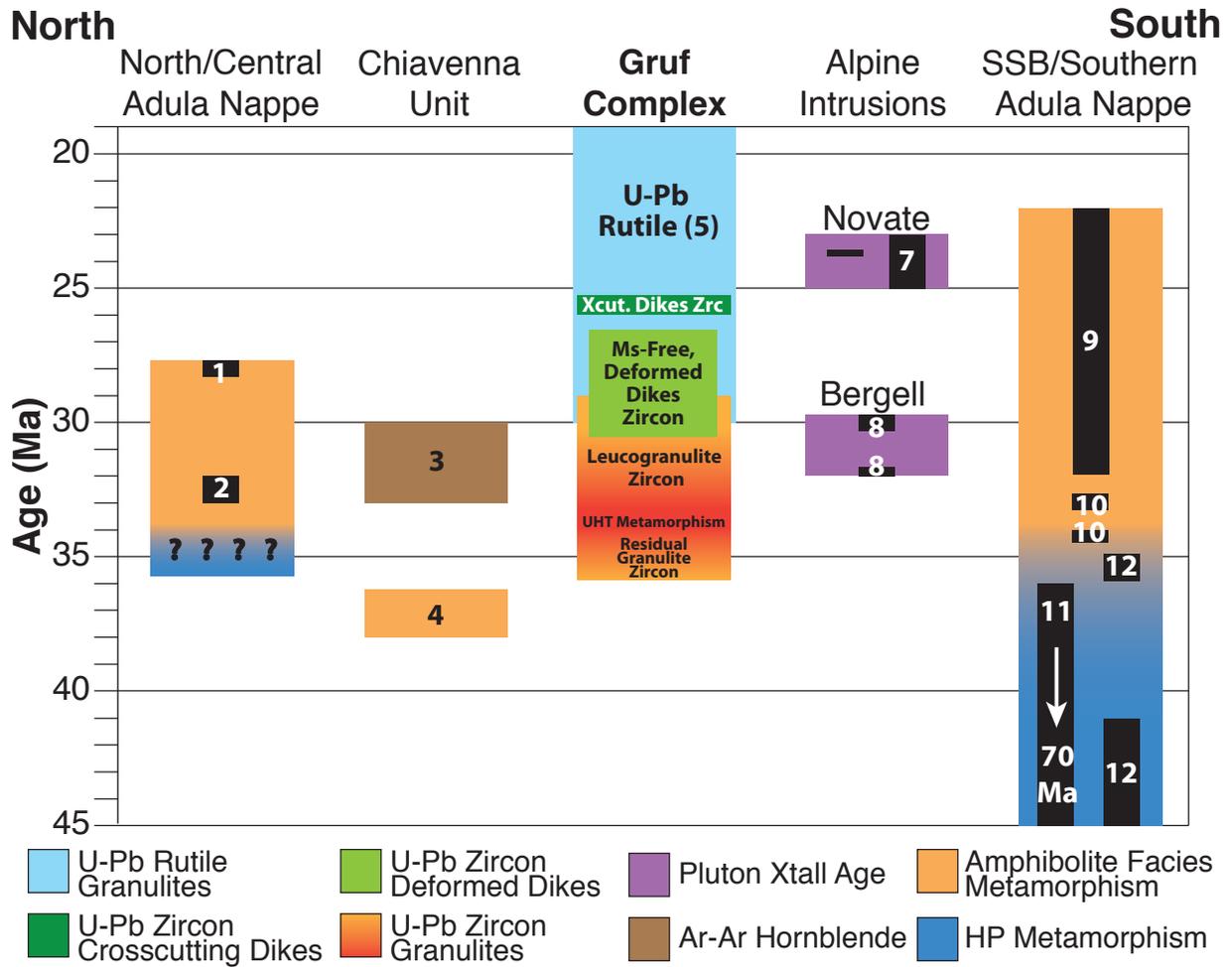


Figure 3.14



Figure 3.A1

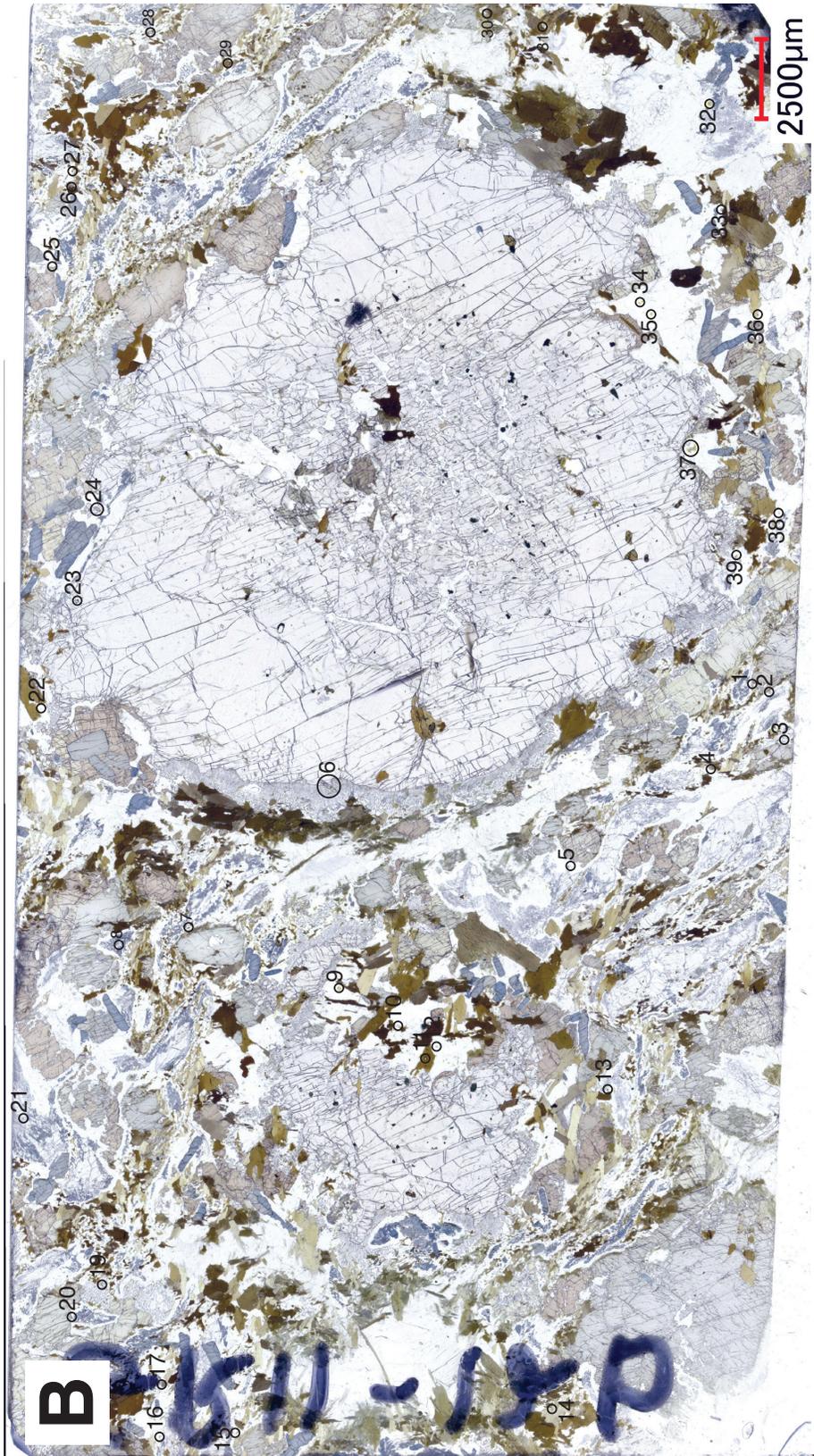


Figure 3.A1 (continued)

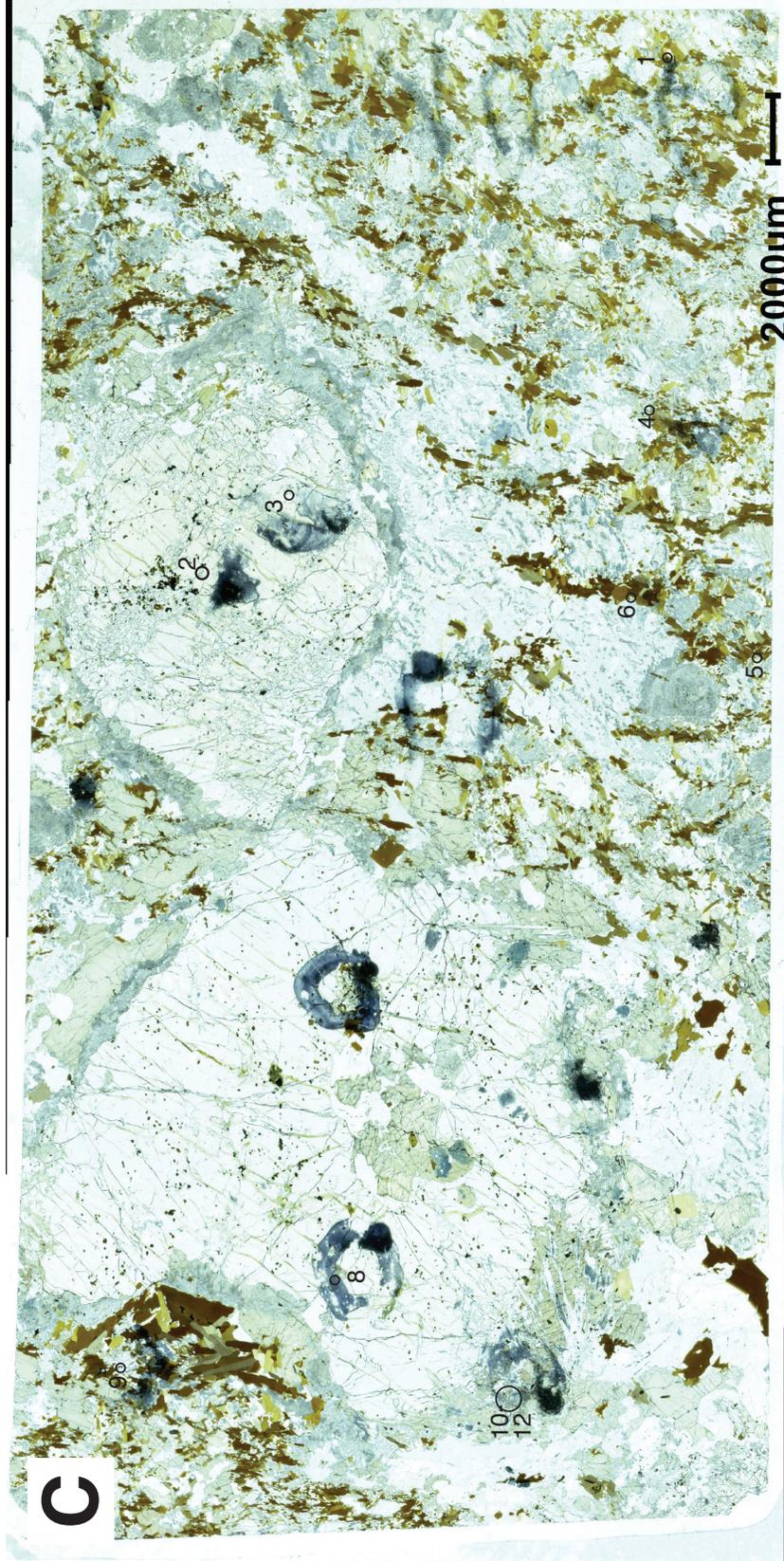


Figure 3.A1 (continued)

Table 3.1 LA-ICP-MS Parameters for Zircon Analyses from Granulite Samples

Laboratory & Sample Preparation	
Laboratory name	The University of Kansas, Dept. of Geology, Isotope Geochemistry Lab
Sample type/mineral	Zircon
Sample preparation	40 µm-thick, polished sections
Imaging	Microscope CL (color images) and SEM CL (black & white images)
Laser ablation system	
Make, model & type	ATL Arf excimer laser (193 nm), Photon Machines Analyte G2
Ablation cell & volume	HeLex
Laser wavelength	193 nm
Pulse width (ns)	5 ns
Fluence	2.5 J/cm ²
Repetition rate	10 Hz
Spot size (µm)	15 µm
Sampling mode / pattern	Single spots
Carrier gas	He, 1.1 l/min, Ar, 1.1 l/min
Ablation duration	20 s
Cell carrier gas flow	He
ICP-MS Instrument	
Make, Model & type	Thermo Element2 magnetic sector field ICP-MS (single collector)
Sample introduction	Aerosol with sample + He was mixed with Ar using a T-connector 15 cm upstream from torch.
RF power	1400 W
Make-up gas flow	Ar, 16 l/min
Sampling depth	20 µm
Detection system	single detector (SEM), counting & analog modes
Elements/ isotopes analyzed	29Si, 49Ti, 89Y, 94Zr, 139La, 140Ce, 141Pr, 146Nd, 147Sm, 153Eu, 157Gd, 163Dy, 166Er, 172Yb, 178Hf, 206Pb, 207Pb, 208Pb, 232Th, 238U
Integration time per peak (Sample Time in miliseconds)	29Si=5, 49Ti=15, 89Y=5, 94Zr=3, 139La=10, 140Ce=10, 141Pr=10, 146Nd=10, 147Sm=10, 153Eu=10, 157Gd=5, 163Dy=5, 166Er=5, 172Yb=5, 178Hf=3, 206Pb=6, 207Pb=10, 208Pb=1, 232Th=1, 238U=2
Total integration time (Segment Duration in miliseconds)	29Si=40, 49Ti=60, 89Y=5, 94Zr=3, 139La=10, 140Ce=10, 141Pr=10, 146Nd=10, 147Sm=10, 153Eu=10, 157Gd=5, 163Dy=5, 166Er=5, 172Yb=5, 178Hf=3, 206Pb=24, 207Pb=40, 208Pb=4, 232Th=4, 238U=8

Table 3.1

Total method time	57 seconds (112 runs, 1 pass)
Sensitivity (cps/ppm)	29Si=21.5cps/ppm, 49Ti=62.3cps/ppm, 51V=1387cps/ppm, 53Cr=bdl, 90Zr=540cps/ppm, 93Nb=1391cps/ppm, 95Mo=348cps/ppm, 118Sn=1448cps/ppm, 120Sn=1992cps/ppm, 121Sb=1275cps/ppm, 178Hf= 563cps/ppm, 181Ta=1743cps/ppm, 184W=770cps/ppm, 206Pb=1064cps/ppm, 208Pb=2127cps/ppm, 232Th=2278cps/ppm, 238U=5099cps/ppm
ICP Dead time	4 ns
UO+/U+	≤0.2%
238U+/232Th+	≤1.6
Data Processing	
Gas blank	28 s
Calibration strategy	Trace elements: standard-sampling bracketing + internal standardization assuming 15% Si (stoichiometric); U-Pb: standard-sampling bracketing
Reference material info	Trace elements: GJ1 (Jackson et al., 2004); NIST612 (Jochum et al., 2005)
Internal std for trace elements	29Si
Data processing package used / Correction for laser induced elemental fractionation (LIEF)	IGOR PRO, Iolite 2.5 (Patton et al. 2011); Trace elements: TraceElementsIS data reduction scheme; U-Pb: U_Pb_Geochronology (Paton, et al. 2010) and VizualAge data reduction schemes (Petrus & Kamber, 2011), exponential LIEF correction for U-Pb ratios.
Common-Pb correction, composition and uncertainty	Not performed because of high Hg backgrounds
Uncertainty level & propagation	Trace elements: 2se internal uncertainty; U:Pb: 2se propagated uncertainty from VizualAge data reduction scheme. Concordia diagrams were plotted using ET_Redux (McLean et al. 2016) with 2s uncertainty ellipses.
Reproducibility	Trace elements: 5–7%. U-Pb ratios: 2–4%
Quality control / Validation	Plesovice zircon (Sláma et al., 2008) for U-Pb isotopes and trace elements; GJ1 (Jackson et al., 2004) for trace elements only

Table 3.2 LA-ICP-MS Parameters for Zircon Analyses from Dike and Leucosome Samples

Laboratory & Sample Preparation	
Laboratory name	The University of Kansas, Dept. of Geology, Isotope Geochemistry Lab
Sample type/mineral	Zircon
Sample preparation	Grain mounts and tape mounts
Imaging	Cathodoluminescence
Laser ablation system	
Make, model & type	ATL Arf excimer laser (193 nm), Photon Machines Analyte G1
Ablation cell & volume	Frames Cell with teardrop-shaped insert
Laser wavelength	193 nm
Pulse width (ns)	5 ns
Fluence	2.5 J/cm ²
Repetition rate	10 Hz
Spot size (um)	30 µm
Sampling mode / pattern	Single spots
Carrier gas	He, 0.4 l/min, Ar, 1.2 l/min
Ablation duration	30 s
Cell carrier gas flow	He, 0.4 l/min, Ar, 1.2 l/min
ICP-MS Instrument	
Make, Model & type	Thermo Element2 magnetic sector field ICP-MS (single collector)
Sample introduction	Aerosol with sample + He was mixed with Ar using a Y-connector 15 cm upstream from torch.
RF power	1200 W
Make-up gas flow	Ar, 16 l/min
Sampling depth	25 µm
Detection system	single detector (SEM), counting & analog modes
Elements/ isotopes analyzed	206Pb, 207Pb, 208Pb, 232Th, 235U, 238U
Integration time per peak (Sample Time in miliseconds)	206Pb=1, 207Pb=3, 208Pb=1, 232Th=1, 235U=2, 238U=1
Total integration time (Segment Duration in miliseconds)	206Pb=4, 207Pb=12, 208Pb=4, 232Th=4, 235U=8, 238U=4
Total method time	1 min 18 s (370 runs, 5 pass)
Sensitivity (cps/ppm)	232Th=14600cps/ppm, 235U=240cps/ppm, 238U=34300cps/ppm
ICP Dead time	22 ns
UO+/U+	0.25%
238U+/232Th+	≤1.7
Data Processing	
Gas blank	33 s

Table 3.2

Calibration strategy	Standard-sampling bracketing
Reference material info	GJ1 (Jackson et al., 2004)
Internal std for trace elements	n/a
Data processing package used / Correction for laser induced elemental fractionation (LIEF)	IGOR PRO, Iolite 2.5 (Patton et al. 2011); U_Pb_Geochronology (Paton, et al. 2010) and VizualAge data reduction schemes (Petrus & Kamber, 2011), exponential LIEF correction for U-Pb ratios.
Common-Pb correction, composition and uncertainty	Not performed because of high Hg backgrounds
Uncertainty level & propagation	2SE propagated uncertainty from VizualAge data reduction scheme. Concordia diagrams were plotted using ET_Redux (McLean et al. 2016) with 2s uncertainty ellipses.
Reproducibility	206Pb/238U=1.5-2.0%; 207Pb/235U=2.0-2.5%
Quality control / Validation	Plesovice zircon (Sláma et al., 2008)

Table 3.3. Whole rock compositional data and zircon saturation temperatures for dike and leucosome samples.

Sample	NOV11-1A	NOV11-1B	VER11-1	GR11-34	GR11-32	GR11-27	GR11-54	GR11-55
Rock Type	leucogranite	leucogranite	pegmatite	pegmatite	granitic dike	granitic pod	pegmatite	leucosome
SiO ₂	71.30	74.04	67.94	62.44	68.79	62.95	70.04	70.98
TiO ₂	0.40	0.04	0.21	0.40	0.30	0.67	0.20	0.28
Al ₂ O ₃	14.42	14.70	17.37	17.28	16.18	17.94	16.45	14.11
FeO	2.81	0.48	1.25	3.65	2.00	3.86	1.65	2.40
MnO	0.04	0.03	0.03	0.08	0.04	0.06	0.04	0.06
MgO	0.64	0.07	0.36	3.64	0.76	1.34	0.67	1.56
CaO	1.64	1.41	2.26	5.30	2.36	2.93	2.20	1.91
Na ₂ O	2.80	4.32	4.89	3.84	3.99	4.11	5.10	3.45
K ₂ O	4.92	3.68	3.33	1.89	4.17	4.88	2.93	3.83
P ₂ O ₅	0.17	0.05	0.08	0.44	0.20	0.26	0.18	0.16
LOI	0.50	0.39	0.67	0.70	0.49	0.32	0.67	0.42
Total	99.82	99.33	98.65	99.85	99.50	99.65	100.23	99.33
Na (mol)	0.045	0.070	0.079	0.062	0.064	0.066	0.082	0.056
K (mol)	0.052	0.039	0.035	0.020	0.044	0.052	0.031	0.041
Ca (mol)	0.029	0.025	0.040	0.095	0.042	0.052	0.039	0.034
Al (mol)	0.14	0.14	0.17	0.17	0.16	0.18	0.16	0.14
Si (mol)	1.19	1.23	1.13	1.04	1.14	1.05	1.17	1.18
Zr (ppm)	206	35	108	89	167	405	30	119
ASI	1.11	1.04	1.10	0.96	1.07	1.06	1.05	1.07
M	0.93	0.90	1.01	1.54	1.06	1.21	1.02	1.01
Zre Sat T (°C)	846	699	780	728	815	890	680	789
Age (Ma)	23.9	23.9	25.6	29.5	27.2	29.9	29.2	32.4
Age ±2s (Ma)	0.1	0.1	0.3	0.3	0.6	0.4	1.0	0.1

Notes:

LOI = loss on ignition

ASI = alumina saturation index = molar Al/(Na+K+Ca)

M = (Na+K+2·Ca)/(Al·Si) (Watson and Harrison, 1983))

Zre Sat T (°C) = zircon saturation temperature = 12,900/[2.95 + 0.85M + ln(496,000/Zr)] (Watson and Harrison, 1983)

Table 3.3

C ₁		C ₂		C ₃		C ₄		C ₅		C ₆		C ₇		C ₈		C ₉		C ₁₀		C ₁₁		C ₁₂		C ₁₃		C ₁₄		C ₁₅		C ₁₆		C ₁₇		C ₁₈		C ₁₉		C ₂₀		C ₂₁		C ₂₂		C ₂₃		C ₂₄		C ₂₅		C ₂₆		C ₂₇		C ₂₈		C ₂₉		C ₃₀		C ₃₁		C ₃₂		C ₃₃		C ₃₄		C ₃₅		C ₃₆		C ₃₇		C ₃₈		C ₃₉		C ₄₀		C ₄₁		C ₄₂		C ₄₃		C ₄₄		C ₄₅		C ₄₆		C ₄₇		C ₄₈		C ₄₉		C ₅₀		C ₅₁		C ₅₂		C ₅₃		C ₅₄		C ₅₅		C ₅₆		C ₅₇		C ₅₈		C ₅₉		C ₆₀		C ₆₁		C ₆₂		C ₆₃		C ₆₄		C ₆₅		C ₆₆		C ₆₇		C ₆₈		C ₆₉		C ₇₀		C ₇₁		C ₇₂		C ₇₃		C ₇₄		C ₇₅		C ₇₆		C ₇₇		C ₇₈		C ₇₉		C ₈₀		C ₈₁		C ₈₂		C ₈₃		C ₈₄		C ₈₅		C ₈₆		C ₈₇		C ₈₈		C ₈₉		C ₉₀		C ₉₁		C ₉₂		C ₉₃		C ₉₄		C ₉₅		C ₉₆		C ₉₇		C ₉₈		C ₉₉		C ₁₀₀	
0.00	0.01	0.02	0.03	0.04	0.05	0.06	0.07	0.08	0.09	0.10	0.11	0.12	0.13	0.14	0.15	0.16	0.17	0.18	0.19	0.20	0.21	0.22	0.23	0.24	0.25	0.26	0.27	0.28	0.29	0.30	0.31	0.32	0.33	0.34	0.35	0.36	0.37	0.38	0.39	0.40	0.41	0.42	0.43	0.44	0.45	0.46	0.47	0.48	0.49	0.50	0.51	0.52	0.53	0.54	0.55	0.56	0.57	0.58	0.59	0.60	0.61	0.62	0.63	0.64	0.65	0.66	0.67	0.68	0.69	0.70	0.71	0.72	0.73	0.74	0.75	0.76	0.77	0.78	0.79	0.80	0.81	0.82	0.83	0.84	0.85	0.86	0.87	0.88	0.89	0.90	0.91	0.92	0.93	0.94	0.95	0.96	0.97	0.98	0.99	1.00																																																																																																			
0.00	0.01	0.02	0.03	0.04	0.05	0.06	0.07	0.08	0.09	0.10	0.11	0.12	0.13	0.14	0.15	0.16	0.17	0.18	0.19	0.20	0.21	0.22	0.23	0.24	0.25	0.26	0.27	0.28	0.29	0.30	0.31	0.32	0.33	0.34	0.35	0.36	0.37	0.38	0.39	0.40	0.41	0.42	0.43	0.44	0.45	0.46	0.47	0.48	0.49	0.50	0.51	0.52	0.53	0.54	0.55	0.56	0.57	0.58	0.59	0.60	0.61	0.62	0.63	0.64	0.65	0.66	0.67	0.68	0.69	0.70	0.71	0.72	0.73	0.74	0.75	0.76	0.77	0.78	0.79	0.80	0.81	0.82	0.83	0.84	0.85	0.86	0.87	0.88	0.89	0.90	0.91	0.92	0.93	0.94	0.95	0.96	0.97	0.98	0.99	1.00																																																																																																			

Table 3. 5. U-Pb Isotope Data for Dike and Leucosome Samples

Sample and Analysis Number	Isotopic Ratios					Dates (Ma)					Concentration					
	207Pb/235U (abs)	± 2SE (abs)	206Pb/238U (abs)	± 2SE (abs)	207Pb/206Pb (abs)	± 2SE (abs)	Rho: 206/238 vs 207/235	238U (abs)	± 2SE (abs)	206Pb/238U (abs)	± 2SE (abs)	207Pb/206Pb (abs)	% Discord. 6/38-7/35	U (ppm)	Th/U	
Leucosome of Magmatic Breccia (GR11-55)																
GR11-55-4g.FIN2	0.031	0.002	0.0048	± 0.0003	0.0468	0.0022	0.82141	30.7	1.6	30.8	± 2.2	0.3	104	34	904	0.06
GR11-55-1f.FIN2	0.031	0.002	0.0049	0.0002	0.0472	0.0022	0.49062	31.34	1.2	31.39	1.8	0.2	166	30	236.6	0.32
GR11-55-3c.FIN2	0.033	0.002	0.0049	0.0002	0.0472	0.0022	0.40884	31.52	1	32.65	1.8	3.5	198	33	311	0.56
GR11-55-7f.FIN2	0.039	0.003	0.0049	0.0002	0.057	0.0028	0.38373	31.55	1.2	38.6	2.4	18.3	473	47	233	0.54
GR11-55-8e.FIN2	0.031	0.002	0.0049	0.0002	0.0454	0.0022	0.36051	31.58	1	30.85	1.7	-2.4	160	31	168	0.62
GR11-55-8h.FIN2	0.032	0.002	0.0049	0.0002	0.0461	0.0020	0.58466	31.6	1	31.54	1.8	-0.2	141	25	266.3	0.48
GR11-55-8k.FIN2	0.032	0.002	0.0049	0.0002	0.049	0.0022	0.52594	31.7	1	32.46	1.8	2.3	213	27	245.3	0.53
GR11-55-7i.FIN2	0.032	0.002	0.0049	0.0002	0.0478	0.0021	0.60017	31.78	1.2	31.68	1.7	-0.3	157	21	292	0.63
GR11-55-4k.FIN2	0.032	0.002	0.0049	0.0002	0.0472	0.0020	0.57197	31.79	1.1	31.82	1.8	0.1	159	24	319.2	0.34
GR11-55-8i.FIN2	0.033	0.002	0.0049	0.0002	0.0474	0.0021	0.61468	31.79	1.1	32.49	1.8	2.2	186	23	218.7	0.47
GR11-55-3c.FIN2	0.033	0.002	0.0049	0.0002	0.0478	0.0022	0.63652	31.8	1.3	32.6	2	2.5	174	43	314	0.59
GR11-55-1e.FIN2	0.032	0.002	0.005	0.0002	0.0465	0.0022	0.54625	31.84	1.1	32.1	1.9	0.8	167	30	245.7	0.40
GR11-55-8b.FIN2	0.033	0.002	0.005	0.0002	0.0475	0.0019	0.76351	31.86	1	32.63	1.7	2.4	143	20	256.3	0.00
GR11-55-5a.FIN2	0.031	0.002	0.005	0.0002	0.0455	0.0020	0.63558	31.87	1.2	30.9	1.8	-3.1	138	29	289.4	0.34
GR11-55-2h.FIN2	0.032	0.002	0.005	0.0002	0.0458	0.0020	0.49433	31.88	1.1	32.08	1.8	0.6	151	28	263.3	0.36
GR11-55-5c.FIN2	0.031	0.002	0.005	0.0002	0.0456	0.0021	0.46275	31.9	1	31.26	1.8	-2.0	174	30	201.4	0.39
GR11-55-1g.FIN2	0.032	0.002	0.005	0.0002	0.0461	0.0020	0.62535	31.94	1.1	31.93	1.8	0.0	145	27	225.2	0.38
GR11-55-8g.FIN2	0.031	0.002	0.005	0.0002	0.0451	0.0020	0.50557	31.97	1.1	30.62	1.7	-4.4	114	30	223	0.46
GR11-55-3b.FIN2	0.032	0.002	0.005	0.0003	0.0463	0.0025	0.75353	32	1.6	31.8	2.3	-0.6	120	40	341	0.72
GR11-55-5b.FIN2	0.033	0.002	0.005	0.0002	0.048	0.0022	0.4105	32.01	1.1	32.71	1.8	2.1	182	25	223.8	0.34
GR11-55-1i.FIN2	0.033	0.002	0.005	0.0002	0.0473	0.0019	0.75469	32.05	1.1	32.88	1.7	2.5	116	15	4090	0.01
GR11-55-8d.FIN2	0.033	0.002	0.005	0.0002	0.0482	0.0027	0.24902	32.06	1.2	33.1	2	3.1	252	38	79.5	0.46
GR11-55-7k.FIN2	0.032	0.002	0.005	0.0002	0.0469	0.0019	0.78505	32.12	1.1	31.82	1.7	-0.9	115	19	2730	0.01
GR11-55-8c.FIN2	0.037	0.002	0.005	0.0002	0.0528	0.0022	0.72674	32.12	1.2	36.88	2	12.9	345	26	1945	0.00
GR11-55-1n.FIN2	0.032	0.002	0.005	0.0002	0.0465	0.0018	0.82794	32.15	1.1	32.22	1.7	0.2	91	16	3447	0.02
GR11-55-3a.FIN2	0.032	0.002	0.005	0.0002	0.0468	0.0021	0.61312	32.16	1.2	32.13	1.8	-0.1	181	31	262	0.69
GR11-55-8j.FIN2	0.032	0.002	0.005	0.0002	0.0478	0.0022	0.40981	32.16	1.1	32.19	1.9	0.1	164	31	219	0.53
GR11-55-4c.FIN2	0.032	0.002	0.005	0.0002	0.0468	0.0022	0.4478	32.18	1.2	31.77	1.8	-1.3	166	22	215.3	0.33
GR11-55-7e.FIN2	0.034	0.002	0.005	0.0002	0.0492	0.0021	0.56703	32.19	1	34.26	1.9	6.0	197	33	1321	0.06
GR11-55-5e.FIN2	0.032	0.002	0.005	0.0002	0.0463	0.0021	0.60576	32.21	1.1	32.46	1.9	0.8	145	28	189.4	0.73
GR11-55-7j.FIN2	0.033	0.002	0.005	0.0002	0.048	0.0021	0.6364	32.21	1.1	32.9	1.9	2.1	178	24	445	0.32

Table 3.5

Sample and Analysis Number	Isotopic Ratios				Dates (Ma)				Concentrator							
	207Pb/ ± 2SE		206Pb/ ± 2SE		207Pb/ ± 2SE		Rho: 206/238 vs 207/235		206Pb/ ± 2SE		207Pb/ ± 2SE		U			
	235U (abs)	238U (abs)	206Pb/ (abs)	206Pb/ (abs)	207Pb/ (abs)	206Pb/ (abs)	207Pb/ (abs)	206Pb/ (abs)	235U (abs)	238U (abs)	207Pb/ (abs)	206Pb/ (abs)	(ppm)	Th/U		
GR11-55-8a.FIN2	0.033	0.002	0.005	0.0002	0.0476	0.0019	0.88715	32.27	1.3	32.89	1.8	1.9	100	16	2070	0.00
GR11-55-5f.FIN2	0.032	0.002	0.005	0.0002	0.0477	0.0022	0.55418	32.32	1.2	32.4	1.9	0.2	171	24	236	0.22
GR11-55-2i.FIN2	0.033	0.002	0.005	0.0002	0.0469	0.0021	0.52677	32.36	1.1	33.13	1.9	2.3	196	34	237.7	0.34
GR11-55-2f.FIN2	0.032	0.002	0.005	0.0002	0.0456	0.0023	0.32975	32.38	1.1	31.9	1.8	-1.5	196	30	123.8	0.43
GR11-55-4j.FIN2	0.032	0.002	0.005	0.0002	0.0469	0.0021	0.51214	32.38	1.1	32.33	1.9	-0.2	200	33	232.6	0.29
GR11-55-2g.FIN2	0.032	0.002	0.005	0.0002	0.0455	0.0019	0.76086	32.39	1.1	32.3	1.9	-0.3	127	25	384	0.61
GR11-55-3d.FIN2	0.034	0.002	0.005	0.0002	0.048	0.0021	0.53123	32.42	1.1	33.59	1.8	3.5	178	24	321	0.53
GR11-55-5j.FIN2	0.034	0.002	0.005	0.0002	0.0495	0.0022	0.55081	32.42	1.1	34.2	2	5.2	239	30	218	0.47
GR11-55-7m.FIN2	0.032	0.002	0.0051	0.0002	0.0464	0.0019	0.76996	32.48	1.1	31.96	1.7	-1.6	91	14	1570	0.02
GR11-55-6a.FIN2	0.033	0.002	0.0051	0.0002	0.0471	0.0019	0.72697	32.5	1.1	33.12	1.8	1.9	113	21	1403	0.03
GR11-55-1k.FIN2	0.034	0.002	0.0051	0.0002	0.0493	0.0020	0.68879	32.52	1.1	34.22	1.8	5.0	180	19	3240	0.03
GR11-55-1a.FIN2	0.032	0.002	0.0051	0.0002	0.0466	0.0019	0.80506	32.54	1.2	32.22	1.8	-1.0	90	16	5110	0.02
GR11-55-7a.FIN2	0.034	0.002	0.0051	0.0002	0.0472	0.0023	0.48728	32.58	1.1	34.1	2	4.5	260	30	112.9	0.65
GR11-55-9a.FIN2	0.033	0.002	0.0051	0.0002	0.0483	0.0021	0.57467	32.62	1.1	32.83	1.8	0.6	181	24	268.8	0.50
GR11-55-5i.FIN2	0.033	0.002	0.0051	0.0002	0.0475	0.0023	0.49048	32.62	1.1	32.9	2	0.9	245	32	188.1	0.42
GR11-55-1b.FIN2	0.033	0.002	0.0051	0.0002	0.0472	0.0019	0.81419	32.62	1.2	33.1	1.8	1.5	125	16	4630	0.01
GR11-55-4f.FIN2	0.033	0.002	0.0051	0.0002	0.0479	0.0022	0.57968	32.65	1.1	33.32	1.9	2.0	205	35	253.5	0.43
GR11-55-1l.FIN2	0.033	0.002	0.0051	0.0002	0.0476	0.0019	0.69215	32.67	1.1	33.43	1.8	2.3	126	18	3270	0.02
GR11-55-4i.FIN2	0.033	0.002	0.0051	0.0002	0.0477	0.0021	0.67903	32.7	1.3	33.2	1.9	1.5	96	22	926	0.05
GR11-55-9b.FIN2	0.033	0.002	0.0051	0.0002	0.0467	0.0020	0.62661	32.74	1.1	32.64	1.8	-0.3	155	24	237	0.38
GR11-55-1d.FIN2	0.033	0.002	0.0051	0.0002	0.0467	0.0019	0.75331	32.74	1.1	32.85	1.7	0.3	106	19	6760	0.02
GR11-55-3h.FIN2	0.034	0.002	0.0051	0.0002	0.0481	0.0020	0.71621	32.77	1.1	34	1.8	3.6	148	20	1282	0.00
GR11-55-4h.FIN2	0.032	0.002	0.0051	0.0002	0.0463	0.0028	0.42141	32.8	1.4	32.2	2.1	-1.9	163	53	236.3	0.71
GR11-55-7g.FIN2	0.033	0.002	0.0051	0.0002	0.0466	0.0021	0.50913	32.8	1.1	32.5	1.9	-0.9	141	23	224	0.48
GR11-55-3g.FIN2	0.033	0.003	0.0051	0.0004	0.0475	0.0023	0.84479	32.8	2.3	32.9	2.6	0.3	139	63	3040	0.00
GR11-55-5d.FIN2	0.032	0.002	0.0051	0.0002	0.0469	0.0024	0.40338	32.85	1.1	32.2	1.9	-2.0	202	34	137.1	0.68
GR11-55-4l.FIN2	0.033	0.002	0.0051	0.0002	0.0461	0.0019	0.70144	32.92	1.1	32.54	1.8	-1.2	113	22	1040	0.03
GR11-55-4a.FIN2	0.034	0.002	0.0051	0.0002	0.0484	0.0022	0.58778	33	1.4	33.84	1.9	2.5	162	35	1084	0.05
GR11-55-4b.FIN2	0.033	0.002	0.0051	0.0002	0.0468	0.0019	0.73822	33.03	1.1	32.9	1.8	-0.4	115	20	2257	0.03
GR11-55-7c.FIN2	0.034	0.002	0.0051	0.0002	0.0477	0.0020	0.64394	33.03	1.1	34.2	1.9	3.4	155	24	480	0.18
GR11-55-6c.FIN2	0.034	0.002	0.0051	0.0002	0.0475	0.0019	0.67306	33.08	1.1	34.12	1.8	3.0	137	19	2221	0.01
GR11-55-1h.FIN2	0.033	0.002	0.0052	0.0002	0.0462	0.0020	0.55593	33.25	1.1	33.14	1.8	-0.3	137	28	293	0.35
GR11-55-7d.FIN2	0.037	0.002	0.0052	0.0002	0.0505	0.0027	0.46123	33.3	1.4	37.3	2.4	10.7	341	67	213	0.53

Sample and Analysis Number	Isotopic Ratios										Dates (Ma)				Concentrator			
	207Pb/ ± 2SE		206Pb/ ± 2SE		207Pb/ ± 2SE		207Pb/ ± 2SE		Rho: 206/238 vs 207/235		206Pb/ ± 2SE		207Pb/ ± 2SE		207Pb/ ± 2SE		U (ppm)	Th/U
	(abs)	235U	(abs)	238U	(abs)	206Pb	(abs)	206Pb	(abs)	206Pb	(abs)	235U	(abs)	235U	(abs)	206Pb		
GR11-55-8f.FIN2	0.034	0.002	0.0052	0.0002	0.0463	0.0024	0.28394	33.35	1.1	33.4	2	0.1	188	34	106.2	0.48		
GR11-55-3f.FIN2	0.035	0.002	0.0052	0.0002	0.0495	0.0020	0.76558	33.37	1.2	35.07	1.9	4.8	190	21	1750	0.00		
GR11-55-1j.FIN2	0.035	0.002	0.0052	0.0002	0.0487	0.0020	0.77873	33.59	1.2	35.05	1.9	4.2	152	20	2424	0.02		
GR11-55-5h.FIN2	0.033	0.002	0.0052	0.0002	0.046	0.0020	0.64036	33.65	1.2	33.23	1.9	-1.3	175	33	254.7	0.35		
GR11-55-3i.FIN2	0.045	0.003	0.0067	0.0003	0.0491	0.0020	0.90172	49.3	2	44.6	2.7	2.9	177	19	640	0.02		
GR11-55-7h.FIN2	0.053	0.003	0.0076	0.0004	0.0503	0.0025	0.74182	49	2.4	52	3.2	5.8	181	38	1805	0.09		
GR11-55-1m.FIN2	0.114	0.010	0.016	0.0011	0.0521	0.0021	0.96663	102.1	7.2	109.8	9	7.0	294	23	1256	0.16		
GR11-55-1c.FIN2	0.139	0.008	0.0194	0.0009	0.0524	0.0022	0.80069	123.7	5.4	131.7	7.5	6.1	308	31	2239	0.41		
GR11-55-6b.FIN2	0.138	0.009	0.0196	0.0009	0.0511	0.0020	0.92023	125.1	6	130.9	7.7	4.4	250	25	1300	0.05		
GR11-55-7b.FIN2	0.146	0.009	0.0199	0.0008	0.0529	0.0022	0.8263	127.3	4.8	138.3	7.6	8.0	317	23	1930	0.36		
GR11-55-4d.FIN2	0.156	0.014	0.0228	0.0018	0.0492	0.0020	0.96708	145	11	145	12	0.0	188	22	512	0.11		
GR11-55-2a.FIN2	0.163	0.013	0.0229	0.0015	0.0518	0.0021	0.96607	146	9.7	152.6	12	4.3	281	17	2407	0.00		
GR11-55-7l.FIN2	0.212	0.012	0.031	0.0013	0.0511	0.0021	0.86277	196.4	7.9	195.2	10	-0.6	246	21	478	0.21		
GR11-55-2d.FIN2	0.225	0.013	0.0312	0.0012	0.0524	0.0022	0.79582	197.7	7.8	205.4	11	3.7	303	30	614	0.06		
GR11-55-2c.FIN2	0.294	0.016	0.0402	0.0014	0.0527	0.0021	0.79359	254.2	8.6	261.7	12	2.9	337	18	720	0.09		
GR11-55-3k.FIN2	0.298	0.016	0.042	0.0014	0.0518	0.0021	0.77887	265	8.4	264.5	13	-0.2	272	18	308	0.14		
GR11-55-4c.FIN2	0.304	0.017	0.0433	0.0016	0.051	0.0020	0.8448	273.2	10	269.4	13	-1.4	231	18	221	0.21		
GR11-55-5g.FIN2	0.43	0.027	0.0444	0.0018	0.0708	0.0030	0.781	279.8	11	364	19	23.1	942	35	646	0.14		
GR11-55-2e.FIN2	0.32	0.018	0.0447	0.0015	0.0513	0.0020	0.83746	281.9	9	281.2	13	-0.2	264	17	359	0.06		
GR11-55-2b.FIN2	0.326	0.018	0.0455	0.0016	0.052	0.0021	0.79009	286.5	10	285.8	14	-0.2	303	20	242	0.09		
GR11-55-3j.FIN2	0.341	0.018	0.0464	0.0015	0.0532	0.0021	0.75392	292.1	9	297.9	14	1.9	326	17	353	0.38		
Deformed Biotite Pegmatite Dike (GR11-54)																		
GR11-54-1.FIN2	0.035	0.008	0.0039	0.0005	0.067	0.0190	0.001	24.8	3.6	34.4	7.9	27.9	1350	580	143	0.25		
GR11-54-26.FIN2	0.024	0.004	0.004	0.0007	0.0487	0.0098	0.079843	25.5	4.5	24.4	3.9	-4.5	220	240	416	0.07		
GR11-54-29.FIN2	0.028	0.006	0.004	0.0007	0.0504	0.0092	0.4819	25.8	4.5	28	5.5	7.9	530	360	384	0.09		
GR11-54-35.FIN2	0.022	0.004	0.004	0.0007	0.0478	0.0056	0.50305	25.9	4.7	22.3	4.1	-16.1	156	80	319	0.05		
GR11-54-17.FIN2	0.027	0.004	0.0044	0.0006	0.0468	0.0050	0.75876	28.2	3.6	26.9	4.3	-4.8	280	100	243	0.11		
GR11-54-33.FIN2	0.032	0.003	0.0045	0.0003	0.0522	0.0043	0.47598	29.2	2.1	32.1	3.1	9.7	546	89	277	0.12		
GR11-54-37.FIN2	0.031	0.003	0.0045	0.0004	0.0478	0.0028	0.69071	29.2	2.4	30.7	2.9	4.9	406	92	4350	0.02		
GR11-54-43.FIN2	0.034	0.006	0.0046	0.0006	0.0579	0.0085	0.48021	29.3	3.7	34.1	5.6	14.1	820	230	274	0.10		
GR11-54-22.FIN2	0.026	0.002	0.0046	0.0009	0.0418	0.0088	0.30396	29.5	5.9	25.6	2.4	-15.2	406	92	4350	0.02		
GR11-54-31.FIN2	0.033	0.002	0.0047	0.0003	0.0533	0.0029	0.53044	30.4	1.9	33.2	2.1	8.4	437	54	222.1	0.00		
GR11-54-19.FIN2	0.067	0.014	0.0047	0.0005	0.106	0.0300	0.001	30.4	3.3	66	13	53.9	1890	240	50.6	0.03		

Sample and Analysis Number	Isotopic Ratios										Dates (Ma)					Concentrator		
	207Pb/ ± 2SE		206Pb/ ± 2SE		207Pb/ ± 2SE		Rho: 206/238 vs 207/235		206Pb/ ± 2SE		207Pb/ ± 2SE		% Discord.		207Pb/ ± 2SE		U	
	235U (abs)	238U (abs)	235U (abs)	238U (abs)	206Pb (abs)	206Pb (abs)	207Pb (abs)	207Pb (abs)	206Pb (abs)	235U (abs)	235U (abs)	238U (abs)	238U (abs)	6/38-7/35	206Pb (abs)	206Pb (abs)	Th/U (ppm)	
GR11-54-39.FIN2	0.031	0.005	0.0048	0.0012	0.0466	0.0072	0.62039	30.7	7.6	31.1	5.1	1.3	620	250	393	0.12		
GR11-54-11.FIN2	0.045	0.013	0.0048	0.0014	0.0704	0.0080	0.99796	31.1	9.3	45	13	30.9	830	170	290	0.10		
GR11-54-4.FIN2	0.055	0.005	0.0049	0.0004	0.0815	0.0086	0.47929	31.5	2.6	55.6	4.9	43.3	1230	150	161	0.23		
GR11-54-34.FIN2	0.027	0.007	0.0049	0.0012	0.0397	0.0056	0.80792	31.7	7.6	26.7	6.5	-18.7			360	0.09		
GR11-54-21.FIN2	0.036	0.006	0.005	0.0006	0.0514	0.0067	0.58238	32.2	3.9	35.9	5.8	10.3	640	210	204	0.12		
GR11-54-12.FIN2	0.042	0.006	0.0051	0.0007	0.062	0.0120	0.1329	32.5	4.7	41.2	5.8	21.1	660	150	175	0.21		
GR11-54-8.FIN2	0.056	0.014	0.0051	0.0010	0.084	0.0190	0.30971	32.7	6.1	55	13	40.5	1500	190	55	0.10		
GR11-54-13.FIN2	0.037	0.004	0.0052	0.0004	0.0518	0.0068	0.27676	33.4	2.5	37	4.2	9.7	520	100	191	0.20		
GR11-54-40.FIN2	0.054	0.007	0.0052	0.0005	0.0788	0.0092	0.42016	33.4	2.9	53.6	6.5	37.7	1200	130	103	0.24		
GR11-54-7.FIN2	0.037	0.004	0.0052	0.0004	0.0522	0.0034	0.80565	33.6	2.5	37	4	9.2	500	97	670	0.09		
GR11-54-16.FIN2	0.046	0.011	0.0052	0.0008	0.062	0.0110	0.68516	33.6	5.2	46	10	27.0	660	150	251	0.13		
GR11-54-14.FIN2	0.036	0.005	0.0053	0.0007	0.0498	0.0036	0.92384	33.9	4.5	35.9	5.2	5.6	215	54	939	0.01		
GR11-54-3.FIN2	0.059	0.022	0.0057	0.0016	0.075	0.0130	0.93293	36	10	58	21	37.9	910	250	213	0.03		
GR11-54-20.FIN2	0.046	0.007	0.0057	0.0003	0.0589	0.0074	0.001	36.3	2.2	45.7	6.3	20.6	720	170	114	0.18		
GR11-54-18.FIN2	0.072	0.012	0.0057	0.0005	0.093	0.0180	0.040078	36.9	3.3	71	11	48.0	1480	190	219	0.15		
GR11-54-32.FIN2	0.041	0.007	0.0062	0.0009	0.0509	0.0054	0.78126	39.9	5.6	41.1	6.8	2.9	310	94	292	0.15		
GR11-54-27.FIN2	0.082	0.010	0.0075	0.0009	0.082	0.0100	0.63657	47.9	5.7	79.5	9.8	39.7	1290	160	168	0.21		
GR11-54-25.FIN2	0.108	0.007	0.0159	0.0009	0.0487	0.0020	0.73821	101.6	5.6	103.3	6.5	1.6	363	53	375	0.23		
GR11-54-15.FIN2	0.191	0.012	0.0268	0.0019	0.0524	0.0025	0.64937	171	12	176.8	10	3.3	470	72	164	0.32		
GR11-54-30.FIN2	0.264	0.013	0.0359	0.0018	0.0529	0.0019	0.68547	227	11	238	11	4.6	393	48	134.7	0.37		
GR11-54-6.FIN2	4.08	0.290	0.0399	0.0025	0.734	0.0350	0.69398	252	16	1646	61	84.7	4628	40	17.67	0.04		
GR11-54-23.FIN2	0.309	0.017	0.0402	0.0021	0.0549	0.0019	0.74685	254	13	272	13	6.6	458	52	177	0.24		
GR11-54-9.FIN2	0.292	0.023	0.0404	0.0028	0.0538	0.0034	0.65805	255	18	261	18	2.3	550	81	122.6	0.42		
GR11-54-42.FIN2	0.308	0.021	0.0415	0.0025	0.0526	0.0021	0.81058	262	16	272	16	3.7	482	65	504	0.05		
GR11-54-41.FIN2	0.327	0.021	0.0447	0.0025	0.0533	0.0020	0.75359	282	16	287	16	1.7	517	60	1780	0.06		
GR11-54-2.FIN2	0.342	0.027	0.047	0.0028	0.0526	0.0025	0.80496	296	17	294	20	-0.7	581	68	600	0.22		
GR11-54-24.FIN2	0.356	0.021	0.0491	0.0030	0.053	0.0020	0.77454	308	18	309	15	0.3	376	54	405	0.30		
GR11-54-28.FIN2	0.37	0.029	0.0524	0.0036	0.0515	0.0025	0.77565	329	22	323	21	-1.9	342	61	227	0.31		
GR11-54-5.FIN2	0.813	0.047	0.0935	0.0054	0.0629	0.0026	0.70013	575	32	600	26	4.2	724	70	154	1.34		
GR11-54-36.FIN2	4.72	0.300	0.308	0.0170	0.1073	0.0041	0.74808	1722	84	1744	55	1.3	1775	64	220	0.89		
GR11-54-10.FIN2	4.99	0.360	0.312	0.0190	0.115	0.0049	0.79034	1749	94	1803	64	3.0	1873	68	243	0.26		
Biotite Granite Pod (GR11-27)																		
GR11-27-9.FIN2	0.031	0.002	0.0045	0.0003	0.0505	0.0029	0.68332	29.1	2	31.2	2.4	6.7	404	73	1108	0.49		

Sample and Analysis Number	Isotopic Ratios				Dates (Ma)				Concentrator							
	207Pb/ ± 2SE 235U (abs)	206Pb/ ± 2SE 238U (abs)	207Pb/ ± 2SE 206Pb (abs)	Rho: 206/238 vs 207/235	206Pb/ ± 2SE 238U (abs)	207Pb/ ± 2SE 235U (abs)	207Pb/ ± 2SE 206Pb (abs)	% Discord. 6/38-7/35	207Pb/ ± 2SE 206Pb (abs)	U (ppm)	Th/U					
GR11-27-13.FIN2	0.03	0.001	0.0045	0.0002	0.0478	0.0017	0.55157	29.11	1.1	29.9	1.3	2.6	228	36	715	0.17
GR11-27-5.FIN2	0.031	0.002	0.0046	0.0002	0.0504	0.0021	0.48312	29.3	1.2	31.2	1.5	6.1	277	51	1250	0.45
GR11-27-15.FIN2	0.03	0.002	0.0046	0.0002	0.0475	0.0021	0.70369	29.5	1.6	29.8	1.9	1.0	334	59	787	0.27
GR11-27-41.FIN2	0.03	0.001	0.0046	0.0002	0.0476	0.0016	0.41239	29.63	1.1	30.08	1.1	1.5	210	32	473	0.29
GR11-27-26.FIN2	0.029	0.002	0.0046	0.0003	0.0455	0.0019	0.70938	29.8	1.6	28.8	1.7	-3.5	224	37	719	0.23
GR11-27-32.FIN2	0.032	0.001	0.0046	0.0002	0.0499	0.0016	0.6556	29.8	1.4	31.6	1.3	5.7	219	36	1357	0.18
GR11-27-34.FIN2	0.03	0.002	0.0047	0.0002	0.0459	0.0012	0.83078	29.9	1.5	29.6	1.6	-1.0	178	30	1970	0.43
GR11-27-16.FIN2	0.03	0.003	0.0047	0.0003	0.0454	0.0024	0.79664	30	1.8	29.5	2.5	-1.7	190	59	321	0.08
GR11-27-43.FIN2	0.033	0.002	0.0047	0.0003	0.0501	0.0038	0.2824	30.3	1.8	32.7	2.5	7.3	290	110	354	0.30
GR11-27-3.FIN2	0.033	0.002	0.0048	0.0003	0.0504	0.0024	0.62294	30.5	1.9	32.7	2.1	6.7	265	62	1260	0.20
GR11-27-39.FIN2	0.03	0.002	0.0047	0.0003	0.048	0.0037	0.4319	30.5	2	30.3	2.4	-0.7	200	130	2210	0.14
GR11-27-44.FIN2	0.032	0.001	0.0047	0.0002	0.0499	0.0016	0.55095	30.51	1.1	32	1.3	4.7	252	34	532	0.45
GR11-27-36.FIN2	0.039	0.003	0.0048	0.0003	0.0633	0.0041	0.47503	30.6	1.8	39.2	2.8	21.9	781	83	256	0.01
GR11-27-35.FIN2	0.03	0.002	0.0048	0.0004	0.0462	0.0026	0.68344	30.8	2.2	29.8	2.2	-3.4	218	48	439	0.17
GR11-27-8.FIN2	0.026	0.004	0.0049	0.0006	0.0401	0.0042	0.65746	31.4	3.9	26.2	3.5	-19.8	160	190	259	0.13
GR11-27-25.FIN2	0.04	0.002	0.005	0.0003	0.0596	0.0042	0.25044	31.9	1.8	39.4	2.2	19.0	545	89	981	0.28
GR11-27-1.FIN2	0.034	0.004	0.005	0.0007	0.0487	0.0025	0.90554	32.1	4.3	34.1	4.2	5.9	193	96	1300	0.01
GR11-27-12.FIN2	0.04	0.003	0.0051	0.0003	0.0566	0.0036	0.35541	32.7	2	40.2	3	18.7	517	77	292	0.07
GR11-27-42.FIN2	0.033	0.003	0.0051	0.0003	0.0476	0.0028	0.65714	32.7	2.2	32.4	2.7	-0.9	226	59	566	0.07
GR11-27-23.FIN2	0.035	0.009	0.0052	0.0011	0.0512	0.0053	0.8651	33.3	7.1	34.8	8.6	4.3	490	150	750	0.16
GR11-27-6.FIN2	0.032	0.006	0.0052	0.0007	0.0445	0.0046	0.81704	33.3	4.5	32.2	5.4	-3.4	357	93	506	0.01
GR11-27-40.FIN2	0.037	0.002	0.0053	0.0003	0.0499	0.0024	0.64047	33.8	2	36.8	2.4	8.2	378	55	458	0.00
GR11-27-31.FIN2	0.031	0.004	0.0053	0.0007	0.0437	0.0042	0.71094	34.1	4.3	31	4	-10.0	302	55	253	0.10
GR11-27-17.FIN2	0.036	0.003	0.0057	0.0004	0.0468	0.0050	0.10761	36.6	2.7	36.3	3	-0.8	160	180	510	0.06
GR11-27-30.FIN2	0.053	0.005	0.0079	0.0007	0.0517	0.0035	0.77563	50.6	4.3	52.7	5.1	4.0	427	85	312	0.39
GR11-27-27.FIN2	0.092	0.011	0.0126	0.0015	0.0531	0.0030	0.75541	80.8	9.7	88.9	10	9.1	334	80	480	0.13
GR11-27-33.FIN2	0.166	0.013	0.0223	0.0015	0.0532	0.0023	0.78851	142.9	9.3	156	12	8.4	468	63	92.8	0.36
GR11-27-38.FIN2	0.211	0.013	0.028	0.0016	0.0542	0.0016	0.85753	178.7	9.8	194	11	7.9	420	41	319	0.12
GR11-27-18.FIN2	0.246	0.010	0.0327	0.0013	0.0552	0.0014	0.68268	207	7.9	222.8	8.1	7.1	405	35	450	0.31
GR11-27-7.FIN2	0.268	0.022	0.0375	0.0026	0.0514	0.0026	0.75175	237	16	239	17	0.8	428	87	433	0.33
GR11-27-10.FIN2	0.28	0.025	0.0399	0.0028	0.0507	0.0026	0.77635	252	18	251	20	-0.4	442	76	2130	0.12
GR11-27-37.FIN2	0.297	0.012	0.04	0.0019	0.0534	0.0014	0.77691	253	12	264.3	9	4.3	381	35	990	0.22
GR11-27-29.FIN2	0.315	0.012	0.0434	0.0017	0.053	0.0011	0.81564	273.5	11	277.7	9	1.5	324	26	354	0.25

Sample and Analysis Number	Isotopic Ratios				Dates (Ma)				Concentrator							
	207Pb/ ± 2SE 235U (abs)	206Pb/ ± 2SE 238U (abs)	207Pb/ ± 2SE 206Pb (abs)	Rho: 206/238 vs 207/235	206Pb/ ± 2SE 238U (abs)	207Pb/ ± 2SE 235U (abs)	207Pb/ ± 2SE 206Pb (abs)	% Discord. 6/38-7/35	207Pb/ ± 2SE 206Pb (abs)	U (ppm)	Th/U	U (ppm)				
GR11-27-22.FIN2	0.355	0.022	0.0492	0.0026	0.0523	0.0020	0.75413	309	16	306	16	-1.0	421	50	460	0.32
GR11-27-20.FIN2	0.569	0.036	0.0716	0.0041	0.0566	0.0024	0.70899	445	25	452	23	1.5	499	57	168	0.24
Biotite Granite Dike (GR11-32)																
GR11-32_4_e.FIN2	0.025	0.001	0.0039	0.0001	0.0448	0.0016	0.73204	25.22	0.96	24.5	1.2	-2.9	236	50	1910	0.02
GR11-32_5_a.FIN2	0.027	0.001	0.0041	0.0002	0.0478	0.0017	0.67075	26.2	1.1	27	1.1	3.0	231	41	4030	0.11
GR11-32_6_d.FIN2	0.027	0.001	0.0042	0.0002	0.0468	0.0014	0.75536	27	1.1	27.1	1.2	0.4	215	42	4140	0.12
GR11-32_5_d.FIN2	0.027	0.001	0.0042	0.0002	0.0467	0.0015	0.63761	27.2	1	27.1	1.2	-0.4	214	41	881	0.28
GR11-32_7_b.FIN2	0.027	0.001	0.0043	0.0002	0.0474	0.0015	0.6693	27.3	1.1	27	0.98	-1.1	173	28	1570	0.17
GR11-32_5_c.FIN2	0.028	0.002	0.0043	0.0002	0.0471	0.0024	0.49825	27.5	1.2	28.2	1.5	2.5	297	93	568	0.16
GR11-32_4_a.FIN2	0.027	0.001	0.0043	0.0002	0.0465	0.0015	0.71193	27.6	1.1	27.2	1.1	-1.5	194	33	730	0.17
GR11-32_8_a.FIN2	0.028	0.001	0.0043	0.0002	0.0477	0.0016	0.64746	27.6	1.2	28.6	1.2	3.5	221	51	7460	0.09
GR11-32_7_a.FIN2	0.029	0.001	0.0043	0.0001	0.0475	0.0016	0.52614	27.83	0.98	28.7	1.1	3.0	273	38	1178	0.15
GR11-32_2_g.FIN2	0.028	0.001	0.0044	0.0002	0.046	0.0013	0.76627	28	1.2	27.9	1.2	-0.4	172	35	2280	0.37
GR11-32_7_c.FIN2	0.035	0.002	0.0044	0.0002	0.0582	0.0020	0.64927	28.03	0.99	35	1.6	19.9	574	59	1910	0.15
GR11-32_2_c.FIN2	0.028	0.001	0.0044	0.0002	0.0461	0.0011	0.81587	28.09	1	28.3	1.2	0.7	152	30	5430	0.10
GR11-32_6_b.FIN2	0.028	0.001	0.0044	0.0002	0.047	0.0014	0.78912	28.2	1.5	28.4	1.5	0.7	128	33	4000	0.09
GR11-32_6_e.FIN2	0.028	0.001	0.0044	0.0002	0.046	0.0014	0.69759	28.4	1.1	28.3	1.2	-0.4	237	38	2600	0.05
GR11-32_6_c.FIN2	0.029	0.001	0.0044	0.0002	0.0483	0.0017	0.75169	28.4	1.3	29.1	1.3	2.4	310	43	1690	0.16
GR11-32_3_a.FIN2	0.031	0.001	0.0044	0.0002	0.0494	0.0019	0.635	28.5	1.3	30.7	1.5	7.2	327	47	1009	0.21
GR11-32_4_c.FIN2	0.029	0.001	0.0045	0.0002	0.0478	0.0015	0.69863	28.6	1.2	29.2	1.2	2.1	202	32	2520	0.14
GR11-32_1_d.FIN2	0.028	0.001	0.0045	0.0002	0.0459	0.0016	0.61012	28.85	1.1	28.4	1.1	-1.6	202	31	728	0.39
GR11-32_4_f.FIN2	0.031	0.001	0.0045	0.0002	0.049	0.0014	0.6833	29	1.1	30.8	1.2	5.8	227	31	3430	0.09
GR11-32_1_e.FIN2	0.031	0.001	0.0045	0.0002	0.0488	0.0014	0.69691	29.18	1	31	1.2	5.9	252	32	1474	0.02
GR11-32_6_a.FIN2	0.03	0.002	0.0046	0.0002	0.048	0.0019	0.64654	29.3	1.2	30.3	1.5	3.3	331	54	886	0.16
GR11-32_2_b.FIN2	0.032	0.002	0.0046	0.0002	0.0494	0.0023	0.64494	29.3	1.4	31.5	2	7.0	454	80	3810	0.09
GR11-32_4_d.FIN2	0.035	0.002	0.0046	0.0002	0.056	0.0026	0.52002	29.3	1.1	34.8	1.8	15.8	513	54	456	0.32
GR11-32_4_g.FIN2	0.031	0.001	0.0046	0.0002	0.048	0.0017	0.686	29.4	1.1	30.7	1.5	4.2	327	52	1262	0.01
GR11-32_4_b.FIN2	0.033	0.002	0.0046	0.0002	0.0513	0.0029	0.56148	29.6	1.5	33	2.3	10.3	480	91	913	0.27
GR11-32_1_a.FIN2	0.032	0.001	0.0046	0.0002	0.0492	0.0017	0.7015	29.78	1	31.7	1.5	6.1	280	38	1820	0.02
GR11-32_2_d.FIN2	0.032	0.002	0.0046	0.0002	0.0512	0.0026	0.66885	29.8	1.4	32	2.1	6.9	533	67	1274	0.32
GR11-32_5_b.FIN2	0.032	0.001	0.0047	0.0002	0.0494	0.0014	0.68289	29.91	0.95	31.6	1.2	5.3	256	34	3750	0.18
GR11-32_7_d.FIN2	0.033	0.003	0.0047	0.0003	0.0499	0.0034	0.58322	30.2	1.7	32.8	2.6	7.9	259	64	728	0.20
GR11-32_2_f.FIN2	0.032	0.001	0.0047	0.0002	0.0489	0.0015	0.72725	30.3	1.4	31.8	1.5	4.7	187	32	1460	0.18

Sample and Analysis Number	Isotopic Ratios				Dates (Ma)				Concentrator							
	207Pb/235U (abs)	206Pb/238U (abs)	207Pb/206Pb (abs)	Rho: 206/238 vs 207/235	206Pb/238U (abs)	207Pb/235U (abs)	207Pb/206Pb (abs)	% Discord. 6/38-7/35	207Pb/206Pb (abs)	±2SE	U (ppm)	Th/U				
GR11-32_2_a.FIN2	0.033	0.002	0.0047	0.0002	0.0479	0.0023	0.41349	30.4	1.3	32.5	1.9	6.5	198	76	1561	0.02
GR11-32_6_f.FIN2	0.051	0.004	0.0048	0.0002	0.0787	0.0053	0.48795	30.5	1.2	50	3.7	39.0	1120	110	392	0.26
GR11-32_1_b.FIN2	0.034	0.003	0.0048	0.0002	0.0512	0.0029	0.73954	31.1	1.5	33.5	2.7	7.2	382	61	1490	0.08
GR11-32_1_c.FIN2	0.036	0.002	0.0051	0.0002	0.0508	0.0019	0.6155	32.5	1.3	36.2	1.7	10.2	391	45	1026	0.03
GR11-32_7_e.FIN2	0.078	0.004	0.0111	0.0005	0.0506	0.0015	0.85187	70.9	3.2	76.2	3.5	7.0	263	41	3190	0.18
GR11-32_2_e.FIN2	0.364	0.015	0.0475	0.0019	0.0563	0.0016	0.77125	299	12	315	11	5.1	471	43	917	0.07
Garnet-Biotite-Muscovite Pegmatite Dike (GR11-34)																
GR11-34_7_a.FIN2	0.029	0.001	0.0044	0.0002	0.047	0.0017	0.59435	28.2	1.1	28.6	1.4	1.4	207	34	578	0.09
GR11-34_4_d.FIN2	0.031	0.002	0.0045	0.0002	0.0498	0.0029	0.14557	28.82	1	31.3	1.6	7.9	404	57	220	0.01
GR11-34_2_c.FIN2	0.029	0.001	0.0045	0.0002	0.0457	0.0018	0.54587	28.99	1	28.5	1.2	-1.7	248	50	503	0.09
GR11-34_6_e.FIN2	0.028	0.002	0.0045	0.0002	0.0452	0.0023	0.62617	29	1.4	27.7	1.6	-4.7	159	38	637	0.11
GR11-34_2_b.FIN2	0.027	0.001	0.0045	0.0002	0.0443	0.0016	0.61295	29.2	1.1	26.9	1.1	-8.6	172	52	680	0.32
GR11-34_3_e.FIN2	0.029	0.001	0.0046	0.0002	0.0451	0.0014	0.7194	29.4	1.2	29.1	1.3	-1.0	181	39	1270	0.11
GR11-34_3_b.FIN2	0.032	0.001	0.0046	0.0002	0.0494	0.0018	0.55185	29.64	1	31.5	1.4	5.9	303	41	506	0.10
GR11-34_5_b.FIN2	0.029	0.001	0.0046	0.0002	0.0468	0.0019	0.51679	29.7	1.2	29	1.3	-2.4	196	36	429	0.09
GR11-34_10_a.FIN2	0.029	0.001	0.0046	0.0002	0.0459	0.0019	0.49661	29.7	1.1	29.1	1.3	-2.1	233	37	387	0.34
GR11-34_2_a.FIN2	0.033	0.001	0.0046	0.0001	0.0514	0.0017	0.48819	29.76	0.96	32.5	1.1	8.4	311	39	557	0.09
GR11-34_10_c.FIN2	0.033	0.003	0.0046	0.0002	0.048	0.0039	0.45613	29.8	1.5	33.1	3.3	10.0	258	84	257	0.09
GR11-34_3_c.FIN2	0.031	0.001	0.0047	0.0002	0.0479	0.0016	0.69789	30.1	1.2	31.1	1.5	3.2	265	44	1130	0.00
GR11-34_9_d.FIN2	0.03	0.001	0.0047	0.0002	0.0464	0.0017	0.45339	30.2	1.1	29.9	1.1	-1.0	197	43	584	0.10
GR11-34_9_c.FIN2	0.035	0.002	0.0047	0.0002	0.0548	0.0024	0.37894	30.31	1.1	35.1	1.6	13.6	449	48	316	0.08
GR11-34_9_b.FIN2	0.03	0.001	0.0047	0.0002	0.0457	0.0017	0.56071	30.4	1.1	30	1.2	-1.3	92	32	1917	0.12
GR11-34_8_b.FIN2	0.032	0.002	0.0047	0.0002	0.049	0.0030	0.26316	30.4	1.3	31.9	1.9	4.7	280	66	384	0.43
GR11-34_8_a.FIN2	0.031	0.002	0.0048	0.0002	0.0463	0.0025	0.59317	30.6	1.4	30.5	2.1	-0.3	257	59	314	0.53
GR11-34_8_c.FIN2	0.116	0.030	0.0058	0.0005	0.149	0.0360	0.51173	37.5	3.1	115	29	67.4	2170	500	253	0.08
GR11-34_10_b.FIN2	0.044	0.006	0.0064	0.0007	0.0519	0.0049	0.74831	40.9	4.6	43.6	6.2	6.2	320	110	331	0.18
GR11-34_9_e.FIN2	0.072	0.005	0.0101	0.0006	0.0507	0.0022	0.77799	64.5	4	70.9	5.1	9.0	271	48	538	0.08
GR11-34_6_f.FIN2	0.163	0.012	0.0122	0.0006	0.0929	0.0069	0.41587	78.4	4.1	152	11	48.4	1480	130	299	0.44
GR11-34_4_a.FIN2	0.185	0.008	0.0239	0.0010	0.0539	0.0015	0.66669	152.4	6.5	171.8	7	11.3	361	34	1717	0.12
GR11-34_7_b.FIN2	0.202	0.010	0.0286	0.0013	0.0511	0.0017	0.74429	181.9	7.9	186	8.1	2.2	294	47	4410	0.08
GR11-34_5_a.FIN2	0.216	0.012	0.0295	0.0015	0.0509	0.0016	0.83488	187.1	9.5	197.8	10	5.4	291	47	883	0.77
GR11-34_3_a.FIN2	0.265	0.012	0.036	0.0014	0.0535	0.0019	0.67902	229	8.6	239.1	9.5	4.2	394	49	723	0.23
GR11-34_5_d.FIN2	0.264	0.012	0.0377	0.0017	0.0511	0.0015	0.78442	238.4	11	236.8	9.3	-0.7	265	32	1580	0.52

Sample and Analysis Number	Isotopic Ratios				Dates (Ma)				Concentrator								
	207Pb/235U (abs)	± 2SE	206Pb/238U (abs)	± 2SE	206Pb/238U (abs)	± 2SE	207Pb/235U (abs)	± 2SE	207Pb/206Pb (abs)	± 2SE	U (ppm)	Th/U					
GR11-34_5_c.FIN2	0.273	0.009	0.038	0.0013	0.0528	0.0014	0.72408	0.72408	240.3	8.3	245.6	7.3	2.2	362	30	944	0.28
GR11-34_4_c.FIN2	0.287	0.012	0.0395	0.0016	0.0526	0.0013	0.7999	0.7999	249.6	10	256.6	8.9	2.7	328	39	1133	0.18
GR11-34_9_a.FIN2	0.285	0.010	0.0396	0.0013	0.0513	0.0013	0.73047	0.73047	250.1	8.3	253.7	8.1	1.4	283	38	446	0.15
GR11-34_3_f.FIN2	0.298	0.012	0.0428	0.0016	0.0511	0.0013	0.80542	0.80542	270.1	10	265.6	9.5	-1.7	273	37	297	0.29
GR11-34_4_b.FIN2	0.314	0.011	0.0438	0.0016	0.0511	0.0014	0.70479	0.70479	275.9	9.7	276.8	8.5	0.3	285	39	1084	0.20
GR11-34_2_d.FIN2	0.317	0.012	0.0442	0.0016	0.0525	0.0013	0.80934	0.80934	278.5	9.6	278.5	8.9	0.0	330	29	841	0.23
GR11-34_3_d.FIN2	0.314	0.013	0.0444	0.0017	0.0502	0.0012	0.83045	0.83045	279.6	11	276.1	10	-1.3	284	37	444	0.27
Two-mica Pegmatite Dike (VER11-1)																	
VER11-1_1_i.FIN2	0.031	0.003	0.0038	0.0003	0.0584	0.0039	0.60358	0.60358	24.2	1.7	31.1	2.6	22.2	622	85	383	0.55
VER11-1_1_h.FIN2	0.026	0.002	0.0038	0.0002	0.0477	0.0021	0.74653	0.74653	24.68	1.4	26.24	1.8	5.9	223	31	634	0.37
VER11-1_1_7_b.FIN2	0.027	0.002	0.0039	0.0002	0.0508	0.0024	0.62377	0.62377	24.86	1.4	27.4	1.9	9.3	272	33	181.8	0.44
VER11-1_1_j.FIN2	0.027	0.002	0.0039	0.0002	0.0478	0.0021	0.78338	0.78338	24.94	1.4	26.65	1.8	6.4	224	28	5560	0.08
VER11-1_3_g.FIN2	0.025	0.002	0.0039	0.0002	0.0455	0.0020	0.73889	0.73889	25	1.4	24.85	1.7	-0.6	99	22	361	0.13
VER11-1_1_e.FIN2	0.026	0.002	0.0039	0.0003	0.047	0.0033	0.44573	0.44573	25	1.7	25.6	2.3	2.3	346	90	136	0.24
VER11-1_3_e.FIN2	0.025	0.002	0.0039	0.0003	0.0463	0.0025	0.73216	0.73216	25.1	1.7	25.3	2.2	0.8	192	50	234.7	0.47
VER11-1_4_b.FIN2	0.025	0.002	0.0039	0.0002	0.0465	0.0020	0.78981	0.78981	25.12	1.4	24.73	1.7	-1.6	186	30	892	0.11
VER11-1_4_i.FIN2	0.029	0.003	0.0039	0.0003	0.0535	0.0032	0.74631	0.74631	25.3	1.9	29.3	2.5	13.7	447	64	154.1	0.40
VER11-1_4_f.FIN2	0.025	0.002	0.0039	0.0002	0.0446	0.0024	0.61329	0.61329	25.4	1.6	25.1	2	-1.2	213	49	137.9	0.36
VER11-1_3_c.FIN2	0.024	0.002	0.004	0.0003	0.0448	0.0024	0.7572	0.7572	25.6	1.9	24.4	2	-4.9	145	42	243	0.15
VER11-1_1_b.FIN2	0.024	0.002	0.004	0.0002	0.0458	0.0022	0.67358	0.67358	25.6	1.5	24.52	1.7	-4.4	195	38	304	0.50
VER11-1_5_a.FIN2	0.026	0.002	0.004	0.0003	0.0469	0.0021	0.88578	0.88578	25.6	1.9	25.8	2	0.8	163	28	4520	0.02
VER11-1_6_h.FIN2	0.027	0.002	0.004	0.0002	0.0477	0.0023	0.59556	0.59556	25.61	1.4	26.7	1.9	4.1	229	42	245.1	0.31
VER11-1_4_e.FIN2	0.026	0.002	0.004	0.0002	0.0474	0.0024	0.67981	0.67981	25.8	1.5	26	1.8	0.8	295	49	268	0.38
VER11-1_2_c.FIN2	0.028	0.002	0.004	0.0002	0.05	0.0024	0.64675	0.64675	25.84	1.4	27.8	2	7.1	304	46	274.5	0.47
VER11-1_4_j.FIN2	0.032	0.003	0.004	0.0002	0.0553	0.0033	0.43324	0.43324	25.84	1.4	31.8	2.6	18.7	500	85	269.1	0.52
VER11-1_2_b.FIN2	0.03	0.002	0.004	0.0002	0.0545	0.0031	0.44808	0.44808	25.96	1.5	30	2.4	13.5	512	70	250	0.50
VER11-1_6_c.FIN2	0.032	0.003	0.004	0.0004	0.0593	0.0040	0.71283	0.71283	26	2.6	31.6	3	17.7	562	89	1011	0.01
VER11-1_1_a.FIN2	0.028	0.002	0.0041	0.0002	0.0505	0.0026	0.4842	0.4842	26.07	1.3	27.6	1.9	5.5	292	44	266.7	0.54
VER11-1_1_k.FIN2	0.026	0.002	0.0041	0.0002	0.0457	0.0020	0.80802	0.80802	26.1	1.5	26.3	1.9	0.8	167	29	2390	0.11
VER11-1_6_f.FIN2	0.027	0.002	0.0041	0.0003	0.0464	0.0027	0.62436	0.62436	26.1	1.9	26.8	2.3	2.6	222	32	107	0.32
VER11-1_6_b.FIN2	0.036	0.006	0.0041	0.0002	0.0599	0.0059	0.85738	0.85738	26.2	1.6	36	5.5	27.2	600	120	213	0.35
VER11-1_1_d.FIN2	0.025	0.002	0.0041	0.0003	0.0453	0.0023	0.69494	0.69494	26.3	1.7	24.8	1.9	-6.0	188	47	256	0.23
VER11-1_4_a.FIN2	0.028	0.002	0.0041	0.0002	0.0502	0.0022	0.74975	0.74975	26.45	1.4	27.9	1.9	5.2	244	31	930	0.22
VER11-1_3_f.FIN2	0.027	0.002	0.0041	0.0002	0.047	0.0022	0.7668	0.7668	26.5	1.5	27.2	2	2.6	160	32	397	0.17

Sample and Analysis Number	Isotopic Ratios					Dates (Ma)					Concentrator					
	207Pb/235U (abs)	± 2SE	206Pb/238U (abs)	± 2SE	207Pb/206Pb (abs)	± 2SE	206Pb/238U (abs)	± 2SE	207Pb/206Pb (abs)	± 2SE	% Discord.	207Pb/206Pb (abs)	± 2SE	U (ppm)	Th/U	
VER11-1_6_a.FIN2	0.043	0.004	0.0042	0.0003	0.073	0.0050	0.67552	26.7	1.7	42.2	4.1	36.7	960	110	560	0.13
VER11-1_2_a.FIN2	0.031	0.002	0.0042	0.0002	0.0548	0.0025	0.57196	26.72	1.3	31.3	2.1	14.6	402	41	276.3	0.61
VER11-1_4_c.FIN2	0.028	0.004	0.0042	0.0005	0.048	0.0024	0.95351	26.9	3.1	28.4	4.3	5.3	149	58	1650	0.13
VER11-1_4_g.FIN2	0.074	0.007	0.0043	0.0002	0.1238	0.0080	0.57228	27.5	1.6	72.2	6.1	61.9	1999	85	373	0.27
VER11-1_3_j.FIN2	0.037	0.004	0.0051	0.0005	0.0504	0.0039	0.76595	32.9	3.1	36.6	4.1	10.1	297	56	339	0.06
VER11-1_3_b.FIN2	0.199	0.016	0.0274	0.0019	0.0528	0.0022	0.9063	174.3	12	183.6	14	5.1	332	35	409	0.14
VER11-1_3_a.FIN2	0.21	0.025	0.03	0.0035	0.0505	0.0040	0.68816	190	22	192	21	1.0	350	160	1880	0.05
VER11-1_4_d.FIN2	0.226	0.017	0.0326	0.0020	0.0501	0.0020	0.89362	206.6	12	207.4	14	0.4	256	30	775	0.54
VER11-1_7_a.FIN2	0.309	0.023	0.0339	0.0022	0.0664	0.0030	0.80898	215	14	272	18	21.0	834	45	534	1.19
VER11-1_4_h.FIN2	0.279	0.020	0.0374	0.0022	0.0526	0.0021	0.88342	236.4	14	249.3	16	5.2	350	27	830	0.23
VER11-1_2_f.FIN2	0.283	0.020	0.0387	0.0023	0.0509	0.0021	0.83657	244.9	14	254.3	17	3.7	276	35	369	0.28
VER11-1_6_d.FIN2	0.301	0.022	0.0409	0.0026	0.0519	0.0023	0.80008	258	16	268	17	3.7	275	45	865	0.06
VER11-1_3_h.FIN2	0.318	0.022	0.0433	0.0025	0.0512	0.0021	0.83276	273	16	279.7	17	2.4	292	35	326	0.23
VER11-1_6_g.FIN2	0.318	0.023	0.0435	0.0026	0.0524	0.0023	0.799	275	16	282.4	17	2.6	356	34	930	0.00
VER11-1_1_e.FIN2	0.318	0.024	0.0474	0.0027	0.0494	0.0021	0.81034	298	17	279	18	-6.8	240	38	991	0.38
VER11-1_3_i.FIN2	0.375	0.026	0.0473	0.0026	0.0556	0.0027	0.53484	298	16	322.6	19	7.6	456	33	37.7	0.31
VER11-1_2_g.FIN2	0.473	0.033	0.0541	0.0032	0.0595	0.0024	0.83716	339	20	394	23	14.0	582	33	632	0.04
VER11-1_3_d.FIN2	0.602	0.066	0.0737	0.0072	0.0595	0.0024	0.97555	455	43	467	43	2.6	603	28	286	0.18
VER11-1_1_f.FIN2	0.605	0.041	0.0771	0.0041	0.0574	0.0025	0.68621	479	24	483	26	0.8	533	33	770	0.18
VER11-1_5_b.FIN2	0.736	0.056	0.0899	0.0061	0.0593	0.0027	0.83169	554	36	560	32	1.1	591	43	156	0.32
VER11-1_2_e.FIN2	0.797	0.056	0.0906	0.0049	0.0612	0.0025	0.77352	558	29	593	32	5.9	641	37	209	0.44
VER11-1_6_e.FIN2	0.745	0.053	0.0917	0.0053	0.0575	0.0023	0.88667	565	32	562	31	-0.5	518	27	365	0.34
VER11-1_1_g.FIN2	0.862	0.063	0.0979	0.0056	0.0642	0.0029	0.80977	601	33	633	33	5.1	763	37	162	0.69
VER11-1_2_d.FIN2	1.214	0.097	0.1357	0.0088	0.0641	0.0025	0.91385	820	50	803	44	-2.1	760	31	169.5	0.52
Leucogranite NOV11-1A																
NOV11-1A_6j.FIN2	0.484	0.049	0.0397	0.0017	0.09	0.0083	0.42854	250.9	10	395	33	36.5	1300	170	868	0.22
NOV11-1A_2d.FIN2	0.285	0.010	0.04	0.0014	0.052	0.0012	0.84219	252.5	8.6	255.2	8.2	1.1	283	16	563	0.67
NOV11-1A_5i.FIN2	0.295	0.011	0.0401	0.0015	0.0527	0.0014	0.75327	253.4	9.5	262.7	8.7	3.5	338	25	675	0.08
NOV11-1A_5r.FIN2	0.295	0.011	0.0403	0.0014	0.0523	0.0013	0.7611	254.5	8.7	262.6	8.8	3.1	313	22	577	0.08
NOV11-1A_4b.FIN2	0.289	0.011	0.0403	0.0017	0.0513	0.0015	0.67636	255.9	11	257.3	8.8	0.5	263	25	213	0.17
NOV11-1A_4d.FIN2	0.295	0.011	0.0406	0.0014	0.0518	0.0012	0.76832	256.5	8.5	262.4	8.4	2.2	268	18	389	0.63
NOV11-1A_7l.FIN2	0.286	0.012	0.0408	0.0016	0.0508	0.0015	0.65817	257.9	9.8	255.2	9.8	-1.1	221	30	1065	0.09
NOV11-1A_8b.FIN2	0.287	0.010	0.0409	0.0016	0.0516	0.0015	0.75353	258.5	9.8	255.9	8	-1.0	285	25	686	0.75
NOV11-1A_8d.FIN2	0.284	0.012	0.0411	0.0017	0.051	0.0015	0.18324	259.6	11	253.7	9.6	-2.3	256	27	162.1	0.33

Sample and Analysis Number	Isotopic Ratios				Rho: 206/238 vs 207/235	Dates (Ma)				Concentrator						
	207Pb/235U (abs)	± 2SE	206Pb/238U (abs)	± 2SE		207Pb/206Pb (abs)	± 2SE	206Pb/238U (abs)	± 2SE	207Pb/206Pb (abs)	± 2SE	U (ppm)	Th/U			
NOV11-1A_6l.FIN2	0.302	0.013	0.0412	0.0017	0.0528	0.0018	0.64605	260	11	269.1	10	3.4	325	32	1400	0.08
NOV11-1A_4m.FIN: 0.299	0.299	0.011	0.0412	0.0015	0.0514	0.0013	0.75113	260.3	9.2	265.6	8.4	2.0	264	16	287.5	0.42
NOV11-1A_5q.FIN2	0.299	0.011	0.0412	0.0015	0.052	0.0013	0.82778	260.3	9.4	265.4	8.8	1.9	277	19	1004	0.08
NOV11-1A_7d.FIN2	0.293	0.011	0.0412	0.0016	0.0515	0.0013	0.80053	260.3	9.6	261.2	8.6	0.3	265	19	402	0.38
NOV11-1A_6h.FIN2	0.303	0.011	0.0413	0.0014	0.053	0.0013	0.73235	260.5	8.7	268.6	8.3	3.0	337	17	526	0.20
NOV11-1A_3l.FIN2	0.305	0.011	0.0413	0.0016	0.0528	0.0015	0.70048	260.6	9.8	270.2	8.8	3.6	313	22	515	0.06
NOV11-1A_1b.FIN2	0.298	0.012	0.0414	0.0016	0.0522	0.0014	0.73094	261.4	9.9	264.1	9.1	1.0	314	20	89	0.33
NOV11-1A_3m.FIN: 0.296	0.296	0.010	0.0414	0.0014	0.0518	0.0012	0.79599	261.4	8.6	263.1	8.1	0.6	283	16	632	0.07
NOV11-1A_3n.FIN2	0.302	0.011	0.0414	0.0015	0.0521	0.0013	0.78848	261.4	9	267.8	8.5	2.4	288	20	365	0.17
NOV11-1A_3g.FIN2	0.305	0.011	0.0414	0.0014	0.0521	0.0014	0.71545	261.5	8.9	270.4	8.7	3.3	293	21	371	0.44
NOV11-1A_3i.FIN2	0.297	0.011	0.0414	0.0014	0.0512	0.0013	0.69232	261.5	8.5	264	8.4	0.9	248	22	179.8	0.56
NOV11-1A_8j.FIN2	0.296	0.012	0.0414	0.0018	0.051	0.0015	0.80267	261.7	11	262.6	9.5	0.3	252	33	310.2	0.25
NOV11-1A_6d.FIN2	0.301	0.011	0.0415	0.0014	0.0526	0.0014	0.74674	261.9	8.8	266.8	9	1.8	309	21	420.3	0.45
NOV11-1A_6o.FIN2	0.316	0.013	0.0415	0.0016	0.0554	0.0015	0.62752	262.2	9.7	278.5	10	5.9	423	28	194.5	0.48
NOV11-1A_5d.FIN2	0.304	0.011	0.0416	0.0014	0.0522	0.0013	0.65099	262.4	8.8	269.8	8.3	2.7	294	20	218.4	0.32
NOV11-1A_8l.FIN2	0.304	0.014	0.0416	0.0017	0.0547	0.0015	0.7843	262.5	11	268.7	11	2.3	399	31	1864	0.55
NOV11-1A_3d.FIN2	0.299	0.011	0.0417	0.0015	0.0515	0.0013	0.78721	263.1	9	265.7	8.2	1.0	262	15	430	0.42
NOV11-1A_3a.FIN2	0.298	0.010	0.0417	0.0014	0.0512	0.0012	0.72426	263.5	8.5	264.4	8	0.3	254	17	1470	0.03
NOV11-1A_8e.FIN2	0.296	0.010	0.0418	0.0015	0.0524	0.0013	0.78224	263.6	9	263.4	8	-0.1	309	16	595	0.12
NOV11-1A_7o.FIN2	0.3	0.013	0.0418	0.0017	0.0528	0.0014	0.929	263.7	11	266	10	0.9	330	22	808	0.24
NOV11-1A_6q.FIN2	0.299	0.011	0.0418	0.0016	0.0516	0.0014	0.7302	263.8	9.7	265.6	8.7	0.7	272	25	551	0.28
NOV11-1A_1f.FIN2	0.292	0.011	0.0418	0.0015	0.0506	0.0014	0.65441	263.9	9.2	259.9	8.5	-1.5	236	26	823	0.08
NOV11-1A_4j.FIN2	0.301	0.011	0.0418	0.0015	0.0513	0.0013	0.7686	263.9	9.3	267	8.2	1.2	275	22	724	0.18
NOV11-1A_5c.FIN2	0.305	0.011	0.0418	0.0014	0.0526	0.0013	0.67482	264.2	8.6	270.6	8.6	2.4	319	20	277.9	0.50
NOV11-1A_3o.FIN2	0.665	0.054	0.0419	0.0018	0.1126	0.0076	0.58835	264.6	11	513	33	48.4	1850	110	281	0.60
NOV11-1A_5n.FIN2	0.299	0.013	0.0419	0.0018	0.051	0.0015	0.76928	264.7	11	265.5	10	0.3	247	29	331.3	0.43
NOV11-1A_7e.FIN2	0.301	0.011	0.0419	0.0015	0.0522	0.0013	0.72799	264.8	9.1	267.3	8.4	0.9	290	17	355	0.43
NOV11-1A_6t.FIN2	0.299	0.011	0.042	0.0015	0.0517	0.0012	0.81321	265.2	9.1	265	8.4	-0.1	284	17	602	0.16
NOV11-1A_8c.FIN2	0.296	0.011	0.042	0.0016	0.0512	0.0013	0.78584	265.2	9.8	263.2	8.5	-0.8	267	19	429	0.06
NOV11-1A_1g.FIN2	0.305	0.011	0.0421	0.0015	0.0532	0.0013	0.75702	265.5	9.2	270.3	8.3	1.8	334	21	842	0.06
NOV11-1A_3j.FIN2	0.303	0.011	0.0421	0.0015	0.0523	0.0013	0.7523	265.6	9	269.1	8.2	1.3	297	19	828	0.10
NOV11-1A_5j.FIN2	0.306	0.011	0.0421	0.0016	0.0531	0.0017	0.56699	265.6	9.9	270.8	8.4	1.9	334	34	436	0.16
NOV11-1A_6m.FIN: 0.3	0.3	0.011	0.0421	0.0015	0.0509	0.0012	0.80385	265.9	9.3	266.1	8.7	0.1	239	18	1097	0.09
NOV11-1A_3b.FIN2	0.303	0.010	0.0422	0.0015	0.0517	0.0014	0.6182	266.1	9.1	268.5	7.9	0.9	276	20	174.8	0.44

Sample and Analysis Number	Isotopic Ratios				Rho: 206/238 vs 207/235	Dates (Ma)				Concentrator						
	207Pb/235U	± 2SE	206Pb/238U	± 2SE		207Pb/206Pb	± 2SE	207Pb/235U	± 2SE	207Pb/206Pb	± 2SE	U (ppm)	Th/U			
NOV11-1A_3p.FIN2	0.309	0.011	0.0422	0.0014	0.0525	0.0013	0.70876	266.4	8.7	273.1	8.2	2.5	302	23	336	0.28
NOV11-1A_5a.FIN2	0.306	0.012	0.0422	0.0015	0.0515	0.0013	0.77761	266.6	9.5	270.4	9.1	1.4	289	22	1310	0.12
NOV11-1A_5o.FIN2	0.313	0.011	0.0423	0.0017	0.0535	0.0014	0.73921	266.7	10	275.9	8.8	3.3	364	20	492	0.21
NOV11-1A_7a.FIN2	0.303	0.011	0.0423	0.0015	0.052	0.0013	0.73777	266.7	9.5	268.8	8.4	0.8	287	21	1240	0.06
NOV11-1A_6n.FIN2	0.306	0.012	0.0421	0.0017	0.0519	0.0014	0.80909	266.8	10	270.4	9.3	1.3	309	23	259	0.46
NOV11-1A_4n.FIN2	0.306	0.010	0.0423	0.0015	0.0517	0.0013	0.73518	266.9	9.2	271.1	8.1	1.5	276	17	486	0.15
NOV11-1A_6g.FIN2	0.315	0.013	0.0423	0.0017	0.0545	0.0016	0.72798	267	10	278	9.7	4.0	389	19	289	0.12
NOV11-1A_5e.FIN2	0.388	0.018	0.0423	0.0016	0.066	0.0023	0.70336	267.2	9.6	332.1	13	19.5	800	50	702	0.13
NOV11-1A_8h.FIN2	0.299	0.011	0.0423	0.0015	0.0523	0.0014	0.73252	267.2	9.4	265.6	8.7	-0.6	289	27	187.5	0.54
NOV11-1A_9j.FIN2	0.305	0.012	0.0423	0.0016	0.0525	0.0016	0.64887	267.2	9.7	269.8	9.5	1.0	308	26	195	0.49
NOV11-1A_5m.FIN2	0.441	0.038	0.0423	0.0017	0.0732	0.0051	0.63251	267.3	11	367	26	27.2	950	130	280	0.23
NOV11-1A_6c.FIN2	0.743	0.075	0.0424	0.0020	0.1234	0.0092	0.74606	267.5	12	547	41	51.1	2030	100	412	0.38
NOV11-1A_4f.FIN2	0.305	0.011	0.0424	0.0015	0.0514	0.0015	0.66058	267.7	9.2	269.8	8.7	0.8	267	22	469	0.24
NOV11-1A_5b.FIN2	0.312	0.012	0.0424	0.0017	0.0529	0.0014	0.82284	267.7	11	275.1	9.4	2.7	321	27	779	0.13
NOV11-1A_3f.FIN2	0.31	0.012	0.0425	0.0016	0.0525	0.0014	0.76451	268	9.8	273.6	9.2	2.0	300	26	493	0.45
NOV11-1A_4p.FIN2	0.312	0.014	0.0425	0.0019	0.0526	0.0013	0.90984	268	12	274.9	11	2.5	301	14	1008	0.08
NOV11-1A_2e.FIN2	0.306	0.011	0.0424	0.0015	0.0521	0.0013	0.76655	268.4	9.2	271.6	8.7	1.2	289	16	589	0.19
NOV11-1A_2a.FIN2	0.324	0.014	0.0426	0.0015	0.055	0.0019	0.4462	268.8	9	284.4	11	5.5	401	46	118	0.26
NOV11-1A_2c.FIN2	0.313	0.012	0.0426	0.0015	0.0536	0.0014	0.72773	268.8	9.5	276.6	8.9	2.8	352	20	750	0.13
NOV11-1A_5p.FIN2	0.303	0.011	0.0426	0.0015	0.0519	0.0015	0.59483	268.9	9.5	269.3	8.5	0.1	277	24	1670	0.07
NOV11-1A_9h.FIN2	0.343	0.027	0.0426	0.0022	0.0586	0.0035	0.73657	269	14	298	20	9.7	560	100	272	0.15
NOV11-1A_7b.FIN2	0.298	0.012	0.0427	0.0020	0.052	0.0016	0.76987	269.2	12	264.6	9.5	-1.7	284	34	943	0.09
NOV11-1A_8a.FIN2	0.316	0.010	0.0427	0.0014	0.0542	0.0014	0.55827	269.2	8.7	279.4	8	3.7	365	18	581	0.28
NOV11-1A_9d.FIN2	0.3	0.012	0.0427	0.0016	0.0525	0.0013	0.79629	269.2	9.8	265.6	9	-1.4	301	20	301	0.51
NOV11-1A_9e.FIN2	0.287	0.011	0.0427	0.0017	0.049	0.0013	0.76844	269.3	11	256.7	9	-4.9	178	22	2145	0.08
NOV11-1A_6r.FIN2	0.313	0.012	0.0428	0.0017	0.0532	0.0015	0.65728	269.9	10	277.4	9.3	2.7	329	28	1027	0.16
NOV11-1A_8f.FIN2	0.302	0.014	0.0427	0.0022	0.0521	0.0016	0.7232	270	14	269.4	11	-0.2	283	31	160.7	0.30
NOV11-1A_8k.FIN2	0.303	0.011	0.0428	0.0015	0.0521	0.0013	0.80586	270.2	9.2	268.4	8.6	-0.7	285	15	989	0.05
NOV11-1A_3c.FIN2	0.306	0.010	0.0428	0.0015	0.052	0.0013	0.72133	270.3	9.3	270.7	8.1	0.1	286	16	371	0.46
NOV11-1A_7s.FIN2	0.298	0.011	0.0427	0.0016	0.0509	0.0013	0.78434	270.3	9.9	264.6	8.9	-2.2	246	20	225	0.32
NOV11-1A_4g.FIN2	0.405	0.025	0.0429	0.0019	0.067	0.0034	0.36758	270.4	12	344	18	21.4	840	100	1036	0.19
NOV11-1A_6e.FIN2	0.304	0.011	0.0429	0.0014	0.0515	0.0014	0.67197	270.4	8.9	269.1	8.4	-0.5	261	24	163.4	0.38
NOV11-1A_8g.FIN2	0.304	0.011	0.0429	0.0015	0.0522	0.0013	0.78188	270.5	9.1	269.2	8.2	-0.5	290	18	528	0.26
NOV11-1A_2b.FIN2	0.322	0.013	0.0429	0.0015	0.0545	0.0016	0.4968	270.7	9	283.2	9.8	4.4	392	34	344	0.30

Sample and Analysis Number	Isotopic Ratios					Rho: 206/238 vs 207/235	Dates (Ma)					Concentrator				
	207Pb/235U	± 2SE	206Pb/238U	± 2SE	207Pb/206Pb		± 2SE	206Pb/238U	± 2SE	207Pb/206Pb	± 2SE	206Pb/238U	± 2SE	207Pb/206Pb	± 2SE	U (ppm)
NOV11-1A_7n.FIN2	0.312	0.011	0.0429	0.0015	0.0528	0.0014	0.70648	270.7	9.6	275.6	8.6	1.8	323	22	445	0.23
NOV11-1A_5g.FIN2	0.315	0.011	0.043	0.0015	0.0531	0.0014	0.73167	271.2	9.4	278.3	8.7	2.6	334	20	1038	0.17
NOV11-1A_4l.FIN2	0.305	0.011	0.043	0.0017	0.0515	0.0013	0.77327	271.3	11	269.9	8.5	-0.5	268	18	476.6	0.24
NOV11-1A_7h.FIN2	0.304	0.011	0.043	0.0016	0.0513	0.0013	0.76733	271.5	10	269	8.8	-0.9	260	22	124	0.34
NOV11-1A_1e.FIN2	0.4	0.023	0.043	0.0015	0.066	0.0032	0.39781	271.6	9.3	339	16	19.9	798	79	581	0.01
NOV11-1A_1c.FIN2	0.318	0.012	0.0431	0.0016	0.0542	0.0015	0.72591	271.7	10	279.8	9.6	2.9	392	28	729	0.11
NOV11-1A_5h.FIN2	0.305	0.012	0.0431	0.0016	0.0508	0.0012	0.80472	271.9	10	271.1	8.9	-0.3	235	22	332.8	0.20
NOV11-1A_1h.FIN2	0.304	0.011	0.0431	0.0015	0.0514	0.0014	0.69965	272	9	269	8.5	-1.1	271	24	312	0.41
NOV11-1A_7g.FIN2	0.309	0.011	0.0432	0.0015	0.0523	0.0013	0.71164	272.2	9.2	273.4	8.4	0.4	308	21	731	0.18
NOV11-1A_4k.FIN2	0.317	0.012	0.0432	0.0016	0.0523	0.0014	0.92032	272.7	9.7	279.2	9.5	2.3	308	27	133.8	0.36
NOV11-1A_8i.FIN2	0.308	0.012	0.0432	0.0016	0.0515	0.0014	0.71257	272.8	10	272.1	9.2	-0.3	269	27	741	0.19
NOV11-1A_4c.FIN2	0.33	0.013	0.0433	0.0015	0.0545	0.0014	0.75788	273	9.4	289.5	9.8	5.7	393	24	697	0.16
NOV11-1A_6s.FIN2	0.31	0.011	0.0433	0.0016	0.0521	0.0013	0.75621	273	9.7	274.6	8.8	0.6	280	18	1676	0.05
NOV11-1A_9c.FIN2	0.31	0.011	0.0433	0.0015	0.0524	0.0014	0.72275	273.2	9.6	274.2	8.6	0.4	316	21	534	0.14
NOV11-1A_7m.FIN2	0.312	0.013	0.0434	0.0017	0.051	0.0014	0.77174	273.9	11	275.1	9.9	0.4	271	23	657	0.13
NOV11-1A_4h.FIN2	0.316	0.011	0.0434	0.0015	0.052	0.0013	0.74858	274	9.1	279.4	8.7	1.9	283	18	226	0.24
NOV11-1A_9f.FIN2	0.302	0.012	0.0434	0.0016	0.0514	0.0014	0.67145	274	9.8	267.9	9	-2.3	275	25	155.1	0.34
NOV11-1A_7k.FIN2	0.315	0.012	0.0435	0.0016	0.0527	0.0014	0.69734	274.4	9.8	278.1	9.3	1.3	316	20	205.9	0.37
NOV11-1A_1a.FIN2	0.308	0.012	0.0436	0.0015	0.052	0.0014	0.67677	274.9	9.3	272.6	8.9	-0.8	294	24	717	0.07
NOV11-1A_6f.FIN2	0.325	0.012	0.0436	0.0015	0.0536	0.0014	0.70404	274.9	9.2	285.4	9.1	3.7	355	27	272	0.35
NOV11-1A_9a.FIN2	0.311	0.011	0.0436	0.0015	0.0521	0.0013	0.71301	275	9.4	275.2	8.7	0.1	303	20	474	0.33
NOV11-1A_4i.FIN2	0.697	0.064	0.0437	0.0017	0.1109	0.0073	0.91669	275.8	11	523	34	47.3	1810	110	143	0.31
NOV11-1A_9b.FIN2	0.31	0.011	0.0438	0.0015	0.052	0.0014	0.58833	276.5	9.1	274.7	8.6	-0.7	303	23	130.8	0.68
NOV11-1A_9g.FIN2	0.306	0.011	0.0439	0.0016	0.051	0.0013	0.76718	277.2	9.9	270.8	8.7	-2.4	254	21	331	0.19
NOV11-1A_6a.FIN2	0.325	0.012	0.0442	0.0017	0.0529	0.0015	0.63482	278.6	10	285.4	8.9	2.4	319	30	1400	0.10
NOV11-1A_3e.FIN2	0.326	0.012	0.0442	0.0014	0.0528	0.0014	0.65305	278.7	8.8	286.3	8.8	2.7	319	20	460	0.24
NOV11-1A_7r.FIN2	0.294	0.011	0.0443	0.0017	0.0489	0.0013	0.78272	279.4	11	261.3	8.9	-6.9	176	24	2916	0.07
NOV11-1A_1i.FIN2	0.355	0.018	0.0443	0.0018	0.0585	0.0024	0.45861	279.5	11	308	14	9.3	578	53	1288	0.04
NOV11-1A_5s.FIN2	0.316	0.012	0.0444	0.0017	0.0512	0.0014	0.65517	279.9	11	279.1	9.2	-0.3	282	26	1810	0.06
NOV11-1A_4a.FIN2	0.328	0.012	0.0447	0.0015	0.0523	0.0013	0.72626	281.7	9.6	287.6	8.9	2.1	305	20	840	0.16
NOV11-1A_7f.FIN2	0.322	0.012	0.0447	0.0015	0.0526	0.0015	0.67771	281.7	9.4	283	9.5	0.5	302	23	1671	0.14
NOV11-1A_4r.FIN2	0.337	0.013	0.0447	0.0018	0.0541	0.0015	0.76738	282	11	295.5	9.9	4.6	377	23	1077	0.05
NOV11-1A_5k.FIN2	0.334	0.012	0.045	0.0016	0.0538	0.0014	0.77025	283.8	9.9	292	9	2.8	353	21	543.4	0.19
NOV11-1A_7j.FIN2	0.32	0.012	0.0452	0.0017	0.0512	0.0014	0.64748	284.9	11	281.3	8.9	-1.3	261	30	2220	0.06

Sample and Analysis Number	Isotopic Ratios				Dates (Ma)				Concentrator							
	207Pb/235U	206Pb/238U	207Pb/206Pb	± 2SE	206Pb/238U	207Pb/206Pb	± 2SE	207Pb/235U	206Pb/238U	± 2SE	U (ppm)	Th/U				
NOV11-1A_6k.FIN2	0.322	0.012	0.0453	0.0018	0.0516	0.0013	0.80125	285.3	11	283.1	9.4	-0.8	269	21	1143	0.21
NOV11-1A_3h.FIN2	0.328	0.013	0.0453	0.0019	0.0509	0.0014	0.76776	285.6	12	287.5	10	0.7	264	23	244	0.27
NOV11-1A_4e.FIN2	0.354	0.012	0.0454	0.0014	0.0558	0.0013	0.71255	286.3	8.8	308	9.3	7.0	430	20	560	0.08
NOV11-1A_3k.FIN2	0.348	0.016	0.0456	0.0018	0.0546	0.0014	0.92264	287.3	11	302.6	12	5.1	397	25	871	0.07
NOV11-1A_9i.FIN2	0.382	0.013	0.0459	0.0016	0.0612	0.0017	0.60297	289.2	10	328.8	9.8	12.0	651	32	242.6	0.93
NOV11-1A_6i.FIN2	0.331	0.012	0.0459	0.0015	0.0518	0.0012	0.86188	289.4	9.4	290.4	9.1	0.3	287	12	775	0.24
NOV11-1A_1j.FIN2	1.42	0.260	0.0465	0.0031	0.204	0.0240	0.92279	293	19	799	87	63.3	2720	170	165.4	0.28
NOV11-1A_7i.FIN2	0.407	0.024	0.0468	0.0019	0.0623	0.0027	0.4014	294.8	12	344	16	14.3	678	70	290	0.29
NOV11-1A_6p.FIN2	0.346	0.017	0.0472	0.0024	0.0522	0.0020	0.75878	297	15	301.5	13	1.5	304	49	1253	0.11
NOV11-1A_7c.FIN2	0.319	0.023	0.0472	0.0037	0.0491	0.0014	0.99438	297	22	279	16	-6.5	182	25	2457	0.03
NOV11-1A_1d.FIN2	0.58	0.066	0.0488	0.0019	0.0831	0.0070	0.72362	307.3	12	436	34	29.5	1160	130	211	0.17
NOV11-1A_7q.FIN2	0.36	0.016	0.0562	0.0026	0.0473	0.0016	0.72764	352	16	312	12	-12.8	125	26	3410	0.17
NOV11-1A_5f.FIN2	0.529	0.021	0.0641	0.0024	0.0592	0.0016	0.68181	400	15	429.9	14	7.0	586	33	1541	0.11
NOV11-1A_6b.FIN2	0.647	0.035	0.0799	0.0045	0.0576	0.0016	0.90484	495	27	509	22	2.8	522	23	992	0.27
NOV11-1A_2f.FIN2	0.674	0.023	0.0811	0.0027	0.0597	0.0015	0.67271	503.9	16	522.5	14	3.6	586	18	439	0.34
NOV11-1A_4o.FIN2	0.777	0.027	0.0904	0.0030	0.0613	0.0015	0.78098	558	18	583.2	16	4.3	649	17	739	0.37
NOV11-1A_2g.FIN2	0.799	0.029	0.0946	0.0034	0.061	0.0015	0.84186	582	20	595.2	16	2.2	645	17	379	0.47
NOV11-1A_4q.FIN2	0.862	0.029	0.099	0.0032	0.0624	0.0015	0.7373	608.1	19	632.3	16	3.8	705	15	429	0.32
NOV11-1A_5l.FIN2	1.398	0.054	0.1369	0.0050	0.0716	0.0018	0.77768	827	28	885	23	6.6	983	19	433	0.39
NOV11-1A_7p.FIN2	3.62	0.160	0.2364	0.0100	0.1113	0.0031	0.84175	1371	53	1550	35	11.5	1832	23	416	0.07
Leucogranite NOV11-1B																
NOV11-1B_1f.FIN2	0.022	0.002	0.0034	0.0003	0.0462	0.0041	0.63172	21.6	2.2	21.8	1.9	0.9	219	71	886	0.12
NOV11-1B_4d.FIN2	0.023	0.002	0.0035	0.0001	0.047	0.0022	0.68359	22.63	0.94	23.23	1.6	2.6	137	24	882	0.25
NOV11-1B_9h.FIN2	0.024	0.002	0.0035	0.0002	0.0495	0.0027	0.67689	22.68	1.2	24.2	1.8	6.3	315	50	3430	0.22
NOV11-1B_9v.FIN2	0.023	0.002	0.0036	0.0002	0.0481	0.0025	0.55679	22.93	1	23.33	1.7	1.7	294	59	433	0.12
NOV11-1B_11s.FIN	0.024	0.002	0.0036	0.0001	0.0494	0.0023	0.70536	22.99	0.99	24.26	1.5	5.2	231	30	547	0.23
NOV11-1B_4e.FIN2	0.024	0.002	0.0036	0.0001	0.0469	0.0024	0.55367	23	0.95	23.77	1.6	3.2	166	32	309	0.12
NOV11-1B_10m.FT	0.024	0.002	0.0036	0.0002	0.0487	0.0023	0.64711	23.02	1	23.75	1.6	3.1	201	40	1414	0.19
NOV11-1B_1c.FIN2	0.025	0.002	0.0036	0.0002	0.05	0.0027	0.48612	23.03	1.1	24.8	1.8	7.1	258	57	466	0.13
NOV11-1B_2n.FIN2	0.024	0.002	0.0036	0.0001	0.0489	0.0022	0.70312	23.15	0.92	24.46	1.6	5.4	193	29	1255	0.15
NOV11-1B_2h.FIN2	0.023	0.001	0.0036	0.0001	0.0457	0.0023	0.48604	23.16	0.9	23.29	1.5	0.6	143	37	462.7	0.14
NOV11-1B_7g.FIN2	0.023	0.002	0.0036	0.0001	0.0465	0.0023	0.52799	23.2	0.92	23.42	1.5	0.9	125	28	584	0.25
NOV11-1B_2d.FIN2	0.024	0.001	0.0036	0.0002	0.0472	0.0023	0.65953	23.24	1	23.66	1.5	1.8	161	25	550	0.16
NOV11-1B_4i.FIN2	0.024	0.002	0.0036	0.0001	0.0482	0.0022	0.75382	23.26	0.92	24.22	1.6	4.0	159	23	793	0.18

Sample and Analysis Number	Isotopic Ratios					Dates (Ma)					Concentrator				
	207Pb/235U	± 2SE	206Pb/238U	± 2SE	207Pb/206Pb	± 2SE	Rho: 206/238 vs 207/235	238U	± 2SE	207Pb/235U	± 2SE	207Pb/206Pb	± 2SE	U (ppm)	Th/U
NOV11-1B_6e.FIN2	0.024	0.002	0.0036	0.0002	0.0492	0.0023	0.72006	23.27	1	24.28	1.6	167	22	1583	0.19
NOV11-1B_6l.FIN2	0.023	0.002	0.0036	0.0002	0.0469	0.0023	0.58338	23.32	0.95	23.26	1.5	199	35	656	0.20
NOV11-1B_2f.FIN2	0.024	0.002	0.0036	0.0002	0.0483	0.0022	0.79874	23.38	1	23.77	1.5	166	26	1908	0.16
NOV11-1B_7e.FIN2	0.025	0.002	0.0036	0.0001	0.0471	0.0023	0.41264	23.39	0.9	24.64	1.5	210	36	482	0.27
NOV11-1B_6i.FIN2	0.032	0.002	0.0036	0.0002	0.065	0.0039	0.32655	23.41	1.2	31.8	2.2	779	77	1200	0.37
NOV11-1B_4f.FIN2	0.026	0.002	0.0036	0.0002	0.0507	0.0024	0.63377	23.43	0.95	25.96	1.7	271	30	692	0.24
NOV11-1B_1In.FIN	0.027	0.002	0.0037	0.0002	0.0514	0.0024	0.67673	23.47	1	26.7	1.7	306	31	1382	0.12
NOV11-1B_8j.FIN2	0.025	0.002	0.0037	0.0002	0.0498	0.0024	0.58325	23.49	1	24.69	1.7	229	32	518	0.27
NOV11-1B_5h.FIN2	0.024	0.002	0.0037	0.0001	0.0484	0.0025	0.53651	23.52	0.88	23.94	1.6	269	44	306.8	0.15
NOV11-1B_9f.FIN2	0.025	0.002	0.0037	0.0002	0.0492	0.0023	0.81018	23.53	1.2	25.4	1.8	246	39	3400	0.18
NOV11-1B_10a.FIN	0.023	0.002	0.0037	0.0002	0.0465	0.0022	0.66865	23.57	1	23.37	1.5	175	26	440	0.15
NOV11-1B_8n.FIN2	0.024	0.002	0.0037	0.0001	0.0477	0.0027	0.40266	23.57	0.99	23.9	1.7	440	110	557	0.18
NOV11-1B_4h.FIN2	0.024	0.002	0.0037	0.0002	0.0474	0.0022	0.64107	23.58	1	24.34	1.6	186	37	918	0.19
NOV11-1B_2b.FIN2	0.031	0.002	0.0037	0.0001	0.0612	0.0029	0.58015	23.59	0.97	31.43	2	658	47	983	0.16
NOV11-1B_7o.FIN2	0.025	0.002	0.0037	0.0001	0.0514	0.0027	0.5264	23.6	0.99	25.5	1.8	380	61	731	0.23
NOV11-1B_4g.FIN2	0.024	0.002	0.0037	0.0002	0.0474	0.0026	0.42651	23.62	1	24.27	1.6	269	40	318	0.11
NOV11-1B_8l.FIN2	0.023	0.002	0.0037	0.0002	0.0472	0.0023	0.69206	23.63	1.2	23.2	1.6	164	29	517	0.14
NOV11-1B_1a.FIN2	0.024	0.002	0.0037	0.0001	0.048	0.0024	0.67776	23.63	0.99	24.3	1.7	223	30	323.3	0.16
NOV11-1B_2i.FIN2	0.024	0.002	0.0037	0.0002	0.0474	0.0021	0.78983	23.63	1	24.3	1.6	151	26	2580	0.09
NOV11-1B_1g.FIN2	0.027	0.002	0.0037	0.0002	0.053	0.0026	0.64023	23.68	1	26.55	1.7	333	40	677	0.13
NOV11-1B_1b.FIN2	0.023	0.002	0.0037	0.0002	0.0464	0.0035	0.57203	23.7	1.5	23.4	2.2	440	100	218	0.12
NOV11-1B_7m.FIN	0.024	0.002	0.0037	0.0002	0.0482	0.0030	0.60748	23.7	1.4	24	2	221	40	417	0.24
NOV11-1B_7d.FIN2	0.054	0.007	0.0037	0.0002	0.1006	0.0100	0.54296	23.7	1.1	53.2	6.5	1570	170	681	0.43
NOV11-1B_2e.FIN2	0.023	0.002	0.0037	0.0001	0.0467	0.0022	0.65722	23.71	0.96	23.57	1.5	177	32	698	0.16
NOV11-1B_2c.FIN2	0.024	0.002	0.0037	0.0002	0.0471	0.0022	0.70013	23.72	1	23.96	1.6	187	28	536	0.15
NOV11-1B_3d.FIN2	0.029	0.002	0.0037	0.0002	0.0551	0.0038	0.463	23.75	1.2	28.5	2.4	481	95	318.3	0.11
NOV11-1B_9n.FIN2	0.027	0.002	0.0037	0.0002	0.0527	0.0027	0.54225	23.76	1.1	26.55	1.8	396	48	537	0.11
NOV11-1B_7f.FIN2	0.024	0.002	0.0037	0.0002	0.0465	0.0024	0.56734	23.81	1	24.04	1.6	121	25	552	0.24
NOV11-1B_3g.FIN2	0.024	0.002	0.0037	0.0002	0.0477	0.0026	0.54658	23.89	1.1	24.46	1.7	237	35	227	0.14
NOV11-1B_10o.FIN	0.025	0.002	0.0037	0.0002	0.0478	0.0022	0.64664	23.89	0.94	24.77	1.6	190	32	667	0.14
NOV11-1B_7h.FIN2	0.024	0.001	0.0037	0.0001	0.0467	0.0020	0.83325	23.9	0.93	23.89	1.5	95	16	1630	0.25
NOV11-1B_2a.FIN2	0.024	0.002	0.0037	0.0002	0.0487	0.0024	0.64967	23.91	1	24.43	1.6	217	31	416	0.21
NOV11-1B_1e.FIN2	0.025	0.002	0.0037	0.0002	0.0487	0.0023	0.72309	23.93	1	24.83	1.6	209	33	593	0.16
NOV11-1B_7a.FIN2	0.025	0.002	0.0037	0.0001	0.048	0.0022	0.66098	23.94	0.94	24.83	1.6	185	24	604	0.16

Sample and Analysis Number	Isotopic Ratios					Dates (Ma)					Concentrator					
	207Pb/235U	206Pb/238U	207Pb/206Pb	± 2SE	Rho: 206/238 vs 207/235	238U	206Pb/238U	207Pb/206Pb	± 2SE	207Pb/206Pb	± 2SE	6/38-7/35	206Pb/207Pb	± 2SE	U (ppm)	Th/U
NOV11-1B_1d.FIN2	0.024	0.002	0.0037	0.0002	0.0478	0.0024	0.59871	23.97	1	24.16	1.6	0.8	258	42	344	0.15
NOV11-1B_5k.FIN2	0.024	0.001	0.0037	0.0001	0.0455	0.0021	0.60251	23.98	0.92	23.65	1.5	-1.4	122	24	774	0.16
NOV11-1B_3l.FIN2	0.025	0.002	0.0037	0.0002	0.0486	0.0025	0.7441	24	1.4	24.99	1.7	4.0	190	37	544	0.14
NOV11-1B_11r.FIN	0.024	0.002	0.0037	0.0002	0.0469	0.0022	0.70465	24.04	0.99	24.49	1.6	1.8	160	27	711	0.28
NOV11-1B_2k.FIN2	0.024	0.002	0.0037	0.0002	0.0463	0.0024	0.60204	24.09	1.1	24.08	1.6	0.0	219	36	213.9	0.10
NOV11-1B_11h.FIN	0.024	0.001	0.0037	0.0002	0.0467	0.0021	0.74409	24.09	1	24.35	1.5	1.1	121	21	1616	0.27
NOV11-1B_6g.FIN2	0.025	0.002	0.0038	0.0002	0.0498	0.0023	0.79456	24.12	1	24.97	1.7	3.4	247	35	1121	0.21
NOV11-1B_10x.FIN	0.025	0.002	0.0038	0.0002	0.048	0.0024	0.54676	24.13	0.99	24.98	1.7	3.4	249	59	855	0.17
NOV11-1B_10n.FIN	0.024	0.002	0.0038	0.0002	0.0472	0.0022	0.70161	24.14	0.98	24.31	1.6	0.7	176	27	1234	0.16
NOV11-1B_5b.FIN2	0.025	0.002	0.0038	0.0002	0.0478	0.0026	0.54764	24.15	1	25.53	1.8	5.4	209	31	171	0.17
NOV11-1B_3e.FIN2	0.034	0.003	0.0038	0.0002	0.064	0.0039	0.61273	24.18	1	33.7	2.7	28.2	716	89	505	0.12
NOV11-1B_4a.FIN2	0.024	0.002	0.0038	0.0004	0.0462	0.0031	0.79942	24.2	2.3	24.3	2.3	0.4	255	31	2120	0.06
NOV11-1B_5c.FIN2	0.028	0.002	0.0038	0.0001	0.0533	0.0036	0.26335	24.2	0.94	27.6	2.1	12.3	343	56	188.1	0.17
NOV11-1B_5i.FIN2	0.024	0.002	0.0038	0.0001	0.0466	0.0021	0.7342	24.21	0.95	23.96	1.5	-1.0	126	26	1876	0.27
NOV11-1B_6f.FIN2	0.024	0.001	0.0038	0.0002	0.0475	0.0027	0.60558	24.21	1.2	24	1.4	-0.9	81	24	1038	0.04
NOV11-1B_10b.FIN	0.025	0.002	0.0038	0.0002	0.0495	0.0027	0.52181	24.21	1	25	1.8	3.2	263	39	236	0.19
NOV11-1B_9t.FIN2	0.024	0.002	0.0038	0.0002	0.0475	0.0023	0.68144	24.23	1	24.27	1.6	0.2	193	26	512	0.18
NOV11-1B_5a.FIN2	0.024	0.002	0.0038	0.0002	0.0467	0.0022	0.62526	24.26	0.99	24.47	1.5	0.9	160	30	645	0.24
NOV11-1B_8c.FIN2	0.025	0.002	0.0038	0.0001	0.0503	0.0022	0.78132	24.29	1	25.5	1.7	4.7	198	24	889	0.12
NOV11-1B_10v.FIN	0.025	0.002	0.0038	0.0001	0.0476	0.0022	0.67411	24.35	0.98	25.1	1.6	3.0	152	31	875	0.15
NOV11-1B_11k.FIN	0.025	0.002	0.0038	0.0002	0.0477	0.0025	0.45556	24.42	1.1	25.03	1.7	2.4	165	36	451	0.13
NOV11-1B_10g.FIN	0.027	0.002	0.0038	0.0002	0.0513	0.0026	0.7075	24.47	1.1	26.7	1.9	8.4	283	34	1181	0.15
NOV11-1B_10f.FIN	0.033	0.003	0.0038	0.0002	0.0636	0.0041	0.35093	24.51	1	33.4	2.6	26.6	733	98	595	0.12
NOV11-1B_7r.FIN2	0.029	0.002	0.0038	0.0001	0.0541	0.0025	0.53591	24.52	0.92	28.82	1.8	14.9	399	38	1664	0.20
NOV11-1B_9u.FIN2	0.035	0.003	0.0038	0.0002	0.0656	0.0049	0.49386	24.59	1.1	34.3	3.3	28.3	870	140	913	0.28
NOV11-1B_9a.FIN2	0.026	0.002	0.0038	0.0002	0.0471	0.0029	0.57112	24.6	1.5	25.5	2	3.5	253	74	315	0.12
NOV11-1B_11j.FIN	0.031	0.002	0.0038	0.0002	0.0579	0.0032	0.61669	24.6	1	30.6	2.2	19.6	559	75	1055	0.20
NOV11-1B_7k.FIN2	0.029	0.002	0.0038	0.0002	0.055	0.0032	0.35891	24.6	1.1	28.5	1.9	13.7	491	79	886	0.19
NOV11-1B_5e.FIN2	0.026	0.002	0.0038	0.0002	0.0491	0.0026	0.49166	24.64	0.99	25.68	1.7	4.0	302	53	419	0.15
NOV11-1B_9s.FIN2	0.025	0.002	0.0039	0.0002	0.046	0.0025	0.59436	24.77	1	24.6	1.7	-0.7	119	41	335	0.15
NOV11-1B_3f.FIN2	0.024	0.002	0.0039	0.0002	0.0463	0.0023	0.68465	24.79	1.1	24.39	1.7	-1.6	193	31	369	0.18
NOV11-1B_5f.FIN2	0.024	0.002	0.0039	0.0002	0.0451	0.0023	0.5526	24.83	1.1	24.18	1.6	-2.7	157	34	425	0.19
NOV11-1B_5d.FIN2	0.034	0.003	0.0039	0.0002	0.0629	0.0037	0.3629	24.83	1	34	2.6	27.0	674	68	382	0.14
NOV11-1B_9l.FIN2	0.025	0.002	0.0039	0.0002	0.0481	0.0026	0.50204	24.88	1	25.47	1.7	2.3	196	34	225	0.08

Sample and Analysis Number	Isotopic Ratios										Dates (Ma)					Concentrator	
	207Pb/ ± 2SE	206Pb/ ± 2SE	207Pb/ ± 2SE	206Pb/ ± 2SE	Rho: 206/238 vs 207/235	206Pb/ ± 2SE	207Pb/ ± 2SE	206Pb/ ± 2SE	207Pb/ ± 2SE	% Discord.	207Pb/ ± 2SE	206Pb/ ± 2SE	U (ppm)	Th/U			
	235U (abs)	238U (abs)	206Pb (abs)	207Pb (abs)		206Pb (abs)	207Pb (abs)	235U (abs)	207Pb (abs)	6/38-7/35	206Pb (abs)						
NOV11-1B_8r.FIN2	0.024	0.002	0.0039	0.0002	0.0457	0.0023	0.62999	24.9	1	24.29	1.6	158	34	269	0.18		
NOV11-1B_2l.FIN2	0.033	0.002	0.0039	0.0001	0.0612	0.0033	0.3186	24.9	0.83	33.3	2.3	651	72	1069	0.13		
NOV11-1B_11p.FIN	0.029	0.002	0.0039	0.0002	0.0533	0.0030	0.58748	24.91	0.97	29	2.2	393	80	804	0.17		
NOV11-1B_3b.FIN2	0.033	0.002	0.0039	0.0002	0.0592	0.0032	0.32398	24.91	0.94	32.7	2.1	595	67	2450	0.24		
NOV11-1B_6k.FIN2	0.037	0.002	0.0039	0.0001	0.0704	0.0031	0.70374	24.91	0.92	36.71	2.3	937	26	938	0.18		
NOV11-1B_7l.FIN2	0.034	0.003	0.0039	0.0002	0.0633	0.0041	0.51364	24.91	1.1	33.7	2.9	684	96	723	0.21		
NOV11-1B_10y.FIN	0.025	0.002	0.0039	0.0003	0.0471	0.0031	0.66935	25	2	24.8	2	146	57	488	0.14		
NOV11-1B_8q.FIN2	0.029	0.002	0.0039	0.0002	0.0548	0.0032	0.43169	25.02	1.2	28.8	2.1	402	65	674	0.24		
NOV11-1B_9r.FIN2	0.031	0.002	0.0039	0.0001	0.0587	0.0033	0.61499	25.07	0.99	30.6	2.3	567	70	527	0.14		
NOV11-1B_8k.FIN2	0.024	0.002	0.0039	0.0002	0.046	0.0022	0.68666	25.16	1	24.43	1.6	157	28	768	0.22		
NOV11-1B_10i.FIN	0.033	0.003	0.0039	0.0002	0.0616	0.0036	0.54606	25.2	1.4	33.4	2.5	709	79	417	0.12		
NOV11-1B_5g.FIN2	0.038	0.003	0.0039	0.0001	0.0703	0.0035	0.44398	25.25	0.92	37.6	2.5	913	52	255.4	0.20		
NOV11-1B_8p.FIN2	0.032	0.003	0.0039	0.0002	0.059	0.0039	0.49963	25.32	1.2	31.7	2.7	537	82	732	0.30		
NOV11-1B_10j.FIN2	0.031	0.003	0.004	0.0002	0.0585	0.0036	0.67374	25.49	1	30.7	2.5	592	92	795	0.30		
NOV11-1B_10c.FIN	0.028	0.002	0.004	0.0002	0.0521	0.0031	0.5848	25.6	1.4	28.3	2.2	369	61	809	0.16		
NOV11-1B_10z.FIN	0.026	0.002	0.004	0.0002	0.0476	0.0022	0.77702	25.65	1.1	26	1.8	245	35	4940	0.07		
NOV11-1B_6a.FIN2	0.026	0.002	0.004	0.0002	0.0459	0.0025	0.51	25.77	1.1	25.6	1.9	230	77	461	0.18		
NOV11-1B_11e.FIN	0.026	0.002	0.004	0.0002	0.0471	0.0023	0.65409	25.79	1.1	25.62	1.7	170	29	415	0.21		
NOV11-1B_11q.FIN	0.039	0.004	0.004	0.0002	0.0703	0.0056	0.5861	25.82	1.1	39.1	3.8	930	130	530	0.24		
NOV11-1B_7c.FIN2	0.046	0.003	0.0041	0.0001	0.0785	0.0046	0.27281	26.15	0.9	45.2	3.3	42.1	1155	509	0.26		
NOV11-1B_8m.FIN2	0.031	0.002	0.0041	0.0002	0.0573	0.0031	0.66824	26.2	1.3	30.9	2.3	500	52	1462	0.28		
NOV11-1B_10p.FIN	0.033	0.002	0.0041	0.0001	0.0586	0.0026	0.65504	26.4	0.95	33.01	2	555	30	1559	0.35		
NOV11-1B_11m.FIN	0.05	0.005	0.0041	0.0002	0.0879	0.0072	0.64144	26.4	1.2	49.4	5.2	46.6	1320	140	876	0.18	
NOV11-1B_11g.FIN	0.04	0.003	0.0041	0.0002	0.0704	0.0046	0.64162	26.55	1.1	40.1	3.4	33.8	940	100	848	0.33	
NOV11-1B_8i.FIN2	0.066	0.005	0.0042	0.0002	0.1183	0.0068	0.44023	26.68	1.2	64.6	4.5	58.7	1926	61	727	0.25	
NOV11-1B_10l.FIN2	0.036	0.002	0.0042	0.0002	0.0619	0.0033	0.53341	26.9	1.2	35.9	2.4	698	42	587	0.36		
NOV11-1B_6c.FIN2	0.038	0.003	0.0042	0.0002	0.0676	0.0039	0.68581	26.97	1.2	37.6	3	28.3	67	643	0.30		
NOV11-1B_9x.FIN2	0.072	0.011	0.0042	0.0002	0.119	0.0140	0.79813	27.09	1.3	69	10	60.7	1720	210	168	0.14	
NOV11-1B_7q.FIN2	0.047	0.003	0.0042	0.0002	0.083	0.0047	0.46864	27.14	1	46.8	3.2	42.0	1272	52	422	0.24	
NOV11-1B_11c.FIN	0.027	0.002	0.0043	0.0002	0.0462	0.0021	0.69387	27.35	1.1	27.3	1.8	154	31	1062	0.04		
NOV11-1B_3h.FIN2	0.054	0.006	0.0044	0.0002	0.0945	0.0088	0.70065	27.96	1.2	53.2	5.9	47.4	1530	150	616	0.23	
NOV11-1B_10u.FIN	0.065	0.009	0.0044	0.0003	0.1014	0.0092	0.41108	28.5	1.9	63.8	8.3	55.3	1700	140	2050	0.23	
NOV11-1B_9q.FIN2	0.057	0.004	0.0046	0.0002	0.0904	0.0045	0.49212	29.52	1.1	55.8	3.6	47.1	1440	43	204	0.19	
NOV11-1B_10d.FIN	0.149	0.017	0.005	0.0002	0.217	0.0190	0.74488	32.09	1.3	139	15	76.9	2900	130	1000	0.20	

Sample and Analysis Number	Isotopic Ratios					Rho: 206/238 vs 207/235	Dates (Ma)					Concentration				
	207Pb/235U (abs)	± 2SE (abs)	206Pb/238U (abs)	± 2SE (abs)	207Pb/206Pb (abs)		± 2SE (abs)	206Pb/238U (abs)	± 2SE (abs)	207Pb/235U (abs)	± 2SE (abs)	% Discord. 6/38-7/35	207Pb/206Pb (abs)	± 2SE (abs)	U (ppm)	Th/U
NOV11-1B_11d.FIN	0.035	0.003	0.0052	0.0003	0.0467	0.0023	0.88147	33.5	2.1	34.9	3	4.0	156	55	2520	0.01
NOV11-1B_10q.FIN	0.079	0.007	0.0054	0.0003	0.103	0.0072	0.50898	34.6	2.1	76.6	6.5	54.8	1709	55	229.4	0.28
NOV11-1B_2m.FIN	0.043	0.003	0.0059	0.0003	0.0512	0.0023	0.85028	37.8	1.6	42.3	2.7	10.6	264	27	1309	0.11
NOV11-1B_11b.FIN	0.096	0.007	0.0063	0.0003	0.1074	0.0068	0.10004	40.6	1.8	92.4	6.4	56.1	1701	67	1820	0.09
NOV11-1B_10s.FIN	0.106	0.009	0.0067	0.0004	0.116	0.0096	0.16673	43.2	2.8	101.8	8.4	57.6	1860	130	1288	0.16
NOV11-1B_8g.FIN	0.059	0.004	0.0078	0.0004	0.0564	0.0025	0.89512	49.8	2.4	57.9	4	14.0	463	23	1346	0.05
NOV11-1B_6b.FIN	0.079	0.006	0.0093	0.0006	0.0616	0.0033	0.76782	59.5	3.6	77.2	5.6	22.9	616	63	349	0.15
NOV11-1B_3a.FIN	0.08	0.006	0.0104	0.0006	0.0537	0.0026	0.79388	66.9	3.7	77.7	5.8	13.9	366	38	473	0.01
NOV11-1B_9o.FIN	0.132	0.017	0.0106	0.0008	0.0831	0.0058	0.95193	68.2	5	124	15	45.0	1190	120	608	0.03
NOV11-1B_11a.FIN	0.099	0.007	0.0108	0.0006	0.0671	0.0042	0.15396	69.4	3.5	96	6.2	27.7	790	88	1073	0.12
NOV11-1B_8d.FIN	0.081	0.007	0.0117	0.0007	0.0506	0.0028	0.73196	75.1	4.7	79	6.3	4.9	339	64	698	0.10
NOV11-1B_8a.FIN	0.099	0.006	0.0127	0.0005	0.0564	0.0025	0.84216	81.4	3.4	95.3	5.7	14.6	470	22	781	0.06
NOV11-1B_7n.FIN	0.128	0.010	0.013	0.0008	0.071	0.0034	0.84694	83	4.9	122	9.1	32.0	970	39	342	0.19
NOV11-1B_9c.FIN	0.163	0.011	0.0202	0.0009	0.0586	0.0027	0.77683	129	6	153.5	9.6	16.0	555	35	2210	0.05
NOV11-1B_9c.FIN	0.163	0.011	0.0202	0.0009	0.0586	0.0027	0.77683	129	6	153.5	9.6	16.0	555	35	2210	0.05
NOV11-1B_7t.FIN	0.217	0.015	0.0216	0.0009	0.0723	0.0034	0.73325	137.7	5.8	200	12	31.2	990	36	333.7	0.06
NOV11-1B_9e.FIN	0.16	0.010	0.0216	0.0008	0.0533	0.0023	0.60742	137.9	5.1	150.7	8.4	8.5	339	26	1095	0.10
NOV11-1B_6h.FIN	0.157	0.011	0.023	0.0013	0.0512	0.0024	0.81118	146.8	8.1	148.8	9.9	1.3	269	25	1058	0.03
NOV11-1B_3j.FIN	0.176	0.012	0.0246	0.0011	0.0533	0.0024	0.84302	156.4	7.1	164.3	10	4.8	347	22	1006	0.10
NOV11-1B_11o.FIN	0.205	0.014	0.0254	0.0011	0.0564	0.0025	0.90591	161.9	7	189.1	12	14.4	471	26	748	0.05
NOV11-1B_7b.FIN	0.202	0.014	0.0269	0.0012	0.0542	0.0024	0.82075	170.8	7.6	186.6	12	8.5	384	32	753	0.08
NOV11-1B_9g.FIN	0.217	0.014	0.0292	0.0013	0.0537	0.0023	0.88101	185.7	7.9	198.7	12	6.5	335	22	757	0.08
NOV11-1B_11f.FIN	0.241	0.016	0.0326	0.0016	0.0536	0.0024	0.86579	206.9	9.7	218.5	13	5.3	362	26	463	0.11
NOV11-1B_5j.FIN	0.237	0.016	0.0332	0.0015	0.0515	0.0023	0.81252	210.6	9.7	217.3	13	3.1	270	26	1057	0.11
NOV11-1B_8e.FIN	0.246	0.016	0.0336	0.0014	0.0541	0.0024	0.63317	212.7	8.5	223.7	13	4.9	366	27	1089	0.32
NOV11-1B_9w.FIN	0.236	0.016	0.0343	0.0017	0.0508	0.0022	0.89941	217	11	214.4	13	-1.2	235	24	1524	0.13
NOV11-1B_10r.FIN	0.24	0.019	0.0343	0.0021	0.0506	0.0027	0.79555	217	13	218	15	0.5	290	47	2560	0.02
NOV11-1B_9b.FIN	0.247	0.017	0.0346	0.0016	0.0515	0.0024	0.81012	219.3	10	223.8	14	2.0	278	31	2039	0.01
NOV11-1B_8f.FIN	0.254	0.017	0.0364	0.0018	0.0519	0.0024	0.85662	230.4	11	229.3	14	-0.5	306	33	630	0.11
NOV11-1B_9i.FIN	0.273	0.018	0.0367	0.0016	0.0531	0.0023	0.83417	232.3	9.9	244.9	14	5.1	331	30	1071	0.39
NOV11-1B_2j.FIN	0.333	0.036	0.0386	0.0034	0.0621	0.0029	0.97098	243	21	290	28	16.2	674	37	883	0.04
NOV11-1B_9k.FIN	0.282	0.019	0.0384	0.0019	0.0521	0.0022	0.91349	243.6	11	251.4	15	3.1	298	22	1215	0.16
NOV11-1B_7l.FIN	0.284	0.017	0.0398	0.0016	0.0516	0.0024	0.51873	251.7	10	253.5	14	0.7	305	25	1774	0.05
NOV11-1B_8h.FIN	0.364	0.030	0.0402	0.0024	0.0682	0.0030	0.92891	254	15	314	22	19.1	856	25	371	0.22

Sample and Analysis Number	Isotopic Ratios					Dates (Ma)					Concentrator							
	207Pb/235U	± 2SE	206Pb/238U	± 2SE	207Pb/206Pb	± 2SE	206Pb/238U	± 2SE	207Pb/206Pb	± 2SE	206Pb/238U	± 2SE	207Pb/206Pb	± 2SE	U (ppm)	Th/U		
NOV11-1B_8o.FIN2	0.274	0.019	0.0407	0.0021	0.0493	0.0022	0.87901	0.0022	0.87901	257	13	246.7	15	-4.2	171	26	1082	0.09
NOV11-1B_9j.FIN2	0.288	0.018	0.041	0.0017	0.0503	0.0022	0.76186	0.0022	0.76186	259.1	11	256.3	14	-1.1	251	31	744	0.32
NOV11-1B_9m.FIN2	0.367	0.030	0.0431	0.0025	0.0601	0.0027	0.91973	0.0027	0.91973	272	16	314	23	13.4	618	29	459	0.13
NOV11-1B_7j.FIN2	0.316	0.019	0.0439	0.0016	0.0526	0.0023	0.70679	0.0023	0.70679	276.8	9.9	279	15	0.8	336	21	766	0.13
NOV11-1B_6d.FIN2	0.372	0.023	0.047	0.0025	0.0587	0.0031	0.60625	0.0031	0.60625	296	15	320.7	17	7.7	544	59	1544	0.06
NOV11-1B_4c.FIN2	0.446	0.031	0.0602	0.0028	0.0536	0.0025	0.77199	0.0025	0.77199	377	17	375	21	-0.5	334	35	1416	0.03
NOV11-1B_6m.FIN2	0.557	0.040	0.0645	0.0033	0.0631	0.0027	0.9269	0.0027	0.9269	402	20	448	27	10.3	705	22	345	0.11
NOV11-1B_8b.FIN2	0.543	0.034	0.067	0.0027	0.0589	0.0026	0.71186	0.0026	0.71186	418	16	439.6	22	4.9	577	32	650	0.07
NOV11-1B_6j.FIN2	0.509	0.040	0.0674	0.0041	0.0568	0.0028	0.84789	0.0028	0.84789	420	24	420	26	0.0	460	56	330	0.11
NOV11-1B_10w.FIN	0.568	0.035	0.0722	0.0031	0.0567	0.0025	0.70439	0.0025	0.70439	449	18	457.4	24	1.8	498	27	192	0.15
NOV11-1B_10t.FIN2	0.628	0.041	0.0761	0.0030	0.0604	0.0028	0.62	0.0028	0.62	472	18	494	26	4.5	601	28	230	0.13
NOV11-1B_4b.FIN2	0.76	0.052	0.0884	0.0040	0.0626	0.0029	0.77098	0.0029	0.77098	546	24	575	30	5.0	687	35	286.7	0.07
NOV11-1B_7s.FIN2	1.135	0.072	0.1277	0.0053	0.0638	0.0028	0.83798	0.0028	0.83798	774	30	770	35	-0.5	744	20	99.2	0.56
NOV11-1B_10e.FIN	2.67	0.190	0.1535	0.0076	0.1274	0.0065	0.78364	0.0065	0.78364	919	43	1320	55	30.4	2073	41	276.9	0.20
NOV11-1B_7p.FIN2	2.762	0.190	0.167	0.0081	0.1216	0.0053	0.86361	0.0053	0.86361	993	45	1342	50	26.0	1979	23	128	0.24
NOV11-1B_11i.FIN2	2.61	0.180	0.1708	0.0081	0.11	0.0048	0.90809	0.0048	0.90809	1014	45	1306	51	22.4	1803	22	259	0.53
NOV11-1B_11l.FIN2	4.44	0.290	0.2855	0.0120	0.1102	0.0050	0.69713	0.0050	0.69713	1616	61	1718	57	5.9	1814	34	301	0.27
NOV11-1B_3c.FIN2	4.36	0.280	0.2858	0.0120	0.1078	0.0048	0.75453	0.0048	0.75453	1617	60	1708	51	5.3	1776	28	48	1.13
NOV11-1B_10h.FIN	7.4	0.520	0.298	0.0160	0.1806	0.0083	0.85506	0.0083	0.85506	1674	78	2153	62	22.2	2669	27	427	0.29
NOV11-1B_10k.FIN	6.28	0.440	0.326	0.0140	0.1379	0.0073	0.71073	0.0073	0.71073	1819	70	2010	63	9.5	2192	52	67.1	0.55
NOV11-1B_9p.FIN2	7.94	0.540	0.36	0.0160	0.161	0.0071	0.84213	0.0071	0.84213	1989	77	2224	60	10.6	2473	19	150.4	0.58

Notes:

abs = absolute

rho is the correlation coefficient between the ratios

% Discord. $6/38-7/35 = 100 - (100 * (206Pb/238U \text{ date}) / (207Pb/235U \text{ date}))$

Bold analyses were used in age calculations.

CHAPTER 4

Texturally controlled U-Pb monazite dating of sapphirine granulites from the Gruf Complex, Central Alps: Evidence for UHT metamorphism in a Cenozoic continent-continent collision

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Abstract

Whether ultra-high temperature (UHT) metamorphism of the Gruf complex occurred during Permian rifting or Alpine orogenesis has been a matter of debate. Here we present texturally controlled U-Pb data of monazite from UHT mineral assemblages within the Gruf granulites. All monazite grains and domains have U-Pb dates between 34 and 30 Ma. We interpret the association of monazite with peak and post-peak UHT minerals and textures and the lack of inherited monazite to indicate that the UHT metamorphic event resulted in complete recrystallization of monazite and thus must have occurred during Cenozoic Alpine orogenesis. Monazite likely formed as garnet was breaking down to produce orthopyroxene, sapphirine, and cordierite in response to decompression at UHT conditions. The range of dates likely indicates that monazite growth

continued for ca. 4 m.y., and Alpine UHT metamorphism is consistent with a model in which advection of mantle heat and material to the lower crust occurred in response to slab breakoff.

Introduction

Ultra-high temperature (UHT) metamorphic rocks are the products of extreme thermal conditions within the lower crust ($>900^{\circ}\text{C}$ and 7–13 kbar; Harley, 1998; Brown, 2007; Harley, 2008). Attaining UHT conditions requires either a hotter-than-normal geothermal gradient (in excess of $20^{\circ}\text{C}/\text{km}$; Kelsey, 2008) or geodynamic processes that bring heat to the lower crust (e.g. Collins, 2002; Clark et al., 2011). Therefore, understanding when UHT metamorphism occurred in a particular geologic setting has important implications for the thermal, tectonic, and rheological evolution of that setting. The evolution of continent-continent orogens, such as the Alps, can influence surface processes such as volcanic activity, earthquakes, and erosion. Therefore, studies of rocks exhumed from the deep parts of orogens can be used to further understand the deep geodynamic processes.

The Gruf Complex contains the only known UHT rocks in the Central Alps (Bousquet et al., 2012a). The maximum pressure and temperature conditions these rocks experienced have been agreed upon by several researchers (Galli et al., 2011; Guevara and Caddick, 2016; Oalmann et al., in preparation-a), but there has not been a consensus as to the timing of the UHT event. The first geochronological studies on the UHT granulites from the Gruf Complex (Liati and Gebauer, 2003; Schmitz et al., 2009), concluded that peak metamorphism occurred during the Alpine orogeny, at ca. 32 Ma. However, based on U-Pb zircon dating of charnockites associated with the UHT granulites, Galli et al. (2012) concluded that UHT metamorphism occurred during a Permian rifting event and that metamorphic conditions only reached ca. $700\text{--}750^{\circ}\text{C}$ and 6–8 kbar during the Alpine orogeny (Galli et al., 2011). Here we present texturally-controlled U-Pb dating of

monazite from the Gruf granulites, thus linking P-T estimates directly to dates, to constrain the timing of UHT metamorphism in the Gruf Complex.

Geologic Setting

Several units in the Central Alps, including the Adula nappe (Figure 4.1), underwent ultra-high pressure (UHP) metamorphism related to subduction until ca. 35 Ma (Becker, 1993; Gebauer, 1996; Brouwer et al., 2005; Hermann et al., 2006; Liati et al., 2009) followed by exhumation and a Barrovian (temperature-dominated) metamorphic overprint, which began at ca. 34 Ma (Gebauer, 1996; Hermann et al., 2006) and continued to ca. 22 Ma (Rubatto et al., 2009). The Barrovian overprint, which affected an area known as the Lepontine dome, is characterized by greenschist facies metamorphism in the northern Central Alps increasing to upper-amphibolite to lower-granulite facies migmatization in the Southern Steep Belt (Todd and Engi, 1997; Figure 4.1; Frey and Ferreiro Mählmann, 1999; Burri et al., 2005). Previously proposed causes of the switch from high-pressure to high-temperature metamorphism include convective heating within an orogenic wedge (Bousquet et al., 2008; Berger et al., 2011), slab breakoff (von Blanckenburg and Davies, 1995; Handy et al., 2010), and lithospheric thinning (Beltrando et al., 2010).

The Gruf Complex is bordered to the north by the Chiavenna ophiolite and Tambo nappe. The normal sense Forcola Fault separates the Gruf Complex from the Adula nappe, which has a different P-T history, to the west (Figure 4.1; Ciancaleoni and Marquer, 2006; Bousquet et al., 2012b; Galli et al., 2013). The 32–30 Ma (von Blanckenburg, 1992; Samperton et al., 2015) Bergell tonalite to granodiorite intrusion structurally overlies the Gruf Complex (Davidson et al., 1996) and crops out to the east and south. The Novate leucogranite intruded the Gruf Complex at 23.8 ± 0.1 Ma (Oalman et al., in preparation-c).

Upper amphibolite facies migmatitic orthogneisses and minor paragneisses and leucogranites are the most common rock types in the Gruf Complex. However, several mappable units of charnockite also occur, which contain scarce schlieren, rafts, and pods of sapphirine-bearing granulites. The granulites underwent simultaneous decompression and heating from ca. 800°C and 10–14 kbar to UHT conditions of 900–1000°C and 7.5–9.5 kbar (Oalman et al., in preparation-a). The ages of variably deformed felsic dikes indicate that the UHT granulite/charnockite units were juxtaposed against the migmatites along mylonitic shear zones between 30 and 26 Ma (Oalman et al., in preparation-c).

Sample Descriptions

Residual granulite samples Br03-55-3 and Br03-56-2, which lack a structural fabric, are very similar in their mineralogy, grain size, and textures. These granulites are dominated by large (up to 1 cm) anhedral garnet porphyroblasts and subhedral to euhedral sapphirine and orthopyroxene porphyroblasts, some of which are intergrown. Cordierite occurs in corona-textures around garnet and forms moats separating sapphirine and orthopyroxene. Quartz + K-feldspar + plagioclase occur in pod-shaped leucosomes that are scattered throughout the samples. Biotite occurs as monomineralic aggregates or fringes garnet and orthopyroxene porphyroblasts. Embayments in garnet are filled with sapphirine + orthopyroxene porphyroblasts, cordierite ± sapphirine ± orthopyroxene symplectites, and/or quartzofeldspathic leucosomes. In the residual granulites, anhedral to subhedral monazite is intergrown with or included in sapphirine, orthopyroxene, cordierite and/or biotite or occurs within leucosomes (Figure 4.2A-D). Monazite most commonly occurs near embayments in garnet.

Retrogressed granulite sample Br03-56-4 is similar in texture and appearance to the residual granulite samples. However, most of the sapphirine, orthopyroxene, and cordierite has

been replaced by secondary chlorite and biotite. Muscovite is also present in this sample. Anhedra monazite grains occur mainly within the retrogressed parts of this sample (Figure 4.2E).

Leucogranulite sample GR11-39 is characterized by alternating light and dark bands that give the sample a gneissic texture. The dark bands are dominated by orthopyroxene, garnet, biotite, and sapphirine, and the light bands comprise deformed quartz, K-feldspar, and plagioclase. Garnet occurs as single, anhedral porphyroblasts and clusters of anhedral grains that were probably originally a single porphyroblast. Garnet embayments are commonly filled with plagioclase, orthopyroxene, sapphirine, biotite, spinel, and magnetite. Monazite occurs exclusively as small, irregular grains on the edge of larger apatite grains (Figure 4.2F), except for one grain located in a leucosome between plagioclase and K-feldspar (grain 2, Table 4.A-2).

Results of U-Pb Monazite Dating

Monazite grains within four granulite samples (two residual granulites, one leucogranulite, and one retrogressed granulite) were dated by laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) at The University of Kansas (see Appendix A for complete analytical procedures). To preserve textural context, small areas were drilled out of thin sections and mounted in epoxy from the residual and retrogressed granulite samples, and monazite was analyzed directly within a thin section from the leucogranulite sample. Discordance is reported as uncertainty weighted age difference (UWAD) (Table 4.A2).

Backscattered electron (BSE) images of selected, representative monazite-bearing textures are shown in Figure 4.2. All analyses have $^{206}\text{Pb}/^{238}\text{U}$ dates between 34.5 and 30 Ma, and there is limited correlation between position within grain (e.g. cores or rims) and date (Figure 4.2A-D). Some of the highest, concordant $^{206}\text{Pb}/^{238}\text{U}$ dates come from monazite intergrown with high-Al orthopyroxene (two analyses from texture GRM2-2 at 33.6 ± 0.8 and 33.2 ± 0.8 Ma) and a euhedral

monazite grain included in plagioclase within a garnet-breakdown texture (33.6 ± 0.7 Ma) (Figure 4.2C), and monazite inclusions in sapphirine have $^{206}\text{Pb}/^{238}\text{U}$ dates older than 32 Ma. In a texture that is typical for monazite textures in the leucogranulite sample, the youngest monazite grains (31–30 Ma) occur on the edge of apatite grains (e.g. Figure 4.2F).

Within the residual and retrogressed granulite samples, discordant analyses are scarce and trend toward common-Pb upper intercepts (Figure 4.3A, 3B). The most concordant grains spread along concordia from 34 to 31 Ma (Figure 4.3A-C). The $^{206}\text{Pb}/^{238}\text{U}$ dates of the analyses with $\text{UWAD} < 1.2$ from the residual granulite sample Br03-56-2 range from 31.06 ± 0.77 to 32.99 ± 0.86 Ma (Table 4.A2; Figure 4.3A). The seven oldest concordant analyses make up a population that has a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 32.5 ± 0.3 Ma ($\text{MSWD}=0.5$, $n=7$), and the lower intercept age of the other analyses is $31.5 +0.2/-0.2$ Ma ($\text{MSWD}=0.4$, $n=17$). Analyses with $\text{UWAD} < 1.2$ from residual granulite sample Br03-55-3 have $^{206}\text{Pb}/^{238}\text{U}$ dates between 31.15 ± 0.59 and 33.61 ± 0.71 Ma (Table 4.2; Figure 4.3B). Excluding the two oldest and one youngest analyses with $\text{UWAD} < 1.2$ and the more discordant analyses, the weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age for this sample is 32.4 ± 0.2 Ma ($\text{MSWD}=1.0$, $n=14$). The three small grains from retrogressed granulite sample Br03-56-4 have $\text{UWAD} \leq 0.7$ and $^{206}\text{Pb}/^{238}\text{U}$ dates ranging from 32.15 ± 0.62 to 34.10 ± 0.77 Ma (Table 4.2; Figure 4.3C). These grains do not encompass a single age population, and therefore, no ages are interpreted from the retrogressed granulite sample. However, these dates agree with the age range of most monazite grains from the residual granulite samples.

All except one of the analyses from leucogranulite sample GR11-39 have UWAD values > 1.2 (Table 4.2; Figure 4.3D), and these grains have the highest U concentrations (4090–12,330 ppm; Table 4.A2) of all the monazite analyzed in this study. The $^{206}\text{Pb}/^{238}\text{U}$ dates of analyses with $\text{UWAD} \leq 3.2$ range from 30.04 ± 0.52 to 31.98 ± 0.56 Ma, which is a younger age range than in the

residual and retrogressed granulites. There are two distinct discordia trends with common-Pb upper intercepts (Figure 4.3D). The older population has a lower intercept age of $31.7 \pm 0.4/-0.5$ Ma (MSWD=0.2, n=4), and the younger discordant population has a lower intercept age of $30.8 \pm 0.3/-0.3$ Ma (MSWD=0.4, n=6). Two analyses that do not plot along either of the discordia trends have a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 30.1 ± 0.4 Ma (MSWD=0.2, n=2).

Discussion and Conclusions

Although many of the U-Pb monazite dates overlap within uncertainty of one another, analyses from single samples and even from single grains do not form age populations with n-appropriate MSWDs without excluding some analyses. This leads us to the interpretation that not all monazite crystallized at the same time. It is more likely that monazite (re)crystallized more or less continuously over the ca. 4 m.y. period recorded by the $^{206}\text{Pb}/^{238}\text{U}$ dates of the concordant analyses, which range from 30.04 ± 0.52 to 34.10 ± 0.77 Ma. This age range is consistent with early post-peak UHT crystallization of zircon in the same rock units (Oalmann et al., in preparation-c).

Based on the close vicinity of most monazite grains to garnet breakdown textures in the residual granulites (Figure 4.2A, C, D) and high Y content of the monazite (up to 4 wt. % Y_2O_3 ; Schmitz et al., 2009), the mostly likely trigger for monazite growth was a release of P and REE during the breakdown of garnet, which contains ca. 0.1 ppm Ce on average (Appendix B). To balance the LREE needed for monazite, it is likely that melt and apatite and xenotime, which is present in the residual granulites but not dated, were involved in the monazite producing reactions. Monazite also occurs in a variety of textures associated with UHT minerals (e.g. high-Al orthopyroxene, sapphirine, and cordierite) (Figure 4.2A–D). The sapphirine granulites underwent simultaneous decompression and heating from ca. 800°C and 9–12 kbar to $900\text{--}1000^\circ\text{C}$ and 7.0–9.5 kbar, resulting in orthopyroxene and sapphirine (and later cordierite) growth at the expense of

garnet (Oalman et al., in preparation-b). Some of the oldest (i.e. >32 Ma) domains are from monazite grains included in sapphirine (Figure 4.2B, D). The texture shown in Figure 4.2B (orthopyroxene + sapphirine + cordierite) equilibrated at 850–1100°C and 7.25–10.5 kbar (Oalman et al., in preparation-a) and contains monazite with $^{206}\text{Pb}/^{238}\text{U}$ dates ranging from 32.0±0.8–33.0±0.9 Ma (weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age = 32.5±0.5, MSWD=1.5, n=3). The euhedral monazite shown in Figure 4.2C has $^{206}\text{Pb}/^{238}\text{U}$ dates ranging from 32.0±0.8–33.6±0.7 Ma (weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age = 32.7±0.4, MSWD=5.3, n=3) and is within a texture that equilibrated at ca. 962°C and 8.7 kbar (Appendix C). The garnet-breakdown texture shown in Figure 4.2A equilibrated at 910°C and 7.6 kbar (Appendix C) contains anhedral monazite ($^{206}\text{Pb}/^{238}\text{U}$ dates = 31.2±0.6–32.4±0.7 Ma) intergrown with high-Al orthopyroxene and plagioclase. Based on the association of monazite with these UHT to slightly post-UHT textural and mineralogical indicators, we conclude that UHT metamorphism occurred simultaneous with or slightly prior to crystallization of the oldest monazite (i.e. during the Alpine orogeny at ca. 34 Ma).

The textural relationships between monazite and apatite in the leucogranulite sample GR11-35 (Figure 4.2F) are interpreted as monazite growth as the result of either LREE expulsion from apatite during decompression (e.g. Finger and Krenn, 2007; Krenn et al., 2009; Rocha et al., 2016) or fluid-assisted dissolution-reprecipitation (Harlov and Förster, 2002; Harlov et al., 2005). The two relatively young ages interpreted from this sample (30.8 +0.3/-0.3 Ma and 30.1±0.4 Ma) compared to the >31.5 Ma ages from the residual granulites (Figure 4.3) and the lack of concordant >32 Ma grains (Table 4.2; Figure 4.3D) indicate that this was a late-stage process. Because there is no evidence in the leucogranulite sample for low temperature fluid infiltration (e.g. abundant hydrous minerals within fractures), we interpret that LREE expulsion from apatite during

decompression and/or cooling led to the monazite growth on apatite. The different age groups (Figure 4.3D, Table 4.A-2) indicate that this process occurred at least three different times in the leucogranulite. Analyses of two different monazite grains (analyses GR11-35-3-1, and 3-2) along the edge of the same apatite grain have $^{206}\text{Pb}/^{238}\text{U}$ dates of 31.3 ± 0.7 Ma and 34.5 ± 0.8 Ma (discordant) and are part of two different age groups shown in Figure 4.3D. We interpret this to indicate that LREE expulsion occurred at two different times from the same apatite grain and thus was a very localized process.

The few monazite dates obtained from the retrogressed granulite sample Br03-56-4, which range from 32.15 ± 0.62 – 34.10 ± 0.77 Ma, overlap with the oldest dates from the residual granulite samples (Table 4.2; Figure 4.3). Because this sample is likely a retrogressed residual granulite sample (see '*Sample Descriptions*' above), we interpret that (re)crystallization or modification of monazite did not occur during greenschist-facies, retrograde metamorphism.

Many (U)HT granulites from other areas contain complexly zoned monazite grains that preserve the ages of multiple plutonic, metamorphic, or metasomatic events (e.g. Kelly et al., 2006; Mahan et al., 2006; Kelly et al., 2012; Dumond et al., 2015; Rocha et al., 2016). However, the Gruf granulites lack inherited monazite (Figure 4.2, 3). Experimental studies have revealed that the closure temperature for Pb diffusion in monazite is similar to that of Pb in zircon (e.g. Cherniak et al., 2004), and thus U-Pb monazite dates should not be affected by post-crystallization diffusive resetting. Therefore, we interpret the lack of inherited monazite to indicate that the Gruf granulites lacked monazite prior to the Alpine metamorphic event or that the rocks completely recrystallized during or slightly earlier than the time monazite (re)crystallized. The lack of inherited monazite precludes an older (Permian) UHT event followed by a moderate temperature overprint during the Alpine orogeny as suggested by Galli et al. (2011; 2012; 2013).

Figure 4.4 summarizes the likely Alpine P-T-t path taken by the Gruf granulites, which is in part based on data and observations presented by Oalman et al. (in preparation-a; in preparation-c; submitted): (1) Prograde metamorphism reached conditions of 800°C and 10–14 kbar. The exact timing and duration of this part of the P-T path is unknown, but minor zircon growth in equilibrium with garnet occurred at $34.8 \pm 1.0/-1.1$ (Oalman et al., in preparation-c). (2) During decompression, the granulites were heated to UHT conditions, resulting in garnet breakdown and the crystallization of the oldest monazite domains in the residual granulites at ca. 34 Ma. (3) Monazite continued to crystallize in the residual granulites until ca. 31 Ma, and the first pulse of monazite growth on apatite in the leucogranulite occurred at $31.7 \pm 0.4/-0.5$ Ma. (4) Between 31 and 30 Ma, monazite in the leucogranulite continued to crystallize along the edge of apatite grains. (5) The lower crustal granulites (and charnockites) were exhumed into the mid-crustal migmatites between 30 and 26 Ma (Oalman et al., in preparation-c).

Attainment of UHT conditions in the lower crust of a continent-continent collision requires an external source of heat. The most likely heat source is the asthenospheric mantle. We postulate that either of two geodynamic processes could have led to heat advection to the lower crust: (1) Slab breakoff resulted in asthenospheric upwelling and mantle melting (von Blanckenburg and Davies, 1995; Handy et al., 2010); or (2) Slab rollback lead to lithospheric thinning of the overthickened orogenic wedge and thus asthenospheric upwelling and magma generation (Beltrando et al., 2010). However, we prefer the slab breakoff model because heating in response to lithospheric thinning would have resulted in a long-lived UHT event and widespread magmatism, both of which are not observed in the Alps.

An Alpine age for UHT metamorphism indicates that the Gruf Complex hosts the second youngest occurrence of exposed UHT rocks in the world. The youngest known exposure of UHT

granulites occur as lenses within an active subduction zone in Indonesia (Pownall et al., 2014), and Hacker et al. (2000) reported young UHT xenoliths in lavas erupted on the Tibetan plateau. Cenozoic UHT metamorphism may have occurred in other regions, but the products have not been yet recognized due to non-exposure, retrogression, or lack of UHT indicator assemblages (e.g. sapphirine + quartz).

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Appendix 4.A: Analytical Methods for U-Pb Monazite Dating

Monazite was analyzed *in situ* (i.e. directly within thin sections or drilled out parts of thin sections) by LA-ICP-MS at The University of Kansas, Department of Geology's Isotope Geochemistry Lab. Table 4.A1 shows the analytical protocols that were used for all measurements. Table 4.A2 shows the U-Pb isotopic data and calculated ages.

Appendix 4.B: Analytical Methods for Garnet Trace Element Analyses

Trace elements concentrations of garnet from one of the monazite bearing textures were measured by LA-ICP-MS at The University of Kansas, Department of Geology's Isotope Geochemistry Lab. Electron microprobe major element data were used for the internal standard values (Ca concentration) for each spot. Table 4.B1 shows the analytical protocols that were used for all measurements. Table 4.B2 shows the trace element concentrations that were measured.

Figure 4.B1 shows chondrite-normalized rare earth element patterns for all of the spots that were analyzed.

Appendix 4.C: Thermobarometry Calculations

Pressure and temperature conditions corrected for retrograde exchange were estimated for textures in the residual granulites using the computer program RCLC of Pattison et al. (2003). Major element compositions used for the input to the calculations were determined by electron probe microanalysis (EMPA) at the Christian-Albrechts-Universität zu Kiel Institute for Geosciences using the methods presented by Oalman et al. (in preparation-b). The compositional data and results of the calculations are presented in Table 4.C1. These calculations assume equilibrium between garnet, orthopyroxene, and plagioclase.

The estimates obtained for textures GRMPx3 (910°C at 7.6 kbar) (Figure 4.2A) and GRM-27 (962°C at 8.7 kbar) (Figure 4.2C) with this method are consistent with the results of thermobarometric modeling (peak conditions = 900–1000°C and 7.5–9.5 kbar) presented by Oalman et al. (in preparation-b), but the estimated temperature for texture GRM-37 (1040°C at 9.1 kbar) (Figure 4.2B) is higher than the peak temperature range modeled by Oalman et al. (in preparation-b). The garnet in this texture occurs as inclusions in orthopyroxene (Figure 4.2B). We interpret this relationship to indicate that garnet is a left over product of a garnet-consuming reaction that produced orthopyroxene + sapphirine + cordierite (Oalman et al., in preparation-b), and thus all the phases are not in equilibrium in this texture.

Figure Captions

Figure 4.1. Tectonic map of part of the Central Alps (modified from Bousquet et al., 2012b) showing the location of the Gruf Complex and surrounding units. Ages for the Bergell and Novate

intrusions are from von Blanckenburg (1992) and (Oalman et al., in preparation-c), respectively.
SSB = Southern Steep Belt.

Figure 4.2. Backscattered electron (BSE) images of representative monazite-bearing textures showing locations of analyses and $^{206}\text{Pb}/^{238}\text{U}$ dates with 2σ propagated uncertainty. (A–B) Textures from residual granulite sample Br03-56-2. (C–D) Textures from residual granulite sample Br03-55-3. (E) A texture from retrogressed granulite Br03-56-4. (F) A texture from leucogranulite sample GR11-39.

Figure 4.3. Wetherill concordia diagrams plotted with ET_Redux (McLean et al., 2016) showing analyses from the different samples and calculated ages (LI = lower intercept; WM = weighted mean). Ellipse areas represent 2σ propagated uncertainty. (A) Residual granulite sample Br03-56-2. (B) Residual granulite sample Br03-55-3. (C) Retrogressed granulite sample Br03-56-4. (D) Leucogranulite sample GR11-39. For panels A and C, the different colored ellipses were used to calculate the different ages with corresponding colored text. In panel B, blue ellipses were used in the age calculations, and grey ellipses were excluded.

Figure 4.4. P-T diagram showing P-T path taken by the Gruf granulites, schematic mineral reactions, conditions and timing of monazite growth, and conditions and timing of exhumation of the lower crustal UHT rocks into the middle crust.

Figure 4.B1. Chondrite-normalized REE patterns for garnet analyses from residual granulite sample Br03-56-2.

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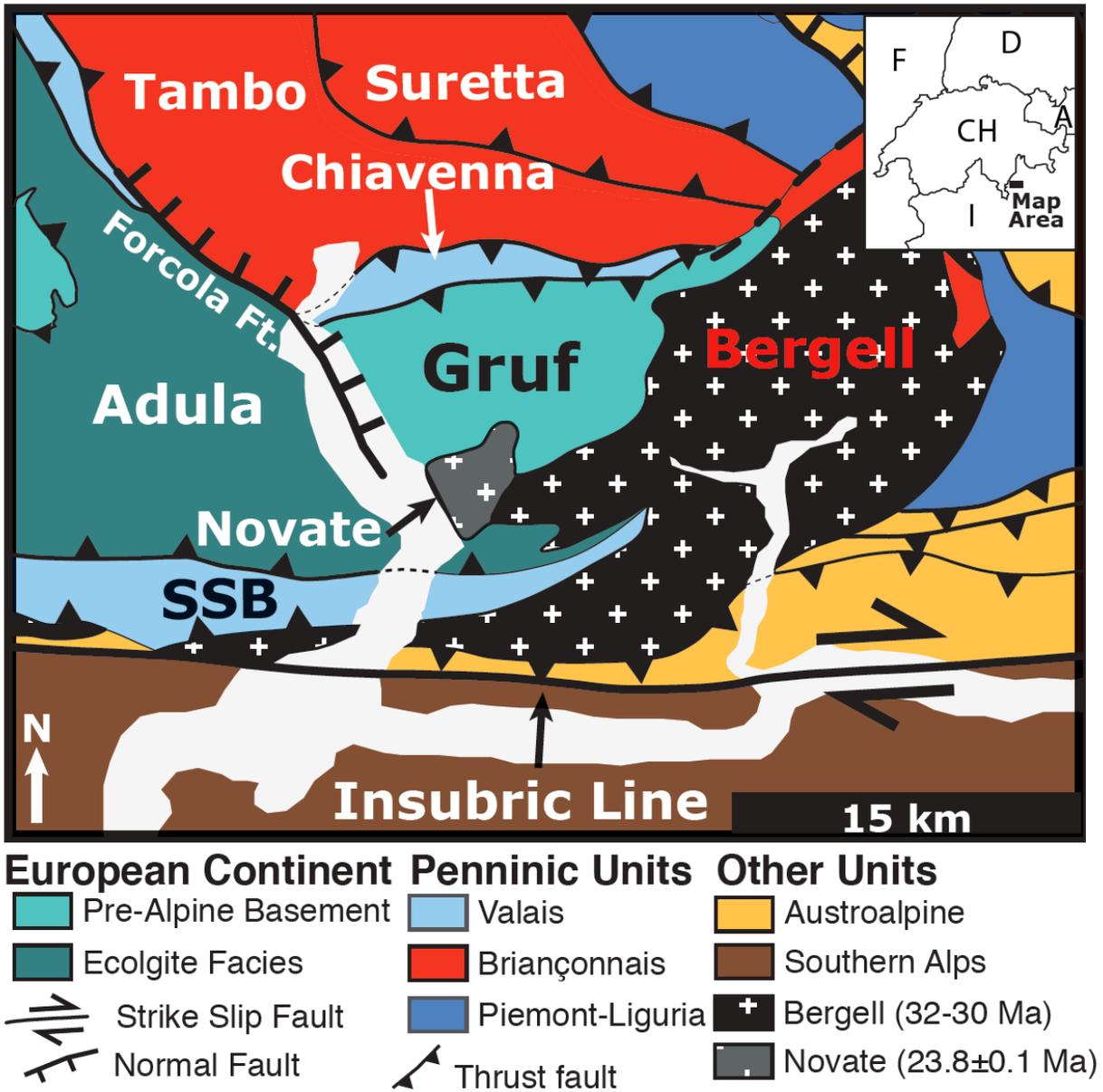


Figure 4.1

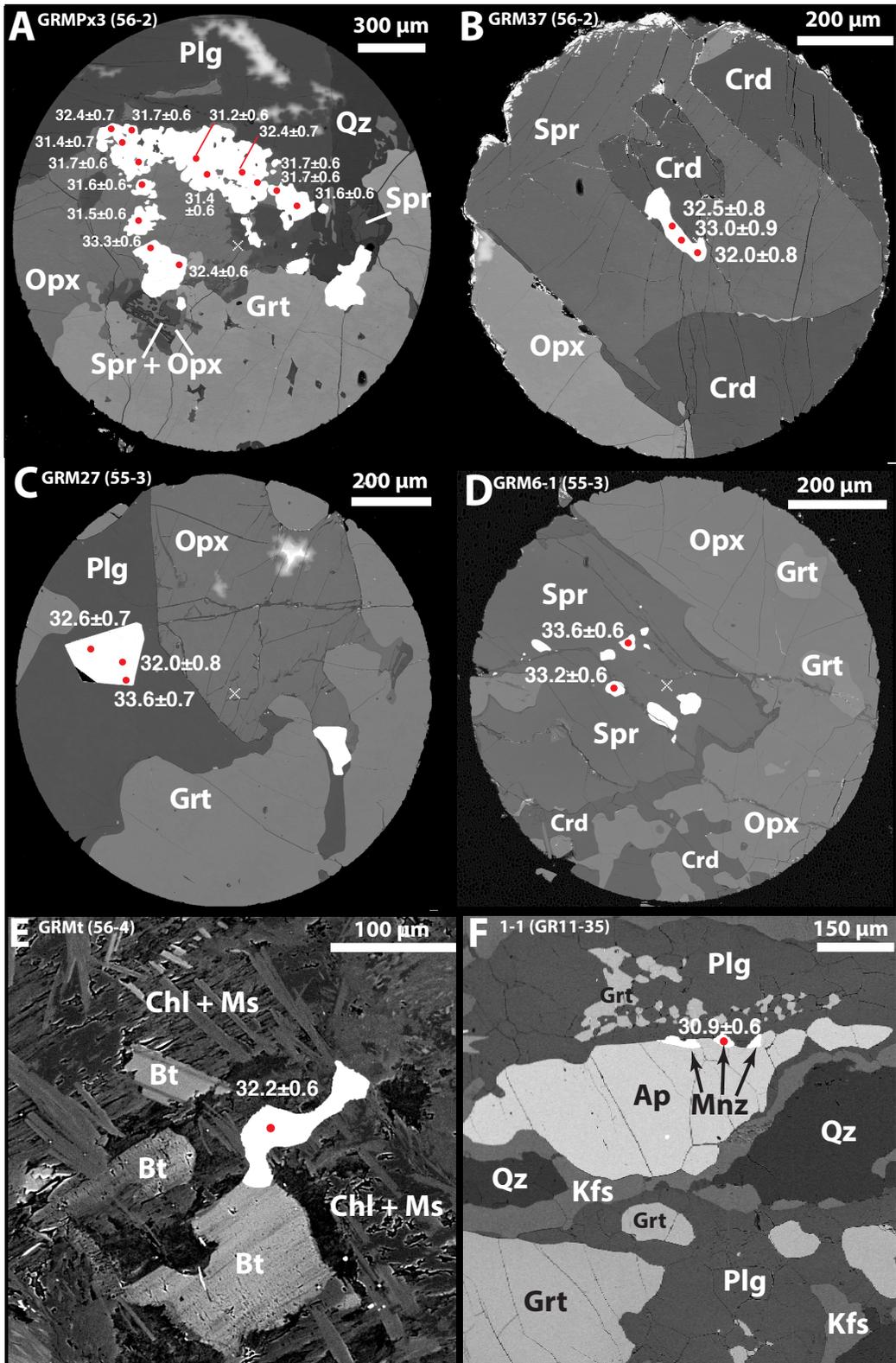


Figure 4.2

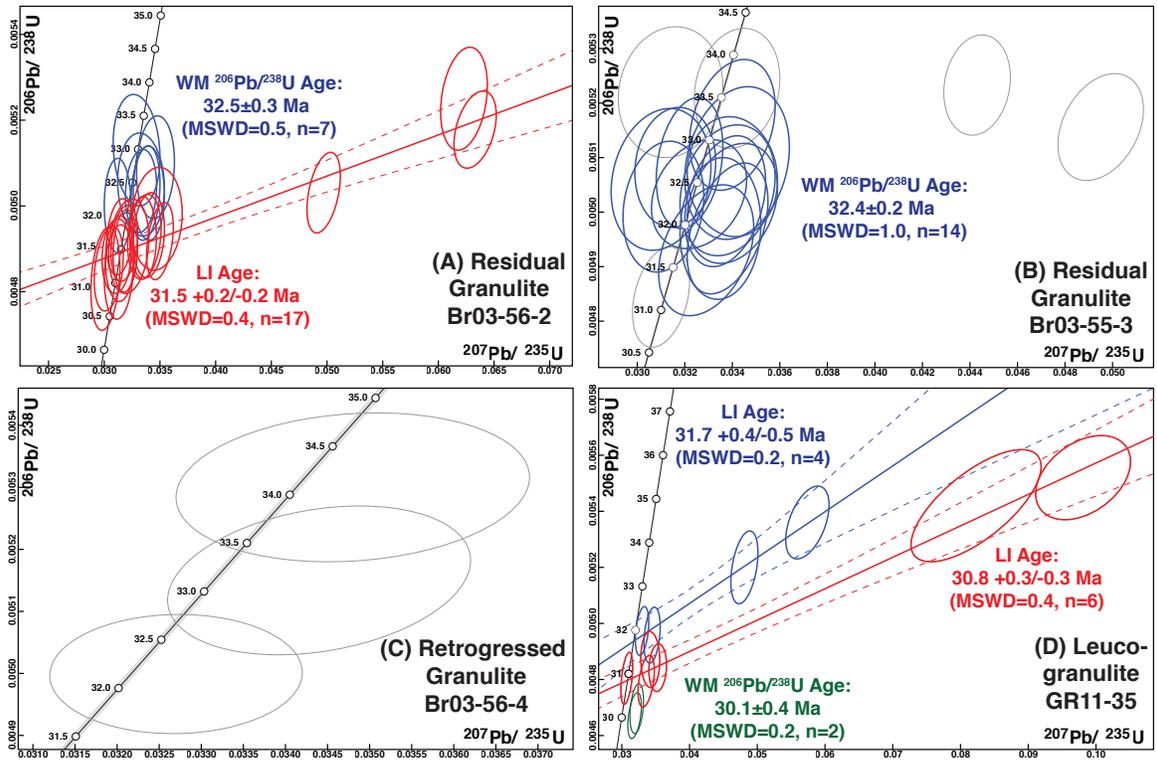


Figure 4.3

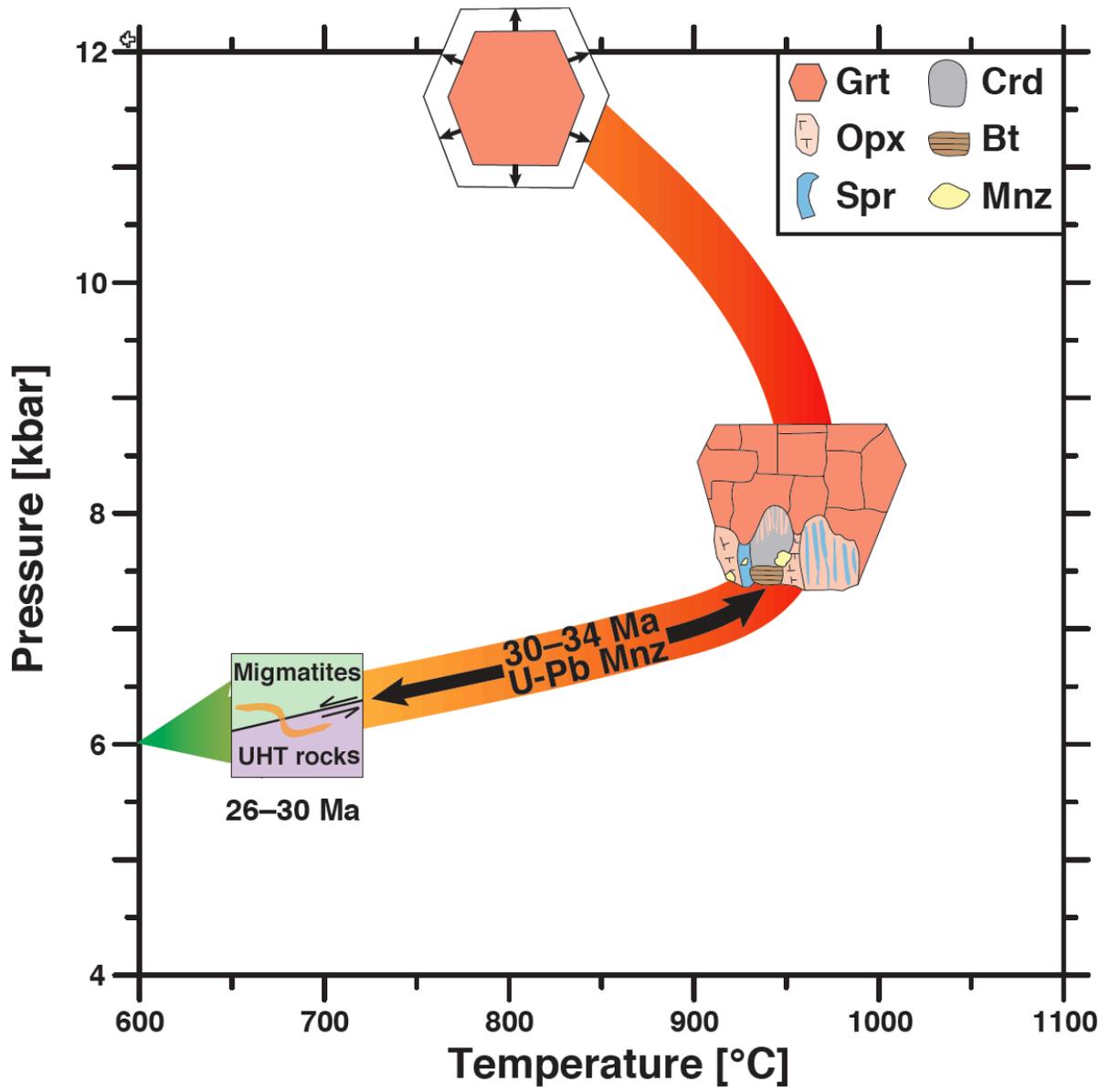


Figure 4.4

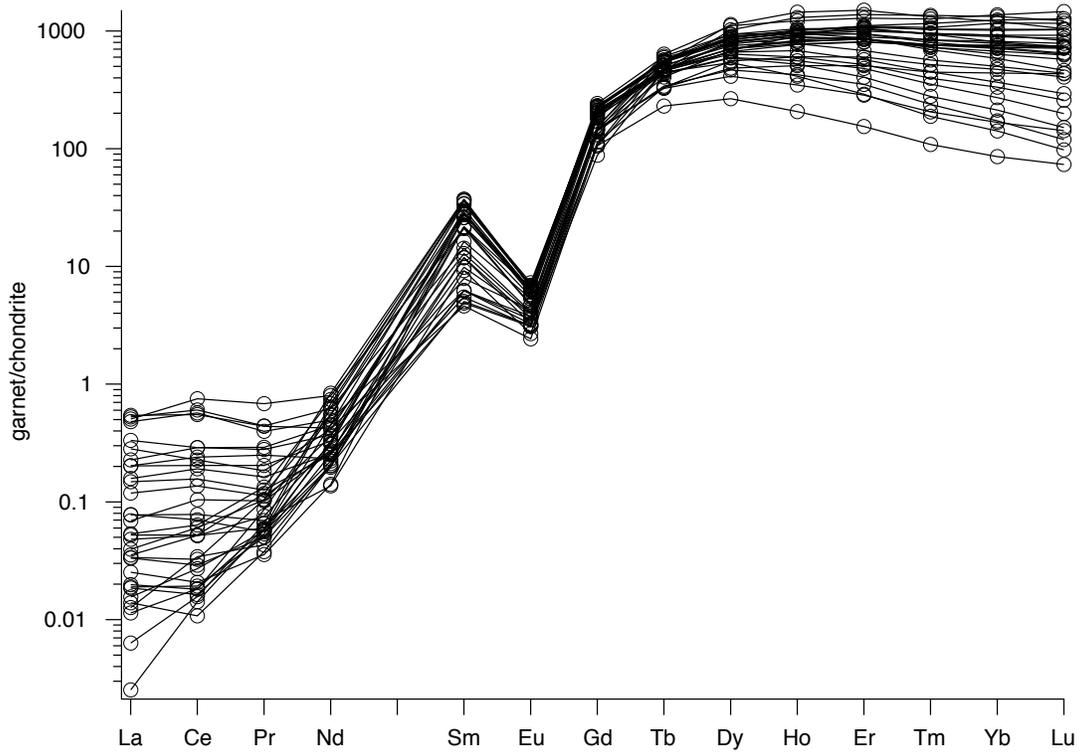


Figure 4.B1

Table 4.A1. Monazite LA-ICP-MS Analytical Parameters.

Laboratory & Sample Preparation	
Laboratory name	The University of Kansas, Dept. of Geology, Isotope Geochemistry Lab
Sample type/mineral	Monazite
Sample preparation	40 µm-thick, polished sections
Imaging	BSE imaging on electron microprobe
Laser ablation system	
Make, model & type	ATL Arf excimer laser (193 nm), Photon Machines Analyte G2
Ablation cell & volume	HeLex
Laser wavelength	193 nm
Pulse width (ns)	5 ns
Fluence	2.4 J/cm ²
Repetition rate	5 Hz
Spot size (µm)	15 µm
Sampling mode / pattern	Single spots
Carrier gas	He, 1.1 l/min, Ar, 1.095 l/min
Ablation duration	22 s
Cell carrier gas flow	He
ICP-MS Instrument	
Make, Model & type	Thermo Element2 magnetic sector field ICP-MS (single collector)
Sample introduction	Aerosol with sample + He was mixed with Ar using a T-connector 30 cm upstream from torch. A 'squid' device was used to attenuate the aerosol between the T-connector and the torch.
RF power	1095 W
Make-up gas flow	Ar, 16 l/min
Sampling depth	20 µm???
Detection system	single detector (SEM), counting & analog modes
Elements/ isotopes analyzed	206Pb, 207Pb, 238U
Integration time per peak (Sample Time in milliseconds)	206Pb = 4, 207Pb = 10, 238U = 2
Total integration time (Segment Duration in milliseconds)	206Pb = 12, 207Pb = 30, 238U = 6
Total method time	37 seconds (176 runs, 4 passes)
Sensitivity (cps/ppm)	238U = 1030 cps/ppm
ICP Dead time	2 ns
UO+/U+	0.10%
238U+/232Th+	1.1

Table 4.A1

Data Processing	
Gas blank	15 s
Calibration strategy	Standard-sampling bracketing
Reference material info	FG Manangotry pass, TIMS age 559.6±1.7 Ma (Möller, unpubl.)
Internal std for trace elements	n/a
Data processing package used / Correction for laser induced elemental fractionation (LIEF)	IGOR PRO, lolite 2.5 (Paton et al. 2011): U_Pb_Geochronology (Paton, et al. 2010) and VizualAge data reduction schemes (Petrus & Kamber, 2011), exponential LIEF correction for U-Pb ratios.
Common-Pb correction, composition and uncertainty	Not performed because of high Hg background
Uncertainty level & propagation	2 σ propagated uncertainty from VizualAge data reduction scheme. Concordia diagrams were plotted using ET_Redux (McLean et al. 2016) with 2 σ uncertainty ellipses.
Reproducibility	206Pb/238U = 1.5-2.0%; 207Pb/235U = 2.5-3.0%
Quality control / Validation	Monazite: Moacyr (Seydoux-Guillaume et al. 2002) and JUP (Kennedy pers. comm)

Sample and Analysis Number	Textural setting	Isotopic Ratios				Dates (Ma)				Disc. a				conc U (ppm)				
		207Pb/235U	±2σ %	206Pb/238U	±2σ %	Rho	±2σ %	207Pb/206Pb	±2σ %	207Pb/235U	±2σ	207Pb/206Pb	±2σ		% disc. a	% disc. b	% disc. c	
GRM28-1.FIN2	in Bt	0.0501	4.3	0.00483	2.5	0.24	0.0452	4.6	31.06	0.77	30.1	1.3	-8	85	-3.2	488.3	-0.7	4420
GRM28-2.FIN2	in Bt	0.0313	4.2	0.00485	2.3	0.14	0.0466	4.3	31.16	0.68	31.3	44	82	0.4	29.2	0.4	29.2	3690
GRM_PX3-6.FIN2	anhedral; with Opx (Fig. 2A)	0.0304	4.3	0.00485	2.0	0.28	0.0466	4.1	31.22	0.63	30.4	1.3	45	85	-2.7	30.6	-0.6	4750
GRM_PX3-8.FIN2	anhedral; with Opx (Fig. 2A)	0.0314	4.1	0.00488	1.8	0.05	0.0450	4.3	31.36	0.66	31.3	63	83	0.0	50.2	0.0	4780	
GRM_PX3-5.FIN2	anhedral; with Opx (Fig. 2A)	0.0311	4.2	0.00488	2.1	0.10	0.0459	4.4	31.40	0.56	31.1	1.3	18	83	-1.0	-74.4	-0.2	5310
GRM_PX3-12.FIN2	anhedral; with Opx (Fig. 2A)	0.0319	3.4	0.00490	2.0	0.12	0.0483	3.9	31.48	0.62	31.9	1.1	124	80	1.2	74.6	0.4	4550
GRM29b1-1.FIN2	in Bt	0.0322	5.6	0.00491	2.2	0.17	0.0485	5.6	31.55	0.72	32.2	1.7	120	110	2.0	73.7	0.4	2237
GRM_PX3-1.FIN2	anhedral; with Opx (Fig. 2A)	0.0318	3.8	0.00491	1.8	0.24	0.0475	4.0	31.59	0.57	31.8	1.2	83	81	0.7	61.9	0.2	4640
GRM_PX3-11.FIN2	anhedral; with Opx (Fig. 2A)	0.0350	3.4	0.00492	1.8	0.36	0.0451	3.5	31.63	0.58	34.9	1.2	251	75	9.4	87.4	2.7	6160
GRM_PX3-9.FIN2	anhedral; with Opx (Fig. 2A)	0.0335	3.6	0.00493	1.9	0.00	0.0501	3.8	31.68	0.62	33.4	1.2	189	81	5.2	83.2	1.5	5560
GRM_PX3-2.FIN2	anhedral; with Opx (Fig. 2A)	0.0336	4.5	0.00493	2.0	0.20	0.0492	4.5	31.69	0.64	33.6	1.5	163	91	5.7	80.6	1.3	4419
GRM_PX3-3.FIN2	anhedral; with Opx (Fig. 2A)	0.0314	3.8	0.00493	1.8	0.24	0.0461	3.7	31.69	0.57	31.4	1.2	28	72	-1.0	-13.2	-0.3	4910
GRM_PX3-7.FIN2	anhedral; with Opx (Fig. 2A)	0.0340	4.1	0.00493	2.0	0.15	0.0494	4.0	31.72	0.64	34.0	1.3	170	86	6.7	81.3	1.8	4190
GRM371-1.FIN2	in Spr, Crd (Fig. 2B)	0.0347	6.6	0.00497	2.4	0.11	0.0516	6.8	31.96	0.80	34.6	2.2	240	140	7.6	86.7	1.2	1467
GRM28-3.FIN2	in Bt	0.0312	3.8	0.00500	2.2	0.00	0.0457	4.2	32.16	0.73	31.3	1.2	9	81	-2.7	-257.3	-0.7	4220
GRM_PX3-14.FIN2	anhedral; with Opx (Fig. 2A)	0.0497	3.0	0.00503	1.9	0.42	0.0714	2.8	32.35	0.60	49.3	1.5	969	58	34.4	96.7	11.3	9340
GRM_PX3-4.FIN2	anhedral; with Opx (Fig. 2A)	0.0337	3.9	0.00503	2.2	0.20	0.0481	4.0	32.35	0.72	33.6	1.3	120	83	3.7	73.0	1.0	3870
GRM_PX3-10.FIN2	anhedral; with Opx (Fig. 2A)	0.0340	4.1	0.00503	2.2	-0.03	0.0487	4.5	32.35	0.72	34.0	1.4	141	91	4.9	77.1	1.2	3760
GRM13-1.FIN2	in Spr, Crd (Fig. 2B)	0.0333	4.2	0.00504	1.9	0.13	0.0481	4.6	32.37	0.62	33.3	1.4	108	90	2.8	70.0	0.7	2903
GRM371-3.FIN2	in Bt	0.0333	6.3	0.00505	2.4	-0.07	0.0463	6.7	32.47	0.77	33.5	2.1	70	120	3.1	53.6	0.5	1559
GRM29b1-2.FIN2	in Bt	0.0346	4.9	0.00509	2.4	0.11	0.0492	4.9	32.70	0.74	34.5	2.1	178	100	5.2	81.6	1.1	2076
GRM371-2.FIN2	in Spr, Crd (Fig. 2B)	0.0329	6.4	0.00513	2.5	-0.13	0.0465	6.7	32.99	0.86	32.8	2.1	40	120	-0.6	17.5	-0.1	1375
GRM_PX3-13.FIN2	anhedral; with Opx (Fig. 2A)	0.0633	3.0	0.00518	1.7	0.26	0.0889	2.8	33.29	0.58	62.3	1.8	1402	54	46.6	97.6	16.1	9910
GRM28-4.FIN2	in Bt	0.0623	3.4	0.00525	2.3	0.36	0.0870	3.3	33.78	0.79	61.3	2.0	1353	67	44.9	97.5	13.8	5740
GRM16-4.FIN2	in Pl	0.0510	3.9	0.00484	1.9	0.31	0.0456	3.7	31.15	0.59	31.0	1.2	-5	72	-0.5	723.0	-0.1	6240
GRM17-2.FIN2	in Kfs	0.0529	6.1	0.00493	2.4	0.35	0.0488	5.7	31.70	0.77	32.9	2.0	150	110	3.6	78.9	0.6	1751
GRM27-3.FIN2	euohedral; in Pl (Fig. 2C)	0.0338	5.0	0.00497	2.4	0.42	0.0492	4.7	31.95	0.75	33.7	1.6	150	92	5.2	78.7	1.1	1971
GRM2_2-3.FIN2	intergrown with Opx	0.0335	5.1	0.00497	2.6	0.07	0.0491	5.5	31.99	0.82	33.4	1.7	150	110	4.2	78.7	0.8	2022
GRM16-1.FIN2	in Pl	0.0311	5.5	0.00501	2.4	0.16	0.0451	5.8	32.20	0.79	31.1	1.7	30	110	-3.5	-7.3	-0.6	2018
GRM2_2-4.FIN2	intergrown with Opx	0.0340	5.6	0.00502	2.2	0.26	0.0500	5.2	32.25	0.73	34.0	1.8	180	110	5.1	82.1	1.0	1612
GRM16-2.FIN2	in Pl	0.0337	5.0	0.00502	2.4	0.30	0.0489	4.9	32.26	0.78	33.6	1.7	151	100	4.0	78.6	0.8	2857
GRM2_1-2.FIN2	in Crd	0.0326	6.1	0.00503	2.0	0.21	0.0489	5.5	32.36	0.67	32.6	1.9	80	110	0.7	59.6	0.1	1862
GRM2_2-6.FIN2	intergrown with Opx	0.0338	5.9	0.00504	2.4	0.21	0.0489	5.7	32.46	0.80	33.8	1.9	160	110	4.0	79.7	0.7	2015
GRM2_1-1.FIN2	in Crd	0.0310	6.8	0.00506	2.6	0.17	0.0447	6.5	32.53	0.80	30.9	2.0	-30	120	-5.3	208.4	-0.8	1988
GRM27-2.FIN2	euohedral; in Pl (Fig. 2C)	0.0340	5.6	0.00507	2.2	-0.06	0.0488	6.1	32.57	0.72	34.0	1.9	150	120	4.2	78.3	0.8	1988
GRM2_2-2.FIN2	intergrown with Opx	0.0316	5.4	0.00508	2.2	0.09	0.0447	5.4	32.68	0.72	31.6	1.6	-39	99	-3.4	183.8	-0.7	2232
GRM17-1.FIN2	in Kfs	0.0342	6.4	0.00509	2.8	0.21	0.0488	6.4	32.73	0.87	34.1	2.1	160	120	4.0	79.5	0.7	1753
GRM16-3.FIN2	in Pl	0.0322	4.3	0.00509	2.4	0.25	0.0465	4.1	32.74	0.76	32.2	1.4	34	79	-1.7	3.7	-0.4	3510
GRM6_1-1.FIN2	in Spr (Fig. 2D)	0.0495	3.6	0.00516	1.9	0.33	0.0689	3.5	33.15	0.63	49.0	1.8	890	67	32.3	96.3	8.8	4390
GRM2_2-1.FIN2	intergrown with Opx	0.0340	5.3	0.00516	2.3	0.28	0.0471	5.1	33.20	0.75	33.9	1.7	67	97	2.1	50.4	0.4	1805
GRM2_2-5.FIN2	intergrown with Opx	0.0314	7.0	0.00522	2.3	0.10	0.0443	7.0	33.55	0.80	31.3	2.2	-30	130	-7.2	211.8	-1.0	1798
GRM6_1-2.FIN2	euohedral; in Pl (Fig. 2C)	0.0342	5.3	0.00523	2.1	0.16	0.0475	5.5	33.61	0.71	34.1	1.8	80	110	1.4	58.0	0.3	1949
Retrogressed Granulite (Bro3-56-4)	in Spr (Fig. 2D)	0.0443	3.2	0.00523	1.8	0.15	0.0614	3.1	33.64	0.59	44.0	1.4	645	64	23.5	94.8	7.4	8010
GRM1-1.FIN2	with Chi, Bt, Ms	0.0327	4.6	0.00500	1.9	0.03	0.0470	4.7	32.15	0.62	32.6	1.5	72	92	1.4	55.3	0.3	3050
GRM33-1.FIN2	with Chi, Bt, Crd	0.0344	5.2	0.00515	2.3	0.27	0.0487	5.1	33.11	0.77	34.3	1.8	160	110	3.5	79.3	0.7	2066
GRM7-1.FIN2	in Crd	0.0348	6.0	0.00530	2.3	0.14	0.0491	6.1	34.10	0.77	34.7	2.0	140	120	1.7	75.6	0.3	1526
Leucogranulite (GR11-35)																		
GR1135_7-1.FIN2	On apatite	0.0319	3.1	0.00467	1.8	0.43	0.0495	2.8	30.04	0.52	31.9	1.0	550	66	5.7	94.5	1.8	12140
GR1135_9-1.FIN2	On apatite	0.0322	3.1	0.00470	1.9	0.31	0.0492	3.0	30.20	0.57	32.2	1.0	50	65	6.1	39.6	2.0	9840
GR1135_2-1.FIN2	Leucosome-b/t pl and Kfs	0.0356	3.6	0.00479	1.9	0.41	0.0502	3.4	30.83	0.59	33.5	1.2	107	71	8.0	71.2	2.2	6420
GR1135_1-1.FIN2	On apatite (Fig. 2F)	0.0307	2.8	0.00480	1.9	0.44	0.0469	2.6	30.89	0.60	30.7	0.9	-	-	-0.9	-	-	12280
GR1135_2-2.FIN2	Leucosome-b/t pl and Kfs	0.0353	3.7	0.00484	1.8	0.11	0.0532	3.9	31.14	0.54	35.3	1.3	-40	86	11.8	177.9	3.2	6680

Table 4.A2. Monazite U-Pb data.

Table 4.A2

Sample and Analysis Number	Textural setting	Isotopic Ratios										Dates (Ma)										207Pb/206Pb ±2σ	207Pb/235U ±2σ	207Pb/238U ±2σ	206Pb/238U ±2σ	Rho	207Pb/206Pb ±2σ	206Pb/238U ±2σ	% disc a	% disc b	disc c	conc U (ppm)
		207Pb/235U	±2σ	%	206Pb/238U	±2σ	%	Rho	207Pb/206Pb	±2σ	%	206Pb/238U	±2σ	%	207Pb/235U	±2σ	abs	206Pb/238U	±2σ	abs	6/38-7/35											
GR1155_3-1.FIN2	On apatite	0.0342	4.1	0.00487	2.1	-0.04	0.0512	4.5	31.32	0.66	34.2	1.4	340	97	8.4	90.8	2.1	4240														
GR1155_16-2.FIN2	On apatite	0.0623	11.7	0.00542	5.7	0.70	0.1100	9.1	34.90	1.90	80.0	9.0	170	190	56.4	79.3	5.0	4610														
GR1155_16-5.FIN2	On apatite	0.0981	7.1	0.00552	2.7	0.38	0.1302	6.9	35.51	0.96	94.8	6.5	170	120	62.5	79.1	5.1	4090														
GR1155_23-1.FIN2	On apatite	0.0345	3.0	0.00496	2.0	0.30	0.0512	3.5	31.88	0.64	34.5	1.2	210	76	7.7	84.8	2.0	7250														
GR1155_18-1.FIN2	On apatite	0.0350	3.5	0.00497	1.7	0.36	0.0489	2.9	31.98	0.56	33.0	1.0	130	61	2.9	75.4	1.0	12350														
GR1155_22-1.FIN2	On apatite, next to spinel	0.0481	4.2	0.00520	2.5	0.40	0.0686	4.1	33.46	0.82	47.6	2.0	1160	89	29.7	97.1	7.1	7660														
GR1155_3-2.FIN2	On apatite	0.0574	5.6	0.00536	2.4	0.48	0.0767	4.8	34.46	0.82	56.6	3.1	230	96	39.1	85.0	7.1	2570														
Calibration Standard F6																																
F6-1.FIN2		0.735	2.6	0.0999	2.2	0.53	0.0598	2.5	555	12	559	11	603	52	0.8	8.0	0.4	1240														
F6-2.FIN2		0.734	2.9	0.0909	2.0	0.60	0.0580	2.4	561	11	558	12	526	55	-0.4	-6.6	0.2	1246														
F6-3.FIN2		0.736	2.6	0.0909	1.9	0.52	0.0582	2.4	561	10	559	12	530	51	-0.3	-5.8	-0.1	1225														
F6-4.FIN2		0.723	2.6	0.0894	1.9	0.39	0.0586	2.6	552	10	552	11	558	56	0.0	1.1	0.0	1227														
F6-5.FIN2		0.753	2.7	0.0922	2.0	0.57	0.0583	2.6	568	11	569	12	542	56	0.2	-4.9	0.1	1192														
F6-6.FIN2		0.735	2.6	0.0901	2.3	0.41	0.0601	2.7	556	15	560	11	601	59	0.7	7.5	0.3	1222														
F6-7.FIN2		0.718	2.6	0.0886	2.1	0.65	0.0582	2.4	547	11	549	11	537	53	0.3	-1.9	0.2	1237														
F6-8.FIN2		0.733	2.6	0.0911	2.4	0.64	0.0595	2.4	562	13	559	11	579	52	-0.6	-2.9	-0.3	1245														
F6-9.FIN2		0.737	2.7	0.0911	2.1	0.55	0.0594	2.5	563	11	561	12	576	55	-0.4	2.3	-0.2	1215														
F6-10.FIN2		0.722	2.8	0.0900	2.0	0.60	0.0585	2.4	555	11	552	12	555	52	-0.7	0.0	-0.3	1228														
F6-13.FIN2		0.726	2.6	0.0915	2.3	0.50	0.0578	2.4	564	12	555	11	517	54	-1.7	-9.1	-0.8	1217														
F6-14.FIN2		0.729	2.7	0.0919	2.2	0.47	0.0573	2.4	567	12	556	11	499	57	-2.1	-13.6	-1.0	1213														
F6-15.FIN2		0.740	2.8	0.0911	2.2	0.57	0.0589	2.4	562	12	563	12	565	54	0.2	0.5	0.1	1198														
F6-16.FIN2		0.737	2.7	0.0908	2.2	0.59	0.0587	2.6	560	12	560	12	566	53	0.0	1.1	0.0	1229														
F6-17.FIN2		0.751	2.8	0.0911	2.1	0.54	0.0596	2.5	562	11	568	12	599	51	1.1	6.2	0.5	1220														
F6-18.FIN2		0.757	2.6	0.0930	2.0	0.67	0.0594	2.2	573	11	572	12	578	48	-0.2	0.9	-0.1	1183														
F6-19.FIN2		0.735	2.7	0.0905	2.2	0.51	0.0587	2.6	558	12	559	12	558	53	0.2	-	0.1	1230														
F6-20.FIN2		0.739	2.8	0.0908	2.1	0.53	0.0585	2.6	560	11	561	12	552	51	0.2	-1.4	0.1	1210														
F6-21.FIN2		0.739	2.7	0.0909	2.1	0.60	0.0592	2.2	561	11	562	12	575	50	0.3	2.5	0.1	1204														
F6-22.FIN2		0.737	2.7	0.0909	2.1	0.58	0.0583	2.4	560	11	561	12	537	52	0.3	-4.4	0.1	1246														
F6-23.FIN2		0.738	2.7	0.0901	2.0	0.52	0.0597	2.5	556	11	562	12	589	52	0.8	5.6	0.4	1230														
F6-24.FIN2		0.743	2.7	0.0910	2.2	0.43	0.0592	2.5	561	12	566	12	577	57	0.8	2.8	0.4	1237														
F6-25.FIN2		0.732	2.9	0.0908	2.1	0.63	0.0588	2.4	560	11	557	12	557	52	-0.6	-0.5	-0.3	1214														
F6-26.FIN2		0.737	3.0	0.0905	2.1	0.49	0.0591	2.7	559	11	560	13	572	58	0.2	2.3	0.1	1229														
F6-27.FIN2		0.732	2.9	0.0905	2.1	0.62	0.0581	2.6	558	11	558	12	535	55	0.0	-4.4	0.0	1229														
F6-28.FIN2		0.729	2.7	0.0901	2.1	0.56	0.0589	2.5	556	11	556	12	565	54	0.1	1.6	0.0	1228														
F6-29.FIN2		0.733	2.9	0.0897	2.2	0.64	0.0591	2.4	554	12	559	13	562	50	0.9	1.4	0.4	1211														
Validation Reference Material JUP																																
JUP-1.FIN2		0.103	2.8	0.0158	1.6	0.48	0.0471	2.8	101	2	100	3	63	56	-1.3	-60.2	-0.5	5260														
JUP-2.FIN2		0.103	2.5	0.0158	1.6	0.37	0.0474	2.5	101	2	100	2	66	51	-1.3	-53.2	-0.5	5360														
JUP-3.FIN2		0.103	2.6	0.0158	1.7	0.35	0.0471	2.5	101	2	100	3	59	53	-1.7	-71.5	-0.7	5340														
JUP-4.FIN2		0.103	2.6	0.0159	1.6	0.51	0.0477	2.3	101	2	100	3	86	90	-1.8	-17.9	-0.7	5280														
JUP-5.FIN2		0.103	2.7	0.0158	1.7	0.22	0.0463	2.7	101	2	102	3	113	96	0.6	10.6	0.2	5940														
JUP-6.FIN2		0.103	2.8	0.0159	1.7	0.41	0.0473	2.7	101	2	100	3	72	57	-1.9	-40.8	-0.7	5500														
JUP-7.FIN2		0.103	2.6	0.0158	1.7	0.23	0.0479	2.5	101	2	99	3	94	55	-1.5	-7.3	-0.6	5450														
JUP-8.FIN2		0.100	2.8	0.0157	1.7	0.28	0.0460	2.6	101	2	97	3	16	52	-3.5	-528.1	-1.3	5400														
JUP-9.FIN2		0.103	2.8	0.0157	1.6	0.21	0.0479	2.7	101	2	100	3	99	60	-0.8	-1.6	-0.3	5270														
JUP-10.FIN2		0.102	2.4	0.0158	1.6	0.21	0.0468	2.6	101	2	99	2	47	53	-2.2	-114.9	-1.0	5360														
JUP-13.FIN2		0.105	2.9	0.0157	1.7	0.39	0.0481	2.5	100	2	101	3	107	57	1.2	6.4	0.4	5380														
JUP-14.FIN2		0.106	2.6	0.0160	1.6	0.24	0.0482	2.5	102	2	102	3	101	55	-0.4	-1.3	-0.2	5260														
JUP-15.FIN2		0.103	2.7	0.0158	1.5	0.25	0.0472	2.5	101	2	99	3	67	54	-1.7	-50.3	-0.7	5400														
JUP-16.FIN2		0.103	2.8	0.0159	1.5	0.25	0.0479	2.7	101	2	98	3	88	58	-1.4	-15.2	-0.5	5240														
JUP-17.FIN2		0.101	2.7	0.0158	1.6	0.44	0.0466	2.6	101	2	100	3	41	50	-3.5	-147.1	-1.4	5180														
JUP-18.FIN2		0.103	2.5	0.0157	1.7	0.42	0.0478	2.5	101	2	99	2	89	53	-1.4	-12.9	-0.6	5330														
JUP-19.FIN2		0.102	2.9	0.0159	1.6	0.31	0.0468	2.8	102	2	98	3	45	56	-3.4	-125.8	-1.2	5560														
Validation Reference Material Moacyr																																
MCR-1.FIN2		0.687	2.9	0.0836	1.7	0.46	0.0591	2.7	518	9	550	12	562	59	2.4	7.9	1.0	1052														

Sample and Analysis Number	Textural setting	Isotopic Ratios				Dates (Ma)				207Pb/206Pb ±2σ	207Pb/235U ±2σ	207Pb/235U abs	207Pb/206Pb ±2σ	% disc a	% disc b	disc c	conc U (ppm)	
		207Pb/235U ±2σ %	206Pb/238U ±2σ %	Rho	207Pb/206Pb ±2σ %	206Pb/238U ±2σ	206Pb/238U	±2σ	abs									abs
MCR-3.FIN2		0.669	2.8	0.0822	1.7	0.48	0.0581	2.5	509	8	52.0	12	586	57	2.1	13.2	0.9	1064
MCR-3.FIN2		0.683	2.8	0.0833	1.7	0.46	0.0595	2.4	517	9	52.6	11	586	52	2.1	11.8	1.0	1052
MCR-4.FIN2		0.674	2.7	0.0831	1.6	0.36	0.0569	2.4	515	8	52.3	11	565	51	1.6	8.9	0.8	1077
MCR-5.FIN2		0.671	2.7	0.0839	1.8	0.39	0.0579	2.4	520	9	52.1	11	526	55	0.3	1.2	0.1	1091
MCR-6.FIN2		0.675	2.8	0.0832	1.8	0.65	0.0586	2.4	515	9	52.3	11	550	50	1.5	6.4	0.7	1092
MCR-7.FIN2		0.658	2.4	0.0826	1.7	0.39	0.0581	2.2	512	9	51.3	10	537	50	0.2	4.7	0.1	1090
MCR-8.FIN2		0.672	2.8	0.0838	1.7	0.60	0.0582	2.4	519	8	52.4	11	541	49	1.0	4.1	0.5	1087
MCR-9.FIN2		0.674	2.5	0.0826	1.9	0.56	0.0589	2.2	512	10	52.3	10	566	49	2.1	9.6	1.1	1084
MCR-10.FIN2		0.673	2.7	0.0828	1.9	0.41	0.0592	2.4	513	10	52.3	11	571	55	1.9	10.2	0.9	1104
MCR-11.FIN2		0.658	2.7	0.0821	1.8	0.40	0.0582	2.6	509	9	51.4	11	539	55	1.0	5.6	0.5	1100
MCR-12.FIN2		0.675	3.0	0.0829	1.8	0.39	0.0590	2.5	513	9	52.3	12	561	58	1.9	8.5	0.8	1085
MCR-13.FIN2		0.661	2.6	0.0831	1.8	0.54	0.0573	2.4	515	9	51.5	11	501	54	0.0	-2.7	0.0	1084
MCR-14.FIN2		0.675	2.4	0.0838	1.8	0.42	0.0582	2.4	518	9	52.3	10	534	52	0.9	2.9	0.5	1077
MCR-15.FIN2		0.679	2.8	0.0835	1.8	0.50	0.0590	2.4	517	9	52.6	11	568	54	1.8	9.0	0.9	1072
MCR-16.FIN2		0.679	2.8	0.0837	1.7	0.22	0.0588	2.6	518	8	52.7	11	562	59	1.6	7.8	0.8	1073
MCR-17.FIN2		0.658	2.7	0.0821	1.8	0.43	0.0582	2.4	509	9	51.3	11	534	55	0.8	4.8	0.4	1053
MCR-18.FIN2		0.678	2.8	0.0829	1.8	0.55	0.0589	2.4	513	9	52.5	11	564	56	2.3	9.0	1.1	1052
MCR-19.FIN2		0.660	2.7	0.0830	1.9	0.59	0.0582	2.4	514	9	51.5	11	547	51	0.2	6.0	0.1	1064
MCR-20.FIN2		0.676	2.8	0.0836	1.8	0.38	0.0585	2.7	518	9	52.5	12	543	59	1.4	4.7	0.6	1048
MCR-21.FIN2		0.657	2.6	0.0832	1.7	0.26	0.0566	2.5	515	9	51.2	10	466	56	-0.6	-10.5	-0.3	1041

Notes:
Rho = correlation coefficient between 206Pb/238U and 207Pb/235U ratios
% discordance a = $100 \cdot (100 \cdot (206\text{Pb}/238\text{U date}) / (207\text{Pb}/235\text{U date}))$
% discordance b = $100 \cdot (100 \cdot (206\text{Pb}/238\text{U date}) / (207\text{Pb}/206\text{Pb date}))$
discordance c = uncertainty weighted age difference (UWAD) = $(206\text{Pb}/238\text{U date}) - (207\text{Pb}/235\text{U date}) / (207\text{Pb}/206\text{Pb date})$
Colored text indicates analyses that were used in age calculations in Fig. 3. Text color corresponds to ellipse color.

Table 4.B1. Garnet LA-ICP-MS analytical methods.

Laboratory & Sample Preparation	
Laboratory name	The University of Kansas, Dept. of Geology, Isotope Geochemistry Lab
Sample type/mineral	Garnet
Sample preparation	40 µm-thick, polished sections
Imaging	BSE images and X-ray compositional maps
Laser ablation system	
Make, model & type	ATL ArF excimer laser (193 nm), Photon Machines Analyte G1
Ablation cell & volume	Frames Cell with teardrop-shaped insert
Laser wavelength	193 nm
Pulse width (ns)	5 ns
Fluence	3.2 J/cm ²
Repetition rate	10 Hz
Spot size (µm)	35 µm circle
Sampling mode / pattern	Single spots
Carrier gas	He, 0.425 l/min, Ar, 1.24 l/min
Ablation duration	25 s
Cell carrier gas flow	He
ICP-MS Instrument	
Make, Model & type	Thermo Element2 magnetic sector field ICP-MS (single collector)
Sample introduction	Aerosol with sample + He was mixed with Ar using a Y-connector 15 cm upstream from torch.
RF power	1100 W
Make-up gas flow	Ar, 16 l/min
Sampling depth	20 µm
Detection system	Single detector (SEM), counting & analog modes
Elements/ isotopes analyzed	29Si, 43Ca, 44Ca, 45Sc, 49Ti, 89Y, 94Zr, 139La, 140Ce, 141Pr, 146Nd, 147Sm, 153Eu, 157Gd, 159Tb, 163Dy, 165Ho, 166Er, 169Tm, 172Yb, 175Lu
Integration time per peak (Sample Time in milliseconds)	29Si=10, 43Ca=10, 44Ca=10, 45Sc=10, 49Ti=10, 89Y=10, 94Zr=10, 139La=30, 140Ce=30, 141Pr=30, 146Nd=20, 147Sm=20, 153Eu=20, 157Gd=20, 159Tb=20, 163Dy=20, 165Ho=20, 166Er=10, 169Tm=10, 172Yb=10, 175Lu=10
Total integration time (Segment Duration in milliseconds)	29Si=10, 43Ca=10, 44Ca=10, 45Sc=10, 49Ti=10, 89Y=10, 94Zr=10, 139La=30, 140Ce=30, 141Pr=30, 146Nd=20, 147Sm=20, 153Eu=20, 157Gd=20, 159Tb=20, 163Dy=20, 165Ho=20, 166Er=10, 169Tm=10, 172Yb=10, 175Lu=10
Total method time	1 min 47 s (100 runs, 2 passes)

Table 4.B1

Sensitivity (cps/ppm)	29Si=122, 43Ca=12, 44Ca=199, 45Sc=5868, 49Ti=447, 89Y=7822, 94Zr=3915, 139La=13000, 140Ce=16723, 141Pr=16659, 146Nd=2615, 147Sm=2153, 153Eu=8914, 157Gd=1791, 159Tb=11996, 163Dy=2985, 165Ho=11585, 166Er=3939, 169Tm=12062, 172Yb=2655, 175Lu=10073
ICP Dead time	22 ns
UO+/U+	0.02%
238U+/232Th+	~1.5
Data Processing	
Gas blank	18 s
Calibration strategy	Standard-sampling bracketing + internal standardization using Ca concentration measured by electron microprobe for each spot
Reference material info	GSD-1G (Guillong et al., 2005; Jochum et al., 2005)
Internal std for trace elements	44Ca
Data processing package used	IGOR PRO, lolite 2.5 (Patton et al. 2011); Trace elements: TraceElementsIS data reduction scheme
Common-Pb correction, composition and uncertainty	n/a
Uncertainty level & propagation	2SE internal uncertainty
Reproducibility	3-8%
Quality control / Validation	GSE-1G (Guillong et al., 2005; Jochum et al., 2005)
Other information	

Table 4.B2. Garnet trace element data.

Analysis	Ca44 (cps)	±2SE (cps)	Si (ppm)	±2SE (ppm)	Ca (ppm)	±2SE (ppm)	Sc (ppm)	±2SE (ppm)	Ti (ppm)	±2SE (ppm)	Y (ppm)	±2SE (ppm)	Zr (ppm)	±2SE (ppm)	La (ppm)	±2SE (ppm)	Ce (ppm)	±2SE (ppm)	Pr (ppm)	±2SE (ppm)
GRMP3-1.FIN2	8.5E+05	2.2E+04	2.5E+05	1.4E+04	8720	300	38.9	0.97	131	4.6	339	14	52.7	2.0	0.003	0.002	0.007	0.006	0.004	0.002
GRMP3-3.FIN2	9.5E+05	4.3E+04	2.2E+05	1.2E+04	9990	620	38.4	2.4	123	9.7	753	89	50.9	4.2	0.125	0.072	0.370	0.210	0.042	0.020
GRMP3-4.FIN2	1.2E+06	4.4E+04	2.2E+05	1.5E+04	10810	660	52.8	2.1	156	8.9	1572	80	75.5	3.6	0.006	0.005	0.013	0.002	0.003	0.001
GRMP3-6.FIN2	1.3E+06	3.0E+04	2.4E+05	9.9E+03	12340	400	68.3	1.5	96.1	4.0	1739	48	53.8	1.6	0.008	0.002	0.018	0.003	0.005	0.002
GRMP3-10.FIN2	1.2E+06	4.2E+04	2.5E+05	1.1E+04	12790	450	108	3.9	14.3	1.0	2400	120	30.9	1.2	0.035	0.006	0.096	0.014	0.012	0.003
GRMP3-14.FIN2	1.2E+06	4.3E+04	2.5E+05	1.3E+04	13140	470	108	2.5	38.7	2.4	1842	63	51.3	1.4	0.028	0.004	0.084	0.007	0.011	0.002
GRMP3-15.FIN2	1.2E+06	9.1E+04	2.6E+05	3.2E+04	12330	850	102	6.8	110	13	1840	130	86.4	5.9	0.067	0.048	0.139	0.089	0.018	0.008
GRMP3-16.FIN2	1.2E+06	4.5E+04	2.2E+05	1.4E+04	10880	480	82.9	3.0	104	2.9	1214	46	72.8	2.5	0.008	0.002	0.032	0.008	0.006	0.002
GRMP3-17.FIN2	1.3E+06	3.0E+04	2.2E+05	8.7E+03	10540	350	52.8	1.6	162	5.1	779	27	79.0	2.3	0.011	0.005	0.032	0.010	0.007	0.002
GRMP3-18.FIN2	1.2E+06	4.3E+04	1.9E+05	1.3E+04	9020	470	53.4	2.1	73.4	4.0	920	52	43.9	2.3	0.079	0.025	0.177	0.039	0.028	0.007
GRMP3-19.FIN2	1.3E+06	3.6E+04	1.7E+05	9.8E+03	8550	280	50.5	1.1	38.9	1.7	1112	36	31.7	0.95	0.018	0.005	0.048	0.011	0.007	0.001
GRMP3-20.FIN2	1.2E+06	5.0E+04	2.6E+05	1.6E+04	12070	390	96.1	1.9	140	7.0	1387	68	62.0	2.9	0.009	0.007	0.037	0.012	0.010	0.003
GRMP3-21.FIN2	1.2E+06	3.3E+04	2.6E+05	9.7E+03	12240	230	85.6	2.7	127	6.6	1510	68	56.0	4.5	0.004	0.002	0.010	0.003	0.005	0.002
GRMP3-22.FIN2	1.2E+06	3.9E+04	2.6E+05	1.0E+04	12390	300	85.3	1.9	149	6.1	1437	35	55.1	1.3	0.012	0.007	0.032	0.016	0.011	0.004
GRMP3-23.FIN2	1.3E+06	3.8E+04	2.5E+05	1.0E+04	12110	340	104	2.7	148	4.4	1766	67	74.6	2.6	0.003	0.001	0.011	0.002	0.006	0.002
GRMP3-24.FIN2	1.3E+06	4.1E+04	2.6E+05	1.1E+04	12880	310	147	4.2	134	5.4	1896	81	59.9	2.1	0.013	0.003	0.039	0.005	0.013	0.002
GRMP3-25.FIN2	1.4E+06	5.4E+04	2.4E+05	1.4E+04	12500	590	163	6.1	102	5.7	1962	87	67.0	2.1	0.008	0.003	0.020	0.004	0.010	0.003
GRMP3-26.FIN2	1.3E+06	4.4E+04	2.4E+05	8.8E+03	12210	320	137	2.4	128	5.0	1425	39	44.8	1.4	0.005	0.002	0.011	0.003	0.008	0.002
GRMP3-27.FIN2	1.3E+06	5.7E+04	2.5E+05	1.5E+04	12480	460	82.9	2.5	163	9.5	1488	82	75.0	4.6	0.001	0.002	0.009	0.004	0.006	0.002
GRMP3-28.FIN2	1.3E+06	4.7E+04	2.2E+05	9.0E+03	10970	380	52.8	0.95	122	4.3	976	30	65.9	1.7	0.003	0.002	0.021	0.004	0.004	0.001
GRMP3-29.FIN2	1.3E+06	4.4E+04	1.9E+05	1.1E+04	9250	290	39.1	1.1	167	7.2	916	35	67.0	2.3	0.005	0.002	0.012	0.005	0.005	0.002
GRMP3-30.FIN2	1.3E+06	4.6E+04	2.5E+05	1.5E+04	12510	340	113	2.8	26.7	2.1	2550	100	44.3	1.6	0.019	0.004	0.043	0.005	0.005	0.002
GRMP3-31.FIN2	1.3E+06	3.5E+04	2.5E+05	1.2E+04	12590	500	105	2.1	38.8	2.1	2148	83	59.6	2.0	0.038	0.007	0.117	0.020	0.016	0.003
GRMP3-32.FIN2	1.3E+06	4.3E+04	2.3E+05	1.0E+04	10730	360	86.6	1.7	32.9	1.8	1798	70	42.4	1.4	0.048	0.010	0.125	0.028	0.019	0.005
GRMP3-33.FIN2	1.4E+06	4.9E+04	1.9E+05	8.3E+03	9340	240	67.5	1.6	240	1.3	1582	64	29.3	1.1	0.129	0.015	0.338	0.035	0.042	0.005
GRMP3-34.FIN2	1.3E+06	4.9E+04	1.9E+05	9.4E+03	9300	260	64.9	1.8	47.0	2.6	1105	40	49.7	2.1	0.048	0.005	0.147	0.017	0.024	0.005
GRMP3-35.FIN2	1.4E+06	4.1E+04	2.4E+05	1.2E+04	11730	500	91.2	2.1	54.7	2.4	1377	60	69.1	2.4	0.054	0.015	0.177	0.041	0.027	0.007
GRMP3-36.FIN2	1.4E+06	4.4E+04	2.4E+05	1.2E+04	12360	380	86.2	2.3	38.7	2.7	1703	58	46.8	1.7	0.114	0.031	0.349	0.058	0.038	0.006
GRMP3-37.FIN2	1.3E+06	3.7E+04	2.3E+05	1.1E+04	11910	380	120	2.7	142	6.2	1689	74	61.7	2.4	0.002	0.002	0.010	0.002	0.006	0.001
GRMP3-38.FIN2	1.3E+06	4.1E+04	2.4E+05	1.4E+04	12140	380	80.0	1.9	146	6.2	1763	69	70.7	3.1	0.004	0.002	0.017	0.006	0.005	0.002
GRMP3-39.FIN2	1.3E+06	6.2E+04	2.2E+05	2.4E+04	10320	580	47.7	0.90	136	12.0	623	39	62.1	2.9	0.120	0.032	0.460	0.170	0.065	0.018
GRMP3-40.FIN2	1.2E+06	5.6E+04	1.9E+05	1.0E+04	8840	480	48.0	1.8	123	9.9	1643	59	48.9	2.0	0.016	0.006	0.064	0.022	0.010	0.003

Table 4.B2

Nd (ppm)	±2SE (ppm)	Sm (ppm)	±2SE (ppm)	Eu (ppm)	±2SE (ppm)	Gd (ppm)	±2SE (ppm)	Tb (ppm)	±2SE (ppm)	Dy (ppm)	±2SE (ppm)	Ho (ppm)	±2SE (ppm)	Er (ppm)	±2SE (ppm)	Tm (ppm)	±2SE (ppm)	Yb (ppm)	±2SE (ppm)	Lu (ppm)	±2SE (ppm)
0.095	0.028	1.82	0.11	0.212	0.021	21.9	0.6	8.6	0.3	67.6	3.0	11.7	0.3	25.6	1.2	2.8	0.2	14.6	0.4	1.9	0.1
0.290	0.200	2.19	0.22	0.239	0.019	29.1	0.9	12.8	0.7	119	5.7	24.1	2.0	58.3	5.9	6.1	0.3	29.4	1.8	3.0	0.1
0.066	0.030	4.21	0.64	0.373	0.043	47.6	1.8	19.0	1.2	197	14.0	49.7	2.8	141	18.7	18.7	0.8	108	5.2	14.6	0.5
0.120	0.022	3.31	0.26	0.322	0.022	37.9	1.6	17.7	0.6	202	7.2	53.0	1.9	165	5.4	24.3	0.8	157	4.6	23.4	0.9
0.118	0.017	0.75	0.06	0.180	0.017	18.1	0.7	17.3	0.6	260	12.0	73.6	3.3	229	9.5	34.6	1.5	226	9.4	31.0	1.3
0.127	0.019	1.51	0.08	0.230	0.016	32.2	0.8	20.7	0.9	229	9.6	57.2	2.3	178	7.4	26.5	0.9	176	5.0	26.2	1.2
0.192	0.074	3.33	0.48	0.376	0.043	49.9	4.1	22.8	4.1	238	33.0	59.2	3.5	175	11	24.0	2.0	137	9.6	18.7	1.4
0.104	0.033	2.49	0.23	0.309	0.021	41.5	1.9	17.7	0.8	168	7.6	38.6	1.5	94.5	3.0	11.5	0.4	62.5	3.0	7.5	0.3
0.146	0.035	4.30	0.15	0.336	0.018	43.2	1.1	16.8	0.5	137	3.6	23.4	0.7	48.1	1.1	4.8	0.2	24.1	0.9	2.5	0.1
0.209	0.056	1.97	0.18	0.231	0.034	30.0	0.7	12.4	0.6	123	7.0	28.8	1.9	84.1	4.4	11.4	0.5	74.9	3.6	11.0	0.6
0.064	0.015	1.41	0.05	0.156	0.012	23.5	0.7	12.5	0.5	141	5.6	34.2	1.3	99.4	2.7	13.2	0.4	80.4	3.7	10.3	0.5
0.169	0.033	3.98	0.27	0.367	0.028	42.3	1.4	17.1	0.9	170	8.4	43.6	1.6	133	6.4	19.5	1.0	128	4.2	18.4	1.1
0.154	0.023	4.42	0.22	0.366	0.022	44.1	1.2	17.5	0.7	180	7.5	47.4	2.1	148	5.7	21.5	0.9	139	6.6	20.0	1.0
0.259	0.042	5.55	0.21	0.391	0.023	43.5	1.2	17.9	0.6	180	7.0	46.2	1.4	141	3.2	19.4	0.5	121	3.8	16.2	0.6
0.297	0.041	5.51	0.22	0.396	0.025	47.0	1.3	20.4	0.8	215	6.9	54.2	1.9	172	5.6	24.0	0.8	152	4.3	22.0	0.9
0.346	0.035	5.69	0.23	0.402	0.021	46.2	1.7	21.3	1.0	223	9.5	57.4	2.5	181	6.7	26.3	0.9	172	6.5	26.1	1.3
0.392	0.045	5.70	0.42	0.385	0.024	47.6	1.5	20.7	1.1	221	13.0	57.3	2.9	183	6.8	27.3	1.1	199	8.0	29.6	1.0
0.255	0.043	4.45	0.18	0.362	0.019	40.4	1.4	18.0	0.7	180	7.3	43.9	1.5	135	3.5	19.2	0.5	129	3.6	19.0	0.7
0.144	0.041	4.71	0.31	0.369	0.034	46.7	1.0	19.1	0.5	191	3.8	46.0	2.2	142	5.3	19.5	0.8	124	4.3	18.1	0.9
0.099	0.019	3.99	0.19	0.351	0.023	41.4	1.3	16.4	0.5	151	6.2	31.7	1.2	79.8	2.2	9.0	0.2	46.3	1.7	5.1	0.2
0.092	0.018	3.31	0.18	0.275	0.020	39.4	1.3	16.1	0.7	150	5.8	29.6	1.0	67.2	2.7	7.1	0.3	36.2	1.4	3.8	0.2
0.095	0.018	0.77	0.09	0.183	0.016	22.4	0.8	20.2	0.7	289	11.0	81.8	3.2	249	7.9	33.7	1.0	205	7.9	26.6	1.0
0.120	0.024	0.96	0.07	0.219	0.019	31.2	0.6	23.8	0.8	282	11.0	69.1	2.6	213	5.5	32.3	0.7	234	6.9	37.2	1.4
0.154	0.052	0.84	0.09	0.180	0.017	27.1	1.3	20.2	0.9	226	9.2	55.2	1.9	156	5.2	20.4	0.6	123	3.6	16.5	0.7
0.198	0.029	0.71	0.07	0.141	0.010	21.8	0.9	17.3	0.7	208	8.3	50.7	1.5	141	5.0	18.6	0.5	115	3.6	15.9	0.7
0.109	0.020	0.95	0.05	0.205	0.015	30.6	1.1	17.3	0.7	162	8.0	34.0	1.3	85.9	2.7	10.1	0.5	56.7	2.6	6.6	0.3
0.174	0.037	1.21	0.10	0.250	0.016	37.6	1.2	22.2	1.1	213	9.4	44.2	2.0	113	2.4	14.5	0.3	85.0	2.4	11.1	0.4
0.232	0.051	0.96	0.08	0.188	0.010	29.6	1.1	21.3	0.9	238	8.4	54.2	1.9	144	5.5	17.7	0.5	100	3.7	11.8	0.5
0.319	0.031	5.25	0.20	0.422	0.020	45.2	1.3	20.3	0.9	206	8.4	52.5	2.0	157	5.4	21.4	0.7	134	4.0	18.3	0.7
0.213	0.025	5.18	0.24	0.391	0.021	46.6	1.6	20.5	0.7	215	8.0	55.1	2.2	172	6.1	24.3	0.9	157	6.0	21.1	1.0
0.374	0.095	3.22	0.28	0.271	0.021	30.6	1.3	12.2	0.2	105	6.5	19.7	1.1	47.2	2.5	5.3	0.4	28.5	1.4	3.6	0.3
0.124	0.040	1.81	0.08	0.189	0.030	29.3	1.3	15.2	0.6	174	8.9	51.7	2.4	184	7.3	29.6	0.8	211	11	32.3	2.0

CHAPTER 5

Constraining the pressure-temperature evolution and geodynamic setting of UHT granulites and migmatitic paragneisses of the Gruf Complex, Central Alps

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Abstract

Thermodynamic modeling of microdomains and whole-rock compositions is used to constrain the pressure-temperature (P-T) evolution of sapphirine granulites and migmatitic paragneisses from the Gruf Complex of the Central Alps. The P-T paths and conditions estimated from granulite microtextures and whole-rock compositions are consistent with one another, indicating that the estimates from both approaches are accurate. The sapphirine granulites were heated to ultra-high temperature (UHT) conditions of 900–1000°C and 7.0–9.5 kbar as they decompressed from ca. 800°C and 9–12 kbar, resulting in garnet breakdown. Nearly isothermal decompression led to the development of cordierite-bearing coronae and symplectites. By ca. 27 Ma, the lower crustal granulites and charnockites had been exhumed to the midcrustal level of migmatitic gneisses, which were undergoing peak metamorphism at ca. 675–750°C and 5–7 kbar. These results are consistent with a geodynamic model that invokes heat advection to the lower

crust closely following the continental-subduction (ultra-high pressure) stage of the Alpine orogeny. The most plausible geodynamic model consistent with the results of this study is breakoff of a southward subducting lithospheric slab, resulting in asthenospheric upwelling.

Introduction

Collisional orogens like the Alps affect global and regional climate and host major earthquakes and magmatism. Deep geodynamic processes that occur in these settings have some control on the surface dynamics such as earthquakes. The thermal structure of orogens also regulates the volume and type of magmatism produced. Metamorphic rocks that resided in the deep crust during orogenesis and were subsequently exhumed to the surface can provide important records of the thermal and geodynamic history of the orogens in which they are found.

Ultra-high temperature (UHT) metamorphism is of particular interest in understanding the thermal evolution of orogens because of the extreme crustal conditions it requires (ca. 900–1100°C and 7–13 kbar; Harley, 1998; Brown, 2007; Harley, 2008). During the Archean and Proterozoic eons, UHT metamorphism was more widespread than it has been during Phanerozoic time due to secular cooling of the Earth (Brown, 2006; Brown, 2007). Therefore, attaining UHT conditions within the Phanerozoic eon requires geodynamic processes that result in either a high geotherm (Kelsey, 2008), advection of heat to the crust from the mantle (Collins, 2002; Sizova et al., 2010; Clark et al., 2011), radiogenic heating (Clark et al., 2011; Kelsey and Hand, 2015), or shear heating (Kelsey, 2008). Constraining the pressure-temperature-time (P-T-t) history of Phanerozoic UHT rocks can be used to elucidate the nature, timing, and intensity of these geodynamic processes. However, determining the peak metamorphic conditions experienced by UHT rocks can be challenging because of retrograde elemental exchange and overprinting (e.g. Carswell and O'Brien, 1993; Fitzsimons and Harley, 1994; Harley, 1998; Pattison et al., 2003). Therefore,

techniques such as equilibrium phase diagram (“pseudosection”) modeling (e.g. Pattison et al., 2003; O'Brien, 2008) and accessory mineral thermometry (e.g. Watson et al., 2006; Luvizotto et al., 2009; Blackburn et al., 2012; Kooijman et al., 2012) must be employed to constrain the P-T evolution of high-grade rocks.

The Gruf Complex is enigmatic in that it hosts the only occurrence of UHT rocks in the Central Alps (Figure 5.1) (Bousquet et al., 2012a). Therefore, understanding whether UHT metamorphism occurred during the Alpine orogeny and what processes led to it has important implications for the geodynamic history of the Central Alps. In this study, we summarize geochronological data and present the results of thermodynamic modeling to constrain the P-T-t history of UHT granulites and upper amphibolite facies migmatites from the Gruf Complex. The granulites are well suited for this study because they contain abundant reaction textures, which are used to constrain the P-T path. These results are used to propose a model for the geodynamic evolution of the Central Alps.

Geological Setting

The European Central Alps comprise nappes that were thrust together during the Alpine orogeny (Figure 5.1), which culminated in the collision of the European and Apulian plates, which began at ca. 40 Ma (Schmid et al., 1996; Pfiffner, 2016). The units that make up the Central Alps were derived from the European continental margin (Adula Nappe and Gruf Complex) and the Penninic units, which comprise the Valais (Chiavenna Ophiolite and Misox Zone) and Piemont-Liguria (Malenco unit) ocean basins and the Briançonnais (Tambo and Suretta Nappes) microcontinent (Figure 5.1). Part of the Central Alps (i.e. the Southern Steep Belt; SSB in Figure 5.1) has been referred to as a lithospheric mélange (Trommsdorff, 1990) or tectonic accretion channel (Engi et al., 2001) in which slivers of crust (nappes) with different paleogeographic

affinities and metamorphic histories were amalgamated along shear zones during the late stages of the Alpine orogeny. Prior to the collisional event and nappe stacking, the Adula nappe and parts of the Penninic units underwent high pressure metamorphism, ranging from blueschist facies conditions in the north to ultra-high pressure (UHP) conditions in the southern Adula nappe (e.g. Gebauer, 1996; Bousquet et al., 2002; Dobrzhinetskaya et al., 2002; Hermann et al., 2006; Berger and Bousquet, 2008), between ca. 45 and 35 Ma (e.g. Becker, 1993; Gebauer, 1996; Brouwer et al., 2005; Hermann et al., 2006; Berger and Bousquet, 2008). This was a result of subduction of the ocean basins, Briançonnais microcontinent, and edge of the European continental margin (e.g. Stampfli et al., 1998; Berger and Bousquet, 2008). After exhumation of the (U)HP units, a large part of the Central Alps (the Lepontine Dome) experienced Barrovian metamorphism between 32 and 22 Ma (Rubatto et al., 2009), ranging from greenschist facies in the north to sillimanite-zone migmatization in the south (Todd and Engi, 1997; Frey and Ferreiro Mählmann, 1999; Burri et al., 2005). In the southeastern Central Alps, the granodiorite to tonalite Bergell Intrusion was emplaced along the Insubric line, which is part of the Periadriatic fault system that represents the suture between the European and Apulian (or Adriatic) plates, between 32 and 30 Ma (von Blanckenburg, 1992; Samperton et al., 2015) (Figure 5.1). Magmatism in the Central Alps waned after the intrusion of the Novate leucogranite at 23.8 ± 0.1 Ma (Oalman et al., in preparation-b).

The Gruf Complex occurs at the eastern extent of the Central Alps (Figure 5.1) where it is bordered to the north by the Chiavenna ophiolite (Valais) and Tambo nappes (Briançonnais). In the south and east, it is structurally overlain by the Bergell Intrusion (Berger et al., 1996; Davidson et al., 1996). The normal-sense Forcola fault separates the Gruf Complex from the Adula nappe, which experienced UHP metamorphism (Heinrich, 1986; Gebauer, 1996; Dobrzhinetskaya et al., 2002; Krogh Ravna and Terry, 2004), in the west (Ciancaleoni and Marquer, 2006; Bousquet et

al., 2012b; Galli et al., 2013). The Novate leucogranite intruded the southwestern part of the Gruf Complex along the Forcola fault at 24 ± 1 Ma (Liati et al., 2000; Oalmann et al., in preparation-b).

The Gruf Complex comprises mostly migmatitic biotite-bearing felsic orthogneisses with more scarce garnet + sillimanite + biotite \pm muscovite metapelitic paragneisses and leucogranite gneisses (detailed mapping by Galli et al., 2013; Figure 5.2). Charnockitic sheets are separated from the migmatites by local, up to 50 cm-wide, mylonitic shear zones with reverse-sinistral shear sense (Galli et al., 2013). The two largest of these charnockite bodies outcrop along a high ridge at the center of the Gruf Complex (Figure 5.2). Scarce Mg- and Al-rich, residual sapphirine-bearing granulites occur as pods, rafts, and schlieren within the charnockites. Variably deformed felsic dikes, ranging from biotite-granite to beryl + garnet + biotite + muscovite pegmatite, occur throughout the Gruf Complex, and the older, muscovite-free and muscovite-poor dikes are deformed whereas the younger, muscovite-rich dikes are undeformed (Oalmann et al., in preparation-b).

In Oalmann et al. (in preparation-a), we used texturally-controlled U-Pb monazite dating to conclude that UHT metamorphism of the granulites occurred during Alpine orogenesis, with monazite dates constraining the timing of peak to post-peak metamorphism to between 34 and 30 Ma. In (Oalmann et al., in preparation-b), we used texturally-controlled zircon petrochronology to conclude that there was minor zircon growth in equilibrium with garnet at 34.8 ± 1.1 Ma, zircon was resorbing during UHT metamorphism, and zircon rims crystallized during early retrograde metamorphism and melt crystallization (ca. 700–800°C) at 32.7 ± 0.7 Ma. We used the U-Pb zircon ages of variably deformed felsic dikes to conclude that the UHT granulite/charnockite units were juxtaposed against the migmatitic ortho- and paragneisses along mylonitic shear zones between 30

and 27 Ma, that contractional deformation ended by 25.6 ± 0.3 Ma, and that decompression led to crustal melting between 26 and 24 Ma (Oalman et al., in preparation-b). In another study (Oalman et al., submitted), we determined that the granulites and charnockites cooled from 700 to 420°C between ca. 30 and 19 Ma using U-Pb rutile thermochronology.

Sample Descriptions

Leucogranulite (GR11-39)

Leucogranulite sample GR11-39 is a felsic granulite characterized by alternating light and dark bands, which define a gneissosity (Figure 5.3A, 4A). Quartz within the leucosomes (light bands) is characterized by a bimodal grain size distribution. Large quartz ribbons (up to 1 cm long) exhibit undulose extinction and are elongate subparallel to the gneissosity. Smaller quartz grains are equant and have lobate grain boundaries. Blocky plagioclase grains are elongate subparallel to the gneissosity and lineation defined by the large quartz porphyroclasts. Lens-like K-feldspar porphyroclasts are also elongate subparallel to the gneissosity.

The melanosomes consist of biotite, garnet, orthopyroxene, sapphirine, and minor spinel and rutile. Biotite occurs as elongate blades that are oriented subparallel to the gneissosity. Anhedral to subhedral garnet porphyroblasts (up to 1 mm in diameter) commonly have embayments filled with individual grains or symplectites of plagioclase \pm orthopyroxene \pm biotite \pm sapphirine \pm spinel (Figure 5.5A). Outside of the garnet embayments, blocky orthopyroxene occurs as isolated porphyroblasts. Up to 3 mm-long aggregates of sapphirine (\pm spinel) are aligned subparallel to the gneissosity and are typically isolated from the other mafic minerals except for biotite and occur within the leucosomes in some cases (Figure 5.4A). The relative proportion of the phases varies from layer to layer. Some melanosome layers consist mainly of elongate biotite with or without sapphirine \pm spinel aggregates, whereas other layers are dominated by blocky

orthopyroxene. Garnet occurs as isolated porphyroblasts within the melanosomes and more scarcely within the leucosomes.

In this study, a garnet embayment filled with symplectites of orthopyroxene, plagioclase, sapphirine, and spinel (Figure 5.5A) is investigated. This garnet-breakdown texture is surrounded by quartz + K-feldspar, which represent crystallized melt. This type of texture is common in the leucogranulite samples and, to a lesser extent, in the charnockites.

Residual Granulite (Gruf100)

Residual granulite sample Gruf100 is a garnet-rich orthopyroxene + sapphirine + biotite + sillimanite + cordierite granulite that lacks a definite fabric (Figure 5.3B, 4B). The distribution of the constituent minerals at hand sample scale is homogeneous (Figure 5.3B).

Garnet porphyroblasts are subhedral and range from 3 mm to 2 cm in diameter (Figure 5.3B). The largest of these porphyroblasts are characterized by inclusion-rich cores and inclusion-poor rims (Figure 5.3B). Embayments in garnet are invariably filled with crystallized leucosome (quartz + feldspar), blocky orthopyroxene porphyroblasts, bladed sapphirine, biotite, cordierite, and/or symplectites consisting of orthopyroxene ± sapphirine ± cordierite (Figure 5.5B, C).

Subhedral orthopyroxene porphyroblasts range from 1–8 mm. In many cases, orthopyroxene and bladed sapphirine are intergrown (Figure 5.5C). In the more biotite-rich parts of the sample, biotite occurs along the edge and within cracks of orthopyroxene (Figure 5.5D).

Bladed sapphirine porphyroblasts range from <1 mm to 5 mm in length. These porphyroblasts are most common near garnet and orthopyroxene (Figure 5.5B, C), but also occur within leucosomes. There are no inclusions of sapphirine within the garnet.

Prismatic sillimanite porphyroblasts reach ca. 3 mm in length and are invariably (at least partly) surrounded by sapphirine + cordierite + minor spinel symplectites, which are separated

from orthopyroxene by cordierite coronae (Figure 5.5D). These textures likely formed from a reaction such as sillimanite + orthopyroxene = sapphirine + cordierite + spinel. In some cases, sillimanite occurs as aggregates made up of numerous small prisms that together form a blade-like shape (Figure 5.5D). This may be the result of sillimanite pseudomorphing kyanite.

Cordierite occurs in orthopyroxene-bearing symplectites or as coronae around and between garnet, orthopyroxene, sapphirine (Figure 5.5B, C) and sillimanite porphyroblasts (Figure 5.5D). In some areas of the thin section, cordierite is so abundant that the other mafic and aluminosilicate phases (garnet, orthopyroxene, sapphirine, and sillimanite) do not touch one another. We interpret these relationships to indicate that cordierite formed slightly after orthopyroxene and sapphirine.

Biotite mostly consists of small masses that surround orthopyroxene and garnet (Figure 5.5D). However, large aggregates of biotite are also present. Based on the relationship of biotite to the other phases, we interpret biotite to be a retrograde phase.

Leucosomes occur as discrete, ≤ 2 cm, white to gray pods comprising quartz + K-feldspar + plagioclase (Figure 5.3B, 4B). The feldspars in these leucosomes lack perthitic texture, and plagioclase and K-feldspar occur in roughly equal proportions. Minor blocky plagioclase also occurs outside of the leucosomes in garnet embayments.

Residual Granulite (Br03-56-2)

Another residual granulite sample with a slightly different bulk composition and mineral assemblage was also modeled. The whole-rock composition that was used for the input of this sample was taken from the Diplom Thesis of Schefer (2005). This residual granulite has a very similar appearance to the other residual granulite sample (Gruf100) in terms of its constituent minerals and their grain sizes, habits, and relationships to one another (Figure 5.4C). The main petrographic differences are that this sample lacks sillimanite (except for sillimanite exsolution

lamellae in orthopyroxene) and contains a higher proportion of spinel and biotite than sample Gruf100. Spinel occurs either as inclusions in or at the fringes of garnet or as fine-grained aggregates that fringe sapphirine (Figure 5.5E).

Residual granulite (GRM-37)

A monazite-bearing texture, which was extracted from a thin section of residual granulite sample Br03-56-2, contains orthopyroxene, sapphirine, and cordierite (Figure 5.5F). The monazite grain occurs with cordierite as part of a two-phase inclusion in sapphirine (Figure 5.5F). This monazite grain was dated by Oalman et al. (in preparation-a) using Laser Ablation Inductively Coupled Plasma Mass Spectrometry (LA-ICP-MS) and yields a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 32.5 ± 0.5 Ma based on three analyses.

Migmatitic paragneiss (GR11-23)

Sample GR11-23 is a garnet + sillimanite + biotite + muscovite gneiss has well-developed gneissic banding, defined by segregations of quartzofeldspathic and garnet-sillimanite-biotite-muscovite-rich layers (Figure 5.3C, 4D). This sample exhibits a well-developed crenulation cleavage defined by folding of the gneissic banding and folds within the mica-rich layers (Figure 5.3C, 4D). Muscovite, biotite, and fibrolitic sillimanite are the most abundant mineral in the dark layers, whereas garnet is scarcer (Figure 5.5G). Euhedral garnet porphyroblasts are elongate subparallel to the gneissosity in some cases, but in most cases garnet is equant (Figure 5.5G). Quartz has a bimodal size distribution in which the largest grains have chessboard-type undulose extinction and are elongate subparallel to the gneissosity. Smaller quartz grains have lobate grain boundaries typically have similar extinction angles. K-feldspar occurs as augen within the leucosome, and plagioclase typically has a blocky habit.

Migmatitic paragneiss (GR11-25)

Sample GR11-25 is a migmatitic garnet + sillimanite + biotite gneiss that has well-developed gneissosity, which is defined by segregation of quartzofeldspathic and garnet-, sillimanite-, and biotite-rich layers (Figure 5.3D, 4E). This sample consists of a higher proportion of leucosome than sample GR11-23 (Figure 5.4D, E). Aggregates of biotite and sillimanite are elongate subparallel to the gneissosity (Figure 5.5H). Euhedral garnet porphyroblasts are mainly restricted to the melanosome layers and are elongate subparallel to the gneissosity in some cases (Figure 5.5H). Quartz has a similar bimodal size distribution and microstructural attributes as the quartz in the other paragneiss sample. The leucosomes of this sample comprise a higher proportion of plagioclase and a lower proportion of K-feldspar than the leucosomes in sample GR11-23.

Methods

Electron Probe Microanalysis (EPMA)

Quantitative major element compositions were collected with a JEOL Superprobe JXA-8900R electron microprobe at Christian Albrechts Universität zu Kiel (Table 5.1). An accelerating voltage of 15–20 kV and a beam current of 20 nA were used for the analyses, and the CITZAF method of Armstrong (1995) was used to correct the data. X-ray compositional maps were obtained using a 15 kV accelerating voltage, 10–40 nA beam current, and 80–200 ms dwell time. Backscattered electron (BSE) images were used with the JEOL Image-Archive software to determine the area percent of the phases comprising reaction textures. The bulk compositions of these textures were then calculated by comparing the intensity (counts) of X-ray compositional maps with using the quantitative EPMA spot analyses (weight percent) (Table 5.2).

X-ray Fluorescence (XRF)

Between 5 and 10 kg of sample (excluding weathered exteriors) were crushed and powdered using a steel plated jaw crusher and ball mill at the University of Kansas, Department of Geology. Splits of the powder were sent to the GeoAnalytical Lab at Washington State University for major and trace element analyses (Table 5.2). Further powdering of the samples, fusing into glass beads, and analysis by XRF were carried out following the methods of Johnson et al. (1999).

LA-ICP-MS

Garnet trace element compositions were measured using laser ablation inductively coupled mass spectrometry (LA-ICP-MS) at the University of Kansas, Department of Geology's Isotope Geochemistry Lab. The analytical protocols that were used for the measurements are shown in Table 5.3.

Pressure-Temperature Modeling

Equilibrium phase diagrams (commonly referred to as pseudosections) were calculated for whole-rock samples (granulites and migmatitic paragneisses; Table 5.2) and microtextural domains (Table 5.2) of the granulites using Theriak-Domino version 04.02.2017 (De Capitani, 1994; De Capitani and Petrakakis, 2010). The whole-rock compositions of the residual granulite samples (Gruf-100 and Br03-56-2) were used as the input for the phase diagram calculations because of the homogeneous distribution of phases in these samples.

In the leucogranulite sample (GR11-39), the effective compositions of reaction textures are not represented by the whole-rock composition because the subsolidus reactions are separated from one another by melt. Therefore, the input compositions for these reaction textures were calculated by combining the BSE image, X-ray composition maps, and EPMA spot analyses of

the area comprising the reaction texture as described above. This method was also used to calculate the composition of a monazite-bearing texture (GRM-37), which was drilled from a thin section of residual granulite sample Br03-56-2.

The Theriak-Domino database *tcds55_p07* (downloaded from dtinkham.net/peq.html), which consists of thermodynamic data from Holland and Powell (1998), was used for all calculations. The following mineral solution models were used: talc, epidote, staurolite, chlorite, chloritoid, and cordierite (Holland and Powell, 1998); white mica (Coggon and Holland, 2002); orthopyroxene and spinel (White et al., 2002); feldspar (Holland and Powell, 2003); sapphirine and osumilite (Kelsey et al., 2004); biotite, garnet, and ilmenite (White et al., 2005); melt for metapelitic compositions (White et al., 2007); amphibole and clinopyroxene (Diener and Powell, 2012).

The H₂O content of all compositions was estimated using temperature versus mole fraction H₂O (M_{H2O}) binary phase diagrams (Figure 5.A1). See Appendix A for details.

Garnet Compositional Zoning

Leucogranulite

The major element zoning patterns of the garnet from the modeled reaction texture in the leucogranulite sample are shown in the X-ray compositional maps in Figure 5.6 (quantitative data in Table 5.1). The grossular component of this garnet is strongly zoned (Figure 5.6A). Mole fraction grossular ($X_{\text{grossular}}$) decreases from 0.1 in the core to 0.03 in the rim. There is a lack of zoning with respect to the other major cations (Mg, Fe, and Mn; Figure 5.6B, C, D) except for a decrease in Mg ($X_{\text{pyrope}} = 0.36$, whereas average $X_{\text{pyrope}} = 0.41$) (Figure 5.6B) and an increase in Fe ($X_{\text{almandine}} = 0.62$, whereas average $X_{\text{almandine}} = 0.51$) (Figure 5.6C) along one edge, which is in contact with orthopyroxene.

Residual Granulite

Figure 5.7 shows a profile of compositional zoning in a large garnet porphyroblast from residual granulite sample Gruf100. This garnet is typical of the residual granulite samples with an inclusion-rich core surrounded by an inclusion-poor rim (Figure 5.4B). The major and trace element zoning profiles described here are typical of garnet porphyroblasts in both residual granulite samples investigated in this study.

There is a sharp contrast in the major element zoning profiles between the core and rim (Figure 5.7). The almandine and grossular components ($X_{\text{almandine}}$ and $X_{\text{grossular}}$, respectively) gradually increase from the center of the inclusion-rich zone toward the contact with the inclusion-poor zone, whereas the pyrope component (X_{pyrope}) decreases. The spessartine ($X_{\text{spessartine}}$) component is invariable across the inclusion-rich zone.

Within the inclusion-poor rim, $X_{\text{almandine}}$ and $X_{\text{grossular}}$ decrease, whereas X_{pyrope} increases, from the core-rim contact toward the edge of the garnet. However, within ca. 0.3 mm of the garnet edge, $X_{\text{almandine}}$ and $X_{\text{grossular}}$ increase, and X_{pyrope} decreases. $X_{\text{spessartine}}$ gradually decreases from the core-rim contact toward the edge of the garnet.

Zoning of trace elements (REE, Y, and HFSE) is markedly different than that of the major elements. The Y concentration decreases from ca. 3500 ppm in the center of the inclusion-rich zone to ca. 400 ppm near the contact with the inclusion-poor zone (Figure 5.7). There is a sharp drop in Y concentration across the core-rim contact, and the Y concentration decreases from ca. 90 ppm to 30 ppm toward the edge of the garnet across the inclusion-poor domain. Zirconium concentration is much higher (ca. 50 ppm Zr on average; Table 5.1) in the inclusion-rich core than in the inclusion-poor rim (decreasing from ca. 7 ppm to 1 ppm; Table 5.1).

Pressure-Temperature Estimates

Leucogranulites

A garnet-breakdown texture (Figure 5.5A), which is representative of garnet textures in the leucogranulite samples, was targeted in sample GR11-39. The observed assemblage of garnet + orthopyroxene + plagioclase + sapphirine + spinel + melt (represented by quartz + K-feldspar) is stable at ca. 900–1050°C and 6–9 kbar on the calculated phase diagram (Figure 5.8A). The garnet compositional isopleths that correspond to the measured composition of the garnet core ($X_{\text{grossular}} = 0.05\text{--}0.10$ and $X_{\text{pyrope}} = 0.4$) intersect at ca. 800°C and 9–12 kbar, which is well outside the stability field of the observed assemblage (Figure 5.8E). The calculated isopleths of garnet volume decrease sharply between the intersection of the garnet compositional isopleths and the stability field of the observed assemblage (Figure 5.8F), indicating that orthopyroxene, plagioclase, sapphirine, and spinel formed at the expense of garnet. The relative positions of the stability fields of sapphirine and orthopyroxene indicate that sapphirine grew earlier than orthopyroxene. This is consistent with the textural relationships between the phases and the anhedral shape of the garnet. Therefore, we interpret that the rock underwent heating to UHT conditions during decompression from near-eclogite facies conditions along the P-T path shown in Figure 5.8F.

Residual Granulites

The peak assemblage of residual granulite sample Gruf100 is interpreted to be K-feldspar + garnet + orthopyroxene + sapphirine + sillimanite + rutile + melt. This assemblage is stable at 925–975°C and 8.25–10.25 kbar on the calculated phase diagram (Figure 5.8B). The presence of sapphirine + orthopyroxene symplectites and coronae around garnet (Figure 5.5B, C) suggest that garnet was breaking down at peak conditions. Cordierite was stable at lower pressure than the interpreted peak assemblage (Figure 5.8B), suggesting that moats of cordierite around garnet, orthopyroxene, sapphirine, and sillimanite formed during continued decompression. Sapphirine +

cordierite symplectites around sillimanite porphyroblasts also suggest decompression to less than ca. 8.5 kbar (the edge of the sillimanite stability field) occurred prior to significant cooling (Figure 5.8B, G). The biotite-out reaction occurs very close to the stability field of the peak assemblage. Therefore, it is likely that some of the biotite did not react out completely at peak conditions and resumed growth shortly after cooling commenced. Garnet-volume percent isopleths (Figure 5.8G) indicate that garnet volume was highest at higher pressures and lower temperatures than the stability field of the peak assemblage, indicating that peak conditions were reached during decompression. Based on these observations, we interpret that this sample followed the P-T path shown in Figure 5.8G in which heating to UHT conditions occurred during decompression, and nearly isothermal decompression continued to <8.5 kbar.

The interpreted peak assemblage of residual granulite sample Br03-56-2 (garnet + orthopyroxene + biotite + sapphirine + spinel + melt) is stable over a large area (925–1075°C and 6 to >11 kbar) of the P-T phase diagram (Figure 5.8C). This sample also shows textural evidence for garnet breaking down to form orthopyroxene, sapphirine, and cordierite (Figure 5.4C) similar to those in residual granulite sample Gruf100, and garnet-volume percent isopleths (Figure 5.8H) also indicate a higher garnet volume percent at higher pressure and lower temperature than the stability field of the peak assemblage. The beginning of the P-T path shown in Figure 5.8H represents a combination of the P-T paths determined for the leucogranulite garnet-breakdown texture (Figure 5.8F) and other residual granulite sample (Figure 5.8G). However, in contrast to sample Gruf100, the cordierite moats in sample Br03-56-2 probably formed in response to cooling rather than decompression according to the phase diagram (Figure 5.8C), indicating a that the slope of the P-T path became shallower (i.e. cooling became more dominant than decompression) at ca. 7 kbar.

Based on the straight grain boundaries between phases and intergrown nature of sapphirine and cordierite (Figure 5.5F), we interpret the texture within microdomain GRM-37 to represent equilibrium. The observed assemblage of orthopyroxene + sapphirine + cordierite was stable at 850–1100°C and 7.25–10.5 kbar (Figure 5.8D). This texture is from the residual granulite sample Br03-56-2 and the P-T estimates obtained using both the whole rock composition and the microdomain composition overlap considerably (Figure 5.8D, C).

Migmatitic paragneisses

The calculated phase diagram for the garnet + sillimanite + biotite + muscovite paragneiss (sample GR11-23) indicates that the observed peak assemblage (plagioclase + garnet + biotite + muscovite + quartz + sillimanite + ilmenite + melt) was stable at ca. 650–750°C and 4.0–7.5 kbar (Figure 5.9A). Sample GR11-25 is mineralogically and texturally similar to sample GR11-23, but it lacks muscovite and has a higher proportion of leucosome. However, the stability field of the observed peak assemblage (675–750°C and 5–7 kbar) overlaps considerably with that of the muscovite-bearing sample (Figure 5.9A, B). We interpret this to indicate that the difference in mineralogy reflects the slight difference in the bulk rock composition of these samples and not differences in peak P-T conditions.

Discussion

Accuracy of P-T Estimates

Determining the effective or equilibrium composition for a rock volume is probably the most challenging part of using thermodynamic software programs like Theriak-Domino to estimate P-T conditions. This is especially true for rocks that have undergone partial melting, which can result in significant changes of the effective composition during prograde and retrograde

metamorphism. However, we believe that our peak P-T estimates for the granulite samples are robust and accurate, because there is considerable overlap in peak P-T estimates and derived P-T paths, whether we used whole rock compositions determined by XRF or compositions of microdomains determined by EPMA (Figure 5.10). The starting point of the P-T paths shown in Figure 5.8E, F, and G is based on the composition of the garnet core in the microtexture from leucogranulite GR11-39. However, the effective bulk composition in which the garnet core formed may be different than the measured composition. Therefore, the starting point of the P-T path may not be accurate, but we believe that it is a close approximation because textural evidence (i.e. sapphirine + orthopyroxene growth at the expense of garnet followed by formation of cordierite coronae) indicates that the residual granulite samples followed similar decompression-heating P-T paths (Figure 5.8G, H) as the one derived from the garnet compositional isopleths for the leucogranulite shown in Figure 5.8E.

For the migmatitic paragneiss samples, we used whole rock compositions to model peak P-T conditions, which have the risk of not representing the equilibrium composition. However, other studies (e.g. Hamilton and Pattison, 2010; Hamilton and Pattison, 2017) have shown that in rocks in which melt segregated from solids, using a non-equilibrium (whole rock) composition only affects the position of the melt-related quartz-, plagioclase-, and K-feldspar-out lines and not the boundaries of the remnant solid phases. Therefore, we did not consider the stability of quartz or feldspars in estimating the peak P-T conditions, and the presented peak P-T estimates should be accurate.

Protoliths of Paragneisses and Granulites

Based on the presence of Al-rich phases such as sillimanite and garnet (and muscovite in sample GR11-23), it is likely that the protoliths of the migmatitic paragneisses were pelitic

sedimentary rocks with various amounts of quartz. The protoliths of the granulites, however, are not as easily discerned. Based on the low abundances of quartz and feldspars and high Al and low Si concentrations (Table 5.2), it is evident that the residual granulites lost melt at some point. All of the granulites contain Permian–Cretaceous, oscillatory-zoned zircon, which is interpreted to have crystallized from melt (Oalman et al., in preparation-b), indicating that there was a higher proportion of melt in the residual granulites prior to Cenozoic time. Therefore, the residual granulites and melanosomes in the leucogranulites likely developed their restitic character, at least in part, prior to UHT metamorphism. The earlier melt-loss events preconditioned the granulites for the subsolidus reactions that took place at UHT conditions as described by Kelsey and Hand (2015). The presence of abundant high-Al phases (e.g. sillimanite, sapphirine, and garnet) and lack of high-Ca phases (e.g. orthopyroxene instead of clinopyroxene or amphibole) leads us to postulate the protoliths of the granulites and charnockites were also pelitic sedimentary rocks.

Composite Pressure-Temperature Path

The conditions of a pre-peak high-pressure stage (ca. 800°C and 9–12 kbar) are constrained by the garnet compositional isopleths from the leucogranulite sample GR11-39 (Figure 5.8E). The enrichment of X_{pyrope} in the inclusion-poor rims of the garnet from the residual granulite sample Gruf 100 (Figure 5.7) is consistent with a high-pressure stage preceding UHT metamorphism. The garnet volume percent isopleths from the leucogranulite and residual granulite samples imply that the granulite/charnockite units were decompressed/exhumed to ca. 7 kbar while being heated to UHT conditions (Figure 5.8F, G), resulting in sapphirine + orthopyroxene growth at the expense of garnet. Qualitatively, the symplectite textures that replace garnet are also suggestive of a decompression P-T path. The rocks were further decompressed into the stability field of cordierite + orthopyroxene for the residual granulites (less than ca. 8.5 kbar at 800–1000°C) (Figure 5.8H).

The lower crustal granulite/charnockite units were then juxtaposed against the mid crustal migmatites, which were undergoing peak metamorphism at ca. 700–750°C and 5–8 kbar. This implies a composite P-T path for the granulites and charnockites characterized by early decompression and heating, an intermediate, nearly isothermal decompression stage at UHT conditions, and finally cooling and decompression to the conditions recorded by the mid crustal migmatites (Figure 5.10).

Figure 5.11 shows the change in modal abundances of minerals in the granulites along the P-T path in Figure 5.10 calculated with the P-T loop function of Theriak-Domino. For the garnet-breakdown texture from leucogranulite sample GR11-39, sapphirine was the first of the symplectite-forming minerals to grow at 950°C and 9.2 kbar as the modal abundance of garnet was decreasing. At ca. 940°C and 7.25 kbar, there was a sharp increase in the modes of orthopyroxene and spinel and decrease in sapphirine, indicating that sapphirine was partially consumed to form orthopyroxene + spinel. Residual granulite sample Gruf100 and Br03-56-2 behaved similarly with orthopyroxene and sapphirine modes increasing and garnet mode decreasing beginning at 950°C and 9.2 kbar. In addition, the modal abundance of sillimanite decreased as cordierite and sapphirine modes increased in the residual granulite sample. These patterns are consistent with the petrographic observations that orthopyroxene + sapphirine (+ cordierite in the residual granulites) grew at the expense of garnet at UHT conditions, indicating that the proposed P-T path is accurate.

The conditions of garnet breakdown and symplectite formation determined in this study contrast the estimates presented by Galli et al. (2011) of 700–750°C and 6–8 kbar. Based in part on these results, the authors concluded that UHT metamorphism did not occur during Alpine orogenesis. However, we interpret our results to indicate that UHT conditions were reached during

the latest metamorphic event, which occurred during Alpine orogenesis, at ca. 34 Ma (Oalman et al., in preparation-b). Furthermore, Galli *et al.* (2011; 2012) attributed UHT metamorphism to post-Variscan, Permian rifting of the European margin. However, isobaric cooling P-T paths are derived for UHT rocks from rift zones in most cases, whereas UHT decompression is more common in orogenic accretion and collision settings (Kelsey and Hand, 2015). Therefore, we interpret the decompression-heating P-T path derived for the Gruf granulites in this study to indicate that UHT conditions were obtained in an orogenic setting (i.e. the Alpine orogeny).

Relating the P-T Path to Ages

Figure 5.12 shows a pressure-temperature-time path for the Gruf Complex based on the U-Pb ages presented by Oalman et al. (in preparation-a); Oalman et al. (in preparation-b); Oalman et al. (submitted) and assuming that UHT metamorphism occurred during a single, Alpine P-T loop as discussed above. Rare earth element patterns of 34.8 ± 1.1 Ma zircon domains in the residual granulite indicate that this zircon grew in equilibrium with garnet and thus prior to garnet breakdown (Oalman et al., in preparation-b). Therefore, we interpret this age to constrain the timing of the pre-peak, high-pressure stage (Figure 5.12). The next youngest zircon population crystallized from post-UHT melts in the leucogranulite at 32.7 ± 0.7 Ma (Oalman et al., in preparation-b), after garnet breakdown as indicated by steep REE-patterns with high HREE concentrations. This leads to the interpretation that UHT metamorphism likely occurred between 34.8 ± 1.1 Ma and 32.7 ± 0.7 Ma (Figure 5.12). A ca. 34 Ma age for UHT metamorphism is consistent with the 32.5 ± 0.5 Ma age (Oalman et al., in preparation-a) of a monazite inclusion in a slightly post-peak sapphirine- + cordierite-bearing texture, which equilibrated at $>850^{\circ}\text{C}$ and >7 kbar (Figure 5.8D).

Because the migmatites were undergoing ductile deformation when the lower crustal charnockites and granulites were exhumed into the middle crust between 30 and 27 Ma (Oalman et al., in preparation-b), we interpret that the P-T conditions estimated for the paragneiss samples (675–750°C and 5–7 kbar) (Figure 5.9) represent the conditions of the middle crust into which the lower crustal rocks were exhumed. The ages of undeformed, muscovite-rich dikes and plutons indicate that contractional deformation ended and the amalgamated Gruf Complex was being exhumed between 26 and 24 Ma (Oalman et al., in preparation-b). U-Pb rutile ages indicate that the Gruf granulites and charnockites underwent cooling from 700–420°C between 30 and 19 Ma (Oalman et al., submitted).

Geodynamic Model for the Central Alps

Any valid tectonometamorphic model for the Gruf Complex must consider the P-T-t path shown in Figure 5.12. One such scenario is that the lower crustal granulites and charnockites were subducted or buried to depths equivalent to 9–12 kbar (ca. 32–42 km) and subsequently exhumed to 7.5–9.5 kbar (ca. 25–33 km depth) while being heated by advected mantle or mantle-derived melts. The switch from nearly isothermal decompression to decompression + cooling (Figure 5.8, 10, 12) may have been the result of the granulites and charnockites retaining heat while being exhumed to depths equivalent to less than 7 kbar within the crust.

The tectonometamorphic evolution of the Gruf Complex discussed above, (U)HP metamorphism at 40–35 Ma (Gebauer, 1996; Brouwer et al., 2005; Hermann et al., 2006) followed by Barrovian metamorphism (beginning at ca. 32 Ma; Rubatto et al., 2009) of the Adula nappe and other units of the Lepontine Dome, and petrogenesis of the magmas that formed the Periadriatic intrusions (e.g. the Biella, Bergell, Adamello, Riesenferner, and Pohorje intrusions), which intruded the Western, Central and Eastern Alps between 34 and 28 Ma (von Blanckenburg and

Davies, 1995; Rosenberg, 2004), (e.g. Bergell and Adamello Intrusions in the Central Alps) must be represented in viable geodynamic models for the Central Alps. Based on U-Pb geochronological constraints, a valid model must also account for partial exhumation of the (U)HP units prior to ca. 34 Ma (Gebauer, 1996; Hermann et al., 2006). Two previously published geodynamic models could possibly produce the observations noted here: 1) Slab breakoff (e.g. von Blanckenburg and Davies, 1995; Handy et al., 2010) lead to buoyant exhumation of the (U)HP units (subducted Penninic units and European margin, i.e. the Adula Nappe) followed by infiltration of asthenospheric mantle into the gap, UHT metamorphism, exhumation of the Gruf lower crustal units, and emplacement of the Bergell magmas. 2) Slab rollback of either the Valaisian or African slabs lead to lithospheric thinning (Beltrando et al., 2010), resulting in exhumation of the (U)HP units along lithospheric-scale faults followed by upwelling of the asthenosphere, which lead to adiabatic melting in the lithospheric mantle and UHT metamorphism in the lower crust, and finally a switch back to contractional deformation. However, we prefer the slab breakoff model because the lithospheric thinning model of Beltrando et al. (2010) does not accurately portray the timing of the Bergell intrusion, and there is no evidence for large scale extension within the Central Alps.

Figure 5.13 schematically shows our proposed model for the geodynamic evolution of the Central Alps based on the results of this study and models presented by Schmid et al. (1996), Bousquet et al. (2002), Handy et al. (2010), and Pfiffner (2016). Figure 5.13A shows the subduction stage of the Alpine orogeny that resulted in (U)HP metamorphism of the Adula and Penninic nappes between 45 and 35 Ma (e.g. Becker, 1993; Gebauer, 1996; Bousquet et al., 2002; Hermann et al., 2006). During this time period the Gruf Complex, interpreted to be of European margin affinity (Bousquet et al., 2012b; Galli et al., 2012), was likely being thrust beneath the Austroalpine nappe stack and became part of the Penninic nappes.

At ca. 34 Ma, the results presented here indicate that the lower crustal part of the Gruf Complex was undergoing UHT metamorphism at 900–1000°C and 7.5–9.5 kbar (25–33 km depth) following decompression from 800°C and 10–13 kbar (35–45 km depth) at ca. 35 Ma (Figure 5.12, 13B). The position of the Gruf Complex at ca. 35 Ma is shown as the transparent unit with dashed outline in Figure 5.13B.

The lack of a UHT overprint on the Adula and Penninic units implies that these units had been exhumed above the European margin and thus slab breakoff had occurred prior to 34 Ma (Figure 5.13B). Between 30 and 27 Ma, the Gruf lower crustal granulites and charnockites were exhumed into the middle crust migmatites (Oalman et al., in preparation-b) (Figure 5.13C). We envisage this to be the result of part of the lower crust of the European margin detaching and exhuming. Amalgamation of the Gruf Complex was coeval with the last stages of crystallization of the Bergell intrusion, indicating that exhumation of the lower crust may have been assisted by the buoyantly intruding Bergell magmas.

Between 26 and 24 Ma, contractional deformation had ceased, and the crustally-derived Novate leucogranite and associated muscovite-rich dikes intruded the Gruf Complex (Oalman et al., in preparation-b). We interpret this to indicate that exhumation of the amalgamated Gruf Complex along the Insubric line occurred then and was likely a result of underplating of Apulian lower crust (Figure 5.13D) at the onset of continent-continent collision, which also cut the supply of mantle-derived magmas to the Gruf Complex. Slip along the normal sense Forcola Fault occurred coevally with intrusion of the Novate leucogranite at ca. 24 Ma, exhuming the Adula nappe relative to the Gruf Complex (Ciancaleoni and Marquer, 2006).

Comparison of the Central Alps with Cenozoic Contractional Settings

Ultra-high temperature rocks have been discovered in other Cenozoic convergent tectonic settings. For example, UHT xenoliths were entrained in lavas that erupted on the Pamir (Gordon et al., 2012) and Tibetan (Hacker et al., 2000) Plateaus, and UHT granulites outcrop on Seram Island within the Banda Arc in Indonesia (Pownall et al., 2014). The Pamir xenoliths were subjected to higher pressures ($\geq 15\text{--}20$ kbar) (Gordon et al., 2012) than the Gruf granulites, whereas the Tibetan xenoliths underwent UHT metamorphism at 8–12 kbar (Hacker et al., 2000). Proposed geodynamic models for UHT metamorphism of the Tibetan lower crust include lithospheric thinning, slab breakoff, and intracontinental subduction. Pownall et al. (2014) concluded that UHT metamorphism in eastern Indonesia resulted from exhumation of lithospheric mantle into the lower crust during a period of extreme extension driven by slab rollback in the Banda Arc. The Alpine orogeny is different from the Pamir-Tibet and Indonesian examples in that the oceanic subduction stage was not long-lived in the Alps.

All these models invoke advection of hot, mantle material into the crust resulting in UHT metamorphism. Therefore, UHT metamorphism can occur in contractional settings if there are episodes of lithospheric-scale extension or mantle upwelling. This may imply that UHT metamorphism is more common in Phanerozoic collisional orogens than previously thought. The products of these extreme thermal processes may not be exposed or not recognized as UHT rocks due to retrograde overprinting or bulk compositions that do not produce obvious UHT assemblages (e.g. sapphirine + quartz).

Conclusions

Thermodynamic modeling of whole-rock and microdomain (reaction texture) compositions from leucocratic as well as restitic granulites yields petrologically reasonable results if the effective reaction composition is determined. The modeled P-T path indicates that the

granulite/charnockite units of the Gruf Complex reached UHT conditions (ca. 900–1000°C and 7.0–9.5 kbar) during decompression from 9–12 kbar and ca. 800°C at ca. 34 Ma, leading to sapphirine + orthopyroxene growth at the expense of garnet in all granulite types. The lower crustal UHT rocks were further decompressed to <7 kbar at (near) UHT temperatures and subsequently juxtaposed against midcrustal migmatitic gneisses, which underwent peak metamorphism at 700–750°C and 5–8 kbar, by between 30 and 27 Ma.

The decompression-heating P-T path experienced by the granulites and charnockites likely was the result of a switch from subduction of oceanic and continental lithosphere to pre-collisional mantle upwelling at ca. 34 Ma within the Central Alps, likely resulting from slab breakoff. The Gruf UHT rocks are most likely somewhat analogous to UHT rocks from other Cenozoic convergent settings including the Tibetan plateau and the Banda Arc in Indonesia.

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Appendix 5.A

The H₂O content of samples and reaction textures were estimated using temperature versus mole fraction hydrogen (M_H) diagrams. Pressure was fixed at 9 kbar for the granulite samples and textures therein (Figure 5.A1A–D) and 6 kbar for the paragneisses (Figure 5.A1E–F) based on the previously published P-T estimates of Galli et al. (2011) and Guevara and Caddick (2016). For all samples, M_H values ranging from 0 (dry conditions) to 20 (H₂O-saturated conditions) were used in the binary plots.

For the garnet breakdown texture from leucogranulite GR11-39 (Figure 5.A1A), the observed peak assemblage consists of plagioclase + garnet + sapphirine + orthopyroxene + spinel + melt ± K-feldspar (red shaded fields in Figure 5.A1A). The presence of coexisting plagioclase, sapphirine, orthopyroxene, and spinel in the garnet-breakdown texture limit M_H to less than 16. It is unclear whether K-feldspar was stable as a solid phase at peak conditions. Therefore, we used a M_H value of 8 to calculate the P-T diagram (Figure 5.8A).

For the residual granulite Gruf100 (Figure 5.A1B), the observed peak assemblage (red-shaded field in Figure 5.A1B) comprises plagioclase + garnet + orthopyroxene + sapphirine + sillimanite + rutile + melt. The sillimanite-out reaction (purple line in Figure 5.A1B) is nearly vertical in the calculated diagram, indicating that the presence or absence of sillimanite is strongly affected by H₂O content for this bulk composition. In this sample, the coexistence of sapphirine, and sillimanite restrict M_H to less than 11 (Figure 5.A1B). Therefore, we used an M_H value of 4, which crosses the field of the observed peak assemblage, to calculate the P-T diagram for this sample.

The field of the observed peak assemblage (garnet + orthopyroxene + sapphirine + biotite + melt) in residual granulite Br03-56-2 is stable over a wide range of M_H values, ranging from 7–19 (Figure 5.A1C). The minerals that were stable at peak conditions (garnet, sapphirine,

orthopyroxene) do not tightly constrain the H₂O content. Therefore, we chose a M_H value of 10, which intersects near the center of field of the observed peak assemblage (Figure 5.A1C).

The observed assemblage in the monazite bearing texture (GRM-37) from residual granulite sample Br03-56-2 is stable over a narrow range of M_H values, ranging from 3–5 (Figure 5.A1D). Assuming that neither free H₂O nor melt were present at peak conditions in this microdomain, we used a M_H value of 5 to calculate the P-T diagram (Figure 5.8D).

The stabilities of quartz and feldspars are not used to estimate the P-T conditions experienced by migmatitic paragneiss samples (see *Accuracy of P-T Estimates* above), and thus are not used to constrain the H₂O contents of these samples. Therefore, the only constraint for H₂O content in the muscovite-bearing paragneiss sample GR11-23 is the presence of muscovite and melt, which coexist at M_H values ranging from 9 to >20 (Figure 5.A1E). The presence of sillimanite and melt and absence of cordierite at peak conditions in the muscovite-free paragneiss sample GR11-25 indicates that M_H was >5.5 at peak conditions (Figure 5.A1F). We used a M_H value of 15 to calculate the P-T diagrams (Figure 5.9) for both paragneiss samples because this value intersects the field of the observed peak assemblage in sample GR11-23 and intersects fields that contain and lack free H₂O for sample GR11-25 (red shaded fields in Figure 5.A1F).

Figure Captions

Figure 5.1. Tectonic Map of the Central Alps (modified from Bousquet et al., 2012b). Ages from Bergell and Novate intrusions are from von Blanckenburg (1992) and Liati et al. (2000), respectively.

Figure 5.2. Geologic map of the Gruf Complex showing locations of samples used in this study (modified from Galli, 2010).

Figure 5.3. Sample photos: (A) Leucogranulite sample GR11-39. (B) Residual granulite sample Gruf100. (C) Muscovite-bearing migmatitic paragneiss sample GR11-23. (D) Muscovite-free migmatitic paragneiss sample GR11-25.

Figure 5.4. Scans of thin sections that were used in this study: (A) Leucogranulite sample GR11-39. (B) Residual granulite sample Gruf100, showing the location of the electron microprobe traverse used in Figure 5.7. (C) Residual granulite sample Br03-56-2. (D) Muscovite-bearing migmatitic paragneiss sample GR11-23. (E) Muscovite-free migmatitic paragneiss sample GR11-25.

Figure 5.5. Photomicrographs (colored) and backscattered electron (BSE) (black and white) images of textures in the granulites and paragneisses: (A) BSE image of garnet breakdown texture from leucogranulite sample GR11-39. (B) Photomicrograph from residual granulite sample Gruf100 showing embayment in garnet filled with prismatic sapphirine, sapphirine + orthopyroxene symplectites and quartz + feldspar, which represents crystallized melt. (C) Photomicrograph from residual granulite sample Gruf100 showing intergrown sapphirine and orthopyroxene porphyroblast separated from garnet by cordierite moat. Garnet is also fringed with orthopyroxene + cordierite symplectites. (D) Photomicrograph from residual granulite sample Gruf100 showing large sillimanite porphyroblast rimmed by sapphirine + spinel + cordierite symplectite, which is separated from orthopyroxene and biotite by a cordierite corona. (E) Photomicrograph from residual granulite sample Br03-56-2 showing a large sapphirine porphyroblast fringed by fine-grained spinel aggregates surrounded by cordierite, biotite, and

orthopyroxene. (F) BSE image of texture GRM-37 from residual granulite sample showing orthopyroxene, sapphirine, cordierite, and monazite. (G) Photomicrograph from muscovite-bearing migmatitic paragneiss sample GR11-23 showing typical leucosomes and melanosome, which consists of biotite, muscovite, sillimanite, and garnet. (H) Photomicrograph from muscovite-free migmatitic paragneiss sample GR11-25 showing typical leucosomes and melanosome, which consists of biotite, sillimanite, and aggregate of garnet porphyroblasts that are aligned subparallel to the foliation/gneissosity.

Figure 5.6. X-ray composition maps of the modeled reaction texture in the leucogranulite sample: (A) Relative Ca concentration. (B) Relative Mg concentration. (C) Relative Fe concentration. (D) Relative Mn concentration.

Figure 5.7. Chemical zoning profile of a large garnet porphyroblast from residual granulite sample Gruf100. The traverse location is shown in Figure 5.4B.

Figure 5.8. (A-D): Equilibrium phase diagrams calculated for granulite samples and reaction textures. (E) P-T diagram showing peak field and garnet compositional isopleths (X_{pyrope} and $X_{\text{grossular}}$) calculated for the garnet breakdown texture in leucogranulite sample GR11-39. The orange bar represents the range of measured compositions in the garnet. (F) P-T diagram with peak field and garnet volume percent isopleths for the garnet breakdown texture in leucogranulite sample GR11-39. The P-T path is drawn from the garnet compositional range from Figure 5.8E to the peak field, indicating a decrease of garnet volume during decompression and heating. (G) P-T diagram with peak field and garnet volume percent isopleths for the residual granulite sample

Gruf100. The starting point of the P-T path coincides with the maximum pressure starting point in Figure 5.8F, and the path is drawn to indicate a decrease in garnet volume during decompression and heating and continued nearly isothermal decompression through the stability field of the peak assemblage into the cordierite stability field. (H) P-T diagram with peak field and garnet volume percent isopleths for the residual granulite sample Br03-56-2. The P-T path is drawn to indicate a decrease in garnet volume during decompression and heating to peak conditions. The starting point coincides with the maximum pressure starting point in Figure 5.8F. The P-T path also considers the growth of spinel after sapphirine and crosses into the cordierite stability field (see Figure 5.8C).

Figure 5.9. Equilibrium phase diagrams calculated for migmatitic paragneiss samples: (A) Muscovite-bearing paragneiss sample GR11-23. (B) Muscovite-free paragneiss sample GR11-25.

Figure 5.10. P-T diagram showing estimates of peak metamorphic conditions for the granulites and paragneisses and a composite P-T path for the granulites.

Figure 5.11. Diagrams showing modal abundance changes of selected minerals along the P-T path shown in Figure 5.10. (A) Leucogranulite GR11-39. (B) Residual granulite Gruf100. (C) Residual granulite Br03-56-2.

Figure 5.12. Pressure-temperature diagram showing the pressure-temperature-time evolution of the Gruf granulites and migmatitic paragneisses with schematic mineral reactions and structural events. Zr-in-rutile temperatures and U-Pb rutile age are from Oalman et al. (submitted).

Monazite and zircon ages are from Oalman et al. (in preparation-a) and Oalman et al. (in preparation-b), respectively.

Figure 5.13. Schematic model for the geodynamic evolution of the Central Alps. (A) 40–36 Ma: subduction of Penninic units and Adula nappe. (B) ca. 34 Ma: UHT metamorphism of the Gruf lower crustal rocks following slab breakoff and exhumation of (U)HP Adula and Penninic nappes. (C) 30–27 Ma: amalgamation of the Gruf Complex closely following intrusion of the Bergell tonalite and granodiorite. (D) ca. 24 Ma: exhumation of the amalgamated Gruf Complex along the Insubric line, intrusion of the Novate leucogranite, slip along the Forcola Fault, and onset of continent-continent collision. The black harpoon symbols represent active slip on thrust faults, and the yellow harpoon symbol in panel D represents active slip along the normal-sense Forcola Fault.

Figure 5.A1: Temperature versus mole fraction hydrogen (M_H) phase diagrams calculated at 9 kbar pressure for granulite microtextures and samples (A–D) and at 6 kbar pressure for migmatitic paragneiss samples (E–F). The red-shaded fields correspond to the observed peak assemblages, and the thick, dashed lines correspond to the M_H values used to compute the P-T diagrams (Figure 8, 9).

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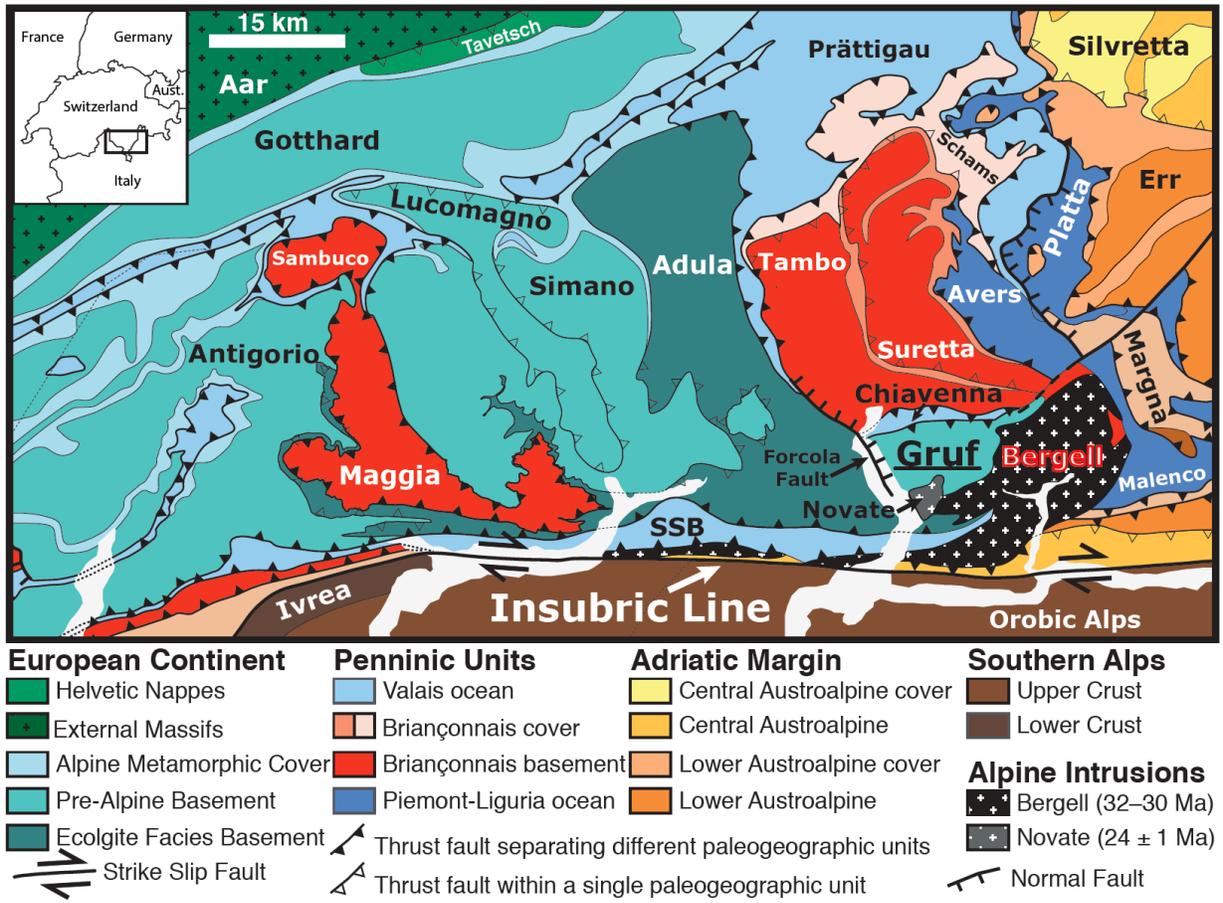


Figure 5.1

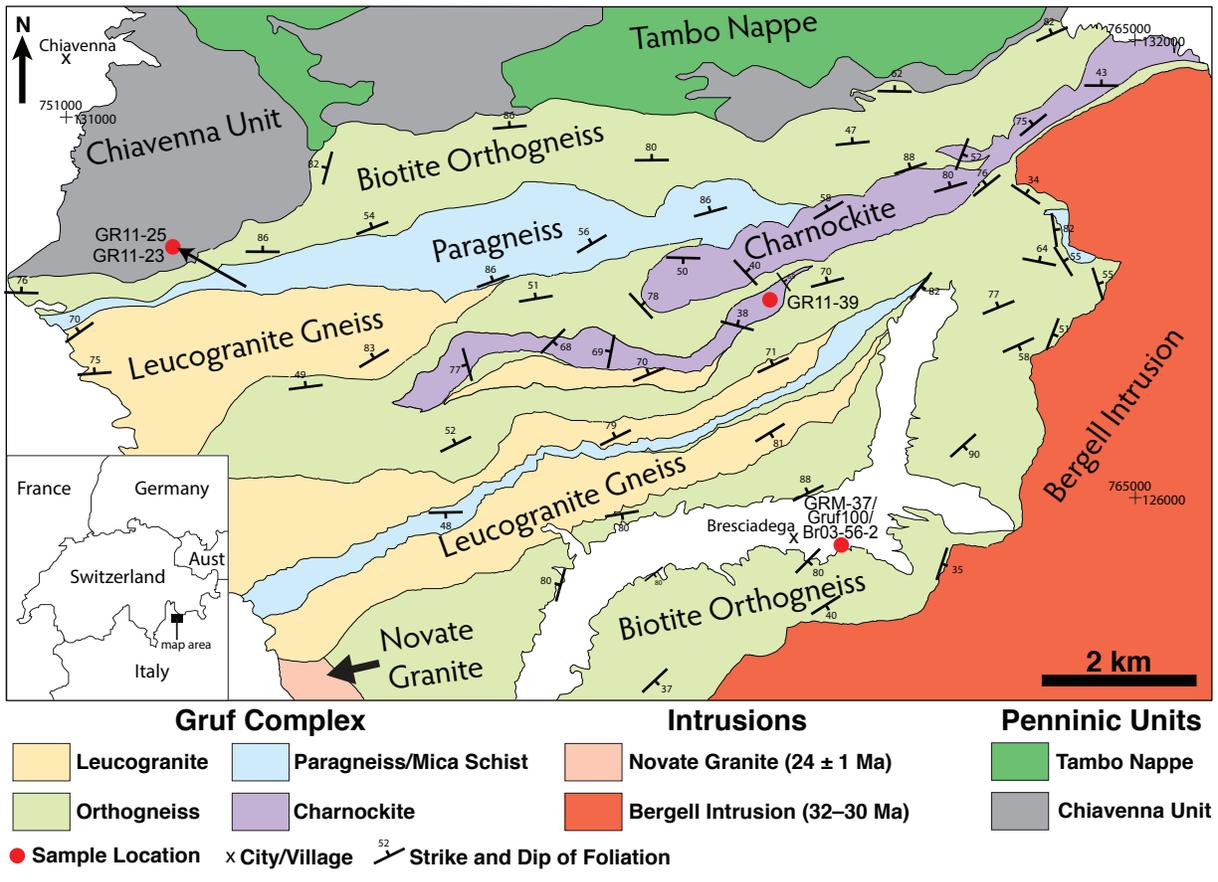


Figure 5.2

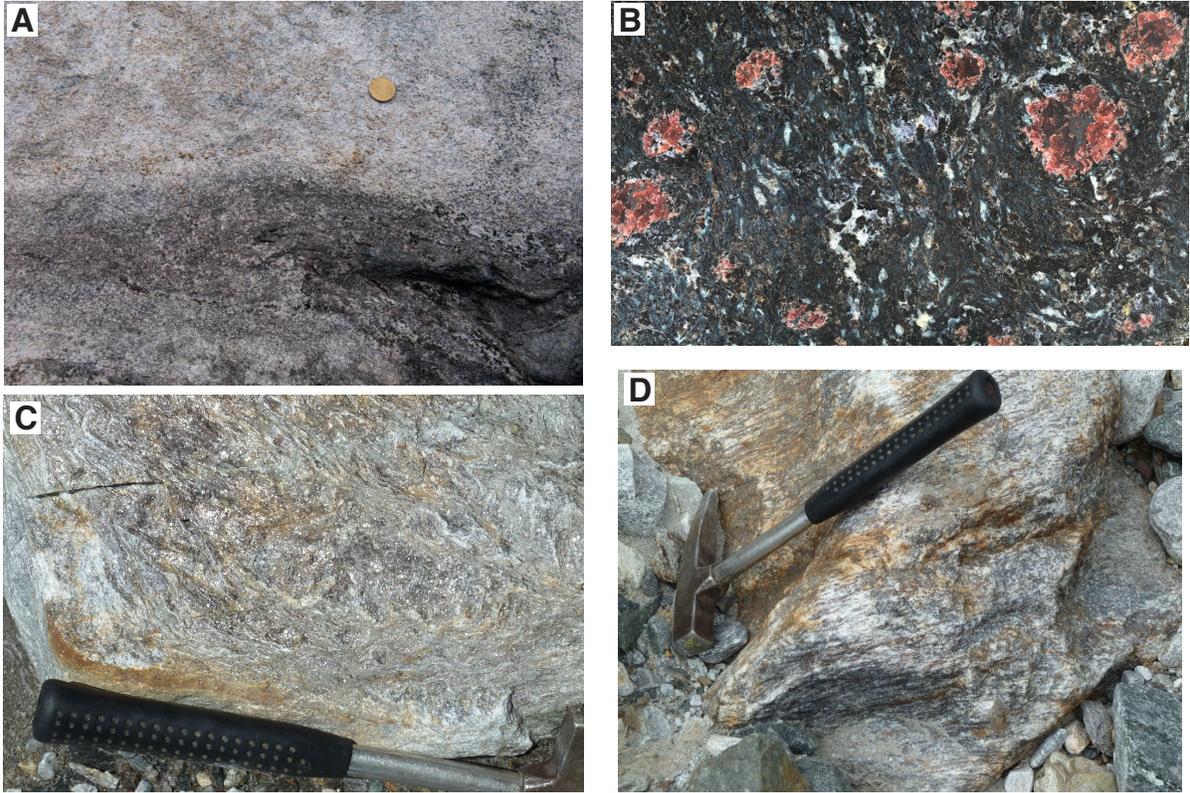


Figure 5.3

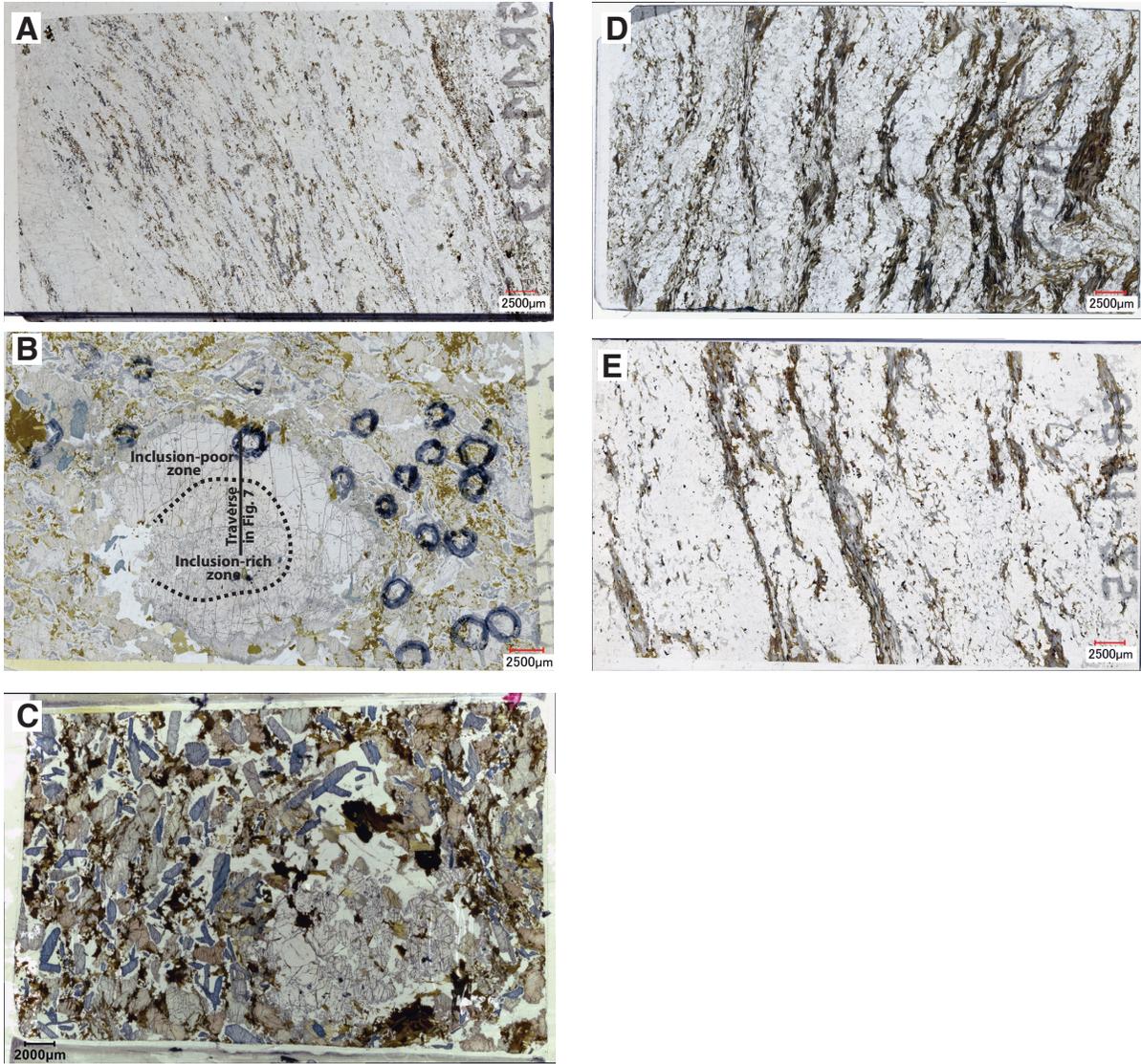


Figure 5.4

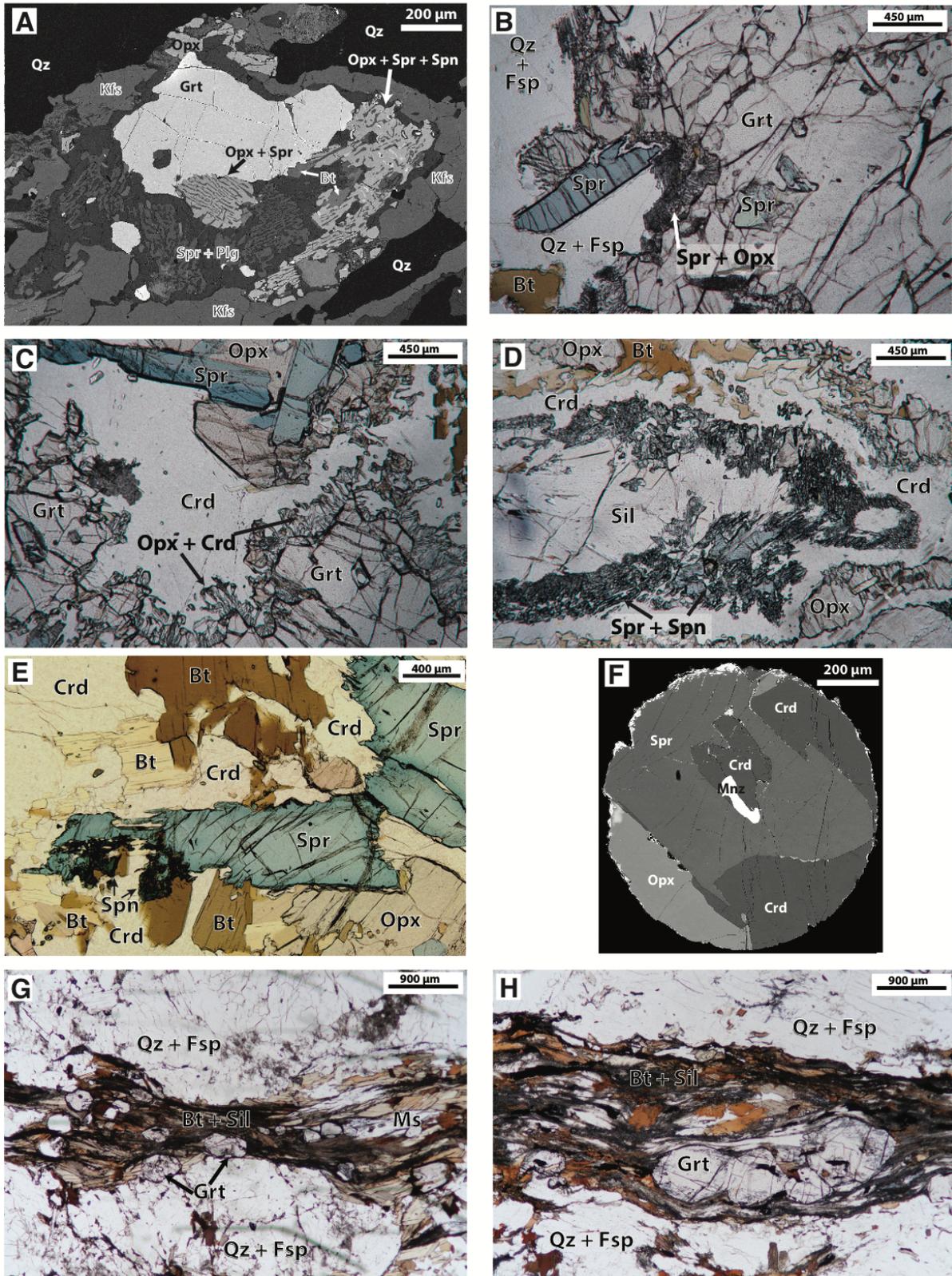


Figure 5.5

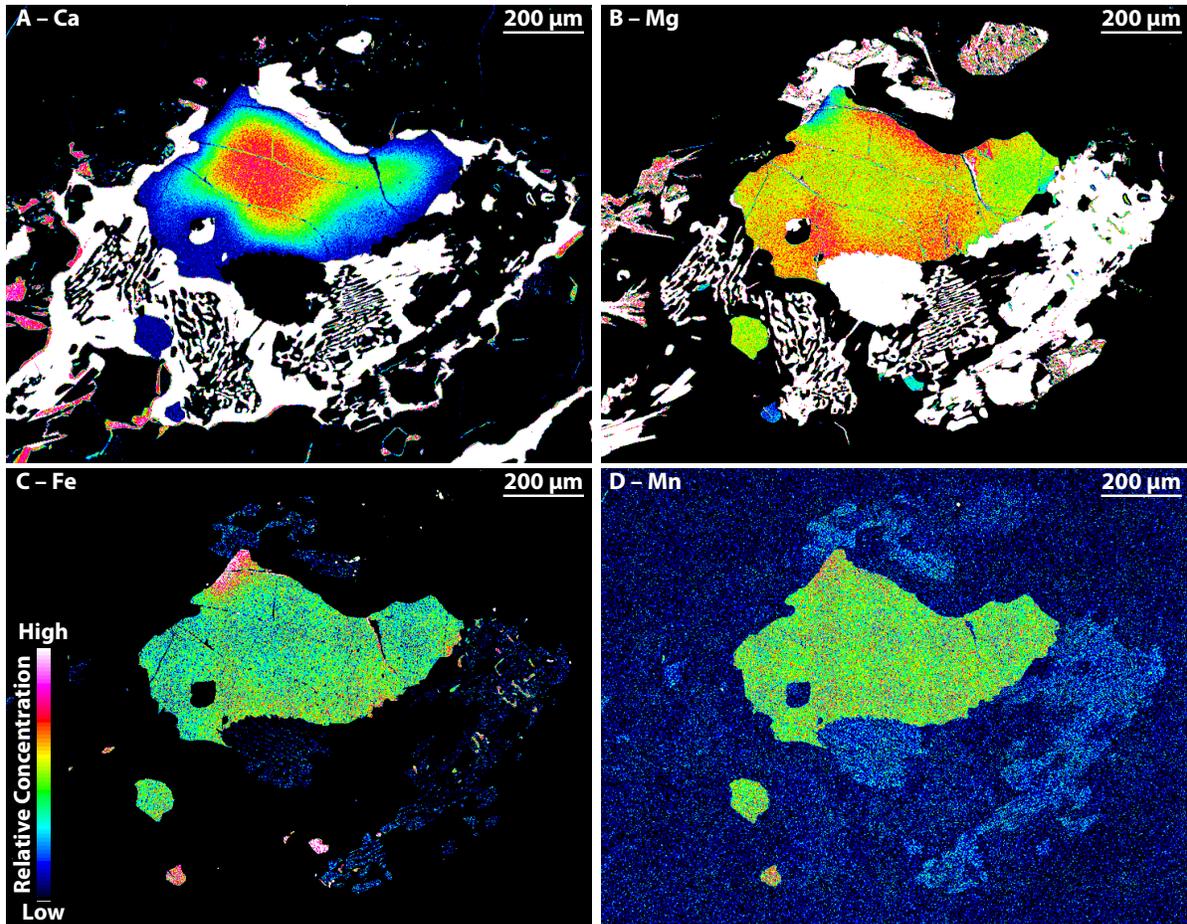


Figure 5.6

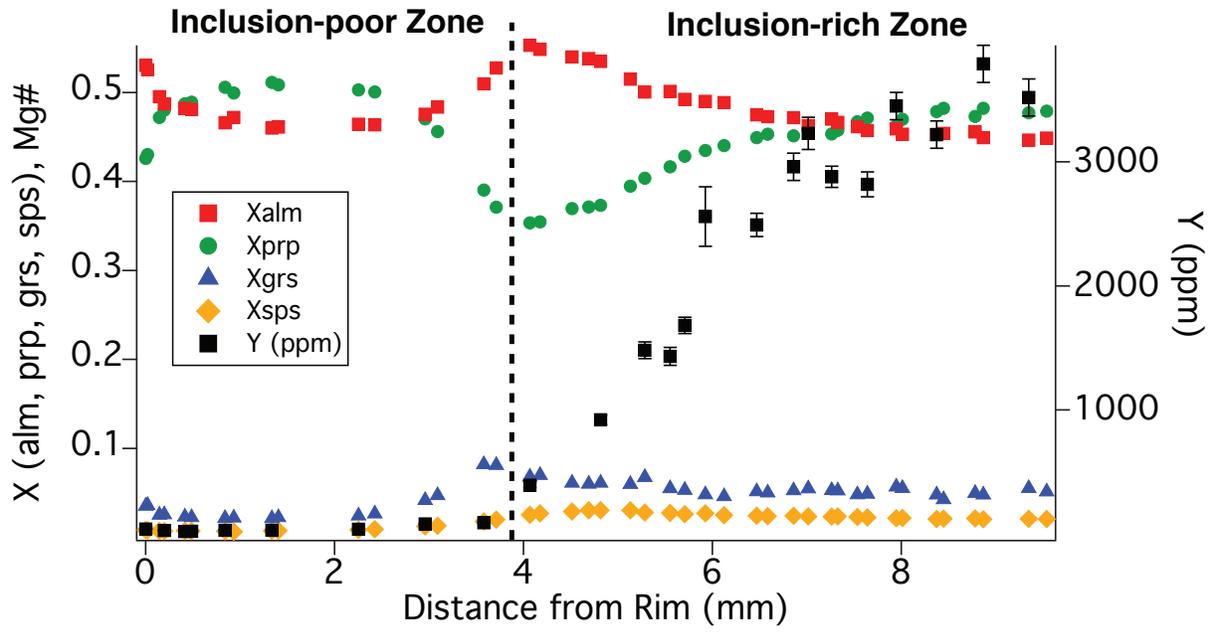
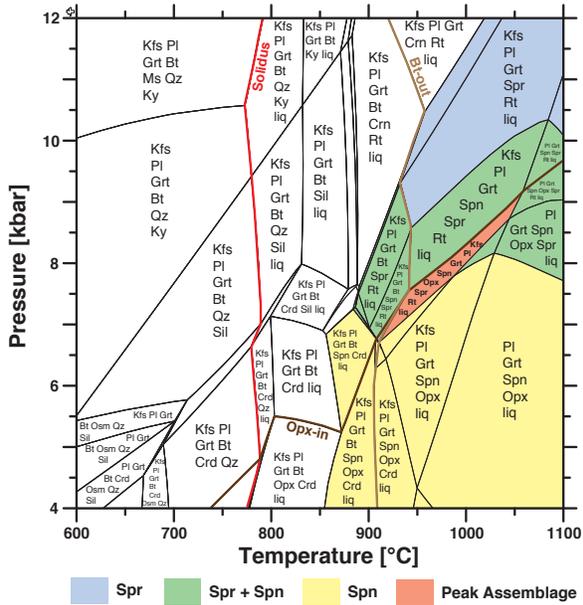
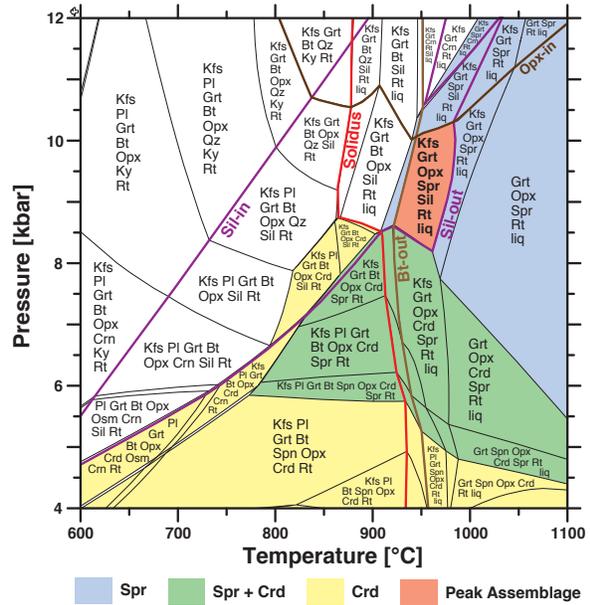


Figure 5.7

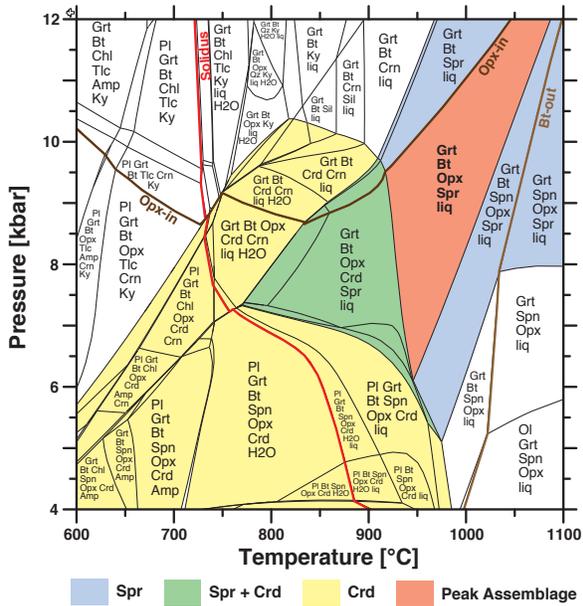
(A) GR11-39 (Texture in Leucogran.)



(B) Gruf100 (Residual Granulite)



(C) Br03-56-2 (Residual Granulite)



(D) GRM-37 (Texture in Resid. Gran.)

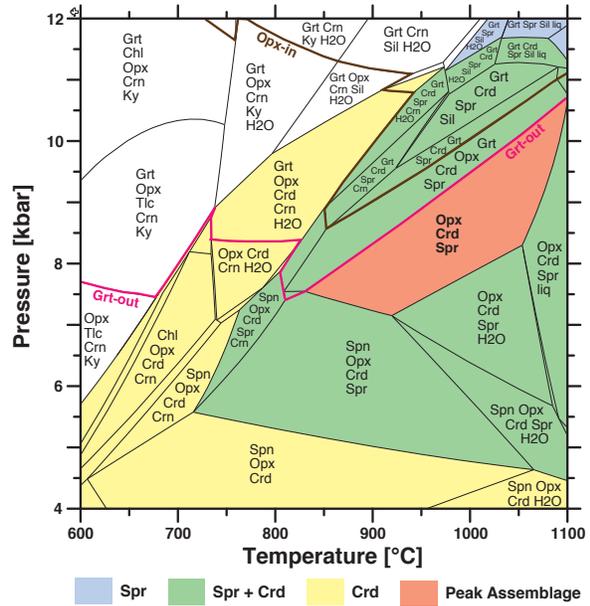


Figure 5.8

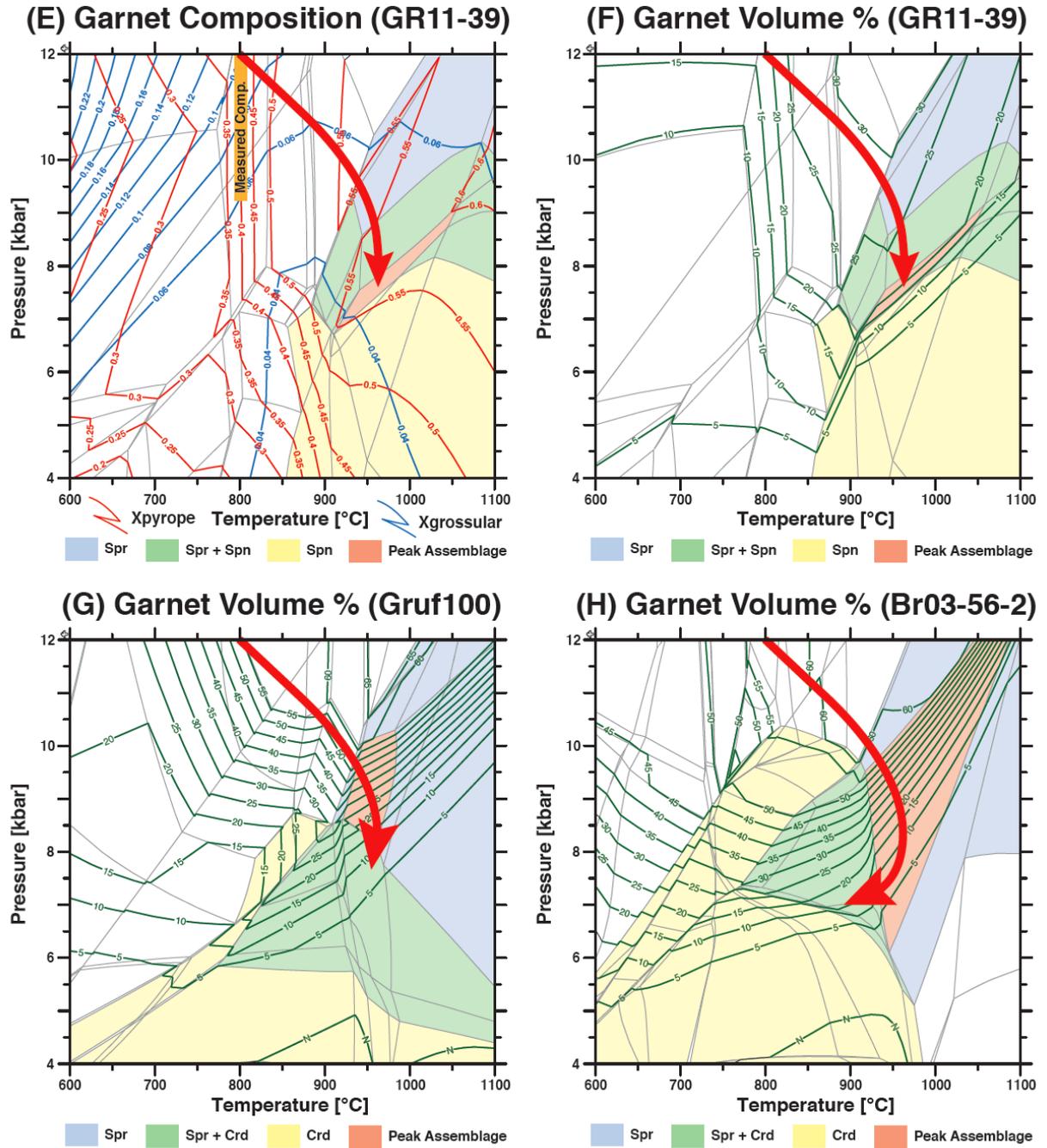


Figure 5.8 (continued)

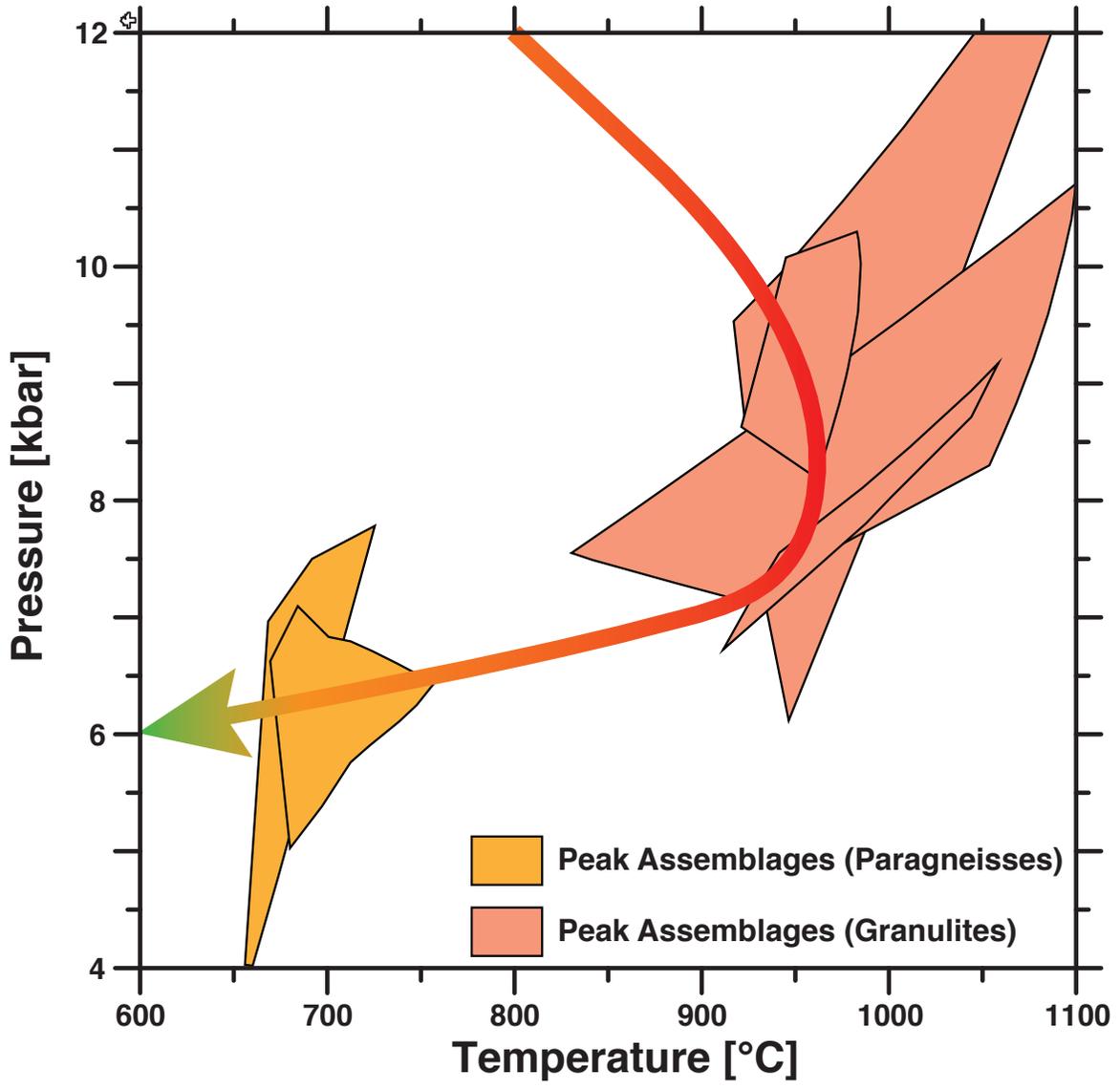


Figure 5.10

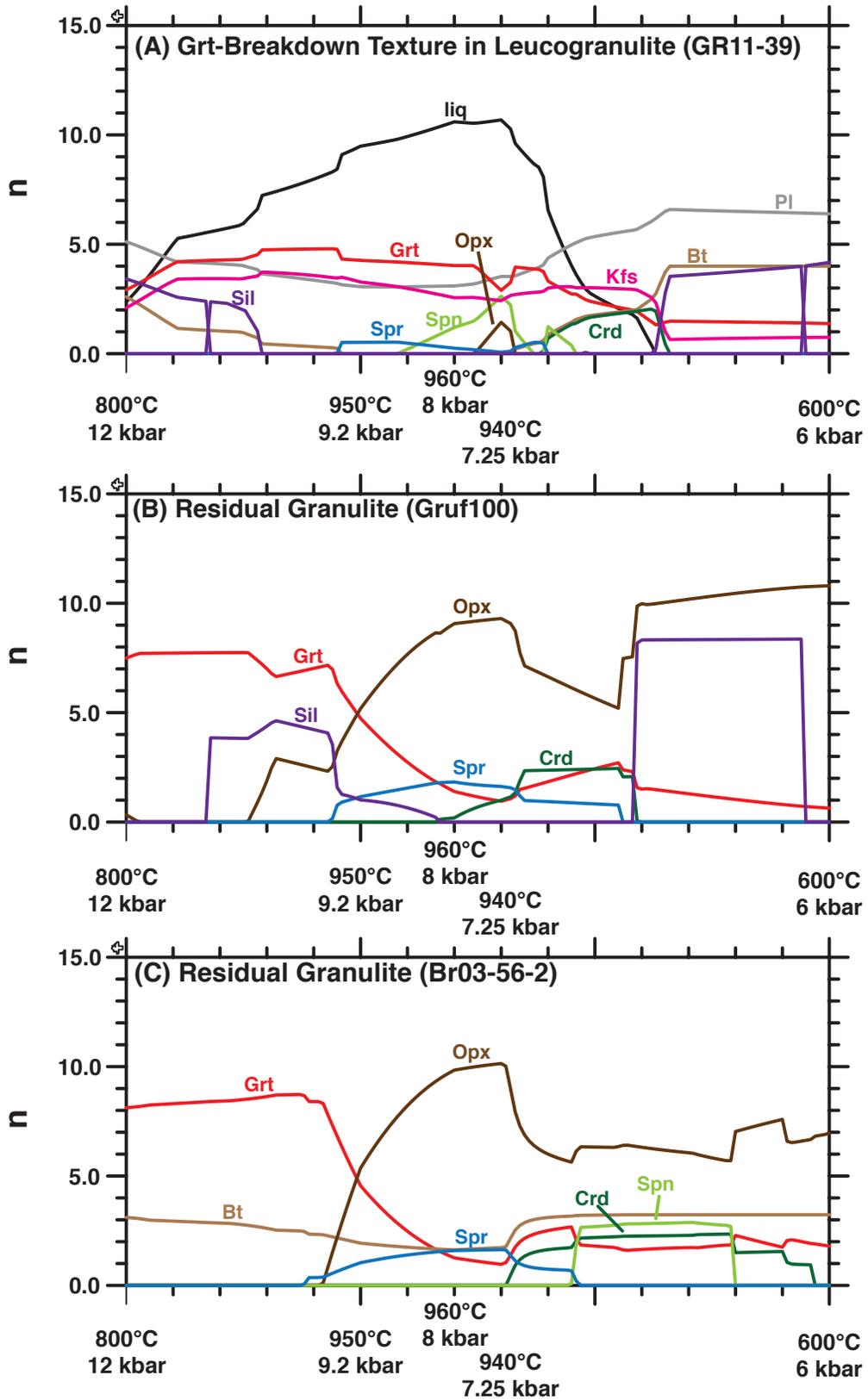


Figure 5.11

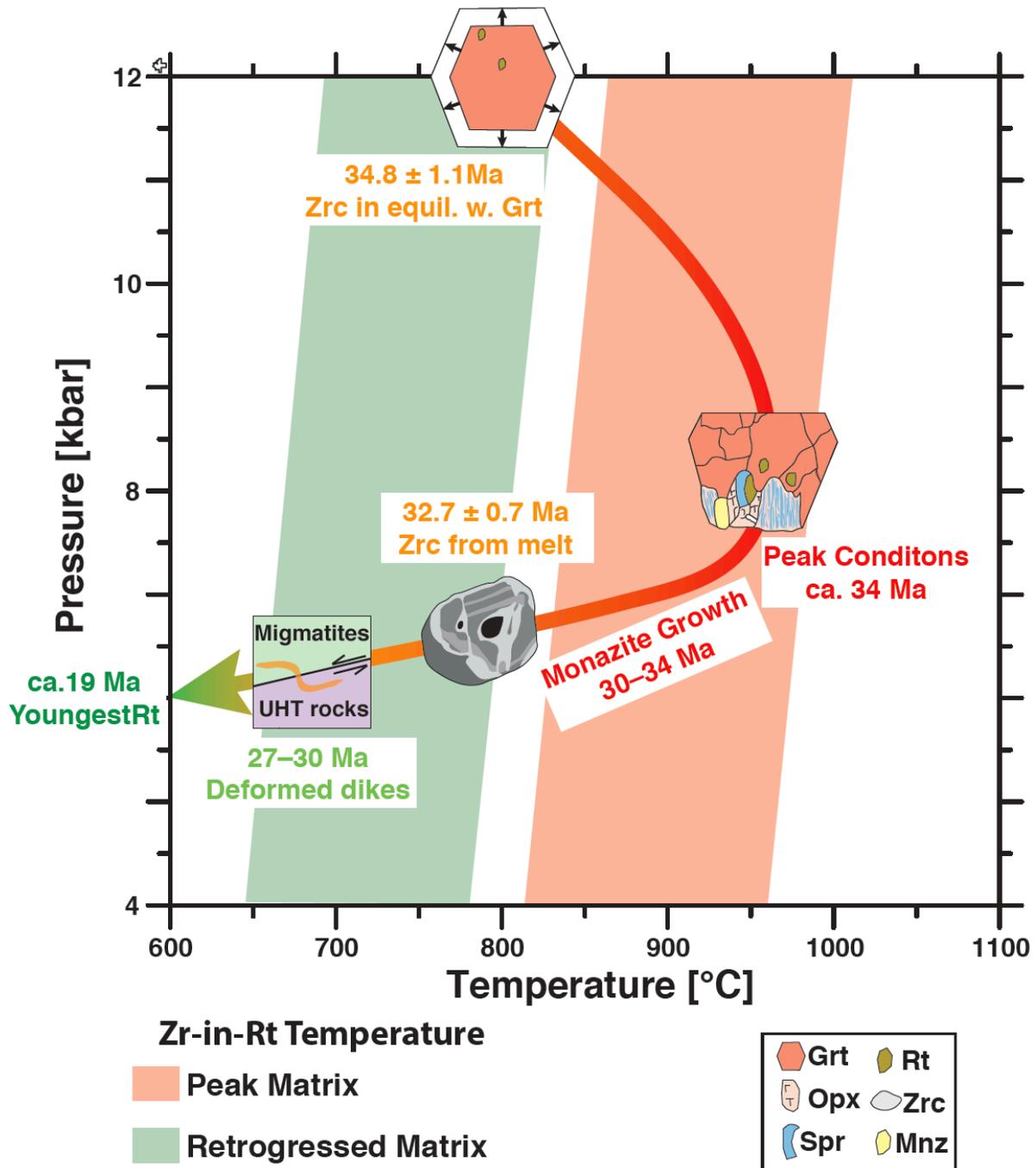
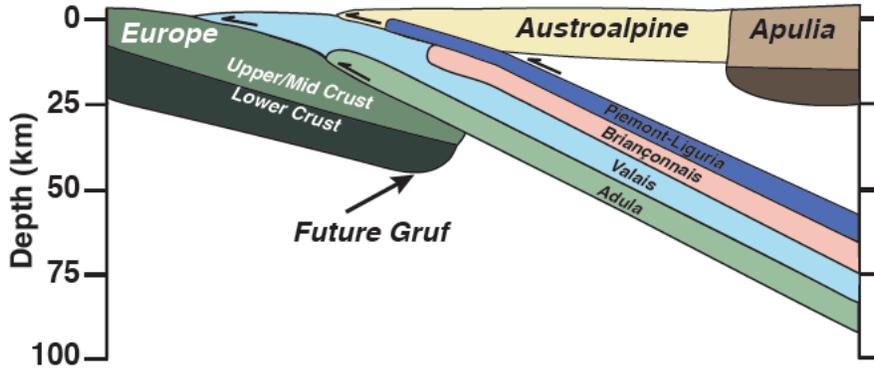
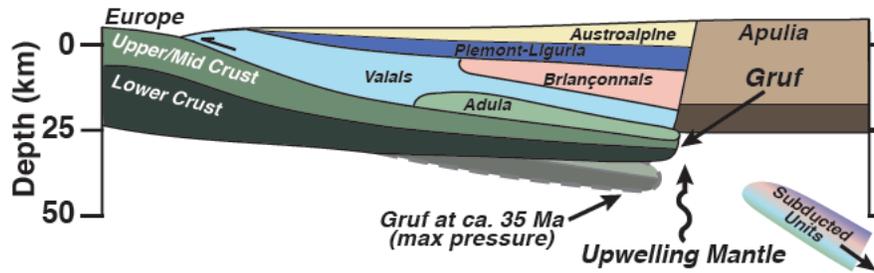


Figure 5.12

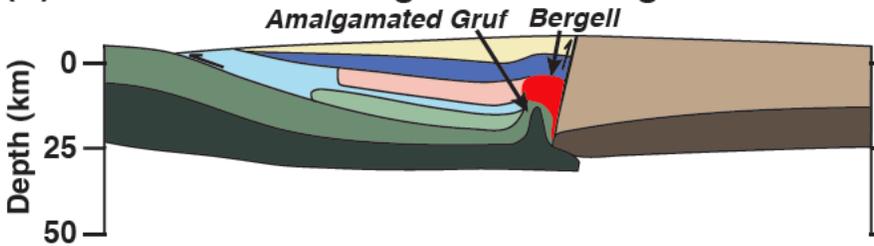
(A) 40–36 Ma: Subduction and (U)HP Metamorphism of Penninic Nappes



(B) ~34 Ma: UHT Metamorphism of Gruf Lower Crust



(C) 30–27 Ma: Post-Bergell Gruf Amalgamation



(D) ~24 Ma: Novate Intrusion and Exhumation

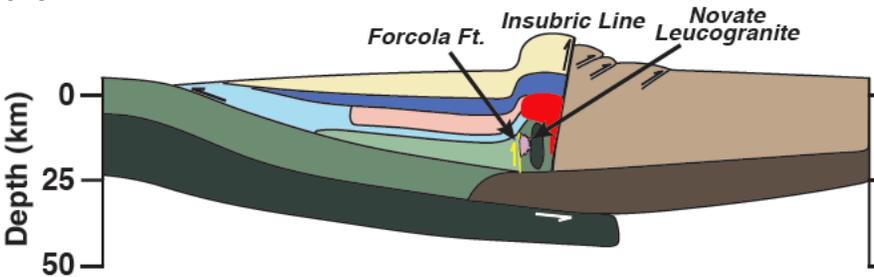
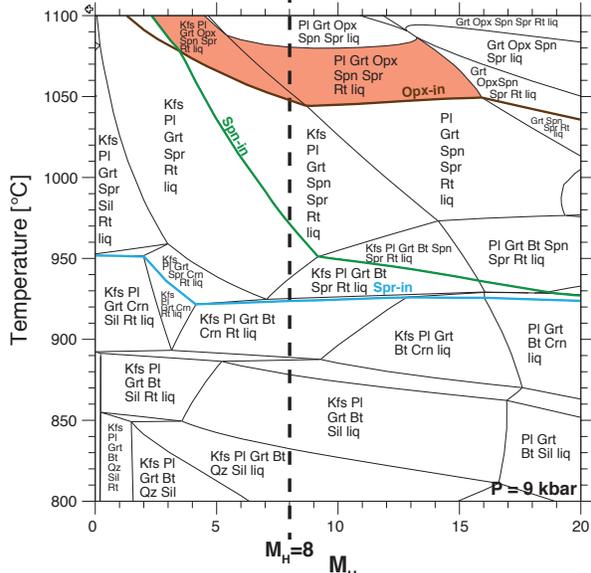
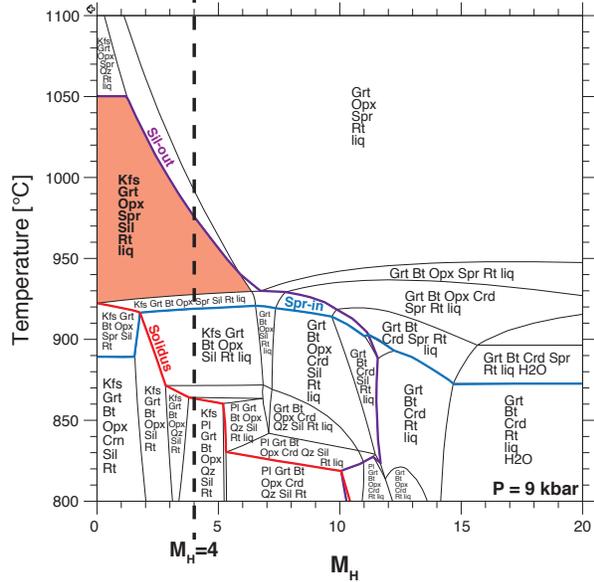


Figure 5.13

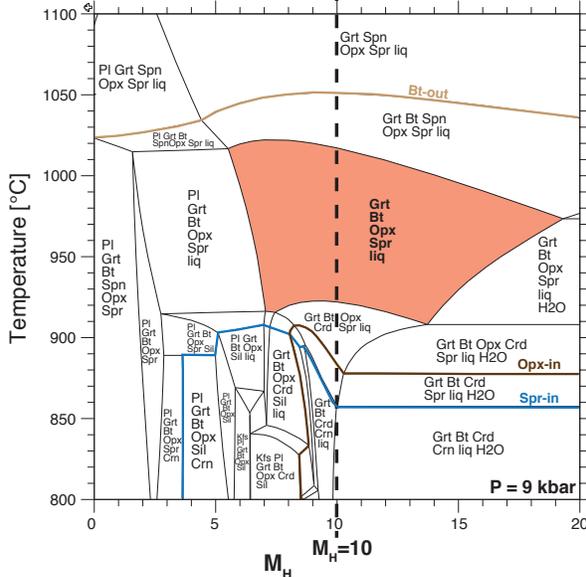
(A) GR11-39 (Texture in Leucogran.)



(B) Gruf100 (Residual Granulite)



(C) Br03-56-2 (Residual Granulite)



(D) GRM-37 (Texture in Resid. Gran.)

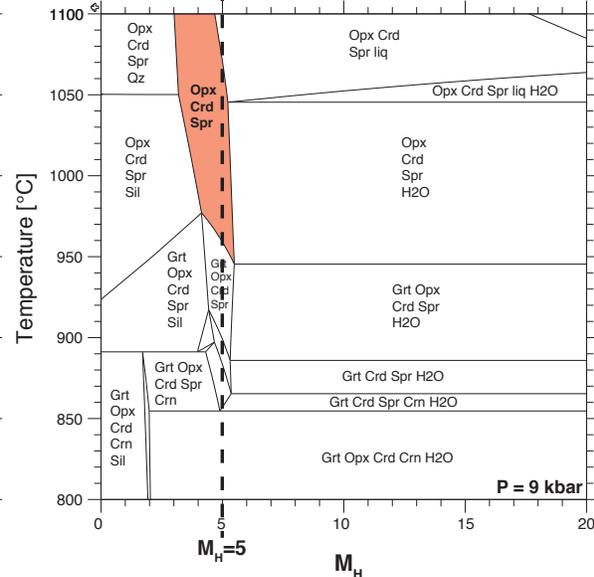
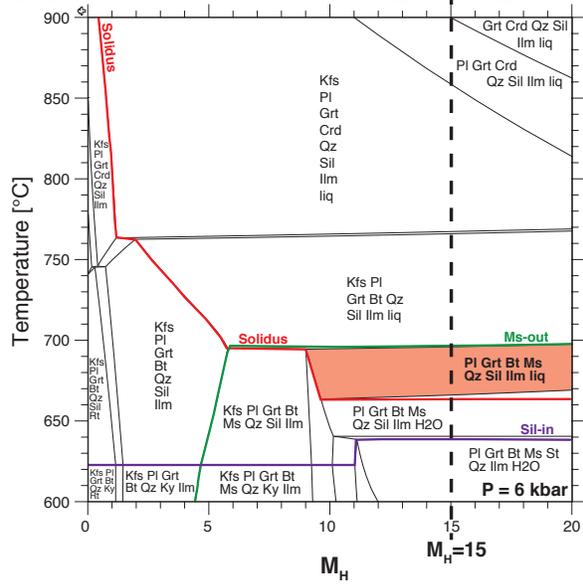


Figure 5.A1

(E) GR11-23 (Ms-Bearing Paragneiss)



(F) GR11-25 (Ms-Free Paragneiss)

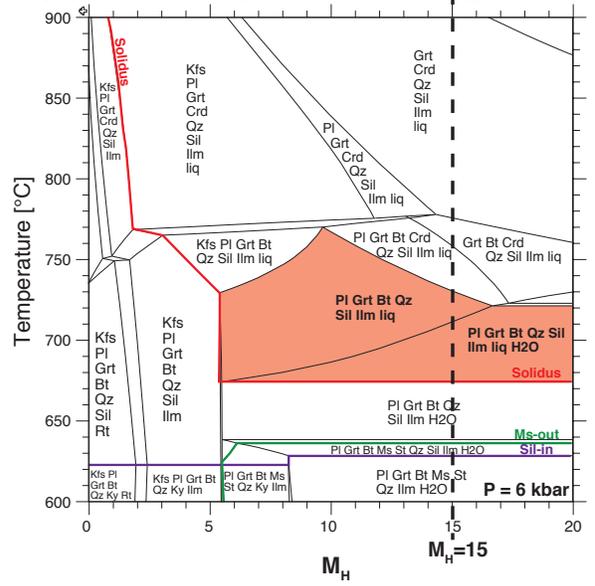


Figure 5.A1 (continued)

Sample and Analysis Number	AlO3	TiO2	K2O	Na2O	Cr2O3	SiO2	FeO	CaO	MgO	MnO	Total Al	Fe	Mg	Ca	Mn	Na	K	Grt	Xpr	Grt	Xpr	Grt	Xpr	Plg	Opx	Mg	Zr	±2SE	Y	±2SE	
	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)	(wt%)											
Residual Garnet (Grt100)																															
Grt100_Grt1	22.62	0.00	0.02	0.03	0.01	39.72	24.57	1.27	11.06	0.38	99.68	0.22	0.34	0.27	0.02	0.01	0.00	0.00	0.53	0.43	0.04	0.01	-	-	0.45	2.3	0.4	41	2		
Grt100_Grt2	22.85	0.00	0.01	0.02	0.00	39.59	24.63	1.32	11.31	0.37	100.49	0.22	0.34	0.28	0.02	0.01	0.00	0.00	0.53	0.43	0.04	0.01	-	-	0.45	-	-	-	-	-	
Grt100_Grt3	23.07	0.01	0.00	0.04	0.00	40.14	23.32	0.92	12.46	0.34	100.32	0.23	0.32	0.31	0.02	0.00	0.00	0.00	0.50	0.47	0.03	0.01	-	-	0.49	1.9	0.2	76	-		
Grt100_Grt4	23.08	0.00	0.00	0.05	0.00	40.30	22.70	0.85	12.69	0.36	100.23	0.23	0.32	0.32	0.02	0.00	0.00	0.00	0.48	0.49	0.02	0.01	-	-	0.50	1.5	0.2	21	1		
Grt100_Grt5	22.93	0.01	0.01	0.04	0.00	40.12	22.93	1.03	13.09	0.34	100.32	0.22	0.32	0.31	0.01	0.00	0.00	0.00	0.48	0.49	0.02	0.01	-	-	0.50	2.2	0.2	23	1		
Grt100_Grt6	23.18	0.01	0.00	0.06	0.00	40.28	22.15	0.80	13.49	0.32	100.30	0.23	0.31	0.33	0.01	0.00	0.00	0.00	0.47	0.51	0.02	0.01	-	-	0.52	1.8	0.2	32	3		
Grt100_Grt7	23.25	0.00	0.00	0.04	0.01	40.42	22.35	0.62	13.27	0.30	100.46	0.23	0.31	0.33	0.01	0.00	0.00	0.00	0.47	0.50	0.02	0.01	-	-	0.51	-	-	-	-		
Grt100_Grt8	23.44	0.01	0.00	0.04	0.00	40.34	21.84	0.74	13.44	0.31	100.34	0.23	0.31	0.33	0.01	0.00	0.00	0.00	0.46	0.51	0.02	0.01	-	-	0.52	0.8	0.1	32	2		
Grt100_Grt9	22.96	0.01	0.04	0.12	0.01	39.39	21.81	0.93	13.51	0.36	99.57	0.23	0.30	0.34	0.01	0.00	0.00	0.00	0.46	0.51	0.02	0.01	-	-	0.52	1.3	0.3	40	3		
Grt100_Grt10	23.17	0.01	0.01	0.09	0.01	40.28	21.97	0.90	13.35	0.40	100.17	0.23	0.31	0.33	0.02	0.01	0.00	0.00	0.46	0.50	0.02	0.01	-	-	0.52	1.8	0.3	40	3		
Grt100_Grt11	23.20	0.00	0.01	0.05	0.00	40.31	22.06	0.93	13.36	0.43	100.41	0.23	0.31	0.33	0.02	0.01	0.00	0.00	0.46	0.50	0.03	0.01	-	-	0.52	-	-	-	-		
Grt100_Grt12	23.08	0.01	0.01	0.06	0.02	39.89	22.42	1.54	12.44	0.58	100.04	0.23	0.31	0.31	0.03	0.01	0.00	0.00	0.48	0.47	0.04	0.01	-	-	0.50	5.0	0.2	82	3		
Grt100_Grt13	22.97	0.05	0.00	0.02	0.00	40.07	22.94	1.75	12.14	0.62	100.28	0.22	0.32	0.30	0.03	0.01	0.00	0.00	0.48	0.46	0.05	0.01	-	-	0.49	-	-	-	-		
Grt100_Grt14	22.69	0.00	0.00	0.03	0.00	40.10	23.00	1.40	12.40	0.44	100.34	0.23	0.32	0.32	0.02	0.00	0.00	0.00	0.48	0.46	0.05	0.01	-	-	0.49	-	-	-	-		
Grt100_Grt15	22.47	0.06	0.01	0.01	0.00	39.52	24.68	2.56	9.74	0.95	100.37	0.22	0.34	0.26	0.05	0.01	0.00	0.00	0.53	0.37	0.08	0.02	-	-	0.41	7.1	0.4	2	-		
Grt100_Grt16	22.31	0.06	0.00	0.01	0.00	39.69	25.64	2.64	9.19	1.17	100.51	0.22	0.36	0.23	0.04	0.02	0.00	0.00	0.55	0.35	0.07	0.03	-	-	0.39	11.0	0.4	395	19		
Grt100_Grt17	22.65	0.03	0.00	0.00	0.00	39.28	25.63	2.55	9.30	1.25	100.68	0.22	0.36	0.23	0.05	0.02	0.00	0.00	0.55	0.35	0.07	0.03	-	-	0.39	-	-	-	-		
Grt100_Grt18	22.60	0.09	0.00	0.03	0.00	39.50	25.06	2.22	9.62	1.34	100.45	0.22	0.35	0.24	0.04	0.02	0.00	0.00	0.54	0.37	0.06	0.03	-	-	0.41	-	-	-	-		
Grt100_Grt19	22.47	0.13	0.00	0.01	0.00	39.24	24.77	2.16	9.59	1.38	99.80	0.22	0.34	0.24	0.04	0.02	0.00	0.00	0.54	0.37	0.06	0.03	-	-	0.41	66.9	3.5	922	7		
Grt100_Grt20	22.58	0.06	0.00	0.00	0.00	39.51	24.65	2.18	10.34	1.41	100.13	0.22	0.33	0.24	0.04	0.02	0.00	0.00	0.54	0.39	0.06	0.03	-	-	0.43	-	-	-	-		
Grt100_Grt21	22.64	0.07	0.01	0.01	0.00	39.66	23.37	2.47	10.58	1.29	100.08	0.22	0.33	0.26	0.04	0.02	0.00	0.00	0.50	0.40	0.07	0.03	-	-	0.45	122.3	3.6	1482	69		
Grt100_Grt22	22.74	0.13	0.01	0.01	0.00	39.50	23.38	2.00	10.90	1.26	98.94	0.22	0.33	0.27	0.04	0.02	0.00	0.00	0.50	0.42	0.05	0.03	-	-	0.45	46.9	2.0	1433	73		
Grt100_Grt23	22.96	0.05	0.00	0.01	0.00	39.88	23.09	1.94	11.27	1.21	100.42	0.23	0.32	0.28	0.03	0.02	0.00	0.00	0.49	0.43	0.05	0.03	-	-	0.47	61.3	2.5	1683	66		
Grt100_Grt24	22.81	0.04	0.00	0.02	0.00	40.02	22.90	1.76	11.40	1.25	100.20	0.22	0.32	0.28	0.03	0.02	0.00	0.00	0.49	0.43	0.05	0.03	-	-	0.47	53.9	6.3	2560	240		
Grt100_Grt25	22.93	0.07	0.01	0.03	0.03	39.57	22.11	1.88	11.72	1.12	99.47	0.22	0.31	0.29	0.03	0.02	0.00	0.00	0.47	0.45	0.05	0.02	-	-	0.49	49.5	1.5	2432	94		
Grt100_Grt26	22.84	0.05	0.00	0.00	0.00	39.82	22.18	1.84	11.92	1.10	99.84	0.22	0.31	0.29	0.03	0.02	0.00	0.00	0.47	0.45	0.05	0.02	-	-	0.49	-	-	-	-		
Grt100_Grt27	22.91	0.05	0.00	0.02	0.03	39.82	22.13	1.93	11.88	1.12	99.90	0.22	0.31	0.29	0.03	0.02	0.00	0.00	0.46	0.47	0.05	0.02	-	-	0.49	48.2	2.1	2960	110		
Grt100_Grt28	22.95	0.04	0.00	0.00	0.00	39.91	21.75	2.01	12.10	1.10	99.86	0.23	0.30	0.30	0.04	0.02	0.00	0.00	0.46	0.46	0.05	0.02	-	-	0.50	46.6	1.8	3230	130		
Grt100_Grt29	23.02	0.07	0.00	0.02	0.00	39.90	21.85	1.94	11.82	1.07	99.69	0.23	0.30	0.29	0.03	0.02	0.00	0.00	0.47	0.45	0.05	0.02	-	-	0.49	62.2	2.6	2861	85		
Grt100_Grt30	22.85	0.03	0.00	0.00	0.00	40.04	21.78	1.78	12.37	1.07	99.92	0.22	0.30	0.31	0.03	0.02	0.00	0.00	0.46	0.47	0.05	0.02	-	-	0.50	-	-	-	-		
Grt100_Grt31	23.06	0.03	0.01	0.00	0.00	40.07	21.56	1.80	12.46	1.05	100.04	0.23	0.30	0.31	0.03	0.01	0.00	0.00	0.46	0.47	0.05	0.02	-	-	0.51	47.9	2.0	2820	100		
Grt100_Grt32	22.60	0.05	0.00	0.02	0.02	39.97	21.55	2.08	12.19	1.00	99.47	0.22	0.30	0.30	0.04	0.01	0.00	0.00	0.46	0.46	0.05	0.02	-	-	0.50	46.3	1.2	3450	110		
Grt100_Grt33	22.92	0.08	0.00	0.03	0.00	39.83	21.57	2.05	12.55	1.02	100.03	0.22	0.30	0.30	0.04	0.01	0.00	0.00	0.45	0.47	0.06	0.02	-	-	0.51	-	-	-	-		
Grt100_Grt34	23.14	0.06	0.01	0.02	0.00	40.01	21.98	1.74	12.71	0.97	100.11	0.23	0.30	0.32	0.03	0.01	0.00	0.00	0.45	0.48	0.05	0.02	-	-	0.52	-	-	-	-		
Grt100_Grt35	22.82	0.05	0.01	0.02	0.00	39.98	21.20	1.80	12.34	0.96	99.20	0.22	0.30	0.31	0.03	0.01	0.00	0.00	0.46	0.47	0.05	0.02	-	-	0.51	-	-	-	-		
Grt100_Grt36	22.97	0.05	0.00	0.02	0.01	40.15	21.08	1.76	12.70	0.95	99.68	0.23	0.29	0.32	0.03	0.01	0.00	0.00	0.45	0.48	0.05	0.02	-	-	0.52	62.7	2.2	3790	150		
Grt100_Grt37	22.87	0.04	0.00	0.04	0.00	40.18	21.02	2.03	12.60	0.97	99.73	0.22	0.29	0.31	0.04	0.01	0.00	0.00	0.45	0.48	0.05	0.02	-	-	0.52	61.8	3.1	3520	150		
Grt100_Grt38	23.22	0.00	0.01	0.03	0.02	39.98	21.22	1.90	12.70	0.96	100.03	0.23	0.30	0.32	0.03	0.01	0.00	0.00	0.45	0.48	0.05	0.02	-	-	0.52	-	-	-	-		
Grt100_Grt39	22.82	0.01	0.02	0.06	0.00	39.56																									

Table 5.2. Bulk compositions used for input to Theriak-Domino.

Rock Type Analysis Type Sample	Residual Granulites			Leucogranulite		Paragneisses	
	XRF Gruf100	XRF Br03-56-2*	EPMA GRM-37	EPMA GR11-39	XRF GR11-23	XRF GR11-25	
SiO ₂	44.85	n.d.	33.17	51.78	62.86	66.88	
TiO ₂	0.63	n.d.	0.05	0.11	0.81	1.18	
Al ₂ O ₃	25.85	n.d.	43.41	24.80	19.35	16.74	
FeO _T	8.65	n.d.	7.20	7.76	6.44	7.12	
MnO	0.11	n.d.	0.05	0.09	0.10	0.10	
MgO	15.74	n.d.	15.39	6.37	1.86	2.33	
CaO	0.356	n.d.	0.03	2.694	0.829	0.501	
Na ₂ O	0.28	n.d.	0.05	2.28	1.72	0.54	
K ₂ O	2.33	n.d.	0.00	4.09	3.73	2.21	
P ₂ O ₅	0.07	n.d.	n.d.	n.d.	0.11	0.05	
LOI (%)	0.82	n.d.	0.61	0.16	1.68	0.85	
Sum (w. LOI)	99.69	n.d.	99.96	100.15	99.50	98.49	
Mole %:							
Si	40.57	37.63	29.22	47.21	60.68	65.93	
Ti	0.43	-	-	0.08	0.59	0.87	
Al	27.56	25.81	45.07	26.65	22.02	19.45	
Fe	6.54	7.53	5.30	5.92	5.20	5.87	
Mn	0.08	0.08	0.08	0.07	0.08	0.08	
Mg	21.23	24.73	20.21	8.66	2.67	3.42	
Ca	0.34	0.54	-	2.63	0.86	0.53	
Na	0.49	0.18	-	4.03	3.22	1.03	
K	2.69	3.23	-	4.76	4.60	2.78	
H	4.00	10.00	5.00	8.00	15.00	15.00	
Excess Si	-	-	-	-	20.00	20.00	

Notes:

Analysis type indicates whether the composition was obtained by XRF of whole-rock powders or by EPMA of microdomains.

*Whole-rock oxide data were not reported in the thesis of Schefer (2005).

LOI = loss on ignition

FeO_T = total iron reported as FeO

n.d. = not determined

Excess Si was added to the quartz-rich paragneiss samples to stabilize quartz in the phase diagrams

Table 5.2

Table 5.3. LA-ICP-MS parameters for garnet trace element analyses.

Laboratory & Sample Preparation	
Laboratory name	The University of Kansas, Dept. of Geology, Isotope Geochemistry Lab
Sample type/mineral	Garnet
Sample preparation	40 µm-thick, polished sections
Imaging	BSE images and X-ray compositional maps
Laser ablation system	
Make, model & type	ATL ArF excimer laser (193 nm), Photon Machines Analyte G1
Ablation cell & volume	Frames Cell with teardrop-shaped insert
Laser wavelength	193 nm
Pulse width (ns)	5 ns
Fluence	3.2 J/cm ²
Repetition rate	10 Hz
Spot size (µm)	35 µm circle
Sampling mode / pattern	Single spots
Carrier gas	He, 0.425 l/min, Ar, 1.24 l/min
Ablation duration	25 s
Cell carrier gas flow	He
ICP-MS Instrument	
Make, Model & type	Thermo Element2 magnetic sector field ICP-MS (single collector)
Sample introduction	Aerosol with sample + He was mixed with Ar using a Y-connector 15 cm upstream from torch.
RF power	1100 W
Make-up gas flow	Ar, 16 l/min
Sampling depth	20 µm
Detection system	Single detector (SEM), counting & analog modes
Elements/ isotopes analyzed	29Si, 43Ca, 44Ca, 45Sc, 49Ti, 89Y, 94Zr, 139La, 140Ce, 141Pr, 146Nd, 147Sm, 153Eu, 157Gd, 159Tb, 163Dy, 165Ho, 166Er, 169Tm, 172Yb, 175Lu
Integration time per peak (Sample Time in milliseconds)	29Si=10, 43Ca=10, 44Ca=10, 45Sc=10, 49Ti=10, 89Y=10, 94Zr=10, 139La=30, 140Ce=30, 141Pr=30, 146Nd=20, 147Sm=20, 153Eu=20, 157Gd=20, 159Tb=20, 163Dy=20, 165Ho=20, 166Er=10, 169Tm=10, 172Yb=10, 175Lu=10
Total integration time (Segment Duration in milliseconds)	29Si=10, 43Ca=10, 44Ca=10, 45Sc=10, 49Ti=10, 89Y=10, 94Zr=10, 139La=30, 140Ce=30, 141Pr=30, 146Nd=20, 147Sm=20, 153Eu=20, 157Gd=20, 159Tb=20, 163Dy=20, 165Ho=20, 166Er=10, 169Tm=10, 172Yb=10, 175Lu=10
Total method time	1 min 47 s (100 runs, 2 passes)

Table 5.3

Sensitivity (cps/ppm)	29Si=122, 43Ca=12, 44Ca=199, 45Sc=5868, 49Ti=447, 89Y=7822, 94Zr=3915, 139La=13000, 140Ce=16723, 141Pr=16659, 146Nd=2615, 147Sm=2153, 153Eu=8914, 157Gd=1791, 159Tb=11996, 163Dy=2985, 165Ho=11585, 166Er=3939, 169Tm=12062, 172Yb=2655, 175Lu=10073
ICP Dead time	22 ns
UO+/U+	0.02%
238U+/232Th+	~1.5
Data Processing	
Gas blank	18 s
Calibration strategy	Standard-sampling bracketing + internal standardization using Ca concentration measured by electron microprobe for each spot
Reference material info	GSD-1G (Guillong et al., 2005; Jochum et al., 2005)
Internal std for trace elements	44Ca
Data processing package used	IGOR PRO, lolite 2.5 (Patton et al. 2011); Trace elements: TraceElementsIS data reduction scheme
Common-Pb correction, composition and uncertainty	n/a
Uncertainty level & propagation	2SE internal uncertainty
Reproducibility	3-8%
Quality control / Validation	GSE-1G (Guillong et al., 2005; Jochum et al., 2005)